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Assessing Uncertainties and Approximations in Solar Heating of the Climate System

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Key Points:

- Known errors and uncertainties in solar heating codes are evaluated within a single framework providing climate-relevant diagnostics
- Spectral resolution of cloud absorption, ice cloud optics, and cloud overlap schemes have large potential errors
- Other significant error sources from numerical or physical approximations, such as 2-stream scattering, can be readily minimized

Supporting Information:

- Supporting Information S1

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Abstract In calculating solar radiation, climate models make many simplifications, in part to reduce computational cost and enable climate modeling, and in part from lack of understanding of critical atmospheric information. Whether known errors or unknown errors, the community's concern is how these could impact the modeled climate. The simplifications are well known and most have published studies evaluating them, but with individual studies it is difficult to compare. Here, we collect a wide range of such simplifications in either radiative transfer modeling or atmospheric conditions and assess potential errors within a consistent framework on climate-relevant scales. We build benchmarking capability around a solar heating code (Solar-J) that doubles as a photolysis code for chemistry and can be readily adapted to consider other errors and uncertainties. The broad classes here include: use of broad wavelength bands to integrate over spectral features; scattering approximations that alter phase function and optical depths for clouds and gases; uncertainty in ice-cloud optics; treatment of fractional cloud cover including overlap; and variability of ocean surface albedo. We geographically map the errors in W m^{-2} using a full climate re-creation for January 2015 from a weather forecasting model. For many approximations assessed here, mean errors are $\sim 2 \text{ W m}^{-2}$ with greater latitudinal biases and are likely to affect a model's ability to match the current climate state. Combining this work with previous studies, we make priority recommendations for fixing these simplifications based on both the magnitude of error and the ease or computational cost of the fix.

Plain Language Summary Solar heating of the climate system—the atmosphere, land surface, and ocean—drives the climate. Accurate numerical calculation of solar heating is a core component of the models we use to project and prepare for climate change. The community has identified many potential sources of error and published studies showing how to improve the solar heating codes used in climate models. Here, we assemble a wide range of these errors, either numerical approximations or uncertainties in defining atmospheric conditions, and put them through the same test: calculating the atmospheric and surface heating over a month of simulated climate conditions. Combining the new calculations here with previous work, we discuss more than a dozen specific areas where improvements could be made and identify high-priority actions.

1. Introduction

The heat from sunlight drives the weather and climate system, the energy in solar photons drives atmospheric chemistry, and the photosynthetically active radiation drives life. For Earth system models (ESMs), one needs to calculate the scattering, absorption, and reflection of solar radiation throughout the atmosphere, ocean, cryosphere, and land surface. The idealized radiative transfer (RT) problem is well known and in many cases has near-exact, but costly solutions; whereas the atmospheric physics of the problem involving gases, clouds, aerosols and surface properties includes unknowns that cause a fundamental uncertainty in the solution. One motivation for this study is to evaluate the potential improvements in solar heating rates if more accurate physics or RT codes are used, but the overriding motivation is to assess a wide range of approximations and uncertainties within a single climate-relevant framework. Where more accurate physics or numerical solutions are known, we can estimate the error in current RT codes, and where there is fundamental uncertainty, we can estimate the potential for error using different methods for framing the problem.

The solar heating module Solar-J (Hsu et al., 2017; hence H2017) is embedded in our chemistry-transport model (CTM) and used to integrate climate-relevant heating rates using the European Center's Integrated

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Forecast System meteorological data that drives the CTM simulation of atmospheric chemistry (M. J. Prather et al., 2017; Szépszó et al., 2019). This CTM + Solar-J model quantified the errors related to spherical geometry of the atmosphere (M. J. Prather & Hsu, 2019; hence P2019). Here, we use that code and diagnostic framework (i.e., January average of hourly global calculations at horizontal resolution of 1.1° including water vapor and ice-/liquid-water clouds) for an extensive set of case studies. In addition, we embed the AER RRTMG-SW version 4.0 (Clough et al., 2005; Mlawer et al., 1997) directly in the CTM, providing a platform for parallel comparison of the two codes in a realistic climate. Combining solar heating and photolysis in one code is obvious as both solutions use the same atmospheric data and we move to ESMs that require both. Nevertheless, compromises between the two requirements, Watts versus photons s^{-1} , are identified here.

In this study, we examine additional classes of approximations or uncertainties beyond P2019: examples of the historical improvement in infrared heating codes and the ongoing work at major climate centers (Section 2); the use of bands to integrate over spectral features (Section 3); multiple-scattering approximations that alter the scattering phase function for clouds, aerosols, and gases (Section 4); uncertainty in ice-cloud optics (Section 5); treatment of fractional cloud cover including cloud overlap (Section 6); and approximation of ocean surface albedo (Section 7). Each one of these sections has been the focus of major research studies that we briefly review. Bringing these together with common climate metric allows for the broad comparison here. Section 8 reviews our findings from Sections 3 through 7 and makes priority recommendations for improving climate models' solar heating codes.

For most cases here, we find global mean absolute errors, or uncertainties, ranging from 0.5 to 5 $W m^{-2}$ with larger systematic latitudinal or root-mean-square errors. Such error levels are likely to shift ESMs into different climate regimes for the current reference period as they are comparable to the changes in climate forcing by greenhouse gases from pre-industrial to present (Myhre et al., 2013). Such shifts are likely to force parametric retuning of other ESM processes to match observed surface temperatures. Either way, errors on this scale are important and must eventually be addressed.

2. Evolving Solar RT Codes

Many of the case studies here follow the ongoing efforts of the scientific community to improve the solar heating codes in current climate models. Most of this work is occurring within the major climate modeling centers. For example, the U.S. Department of Energy's Energy Exascale Earth System Model (E3SM) is implementing a new toolbox that seeks to balance accuracy, efficiency, and flexibility in the solar RT code (Pincus et al., 2019). The Canadian Centre for Climate Modeling and Analysis (CCCma) examined the importance of expanded scattering phase functions for clouds and aerosols possible with a 4-stream RT code (Li et al., 2015, hence LBYY) as well as the use of satellite-derived cloudy atmospheres to compare 3-D RT with the 1-D climate model codes (H. W. Barker et al., 2012; 2015, hence B2015). The Korean Integrated Model has updated their ice cloud treatment (Baek & Bae, 2018), as have also Zhao et al. (2018) for the Community Atmosphere Model Version 5. The European Centre for Medium-Range Weather Forecasts (ECMWF) now includes innovations such as horizontal and vertical cloud structures (R. J. Hogan & Bozzo, 2018; J. K. P. Shonk et al., 2010) based on Hogan et al.'s (2016, hence H2016) solar RT code that combines Tripleclouds and 3D effects in a 1D solution. The Australian Community Climate and Earth System Simulator (ACCESS) has also adopted Tripleclouds (Franklin et al., 2013). The French ARPEGE-Climat model (Séférián et al., 2018) has adopted the Z. H. Jin et al. (2011) ocean surface albedo model. With variants of the Solar-J code we can compare many of these improvements including vertical cloud overlap, but the 3D effects of H2016 are outside of Solar-J's current capability.

To calibrate the climate metric used here to assess errors and uncertainties, we test a sequence of solar RT codes that represents successively improved spectroscopy and modeling of water vapor lines in the infrared. Our reference code, the standard Solar-J version (H2017, denoted *SJ*), uses Cloud-J (M. J. Prather, 2015) spectroscopic data for the ultraviolet and visible wavelengths (0.18–0.78 μm , bins #1–18) with a single value of solar flux (in both photons $cm^{-2} s^{-1}$ and Watts m^{-2}) and cross sections (cm^2). *SJ* adds 9 broad infrared (IR) bands (0.78–12.2 μm , bins #19–27) taken directly from the RRTMG-SW version 4.0 code (Clough et al., 2005; Mlawer et al., 1997). These 9 IR bands plus the overlap bin #18 (0.485–0.778 μm) contain a total

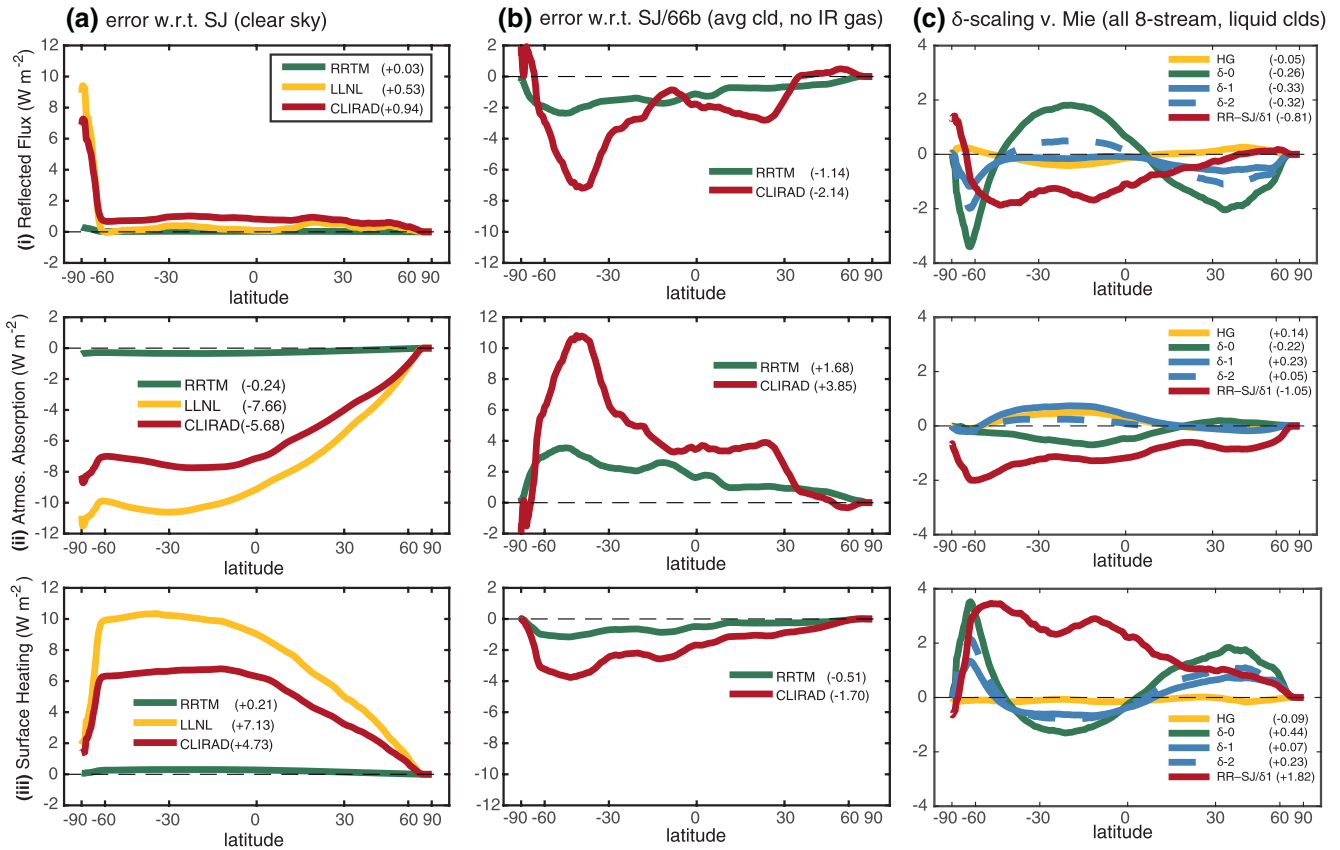


Figure 1. Monthly zonal mean flux differences (W m^{-2}) as a function of latitude for January 2015 with three vertical panels showing reflected, atmospheric absorption, and net surface heating (top down). (a) Case study for H_2O -gas absorption and clear sky, comparing models with different numbers of infrared sub-bins. Differences are relative to standard Solar-J (*SJ*). RRTM refers to the very high-resolution (*SJ/RRX* in Table S1); CLIRAD and LLNL, to the coarser resolutions (*SJ/CLIRAD* and *SJ/LLNL*). (b) All sky with averaged clouds and no infrared (IR) gas absorption, emphasizing the resolution of cloud absorption. Differences are relative to *SJ-66b* (high-resolution infrared bins for clouds). *SJ/noIR* has the standard 9 IR RRTMG bands, and *SJ/CLIRAD/noIR* has 3. (c) Averaged liquid-only clouds shown for a range of re-scalings of the Mie scattering phase function (HG, $\delta-0$, $\delta-1$, $\delta-2$). These are all evaluated within the 8-stream *SJ* code. Also shown is the difference RRTMG minus *SJ/δ1*, where much of the difference, especially in atmospheric heating, is due to the 2-stream minus 8-stream difference. See Table S1 for a complete description of code versions.

of $78 + 5$ g-points (sub-bins) for gas absorption. A 'high-accuracy' *SJ* version in terms of infrared gas absorption is derived from the benchmark code RRTM-SW with the same 9 IR broad bands, but 144 g-points, and is denoted *SJ/RRX*. We made Solar-J variants using water vapor absorption as parameterized by older codes. The oldest and least accurate LLNL code (Grant & Grossman, 1998) has three IR bands ($0.69\text{--}3.85 \mu\text{m}$) with 21 sub-bins (denoted *SJ/LLNL*). The CLIRAD code (Chou & Suarez, 1999, revised with HITRAN 2012 data), used until recently by the Goddard Space Flight Center climate model, has three IR bands ($0.70\text{--}10.0 \mu\text{m}$) including a total of 30 sub-bins (denoted *SJ/CLIRAD*).

We compare these codes by running their *SJ* variants with only water vapor as an absorber in the IR and clear skies (no clouds or aerosols, see Figure 1a). Compared to *SJ*, *SJ/CLIRAD*, and *SJ/LLNL* have about 5 and 7 W m^{-2} less atmospheric absorption, respectively, most of which is absorbed at the surface with small excess fraction ($\sim 1 \text{ W m}^{-2}$) being reflected (see Table 1/Rows 1 & 2, hence designated T1/R1&2). The more accurate *SJ/RRX* code has only $\sim 0.25 \text{ W m}^{-2}$ less atmospheric absorption than *SJ* (T1/R3). Assuming that *SJ/RRX* is the most accurate code for water vapor absorption, we find that in terms of re-partitioning the solar heating between atmosphere and surface ($5\text{--}7 \text{ W m}^{-2}$) these historical improvements are greater than any of the next generation of errors assessed here. In terms of incident or reflected flux errors, however, they are smaller than the cloud uncertainties, spherical geometry, or even issues of wavelength resolution in the visible.

Table 1
Errors in the Three Primary Components of the Solar Radiation Budget (Reflected Sunlight, Atmospheric Absorption, Surface Absorption in $W m^{-2}$) for a Range of Approximations in the Radiative Transfer Models

row	Models	Error being estimated	Mean difference (Wm^{-2})			RMS difference (Wm^{-2})		
			refl.	Atm.	Surf.	refl.	Atm.	Surf.
Spectroscopy errors for infrared and visible gas absorption with clear sky								
R1	B1–B0	SJ/CLIRAD v. SJ	+0.94	−5.68	+4.73	1.4	8.4	7.0
R2	B2–B0	SJ/LLNL v. SJ	+0.53	−7.66	+7.13	1.2	12.0	11.0
R3	B3–B0	SJ/RRX v. SJ	+0.03	−0.24	+0.21	0.0	0.4	0.3
R4	RR0–B0	AER4.0 v. SJ	+1.30	−0.47	−0.82	1.9	0.8	1.4
R5	B4–B0	SJ/hrv v. SJ	+0.34	−0.46	+0.12	-	-	-
R6	B5–B0	SJ/2S versus SJ	+0.31	−0.53	+0.22	0.4	0.4	0.2
Wavelength resolution of cloud absorption (<i>MAX-COR overlap and QCAs, no IR gas absorption</i>)								
R7	C1–C0	SJ/CLIRAD v. SJ/66b	−2.14	+3.85	−1.70	8.7	10.4	3.1
R8	C2–C0	SJ/RRTM v. SJ/66b	−1.14	+1.68	−0.51	2.2	3.2	1.1
Scattering phase function errors for HG and δ -scaling using grid-cell averaged only liquid water clouds. Also differences between RRTMG-SW v4.0 and <i>Solar-J</i> using δ -1 scaling. The sum of the 3 principal components may not equal zero because of round-off or small differences in incident flux (not shown) due to the vertical sub-layers for clouds in <i>SJ</i> . The Rayleigh case below R17 was done with clear sky.								
R9	Mh–M0	SJ/HG v. SJ (Mie)	−0.05	+0.14	−0.09	0.8	0.6	0.6
R10	M1–M0	SJ/ δ_0 v. SJ (Mie)	−0.26	−0.22	+0.44	5.6	1.2	5.1
R11	M2–M0	SJ/ δ_1 v. SJ (Mie)	−0.33	+0.23	+0.07	1.4	1.0	2.2
R12	M3–M0	SJ/ δ_2 v. SJ (Mie)	−0.32	+0.05	+0.23	2.8	0.5	3.1
R13	M4–M0	SJ/ $\delta_1/2S$ v. SJ/ δ_1	−0.68	−1.74	+2.42	3.5	2.7	5.0
R14	RR2-M0	AER4.0 v. SJ	−1.14	−0.82	+1.88	3.0	1.7	4.1
R15	RR2-M2	AER4.0 v. SJ/ δ_1	−0.81	−1.05	+1.82	2.4	1.7	3.6
R16	RR2-M4	AER4.0 v. SJ/ $\delta_1/2S$	−0.13	+0.69	−0.61	3.3	1.4	3.2
R17	MR –B0	Isotropic Rayl. v. SJ	−0.01	+0.01	0.00	0.0	0.1	0.0
R18	MI– D0	SJ/ δ_1 ice v. SJ, QCA, MAX-COR overlap	−0.02	+0.01	+0.01	0.8	0.5	0.5
Cloud Overlap Algorithm. <i>Solar-J</i> with QCAs for both MAX-RAN and MAX-COR overlap. RRTMG-SW v4.0 run with McICA for MAX-RAN cloud overlap. See note above small differences in incident flux.								
R19	D1 – D0	SJ: MAX-RAN v. MAX-COR, both QCA	−1.35	−0.06	+1.41	7.9	1.0	8.6
R20	RR1-M5	AER4.0 v. SJ (avg cld w/ice)	−0.38	−2.67	+2.97	4.4	5.9	6.1
R21	RR3–D1	AER4.0 (McICA) v. SJ (QCA) (both MAX-RAN)	−2.10	−2.36	+4.41	18.5	5.9	20.5
Ocean Surface Albedo (<i>Solar-J</i> with MAX-COR overlap and QCAs)								
R22	O1–O0	SJ:OSA constant v. angle dependent	+0.68	+0.16	−0.84	4.7	0.4	5.1
Spherical errors from M. J. Prather and Hsu (2019) include changes in incident flux, which was not included in the above comparisons that used the same model (spherical or flat) in each comparison. Comparisons R23–R25 use <i>Solar-J</i> with all clouds, MAX-COR overlap and QCAs.								
R23	-	SJ: spherical v. flat	incid. +1.55	refl. +0.68	atm. +0.46	surf. +0.41	-	-
R24	-	SJ: refraction v. flat	+1.94	+0.83	+0.53	+0.58	-	-
R25	-	SJ: geometric v. flat	+2.50	+1.02	+0.86	+0.62	-	-

Note. See methods section Table S2 for description of each of the pairs of simulations noted in column two. Global-mean area-weighted differences are averaged over January 2015 (744 hourly data) with root mean square differences accumulated hourly. All results use the UCI CTM and the T159L60 ($\sim 1.1^\circ \times 1.1^\circ$) ECMWF forecast fields developed by U. Oslo from the Open-IFS system (M. J. Prather et al., 2017; Søvde et al., 2012).

3. Solar-J Spectral Model and Band Resolution

Practical RT solutions for the complex absorption features of atmospheric gases require selecting a limited number of representative wavelengths or wavelength bands for the calculation. Ideally, one picks as broad a band as possible that has nearly the same attenuation of sunlight across all wavelengths and can be accurately represented with a single RT solution. In wavelength regions with many sharp absorption lines, such as the Schumann-Runge bands of O₂ (177–202 nm), the opacity distribution method (Fang et al., 1974) sorts the lines into bins of non-contiguous wavelength micro-bands with similar opacity. In parallel, the calculation of heating rates uses the similar correlated-k distribution (Goody et al., 1989; Lacis & Oinas, 1991) to sort the near infrared and infrared absorption lines of water vapor and other greenhouse gases into bins of similar opacity, designated g-points in RRTMG. Both methods select broad wavelength bands where the opacities of more continuum absorbers like clouds and aerosols can be treated as a constant. The art here lies in selecting the minimum number of total bins, each requiring a separate RT calculation, which can still reproduce, within a specified error, results from spectrally resolved models using thousands of wavelength intervals and calculations.

Solar-J's underlying spectral model and scattering code is Fast-J, which is used for chemical photolysis rates and thus requires a reference solar spectrum in photons cm⁻² s⁻¹ and photon-weighted average cross sections. With the added capabilities for cloud overlap and solar heating, Fast-J has become Cloud-J (M. J. Prather, 2015), which has become the core of Solar-J. Fast-J developed optimized wavelength binning for the stratosphere (175–298 nm, bins #1–12, Bian & Prather, 2002) and the troposphere (298–800 nm, bins #13–18, Wild et al., 2000), and builds upon 4 decades of RT development for atmospheric chemistry (Logan et al., 1978; Olson et al., 1997; PhotoComp, 2010). The 8-stream Feautrier based scattering code (Prather, 1974; Wild et al., 2000) calculates mean specific intensities for the photolysis rates as well as flux divergence across each model layer. The cross sections used for atmospheric attenuation, photolysis rates, and absorption in bins #1–18 are photon-weighted. Heating rates use a reference solar table (Watts m⁻²) for each large bin. Atmospheric attenuation and absorption in IR bands #19–27 is calculated with the RRTMG spectral model; Solar-J extends its reference solar table (Watts m⁻²) to match RRTMG; and photon fluxes are not used since photolysis only occurs in bins #1–18. For the tropospheric wavelengths (#13–18), the bins are contiguous wavelength bands, and only ozone absorption and Rayleigh scattering affect the fluxes, see Figure 2.

3.1. The Wavelength Region 290–778 nm

We expected a simple, or at least explicable comparison of *SJ* with the RRTMG-SW v4.0 standalone code (*RRTMG*) under clear skies because the codes are essentially identical for IR wavelengths >778 nm, and the UV and VIS region <778 nm consists of well quantified, continuous cross sections for ozone absorption and Rayleigh scattering. The primary differences in the codes are the wavelength bands and scattering: *SJ* uses Rayleigh-phase scattering with an 8-stream code while *RRTMG* uses isotropic scattering with a 2-stream code. The difference *RRTMG* minus *SJ* under clear sky gave a surprisingly large difference: (+1.3, -0.5, -0.8 W m⁻²) for the three primary globally averaged components (reflection, atmospheric absorption, and surface absorption), respectively (*T1/R4*). For the UV region <298 nm, the solar flux (13.6 out of 1360.8 W m⁻²) is absorbed in the stratosphere, and both *SJ* and *RRTMG* agree on this. So focusing on the 298–778 nm region that covers the visible, UVA, and part of UVB (for convenience designated as UV-VIS here) we try to understand these differences.

This UV-VIS region has the most solar flux (721 W m⁻²) and for the most part is optically thin. The primary attenuators (molecular Rayleigh, O₃), solar flux, and wavelength bands used in *RRTMG* and *SJ* are shown in Figure 2. For 298–320 nm (13 W m⁻²) the O₃ total optical depth is greater than about 0.5 (0.05 in the troposphere) and much is absorbed in the stratosphere. *SJ* uses 3 narrow bands to resolve the rapid changes in O₃ optical depth and attenuation of sunlight as well as a fourth narrow band to 345 nm where transmission is controlled by Rayleigh scattering. This level of resolution is necessary to calculate photolysis rates. The Rayleigh optical depth (mostly tropospheric) drops rapidly from 1.0 at 300 nm to 0.04 at 700 nm. For wavelengths >345 nm, the optical depths of Rayleigh and ozone are relatively small, and accurate photolysis rates need at most two bands if individual species' absorption cross sections are photon-weighted. O₃

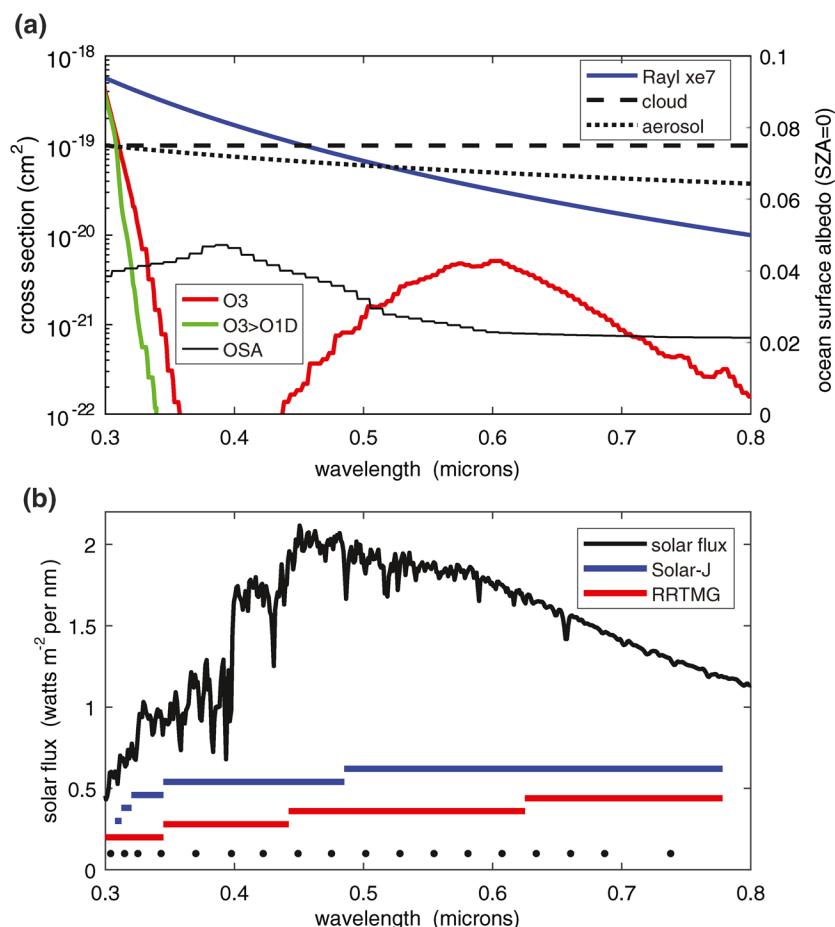


Figure 2. (a) UV-Visible region for Solar-J (298–778 nm, 721 W m⁻²) showing (top) continuum spectra for Rayleigh scattering, O₃ absorption, O¹D production, and ocean surface albedo (OSA, high sun, low wind) plus schematic wavelength dependence (arbitrary cross sections) for clouds (dashed line) and typical aerosol (dotted). Rayleigh cross sections have scaled by 10⁷ to be compared with those of O₃, assuming tropospheric abundance of 100 ppb. Clouds have no wavelength dependence across this region and aerosols are shown for an Angstrom coefficient of 1. (b) The solar spectrum and the wavelength bands used in Solar-J (blue), RRTMG (red) plus the high-resolution-visible version of Solar-J/hrv (black dots).

absorption in the Chappuis bands (450–800 nm) is not important for the photochemistry, but it is an important source of atmospheric heating. Likewise, Rayleigh scattering in the visible plays a minor role in photolysis rates but is an important source of reflected sunlight. A first-order estimate based simply on weighting tropospheric extinction with solar flux has Rayleigh scatter (forward and back) of 123 W m⁻², O₃ absorption of 2.7 W m⁻², and surface reflection (ocean albedo = 0.05) of 36 W m⁻². This should be a relatively easy wavelength region to model accurately.

To cover the 345–778 nm region, RRTMG has three broad bands and 20 total g-points (sub-bins) within them. Each g-point has its own Rayleigh cross section derived from the correlated-k distribution and its own O₃ absorption cross section. Solar-J has only two broad bands with no sub-bins and clearly does not resolve the different patterns of Rayleigh and O₃ extinction, see Figure 2. Thus, we built a high-resolution-visible Solar-J version (*SJ/hrv*), putting 18 bands in the UV-VIS region (black dots in Figure 2) to achieve resolution comparable to RRTMG. The *SJ/hrv* minus *SJ* differences, (+0.3, -0.5, +0.1 W m⁻²) (**T1/R5**), can explain the RRTMG v. *SJ* differences in atmospheric heating, but still leaves a large discrepancy in the reflected flux.

Is this difference caused by Solar-J's use of photon-, not Watt-weighted averages for these UV-VIS bands? In Table S3, we look at the average cross sections (Rayleigh scattering and O₃ absorption) for the three largest

VIS bands in Solar-J using both photon and Watt weighting derived from the same high-resolution solar spectrum. When the bands are narrow (e.g., 320–345 nm) the difference between these two averaging methods is negligible, but for the two broad VIS bands, the Watt-averaged Rayleigh cross sections are 4%–7% larger, and would result in greater reflected flux as seen in *RRTMG* and *SJ/hrv*. Thus Solar-J's requirement for calculating both photolysis and heating rates from the same calculation points out a fundamental conflict in the case of broad bands. For the narrow tropospheric UV (298–345) the current *SJ* bands are fine, but for the VIS we must increase the number of bands beyond 2, and look at Watt-weighting of Rayleigh and O₃ cross sections but not the photolyzed species. We expect that accurate clear-sky solar heating can be computed with 5–10 broad bands and probably does not need the 20 g-points of *RRTMG*. One must remember that clouds and aerosols are also important for the 721 W m⁻² in the UV-VIS. Clouds have little spectral dependence in the UV-VIS and can thus be calculated accurately with very broad bands. Aerosols extinction often scales inversely with wavelength (see Figure 2) and thus aerosol radiative forcing of climate needs to be resolved by the broad band structure and not with g-points.

3.2. Cloud Absorption in the IR

The accuracy of the calculated cloud absorption is determined by the number and spacing of the IR broad bands, not the number of sub-bins or g-points in each, because models assume constant cloud optics across each band. Edwards and Slingo (1996) recognized this problem and devised an ingenious spectral averaging of the single-scattering albedo (SSA) of ice- and liquid-water clouds over the broad infrared bands. Their thick-cloud averaging enabled the UK climate model to run successfully with four broad bands for solar RT, but that model and a 24-band model still had large errors compared to their 220-band reference model. How does this problem look with the *RRTMG* bands? Figures 3a and 3b shows the SSA of typical ice- and liquid-water clouds over the infrared spectrum. The refractive indices for ice and liquid water are similar but with distinct wavelength shifts: the second deep SSA minimum for liquid occurs about 1.9 μm, while that for ice occurs about 2.0 μm. The largest differences in the liquid versus ice curves here are caused by particle size, with smaller particles having larger SSA. The 9 infrared bands (#19–#27 in Figure 3) are denoted with the vertical dashed lines. The average cloud SSA in each bin is shown by the horizontal bar with a circle.

To resolve the variation in cloud absorption within the broad bands, we recalculated a set of narrow bands (0.05–0.10 μm wide), yielding 66 IR bands shown as black squares. The Solar-J cloud optics are calculated using a flux-weighted refractive index for liquid or ice water and Mie theory for a wide range of effective radii (R_{eff}) and a dispersion of spherical particles similar to stratus clouds. The Mie calculation returns SSA, extinction coefficient (Q), and seven Legendre coefficients of the scattering phase function ($g_{1:7}$, see Table 2). For ice clouds, a non-Mie scattering phase function is used and examined in Sections 4 and 5. The Mie calculation of SSA and Q is surprisingly accurate. In the top panel of Figure 3 we compare our narrow-band Mie-ice SSA results with values taken from Ping Yang's library (Bi & Yang, 2017) of ice crystal optics: the four thin colored lines represent four crystal habits (droxtals, 8-columns, 10-plates, small bullet rosettes, all with roughness value of 003) that have a wide range of geometries (maximum to effective diameter ratios of 1.2, 3.3, 7.6 and 2.5 respectively). The SSA is primarily a function of R_{eff} and refractive index, and the Mie approach fits within the range of crystal habits.

The standard Solar-J cloud optics, based on average refractive indices (H2017), should probably be recalculated with the Mie calculation at fine wavelength intervals and then with flux-averaging of SSA, Q , and $g_{1:7}$, as in Edwards and Slingo (1996). Looking forward, however, we should not have bands or sub-bins with such large variations in refractive index. The narrow bands have smaller changes in refractive index across each band and will be minorly affected by the method of averaging. *SJ/66b* is the Solar-J variant using these 66 narrow bands.

Lu et al. (2011: LZL) noted that *RRTM*'s wide infrared bands average over both reflecting and absorbing wavelengths in clouds. LZL use a correlated k-distribution that combines the line-by-line water vapor with the continuum liquid-water cloud optical properties to produce a set of sorted cloud + vapor bins within each wide band. LZL apply this new hybrid model to a sample column of liquid cloud layers and show that including variable cloud absorption within the broad *RRTM* bands results in more scattering, less atmospheric heating with both more reflected flux and more surface heating. They attribute these errors to the

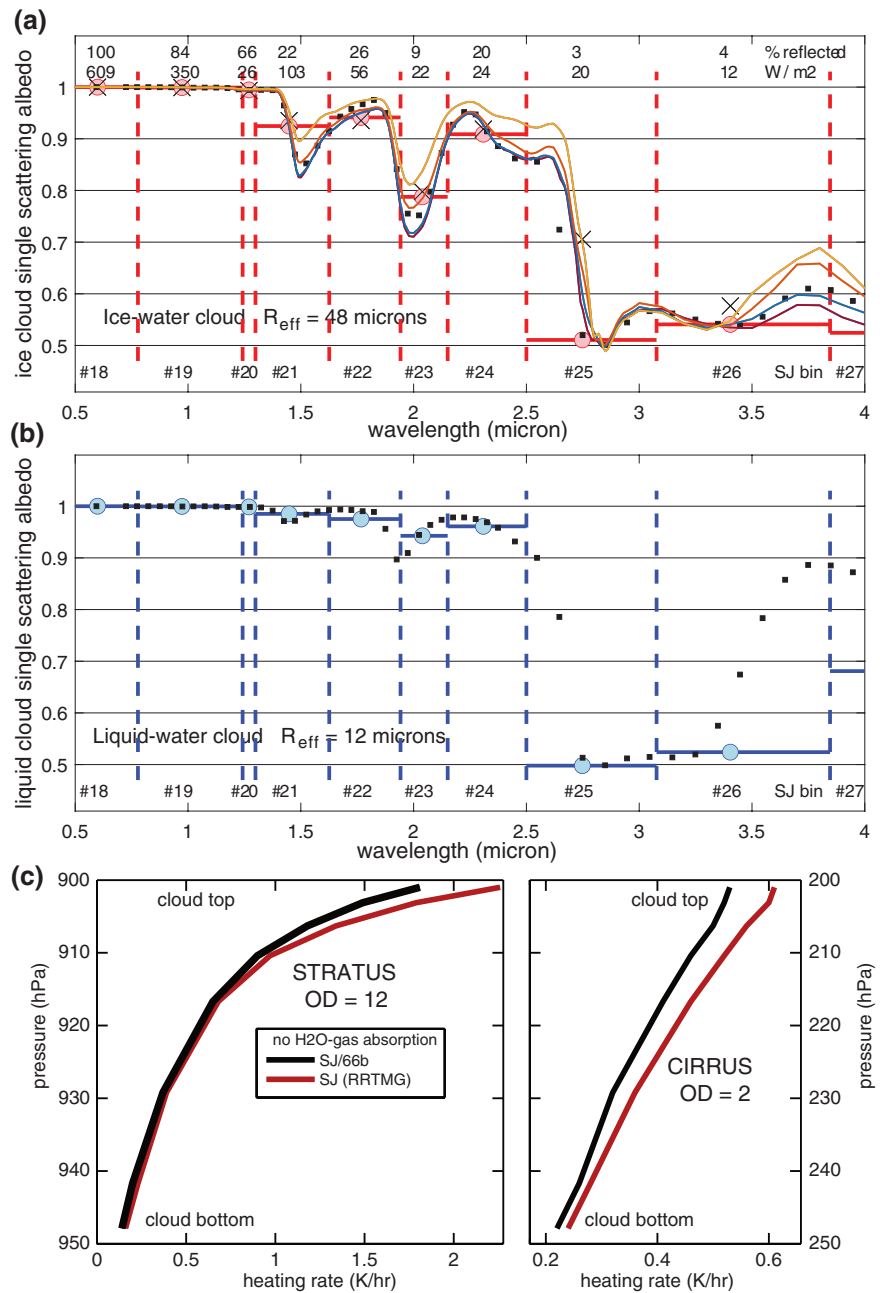


Figure 3. (a and b) Single scattering albedo (SSA, dimensionless) for water clouds versus wavelength (μm), including both (a) ice and (b) liquid water (bottom). The RRTMG infrared (IR) bins, designated #18 through #27 are demarcated by dashed vertical lines with average SSA denoted (large circle with cross bar and light-colored fill). The solar energy (W m^{-2}) in each bin is denoted, and the reflected flux (% , red, above flux) for overhead sun, zero surface albedo, and an optically thick ice-water cloud is also given. Average SSA is calculated from a flux-weighted refractive index for each bin assuming spherical (Mie) particles, including for ice-water. The liquid particles have an effective radius of $12 \mu\text{m}$; and the ice particles, $48 \mu\text{m}$. The black squares show values used in *SJ/66b*, where the resolution ranges from $0.05 \mu\text{m}$ ($0.7\text{--}2.4 \mu\text{m}$) to $0.10 \mu\text{m}$ ($2.5 \mu\text{m}\text{--}4.0 \mu\text{m}$). Four thin colored lines plotted on top of the ice cloud data are taken from Ping Yang's library (Bi & Yang, 2017) of ice crystal optics for droxtals, 8-columns, 10-plates, and small bullet rosettes, all with roughness value of 003). The R_{eff} values of $48 \mu\text{m}$ correspond to maximum diameters of 112, 312, 738, and $238 \mu\text{m}$, respectively. The Zhao et al. (2018) new derivation of ice-cloud single scattering albedo for the RRTMG bands is shown as a light cross (X) to compare with the pink/red circles. (c) Profile of in-cloud rates for stratus (optical depth $\text{OD} = 12$, liquid water) and cirrus ($\text{OD} = 2$, ice water) from *SJ/noIR* and *SJ/66b* (also no IR gas absorption by design) for overhead sun and surface albedo of 0.05.

Table 2
Parameters (Scaling Factor f , Asymmetry Factor g^* , Scattering Optical Depth τ_{sca}^*) for Different Delta-M Scaling Methods

$$P(\Theta) = 1 + g_1 \times 3 P_1(\cos[\Theta]) + g_2 \times 5 P_2(\cos[\Theta]) + g_3 \times 7 P_3(\cos[\Theta]) + \dots + g_7 \times 15 P_7(\cos[\Theta])$$

$$P_{\text{Mie}}(\Theta) = 1 + 0.865 \times 3 P_1(\cos[\Theta]) + 0.795 \times 5 P_2(\cos[\Theta]) + \dots + 0.507 \times 15 P_7(\cos[\Theta])$$

$$P_{\text{HG}}(\Theta) = 1 + 0.865 \times 3 P_1(\cos[\Theta]) + 0.748 \times 5 P_2(\cos[\Theta]) + \dots + 0.362 \times 15 P_7(\cos[\Theta])$$

Method	f	$g^* = (g_1 - f)/(1 - f)$	$\tau_{\text{sca}}^*/\tau_{\text{sca}} = 1 - f$	Notes
δ -0 (isotropic)	$g_1 = 0.865$	0	0.135	$g_n = 0, n \geq 1$
δ -1	$g_1^2 = 0.748$	0.464	0.253	$g_n = 0, n \geq 2$
δ -2	$g_2 = 0.795$	0.342	0.205	$g_n = 0, n \geq 2$

Note. This example assumes liquid water cloud ($R_{\text{eff}} = 12 \mu\text{m}$, wavelength = 600 nm, $\omega_0 = 0.99999$). The Mie phase function is truncated after P_7 . The Henyey-Greenstein phase function is expanded to P_7 is using only the first asymmetry term of the Mie phase function. The δ -0, δ -1 and δ -2 phase functions include at most P_0 and P_1 . All of these SJ variants use 8-stream scattering.

spectral correlation of gas and cloud absorption and make the case for including ice clouds in a three-parameter correlated k-distribution as the path forward.

We extend the LZL work with *SJ/66b* by evaluating the RRTM-band cloud errors in a climate-relevant framework and including ice clouds. We do not have the capability to generate correlated-k sub-bins for the 66 IR bands and thus cannot include infrared (IR) gas absorption. In a single column model like LZL, we calculate cloud heating rate profiles for sample stratus and cirrus clouds at high sun (Figure 3c) using *SJ/66b* and compare it with Solar-J without IR absorption (*SJ/noIR*). For stratus clouds, the excess heating using *SJ/noIR* ranges from +5% (bottom) to +25% (top), with similar values for cirrus ice clouds. At the top of a stratus cloud, *SJ/66b* heating rates are 1.8 K per day, while *SJ/noIR* calculates 2.3 K per day (both are without IR gas absorption). We can include IR gas absorption in *SJ* and find that heating rates in the upper layers of the stratus cloud are reduced by about a third because of the absorption by water vapor above the cloud deck; this difference is much less with high cirrus clouds. The relative differences between 9 bins and 66 infrared bands (i.e., 5%–25%) should remain. Cloud heating errors at this level are likely to affect the lifetime and stability of clouds (Wood, 2012). Overall these results are similar to LZL's liquid clouds and confirm that ice clouds also need to be corrected.

In terms of our climate metrics (e.g., global zonal-mean heating rates for January in Figure 1b), *SJ/noIR* has 1.7 W m^{-2} more atmospheric heating than *SJ/66b*, while *SJ/CLIRAD/noIR*, with only three IR bands, has 3.8 W m^{-2} more (T1/R7&8). Coarse wavelength resolution of cloud absorption clearly results in more in-cloud heating balanced almost equally by less reflection and surface absorption. Our global monthly values are about 20 times lower than LZL's case study of a very thick low cloud with overhead sun, and this is probably not inconsistent. LZL attribute these errors to the correlation of gas and cloud absorption, but our results show that a significant component of the error is simply the failure to separate the highly scattering wavelengths from the partly absorbing wavelengths and is not tied to water vapor absorption. This information may help in the development of a combined gas-liquid-ice model. Improving the accuracy of the gaseous absorption is an ongoing effort (Etminan et al., 2016; Mlawer et al., 2012; Paynter & Ramaswamy, 2014; Pincus et al., 2015; Radcl et al., 2015), but parallel efforts on cloud absorption are limited.

4. Scattering Phase Functions and Multiple Scattering

Ideally, the sunlight scattered by clouds, aerosols and gases is resolved semi-continuously in all directions within the atmosphere, but in practice, RT solutions for solar heating keep track of a limited number of angles (streams) in upward and downward directions and average over the azimuth angle. Solar-J uses eight streams (four up, four down) to resolve multiple scattering and this RT solution is implemented in many global chemistry models (M. J. Prather, 2015). RRTM (Mlawer et al., 1997) uses 16-stream scattering but is not implemented in global models; instead, RRTMG (Clough et al., 2005) with 2-stream scattering is used in many climate models. The Geophysical Fluid Dynamics Laboratory (AM3) and the Canadian Center for

Climate Modeling and Analysis (CanAM4) with 4-stream RT (Li & Ramaswamy, 1996; LBYY) appear to be the most accurate scattering codes currently used in climate models.

The number of scattering angles determines how well the scattering phase function, $P(\Theta)$, is resolved, where Θ is the angle between incident and scattered light. Nominally, these phase functions are calculated using Mie theory for spherical droplets or other approximations for aspherical ice or dust particles (Mishchenko et al., 2016; Yang et al., 2018) and expanded in Legendre polynomials, $P_L(\Theta)$, where the first two terms are $P_0 \equiv 1$ and $P_1 \equiv \cos(\Theta)$. See Figure 1 of LBYY for the Legendre coefficients for water clouds and dust. Delta-Eddington methods for solving the RT problem (Joseph et al., 1976) approximate this phase function as a forward delta function plus a two-term expansion of the phase function, P_0 and P_1 , see Table 2. M-stream RT methods (Lin et al., 2018; W. J. Wiscombe, 1977) extend this to include the first M terms in the Legendre expansion. All of these methods reduce the coefficient of P_1 from $3g_1$ to a lesser value $3g^*$ (g^* is often called the asymmetry factor, see Table 2), and likewise reduce the scattering optical depth. The reduction in cloud optical depth applies only to the scattering optical depth τ_{sca} , while the absorbing τ_{abs} is unchanged. These δ -scaling methods are chosen to avoid the unphysical negative phase functions that result from truncation of the Legendre series. The 8-stream in Solar-J does not use delta-M scaling because early tests showed that simply truncating the phase function for liquid water clouds at P_7 (giving a phase function that oscillates in sign) still produced accurate, non-negative fluxes and mean intensities compared to 160-stream solutions expanded to P_{159} . In these tests, the mean intensity differed by 1% throughout most of the atmosphere, with a worst case of 8% in the uppermost layers of an optically thick cloud and overhead sun (Wild et al., 2000). A major difference between these methods is that Solar-J retains the correct direct flux at the surface while δ -scaling methods can greatly exaggerate it.

Studies to assess the biases of radiative fluxes using 2-stream approximations are historical and extensive (e.g., Chandrasekhar, 1950; Coakley & Chylek, 1975; Harshvardhan & King 1993; King & Harshvardhan, 1986; Liou, 1974; Meador & Weaver, 1980; Räisänen, 2002; W. J. Wiscombe and Grams, 1976). Two modern studies have carried this further with extensive analysis of 2-stream versus exact solutions, using a 128-stream RT code (LBYY) and Monte Carlo simulations (B2015), and we use them in our evaluation of 2-stream versus 8-stream solutions and errors in δ -scaling.

4.1. Phase Function Approximations for 2-Stream Assessed With 8-Stream Scattering

The method designated δ -1 here adopts the commonly used Henyey-Greenstein (HG) phase function (Boucher, 1998) to estimate the Legendre coefficient $5g_2$ (i.e., $g_2 = g_1^2$), see Table 2. This coefficient becomes the scaling factor $f = g_1^2$ and is used to calculate a reduced asymmetry factor $g^* = (g_1 - f)/(1 - f)$ and reduced $\tau_{\text{sca}}^* = (1 - f) \tau_{\text{sca}}$. This scaling, for example, is applied in the Mie expansion for liquid-water clouds of RRTMG code used here. Another alternative, designated δ -2 (δ -M method with $M = 2$), uses the Mie phase function's second term, $f = g_2$ and the revised g^* and τ_{sca}^* are calculated as for δ -1. For comparison, an isotropic-equivalent method, designated δ -0 here, drops the asymmetry factor in the phase function and calculates a reduced τ_{sca}^* using $f = g_1$ and the above formulae. From the example in Table 2, δ -0 is the least forward scattering with the largest reduction in τ_{sca} ($1-f = 0.14$); δ -2 has the next largest reduction ($1-f = 0.20$); while δ -1 has the least ($1-f = 0.25$), reducing τ_{sca} by only a factor of 4. SJ versions have been coded that rewrite the cloud optical depth and scattering phase function in accord with δ scaling and are designated eponymously as $SJ/\delta 0$, $SJ/\delta 1$, and $SJ/\delta 2$. In addition, version SJ/HG uses 8-stream scattering and an 8-term HG expansion based only on the first term of the Mie phase function: $P_{\text{HG}}(\Theta) = 1 + \sum_{L=1:7} (2L + 1) (g_1)^L P_L(\cos[\Theta])$. Rayleigh scattering by air (an important component in the visible region, see above) must also be approximated in 2-stream codes, truncated from $P_{\text{Ray}} = \frac{3}{4}(1 + \cos^2[\Theta])$ to $P_{\text{Ray}} = 1$, removing the forward-backward scattering lobe of the phase function to make it isotropic; this is implemented here as version SJ/Ray . All these versions use the standard 8-stream scattering code. The δ -scaling and HG approximation tests here are run without ice clouds to aid in comparisons with RRTMG.

Approximating Rayleigh scattering as isotropic, required in all 2-stream codes, fortunately has inconsequential errors: all three global mean error metrics (reflection, atmospheric absorption, surface heating) are measurable but within $\pm 0.01 \text{ W m}^{-2}$ (T1/R17). There is little systematic geographic or zenith angle errors since the root-mean-square (rms) errors are $\leq 0.1 \text{ W m}^{-2}$. This test was done with the 8-stream code.

For clouds, the δ -scaling errors are modest in terms of global mean, within $\pm 0.4 \text{ W m}^{-2}$ for any of the three primary components (**T1/R10-11-12**). The pattern is interesting in that all three methods show a similar -0.3 W m^{-2} error in reflected flux, but the surface absorption error shifts from $+0.07$ to $+0.23$ to $+0.44 \text{ W m}^{-2}$ in the order $\delta-1$ to $\delta-2$ to $\delta-0$, being caused by the successively greater reductions in τ_{sca} by the scaling factors of 0.25, 0.20, and 0.14. Thus in $\delta-1$ the reduction in reflected flux goes into atmospheric heating; while in $\delta-2$ and $\delta-0$, it goes into surface heating. Global mean HG errors are $(-0.05, 0.14, -0.09 \text{ W m}^{-2})$ for the three primary components (**T1/R9**). This pattern—less reflection and less forward scattering to the surface—is caused by the weaker forward and backward scattering peaks in the HG phase function. Although these mean errors are modest, there is no basis and no cost advantage to using an 8-term HG phase functions. For a complete analysis on the radiative flux biases due to HG approximations on the clouds and dust aerosols, see LBYY.

The δ -scaling 2-stream models are optimized to give reasonable averages, but with large opposite-sign errors at different SZA (Joseph et al., 1976; W. J. Wiscombe, 1977). Thus, we examine the geographic pattern of δ -scaling errors for fixed sun (00Z) in Figure 4 and find that $\text{SZA} \sim 40^\circ$ (green dashed oval) is the zero-error point for all three δ -scaling methods. Figures 4a–4c columns show the three primary components: reflected flux, atmospheric absorption, and surface absorption, respectively. The first three rows (i)–(iii) show the sequence $\delta-0$ to $\delta-2$ to $\delta-1$, respectively, in order of decreasing scaling factors. Looking at the region with $\text{SZA} < 40^\circ$ (inside the green oval), we see that the error in reflected flux is positive and greatest for $\delta-0$ and decreases along the sequence. Similar dipole patterns as a function of zenith angle can be seen in Figure 2 of LBYY with the relative errors about 20% for thin liquid clouds (optical depth, $\tau < 1$) to about 5% for thicker clouds ($\tau \sim 10$). The zero-error point moves from SZA of about 60° at $\tau \sim 0.01$ toward smaller SZAs with increasing τ up to 10. For thick clouds ($\tau > 10$), the relative errors and the dipole patterns are comparatively minor. Similar patterns are also shown in Figures 2c and 5a of B2015 for the visible band and the near-infrared band respectively. However, the zero-error point in B2015 stays in the $\text{SZA} = 40\text{--}50^\circ$ range.

The vertical profiles of the atmospheric absorption errors in Figure 4 are shown in Figure S1. The error at high sun (i.e., inside the oval $\text{SZA} < 40^\circ$) goes from overall negative for $\delta-0$ (-2.7 W m^{-2}), to small positive for $\delta-2$ ($+0.8 \text{ W m}^{-2}$), to strongly positive for $\delta-1$ ($+2.9 \text{ W m}^{-2}$), and the profiles are consistent with the means. This sequence is similar to the sequence for global mean atmospheric heating, $+0.07$ to $+0.23$ to $+0.44 \text{ W m}^{-2}$. When we mix 2-stream versus 8-stream (*RRTMG* vs. *SJ/Mie*, red line in Figure S1), the profile errors are large and oscillate. While the column mean difference is small in this case, the rms difference is as large as the worst case $\delta-0$. In high solar latitudes (50°S – 70°S and 30°N – 50°N in January) where SZA remains large at all hours, the monthly mean averages have consistent errors of $1\text{--}2 \text{ W m}^{-2}$ at the surface (Figure 1c). The δ -scaling approximations studied here have largest impact on the reflected and surface fluxes rather than atmospheric absorption.

4.2. 2-Stream versus 8-Stream and Exact Scattering Solutions

Here we assess the errors in 2-stream scattering by comparing with 8-stream solutions standard in Solar-J using our metric. Fortunately, two key studies have carried this further with exact solutions, using a 128-stream RT code (LBYY) and Monte Carlo simulations (B2015). Räisänen's (2002) work on tuning 2-stream approximations used comparisons similar to ours with a 16-stream discrete ordinate method. For Solar-J, we coded a 2-stream RT solver based on the Feautrier method using $\delta-1$ scaling, and this version is designated *SJ/2S*. With this we have a clean, bias-free comparison of 2-stream versus 8-stream under clear skies (**T1/R6**) and with liquid clouds using $\delta-1$ scaling for both models (**T1/R13**).

Under clear sky conditions (no clouds, no aerosols) the only scaling in *SJ/2s* is to make Rayleigh scattering isotropic, which has negligible effects as noted above (**T1/R17**). The pattern of 2-stream errors (Figure S2, left column) is unusual, with reflected flux being much larger in regions with high surface albedo (Antarctica, dry sub-tropics). The enhanced reflection is matched by reduced atmospheric absorption. Our surface reflection is Lambertian (isotropic), and it is clear that having only a single zenith angle of 55° for the reflected light reduces the average path length through the atmosphere for the reflected flux and results in a shift from absorbed to reflected. There is also a slight increase in the downward scattered light reaching the surface with only one angle. We looked at the B2015 case studies, but could find no corroboration since

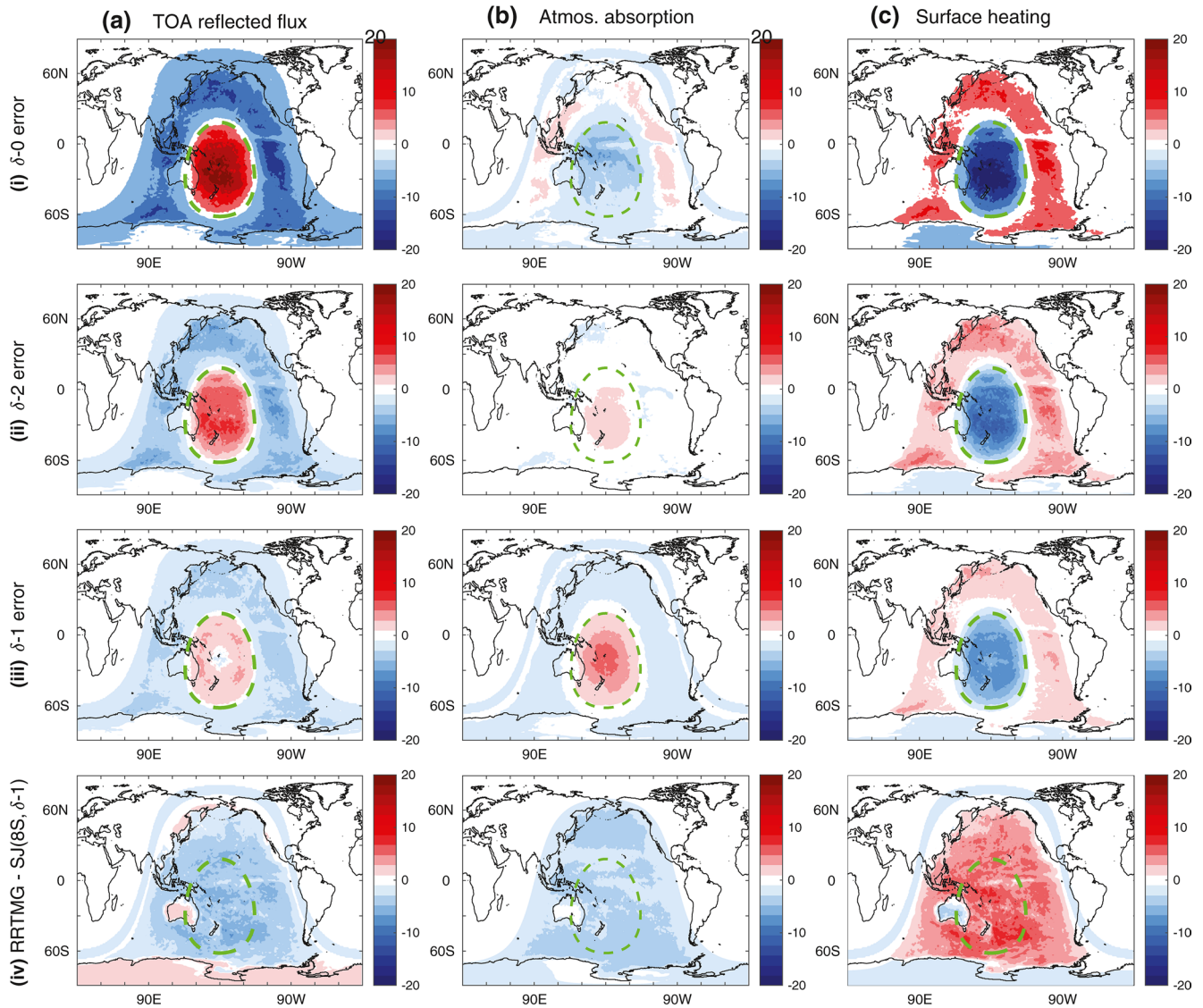


Figure 4. Geographic map of model differences in solar heating terms (W m^{-2}) averaged over 31 days in January at 00Z (sun over the Dateline), with columns show (a) reflected flux, (b) atmospheric absorption and (c) surface heating. Rows (i), (ii), and (iii) show the errors for $\delta-0$, $\delta-2$, and $\delta-1$, respectively, calculated with Solar-J 8-stream scattering relative to the standard Mie phase function, see Table 2. Row (iv) shows the difference, *RRTMG* minus *SJ/δ1*. All calculations use grid-cell averaged liquid clouds only. The green dashed line encloses the region with $\text{SZA} < 40^\circ$.

they used black surfaces. We did find this pattern in the direct comparison of *SJ/δ1* with *RRTMG* (Figure 4, row iv) where *RRTMG*'s 2-stream code also picks out the high albedo land masses with enhanced reflection. Unfortunately, we cannot compare with the LBYY climate model results (their Figure 6) because their 'clear sky' includes, and clearly shows, the importance of northern continental pollution and southern ocean sea salt aerosols in their 2-stream errors.

Trying to separate the effects of δ -scaling errors from 2-stream errors for clouds, *SJ/2S* experiments (*T1/R13*) are carried out extending the phase-function scaling exercises (*T1/R9-R12* and Figure 4). The dipole picture of atmospheric reflection caused by $\delta-1$ scaling of liquid clouds (Figure 4a_{iii}) is consistent with LBYY's Figure 2 row iii, which shows a large shift from increased reflection (high sun) to reduced (low sun), as does (Räisänen's [2002], Figure 2). The LBYY examples in their Figures 2 and 3 are quite informative: For the 4 cases (liquid & ice, reflection & atmospheric absorption), only liquid-reflection shows a distinct pattern change from δ -Eddington (2-stream) to δ -4SHE (4-stream), while the other three show a pattern indic-

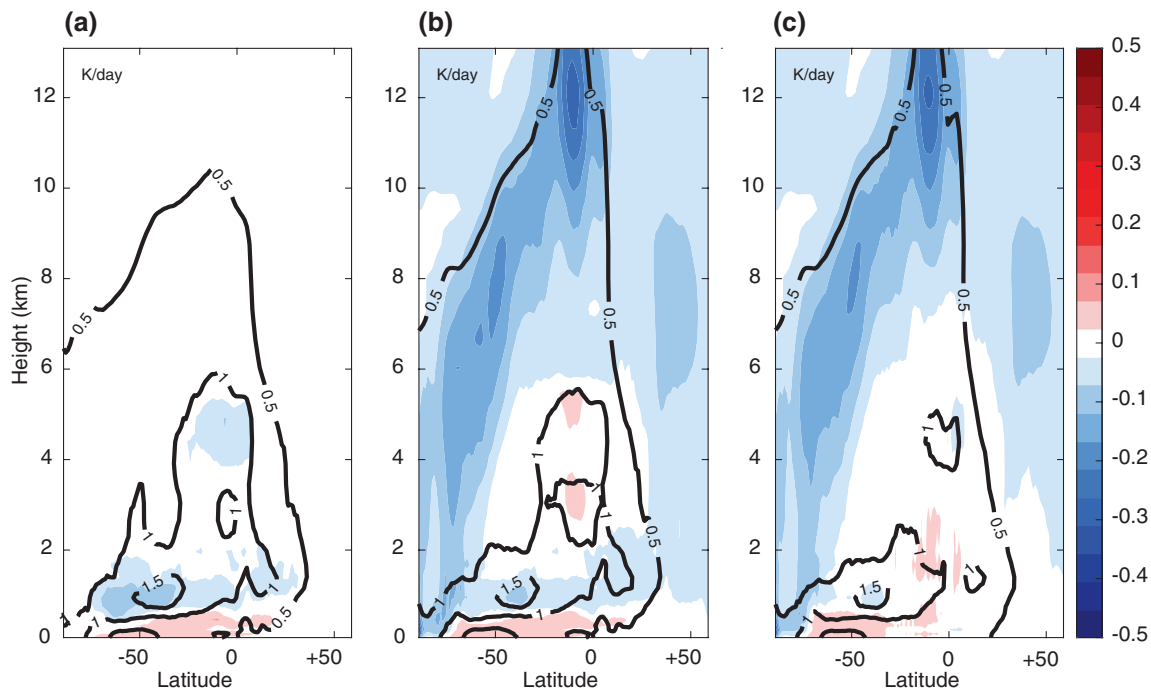


Figure 5. (a) Zonal mean heating rates (black contour lines, 0.5, 1.0, 1.5 K per day) as a function of latitude and height for January 2015 using Solar-J with averaged clouds and no ice water clouds. Color fill (−0.5 to +0.5 K per day color bar, with ± 0.025 as white) show the difference RRTMG-SW v4.0 minus Solar-J. (b) Same as (a) except that ice water clouds are included. (c) Same as (b) except that MAX-RAN cloud overlap is used to generate ICAs; and RRTMG uses McICA to sample the ICAs, while Solar-J uses QCAs (see text).

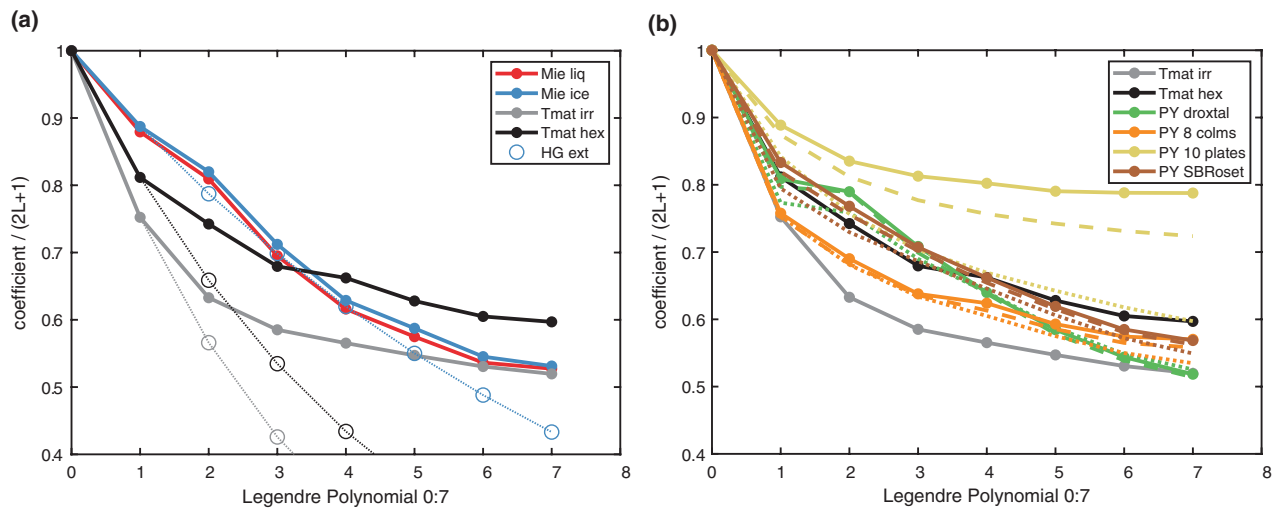


Figure 6. Scattering phase function for ice particles (600 nm, $Reff = 48 \mu\text{m}$) used in Solar-J (T-matrix irregular/warm, T-matrix hexagonal/cold) is compared (a) with Mie calculations for liquid and ice water spheres and (b) with ice crystal habits from the Yang library (droxtals, 8 columns, 10 plates, and solid bullet rosettes) with roughness of 000 (solid), 003 (dashed), and 050 (dotted). The coefficients of the Legendre polynomials (PL, $L = 0:7$, as used in Solar-J) derived from a fit to the full phase function are divided by $2L+1$. The Henyey-Greenstein extension of the phase function using the asymmetry parameter g (value shown at $L = 1$) shows a reasonable approximation to the Mie phase functions, but not those for ice crystals. The T-matrix hexagonal/cold phase function is similar to three of the four habits shown here, but they all differ from the “10 plates” habit (except for roughness 050). The 10 plates habit has the most asymmetric shape with a maximum length to effective diameter ratio of 7.5 (vs. 1.0 for a sphere).

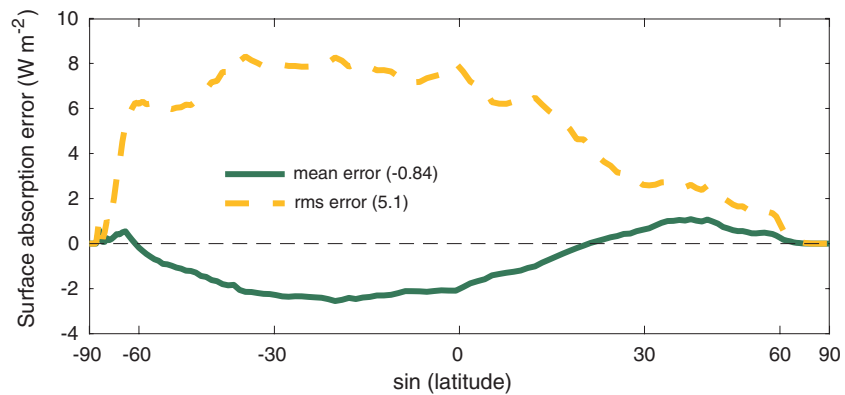


Figure 7. Zonal mean error (mean and rms in W m^{-2}) for fixed ocean surface albedo (OSA) versus interactive OSA that depends on wavelength, incident angle and wind.

ative of the convergent error reduction from 2-stream to 4-stream. If we extrapolate to 8-stream, we would see the errors in δ -4SHE drop by a factor of 2–4. We looked separately at the error caused by δ -1 scaling of ice clouds only and found it to be surprisingly small with global mean differences of $< \pm 0.02 \text{ W m}^{-2}$. So straightforward δ -M scaling of ice clouds, here and in LBYY, does not seem to cause much error, but we cannot easily judge the other more parameterized polynomial fits for 2-stream use (Fu, 1996; Zhao et al., 2018).

In terms of atmospheric heating by clouds, LBYY's parametric study (their Figure 3) shows reduced values over most ranges of optical depth and sun angle, and this error is clearly associated with the number of streams. Similar results are shown in B2015 Figures 5c and 5f. B2015 evaluated 2-stream δ -Eddington errors for a typical mixed-cloud atmosphere derived from high-resolution satellite observations (see their Figures 7 and 8). The atmospheric absorption error is typically 0 to -5 W m^{-2} , while the surface heating error is 0 to $+8 \text{ W m}^{-2}$. These errors are similar to our 2-stream minus 8-stream differences (both δ -1 scaling, **T1/R13**) shown in Figure S2 (middle column).

Atmospheric heating rates by clouds are driven predominantly in the IR region ($>778 \text{ nm}$) where *SJ* and *RRTMG* have identical bands and atmospheric properties (unlike the UV-VIS region discussed above). The geographic pattern of atmospheric absorption at 00Z for *RRTMG* minus *SJ*/ δ 1(liquid-only) (**T1/R15**, Figure 4iv) is consistent and uniformly negative. Since both use δ -1 scaling, this supports the result that 2-stream scattering consistently underestimates the atmospheric absorption by clouds. The *RRTMG*(δ -scaling, 2-stream) minus *SJ*(8-term Mie or ice scattering phase function, 8-stream) comparison clearly show less heating by liquid clouds below 6 km altitude (Figure 5a), Results in Figure 5a used the full-phase-function *SJ* (**T1/R14**), and if we use *SJ*/ δ 1 (**T1/R15**) as the reference, differences are similar. But if we go to our 2-stream, δ 1-scaling code (*SJ*/ δ 1/2S, **T1/R16**), the differences with *RRTMG* are greatly reduced. When we add ice clouds, the difference is much larger, from -0.05 to -0.1 K per day throughout most of the troposphere (Figures 5b and 5c). With ice clouds, there is some reversal of differences in the lower mid-troposphere presumably due to cloud overlap. In terms of global mean, atmospheric heating error jumps from about -1 W m^{-2} to -2 or -3 W m^{-2} (**T1/R14&15** v. **T1/R20&21**).

A clear result from B2015, LBYY, Räisänen (2002) and this study is that 2-stream RT codes systematically underestimate the atmospheric absorption by all clouds, and for ice clouds the error is larger in absolute amount. They also overestimate the surface flux, and in all of Räisänen's efforts to tune a 2-stream code, these fundamental errors remain. The ability to resolve the scattered light across multiple zenith angles is critical in calculating heating rates. With only a 55° zenith angle path, the 2-stream scattered light escapes from liquid clouds more easily. Ice cloud errors, however, either reflection or absorption, are not driven by δ -M scaling but by the more parameterized non- δ -Eddington scaling used in many 2-stream codes. Reflection errors for liquid clouds are different and appear to result primarily from δ -Eddington scaling, and in this case expansion to 4-stream or above significantly improves the accuracy (Räisänen, 2002; LBYY) because a more accurate scattering phase function can be used.

5. Ice Cloud Optics

Ice particles in cirrus or mixed-phase clouds come in a wide range of sizes and crystal habits (Kärcher et al., 2014; Platt & Martin, 1997) with a dizzying array of optical properties (Mishchenko et al., 2016; Yang et al., 2018). In part because of their importance in climate and remote sensing, ice clouds are an intense research area (Baum et al., 2005; Heymsfield et al., 2017; Holz et al., 2016; Platnick et al., 2017). The accurate treatment of ice clouds in climate models remains a fundamental uncertainty because—like the case with aerosol size and chemical composition affecting RT—we can observe, but not accurately characterize the full mix of size and crystal habits within ice clouds (Bailey & Hallett, 2009).

Even when the ice cloud is fully characterized, the RT solution remains difficult and approximations are numerous, for example, RRTMG primarily uses a Fu (1996) parameterization for optical depth, single scattering albedo, and asymmetry parameter. Solar-J inherited the Cloud-J photolysis method for ice clouds including effective diameters (Heymsfield et al., 2003) and T-matrix scattering phase functions (Mishchenko et al., 2004) for the visible (~ 600 nm) using two cloud types (pure hexagonal crystals used for $T \leq -40^\circ\text{C}$, and irregular crystals used for warmer ice clouds). Solar-J's use of a scattering phase function across all wavelengths has minimal error for highly scattering wavelengths ($\text{SSA} > 0.9$) because the ice crystal effective diameters are much larger than the wavelength, and thus the asymmetry parameter g is nearly constant at ~ 0.8 in agreement with Bi and Yang (2017). When we compared atmospheric heating rates between RRTMG and SJ for our January climate simulation (averaged clouds, no overlap model used), the importance of ice clouds stands out. Both models can use the same optical properties for liquid clouds (with RRTMG using δ -1 scaling), and thus the differences in tropospheric heating rates are systematic but modest (± 0.05 K per day, Figure 5a, also see case study Figure 4 in H2017). When ice clouds are included, the two models clearly diverge (Figure 5b). Unlike for liquid clouds, SJ cannot simply match the RRTMG's ice-cloud parameterization (Fu, 1996), because it is not a δ -M scaling. RRTMG heating rates are 0.1–0.2 K per day (10%–20%) less than those in SJ throughout the middle-upper troposphere, with the pattern reversed for liquid water clouds in the tropics (2–6 km) presumably due to less shielding by ice clouds. Running both models with fractional clouds assuming MAX-RAN overlap, gives similar results (Figure 5c). While errors in the parameterization of ice optics will contribute to this error, it is most likely caused by 2-stream models as discussed in Section 4.2.

The ice-cloud phase functions shown in Figure 6a show the coefficients of the Legendre expansion to $L = 7$ scaled by $1/(2L+1)$ for Mie, T-matrix, Bi-Yang and some HG expansions. Where possible, values are calculated for $R_{\text{eff}} = 48$ μm and 600 nm. In the Mie calculation, the difference between liquid and ice water is small as expected, but the T-matrix coefficient drops quickly (becoming more isotropic), but then levels off, becoming more forward peaked than the Mie at $L > 5$ for hexagonal ice and $L > 7$ for irregular ice. The HG phase function matches the Mie out to $L = 5$ before becoming more isotropic. For ice clouds (T-matrix), the HG diverges by $L = 2$ and this is why LBYY show a pronounced error in HG versus their 4-stream RT that uses the $L = 2$ & three terms. The divergence of the ice-cloud phase function from any HG-like extrapolation becomes even greater for $L = 4$ – 7 , which can be included in the 8-stream calculation here. We compare in Figure 6b our T-matrix expansion with those calculated from the Bi-Yang database used in LBYY. Our hexagonal T-matrix fits within the class of ice crystals from Bi-Yang except for the extremely forward peaked 10-plates example, which has the greatest difference from a sphere: the ratio of maximum diameter to effective diameter is 7.5 versus 1.2 for droxtals. As expected, our irregular T-matrix phase function has a larger fraction of isotropic-equivalent scattering, similar to what happens with increased roughness in the Bi-Yang crystals. All of these phase functions have similar asymptotic behavior. The Solar-J treatment of ice clouds certainly needs to be updated (e.g., Bi & Yang, 2017; Heymsfield et al., 2013), but is solidly linked to the optical properties of ice clouds, and is adequate for the comparisons here.

Some recent efforts have updated the RRTMG options for ice clouds. Zhao et al. (2018) combined more realistic mixtures of ice crystals (Baum et al., 2011) with the Yang library of optical properties for a wide range of wavelengths, sizes and habits (Bi & Yang, 2017; Yang et al., 2013) to create a new Baum-Yang ice cloud model for RRTMG(BY). We checked our ice cloud optics against Zhao et al.'s tables and found consistent values in the scattering regions (VIS and IR bands #S19–22) for SSA (Figure 3a) and g , but we calculate about 10% more extinction for the same effective diameter and ice water content. We were not able to implement the new RRTMG(BY) code in Solar-J and can only assess their published results: RRTMG(BY) has weaker

in-cloud heating than *RRTMG(Fu)* (their Figures 2d–f). As shown in LBYY, this is a standing problem with all 2-stream codes, and the new ice optics appear to have made this error worse.

Baek and Bae (2018) updated the Korean Integrated Model (KIM) radiation code based on *RRTMG* using the Yang et al (2013) ice optics library and tested with the KIM forecasting system. They show that use of the updated ice optics versus the *RRTMG(Fu)* parameterization leads to more in-cloud heating by about 0.1 K per day and less reflection by about 3 W m^{-2} (their Figure 8). These results appear to correct much of the error in ice-cloud heating found here and in LBYY and B2015, but they are difficult to understand or consolidate with *RRTMG* simulations here because KIM uses their adaptation of *RRTMG*'s 2-stream solver. We find that the large error in ice cloud heating is due to 2-stream methods and not scaling methods. When we truncate the ice-cloud phase function with δ -1 scaling (*SJ/δ1ice* vs. *SJ, T1/R18*) we find very small differences in reflection or atmospheric heating (Figure S2, right column). The *SJ* ice-cloud phase functions are consistent with the Y2013 library, and so if Baek and Bae are to correct much of the ice-cloud errors within a 2-stream solution then we need to understand how these new ice-cloud optics are implemented.

The treatment of ice-water clouds remains a large source of uncertain error in solar-heating codes. A more accurate treatment of ice water clouds will combine the physics of individual particles (e.g., Mishchenko et al, 2016; Yang et al., 2018) with the actual mix of such particles observed in the atmosphere (Heymssfield et al., 2017; Thornberry et al., 2017). We concur with LBYY that multi-stream scattering codes, such as CCCma's 4-stream or Solar-J's 8-stream, will become more essential for accuracy and also more cost effective in climate models. They are also necessary to assess ice cloud optics since the 2-stream problems obscure such improvements.

6. Cloud overlap

Treatment of overlapping clouds provides a challenge for climate models in solving the RT problem as well as other physical processes such as precipitation and scavenging of trace species (e.g., Neu & Prather, 2012). If we can specify the cloud overlap in terms of separated independent column atmospheres (ICAs, with multiple 1-D RT solutions) or in terms of 3-D cloud fields (a 3-D RT problem) then we can evaluate the errors of different RT methods for case studies using highly accurate methods. One challenge lies in developing an RT solution that works efficiently in climate models, but a greater uncertainty lies the specification of cloud structures. Fortunately, with modern observing systems and cloud resolving models, this problem is becoming less of an unknown (H. W. Barker & Li, 2019; B2015).

6.1. One-Dimensional ICAs

When grid-cell layers specify cloud fraction (presumably in terms of areal coverage), explicit information or an algorithm is needed to describe exactly how the cloud is distributed within the grid cell and how the cells overlap. Most climate models do not resolve the horizontal scale of clouds and simply report a cloud water column and fractional coverage in the cell. A typical algorithm is MAX-RAN (Briegleb, 1992): the cloud is assumed to be uniform in the cloudy fraction of the cell; when two adjacent vertical layers have clouds, they are maximally overlapped; but when two cloudy layers, or two groups of contiguous (maximally) overlapped clouds, are separated by a clear layer, they are randomly overlapped (e.g., see figures in Neu et al., 2007). More realistic cloud-overlap algorithms have been developed based on observations showing that cloud overlap has a vertical decorrelation length (H. W. Barker, 2008; Bergman & Rasch, 2002; Di Giuseppe & Tompkins, 2015; R. J. Hogan & Illingworth, 2000; A. M. Tompkins & Di Giuseppe, 2015). This EXP-RAN method assumes an exponential decorrelation length for connected cloud layers but random overlap across clear layers (A. M. Tompkins & Di Giuseppe, 2007).

In a manner similar to Hogan and Bozzo's (2018) deterministic cloud-cover generator that goes from MAX-RAN to EXP-RAN, Cloud-J developed a deterministic ICA generator for MAX-RAN and then adapted it to use vertical decorrelation lengths in its MAX-COR algorithm (M. J. Prather, 2015). Chemistry models need the selection of ICAs for any overlap method to be deterministic because many critical applications require perturbation-control pairs without stochastic noise (e.g., M. J. Prather & Hsu, 2010). Thus Solar-J cannot

use a stochastic cloud generator (e.g., Räisänen et al., 2004), and this drove the structure of our cloud overlap algorithm. MAX-COR was designed to be (i) deterministic, (ii) linear in cost with increasing numbers of layers, and (iii) robust when cloud data are averaged in time or space, because such averaging tends to eliminate cloud-free layers and revert to MAX overlap. Based on observations of decorrelation length (Kato et al., 2010; Naud et al., 2008; Oreopoulos et al., 2012; Pincus et al., 2005), MAX-COR defines 6-layer groupings by altitude range. Because decorrelation is small across the vertical range of each group, we assume MAX overlap within each group and a decorrelation of the overlap of each MAX group with its neighbor. Adopting terminology of climate community, MAX-COR is effectively a MAX-EXP algorithm. By quantizing the cloud fraction to the nearest 10% and allowing an independent cirrus shield at the top, the absolute maximum number of ICAs under MAX-EXP is $<5 \times 10^6$ and thus ICAs can be rapidly defined and binned with low computational overhead. Deterministic EXP-EXP or EXP-RAN models in our code would have to enumerate up to 2^{33} ICAs for our model that has potentially 33 cloudy layers, which is truly prohibitive and not linearly scalable with resolution. We believe that a MAX-COR or MAX-EXP algorithm is likely the most stable and scalable deterministic ICA generator for vertical cloud decorrelation algorithms. The RRTMG v4.0 code available at the time of this study uses primarily MAX-RAN cloud overlap, but the new v5.0 code includes an EXP-RAN option. Thus, our comparisons of cloud-overlap results with the RRTMG code are limited to MAX-RAN. Within Solar-J we can run both MAX-RAN (*SJ/RAN*) and the standard MAX-COR (*SJ*) and thus compare with J. K. P. Shonk and Hogan (2010), as discussed below.

Let us accept that ICAs generated by cloud overlap algorithms can be solved with 1D RT as horizontally homogeneous plane parallel layers, then the next step is how to solve the RT problem for all ICAs and average the results. The number of ICAs are often numerous enough that no practical climate RT code can solve them all, and most codes do not even count them all (Räisänen et al., 2004). RRTMG randomly selects an ICA for each wavelength bin in the RT solution, a method designated Monte Carlo ICA (McICA, Pincus et al., 2003). McICA has errors at each time step by mixing ICAs across wavelengths and by not accurately sampling the average of ICAs (e.g., average cloud optical depth) in that time step. McICA is intended to deliver the correct mean when averaged enough times over the same cloud system, but it has hourly grid-cell rms errors of 40 W m^{-2} (H. W. Barker et al., 2008; Pincus et al., 2003). A key underlying premise is that solar heating errors propagate symmetrically and linearly in the climate system and average out, as was found for simple forecast models. Assessing net bias errors caused by noisy heating rates would need to examine nonlinear processes in hydrology, cloud systems, ecosystem productivity, and air quality in Earth system models (e.g., Pincus & Stevens, 2013).

With a deterministic ICA generator, we can calculate an "exact" non-stochastic answer as was done for limited test cases in M. J. Prather (2015), but we could not afford to do this for our January climate metric. Solar-J identifies and sorts all ICAs by cloud optical depth and then selects up to four representative quadrature column atmospheres (QCAs) each with a fractional area to represent the distribution of ICAs. The full-wavelength RT solutions are completed for each QCA (Neu et al., 2007). See Figure S3 for a global picture of the average frequency of occurrence of the 4 QCA bins for January 2015. Cloud quadrature does a very good job of averaging over the ICAs with net bias errors of $\sim 1\%$ in solar intensity and rms errors of 2%–4%. To reach equivalent accuracy for a single time step using random selection would require about 50 ICAs each with full wavelength calculation (not as in McICA) versus an average of 2.8 QCAs (many grid cells have less than 4 QCAs).

The binary cloudy-or-clear within a grid cell fails to account for varying cloud thickness, and J. K. P. Shonk and Hogan (2008) invented the Tripleclouds algorithm to represent in-cloud heterogeneity with a thick homogeneous core cloud surrounded by a thinner homogeneous edge cloud. In the binary cloudy-clear approach, two layers can generate 2^2 ICAs from their overlap combinations, but in the Tripleclouds system, it is 3^2 . Tripleclouds address a significant source of error in solar heating calculations, but there is considerable uncertainty in partitioning the cloud water path into thick and thin regions (J. K. P. Shonk et al., 2010). Tripleclouds has not been implemented in Solar-J, and its primary use is in the RT codes of the ECMWF (R. J. Hogan & Bozzo, 2018) and ACCESS (Franklin et al., 2013).

The cloud overlap assumption is clearly an important source of error. Shonk and Hogan (2010, hence SH2010) demonstrated this using 4 months of ERA-40 data and cloud optics similar to the January IFS cloud data used in Solar-J here. SH2010s calculation using Tripleclouds showed that EXP-RAN produced about

+4 W m⁻² more reflected sunlight than MAX-RAN, as expected because EXP-RAN has greater global cloud cover. Similarly, our calculation (using binary clouds) shows MAX-COR has +1.4 W m⁻² more reflection and similar reduction in surface heating versus MAX-RAN (**T1/R19**). This reduced effect is understandable because MAX-COR does not shift to random overlap if there is a gap in one layer. Both models show that increased reflection is balanced by reduced surface heating, with little change in atmospheric heating. A direct comparison is not possible since SH2010 did not calculate EXP-RAN for binary clouds. SH2010s EXP-RAN appears to separate cloud layers more than the MAX-COR sub-groups, but their Tripleclouds may also affect the EXP-RAN results. SH2010 find that the Tripleclouds versus binary clouds is a larger effect (−6 W m²). We support the SH2010 results that cloud overlap is major uncertainty in current models.

Unfortunately, Solar-J coding is not flexible enough to run RRTMG-style, McICA-like calculations. Thus we compare the two codes, each using their best simulation of MAX-RAN clouds: RRTMG v4.0 (MAX-RAN overlap, McICA, 2-stream, and δ -1 scaling) minus SJ/RAN (MAX-RAN, QCA, 8-stream, Mie/T-matrix phase functions). The primary global-mean component differences are (−2.1, −2.4, +4.4 W m⁻²) (**T1/R21**), among the largest found in our range of case studies. If we remove the McICA sampling issues by specifying averaged clouds, the RRTMG minus Solar-J difference is (−0.4, −2.7, +3.0 W m⁻²) (**T1R20**). These differences are due in large part to 2-stream errors (−0.7, −1.7, +2.4 W m⁻²) (**T1R13**) noted in Section 3, but clear-sky differences (**T1/R4**) make the comparison worse. It is encouraging that the profiles of heating rates in Figures 5b and 5c show that McICA-QCA differences are less than the errors we can associate with 2-stream RT. The averaged cloud (Figure 4b) does produce different mean heating rate contours (solid lines) than does MAX-RAN (Figure 4c), and with that there is some shift in the RRTMG-SJ difference pattern in the lower tropical troposphere.

6.2. Three-Dimensional Cloud Fields

How will the next generation of solar heating codes deal with cloud structures, both vertically and horizontally? The vertical coherence of clouds has become standard in RT codes; and the horizontal coherence, including the variations in cloud water path within a cell is now being modeled explicitly with Tripleclouds (J. K. P. Shonk & Hogan, 2008). This work points toward new approaches for understanding how ICAs within a grid cell column interact with each other. As the horizontal resolution in climate models drops to cloud-resolving scales then it is obvious that neighboring column atmospheres are not independent. B2015 used 3D coherent clouds fields from the NASA A-train (MODIS, CloudSat, CALIPSO) with 3D Monte Carlo RT codes to show that 1D minus 3D RT errors in heating and reflection were as large as 2-stream minus exact multi-stream errors. Recent comparison of chemistry models with aircraft measurements (Hall et al., 2018) showed that scattered light from neighboring clouds alters photolysis rates in clear-sky ICAs.

While 3D RT codes are beyond reach for climate models, R. J. Hogan et al. (2016) developed the Speedy Algorithm for Radiative Transfer through Cloud Sides code as a new operational 2-stream RT code for the ECMWF models (R. J. Hogan & Bozzo, 2018). This code includes first-order 3D elements by coupling the 2-stream column solutions across the three elements in each layer. H. W. Barker and Li (2019) investigate how to average the solar heating over a domain consisting of many columns by using more frequent RT calculations but for an objectively selected sub-sample of the atmospheres therein based on the cloud distributions (Partitioned Gauss–Legendre Quadrature). This exciting approach shares many features with Solar-J QCAs (used for ICAs within a single cell), such as producing small rms errors relative to random sampling.

7. Ocean Surface Albedo

Climate models often assume that the ocean surface albedo (OSA) is constant in the visible, typically 0.06 for all incident solar direct and diffuse radiation. OSA varies greatly with incident angle and somewhat with wavelength, wind speed, and chlorophyll concentrations (Z. Jin et al., 2004; 2011; Li et al., 2006; Taylor et al., 1996 and see Figure S4). Recently, this interactive parameterization of OSA (Z. H. Jin et al., 2011) has been implemented in two Earth system models (S  ferian et al., 2018) and shown to better match satellite derived OSA. Here we take the FORTRAN module directly from S  ferian with only minor modifications. Because Solar-J resolves the downward diffuse radiation with four angles, we calculate the albedo specifi-

cally for those four angles plus the direct solar beam, and do not use the OSA averaged albedo for "diffuse radiation," which is necessary when using 2-stream RT codes.

Solar-J lower boundary condition is 2nd-order in finite-difference RT solution and has not changed since the original Fast-J documentation (Wild et al., 2000). The interactive OSA requirement that each angle has a different albedo required a rewrite of the 2nd-order lower boundary condition. The Fast-J Feautrier solver for scattered light (see Equations 9 & 19 of Wild et al., 2000) uses odd-even (leap-frog) first-order finite-difference equations, solving at the lower boundary for j_n .

$$j^n = \frac{1}{2} [I^{\text{up}}(+u_n) + I^{\text{down}}(-u_n)] \equiv \frac{1}{2} [I_n^{\text{up}} + I_n^{\text{down}}], \quad (1)$$

where u_n ($n = 1:4$) are the cosines of the zenith angles for the scattered intensity (I). The angles are Gauss points with weights w_n . We assume a Lambertian reflective surface, and hence $I^{\text{up}}(+u_n)$ is isotropic and denoted simply as I^{up} . The solution requires a linear equation relating I^{up} to the intensities at the four angles $j_{n=1:4}$. For notation below, we use Σ to denote the sum over the quadrature angles $n = 1:4$. The upward flux from the lower boundary is the cosine-weighted sum of the specific intensity

$$F^{\text{up}} = 2 \Sigma I^{\text{up}} u_n w_n = I^{\text{up}}, \text{ where } \Sigma u_n w_n \equiv \frac{1}{2} \quad (2)$$

The upward flux can also be calculated in terms of the downward incident fluxes at the four quadrature angles and direct beam, but with angle-specific albedos A_n and A_0 .

$$F^{\text{up}} = 2 \Sigma A_n I_n^{\text{down}} u_n w_n + \frac{1}{4} A_0 u_0 F^{\text{solar}} \quad (3)$$

Substituting $I_n^{\text{down}} = 2 j_n - I^{\text{up}}$ from Equation (1), we get

$$I^{\text{up}} = F^{\text{up}} = 2 \Sigma A_n (2 j_n - I^{\text{up}}) u_n w_n + \frac{1}{4} A_0 u_0 F^{\text{solar}} = 4 \Sigma j_n A_n u_n w_n - 2 I^{\text{up}} \Sigma A_n u_n w_n + \frac{1}{4} A_0 u_0 F^{\text{solar}} \quad (4)$$

If A is a constant, this reduces to Equation 19 of Wild et al. (2000).

$$I^{\text{up}} = \left[\frac{4A}{1+A} \right] \Sigma j_n u_n w_n + \left[\frac{A}{1+A} \right] A_0 u_0 F^{\text{solar}} \quad (5)$$

With A_n depending on u_n , we derive the new lower boundary condition for I^{up} .

$$I^{\text{up}} = \left[\frac{4}{1 + 2 \Sigma A_n u_n w_n} \right] \left[\Sigma j_n A_n u_n w_n + \frac{1}{4} A_0 u_0 F^{\text{solar}} \right] \quad (6)$$

Evaluation of interactive OSA (SJ minus SJ/OSA) uses the full spherical geometric atmosphere with MAX-COR cloud overlap. The global mean errors with fixed OSA are (+0.7, +0.2, -0.8 W m⁻²) for (reflection, atmospheric absorption, and surface absorption), respectively (**T1/R22**). The global mean error, fixed minus interactive, can be adjusted to near zero by selecting the fixed OSA, but there remains a strong latitudinal error of 3 W m⁻² in ocean heating associated with high sun, see Figure 7. The zonal rms errors are large, 2–8 W m⁻², because of the wide diurnal range of solar zenith angles over the day, but given the thermal inertia of the upper ocean layers, this probably averages out. Overall, these results are similar to those found in S  ferian et al. (2018).

8. Findings and Recommendations

Our goal here is to provide an extensive analysis of the many uncertainties or known errors in our climate model calculations of solar heating rates, and to do this in a consistent framework with climate-relevant diagnostics. We wanted to review as many of the uncertainties/errors as we could, from the mundane (Rayleigh scattering) to the formidable (cloud overlap), so that we could assess priorities for improving the accuracy of solar radiation in climate models. We began this study with the hope of comparing Solar-J directly to AER's RRTMG version 4.0 code, and thus we ran both codes within our chemistry-transport model to

integrate over forecast modeled atmospheres. We became quickly humbled when even the clear-sky differences in the two, at the 1 W m^{-2} level, could not be fully resolved without a complete rewrite of one code or the other. Nevertheless, the parallel simulations with RRTMG were informative (e.g., 2-stream problems) and helped our analysis here. These findings are based primarily on the extensive variants and adaptations made to the Solar-J RT code. In some cases, our comparisons with other published work has provided insight and allowed us to draw insight beyond our own simulations.

Here, we make recommendations based on the magnitude of error and the difficulty or extra computational cost in improving the models. The levels of ranking include 0 (inconsequential errors), 1 (modest errors and easy/cost-effective fix, or significant errors but hard to fix), 2 (significant errors and ready/cost-effective fix).

1. **Spherical, refracting atmosphere, level 2.** Flat-atmosphere models have 1.9 W m^{-2} less incident sunlight and 1.1 W m^{-2} less heating of the climate system (atmosphere and surface) than do spherical, refracting atmospheres (T1/R23&24). Errors can reach 4 W m^{-2} (monthly averages) at high latitudes. Spherical solar ray-tracing with refraction can and should be readily implemented with simple ray tracing code (P2019) and incorporated in standard 2-stream codes (Spurr & Natraj, 2011). There will be minor costs in that about 56% of the Earth, rather than 50%, will require radiation calls every time step.
2. **Geometrical, expanding atmosphere, level 1.** Shifting from geopotential to geometric coordinates is a conceptual change and will need more thought on how to account for the extra mass in the upper layers as well as the extra solar heating (P2019) (T1/R25).
3. **Stratospheric heating, level 1.** Differences in stratospheric heating rates between RRTMG-SW and Solar-J are large ($\sim 10\%$, see Figure 2 of H2017). These are likely caused by the different cross section for O_3 absorption and the lack of O_2 absorption in RRTMG (significant in the mid stratosphere). These differences can and should be readily resolved with some group efforts like PhotoComp (2010) but with diagnosed heating rates, and with the inclusion of O_2 photolysis as heating. (probably no impact on overall cost).
4. **Resolving UV-VIS absorption and scattering, level 1.** The 300–700 nm sunlight that reaches the surface interacts primarily with broad band features of O_3 absorption (Hartley-Huggins and Chappuis bands) and Rayleigh scattering that vary differently and widely across these wavelengths. The RRTM-based codes accurately resolve the O_3 and Rayleigh features with g-points. Solar-J bands are optimized for photolysis but not heating rates. Its number of visible broad bands needs to be expanded beyond 2, and those cross section adjusted to Watts weighting. Aerosol extinction, but not clouds, also varies widely across visible wavelengths, and all solar RT codes may consider testing aerosol wavelength dependence across their broad bands. (important, not too difficult, but limited to Solar-J, hence level 1).
5. **Resolving IR cloud absorption, level 2.** The infrared wavelength bands (#19 through #27 in RRTM) need to be reformulated to more accurately account for the absorption spectrum of liquid and ice water. The error is great in terms of atmospheric absorption (global average of $\sim 2 \text{ W m}^{-2}$ excess heating) and cloud top heating (25% too great), likely affecting the life cycle of clouds. The LZL approach of doing a double correlated-K for water vapor and liquid water within the RRTM bands may be difficult when ice water is added (which it must be) since the absorption features are shifted in wavelength from those of liquid water. Our finding like Edwards and Slingo (1996) is that most of this error is simply failure to resolve the widely different SSA within each broad RRTM band. We suggest it is time to drop the RRTM band structure and re-group a set of non-contiguous narrower bands into a single group with similar cloud SSA, and then redo the correlated-K for water vapor within that new group. This would probably not increase the computational cost of the RT code, but does involve substantial new research.
6. **Rayleigh scattering, level 0.** Forcing Rayleigh scattering to be isotropic, as required in current 2-stream codes, is inconsequential.
7. **δ -scaling of the cloud scattering phase function, no assessment.** The errors caused by δ -Eddington and other scaling methods are significant, causing systematic errors at the $1\text{--}2 \text{ W m}^{-2}$ level across latitudes. Because of the large reduction (factors of 4–7) in liquid cloud optical depths, these scaling methods project much greater solar flux impacting the surface at the SZA rather than as diffuse, scattered light. In terms of reflected light, LBYY show large dipole errors with SZA in reflected flux for δ -Eddington scaling. Fortunately, δ -scaling appears to have little impact on atmospheric heating. There is no assessment here because any solution requires multi-stream scattering, a more formidable task.

8. **Multi-stream scattering, level 2.** From the experiments here and the careful error analysis of B2015 and LBYY, it is clear that 2-stream scattering is a major source of error in many codes, including RRTMG-SW. 2-Stream RT codes uniformly underestimate atmospheric heating rates by more than 2 W m^{-2} . They appear to have bias errors in reflected flux that depend on surface albedo. 2-stream requires some δ -scaling method with its own source of errors (vii above). In terms of solar heating errors, the 2-stream errors are comparable to the errors in moving from 1D to 3D RT (B2015). While full 3D RT is beyond the capabilities of climate models, multi-stream is clearly not. The GFDL and CCCma models can currently use 4-stream RT, and it is likely that 6- or 8-stream RT could be optimized for modern processors and software so as to be affordable. For example, the 8-stream Fast-J code is being implemented in E3SM for chemistry-climate modeling. Based on B2015, any of these options would be a significant improvement for climate modeling.
9. **Ice clouds, level 1.** The range of approaches to parameterizing ice clouds in current solar RT codes reflects both the lack of multi-stream capability to resolve the ice crystal scattering and the fundamental uncertainty in prescribing the mix of crystal habits over the wide atmospheric range of ice clouds. With the recent work by Yang and colleagues (e.g., 2013, 2018) quantifying the scattering, and the better atmospheric data from Heymsfield and colleagues (e.g., 2017), we are in a position to remove ice-cloud parameterizations and apply the ice-cloud physics directly to solar RT codes. Two recent efforts to do this have been restricted to RRTMG-like codes and thus resorted to polynomial fits for a 2-stream code. It is clear that we need a good database for typical ice clouds that includes the basic physics, like the phase function needed for 4-stream and higher codes.
10. **Cloud overlap, level 2.** The representation of sub-grid unresolved cloud overlap is critical in solar heating. The simplistic averaging of clouds over the grid cell produced huge errors ($+20 \text{ W m}^{-2}$ in reflected flux in our case) and was quickly dropped in favor of more realistic cloud overlap models such as MAX-RAN. The latest type of cloud schemes includes the successive decorrelation with altitude separation and include EXP-RAN, EXP, and MAX-COR (used here). All of these decorrelation methods increase the effective cloud cover and the reflected solar flux. In our case, MAX-COR has on average 1.4 W m^{-2} greater reflection and less surface absorption, similar to other approaches, including more exotic approaches (J. K. P. Shonk & Hogan, 2008) that add a second, thinner extension around the primary cloud. The range of approaches reflects the basic uncertainty in mapping climate model cloud data (cloud fraction and cloud water content) into 3D fields of clouds. For the 1D, or quasi-1D (R. J. Hogan & Bozzo, 2018) RT codes in climate models, it would be useful to establish some standard, community-wide, satellite-based cloud overlap models (H. W. Barker, 2008; Bankert et al., 2015; Ham et al., 2015; Kato et al., 2010; Tompkins & Giuseppe, 2015) along with a simple ICA generator (e.g., M. J. Prather, 2015) to provide a basis for comparisons.
11. **Monte Carlo or other ICA averaging, level 1.** Using cloud quadrature QCAs to average over ICAs would greatly reduce the numerical noise generated by McICA random selection. It is easy to implement in any code, but would increase the computational costs by a factor of 2.8. There is a broad interest reducing the stochastic noise in heating rates, particularly at high resolution (H. W. Barker & Li, 2019).
12. **Ocean surface albedo, level 2.** Use of an interactive ocean surface albedo that depends on SZA, wind and wavelength would eliminate a latitudinal mean bias of 3 W m^{-2} in surface absorption. This easy fix, can be used in RRTMG, and has already been implemented in the ARPEGE and LMDZ models (S  ferian et al., 2018).
13. **Photosynthetically Active Radiation, no assessment.** PAR is calculated in Solar-J with 4 downward diffuse streams and no δ -scaling of cloud optical depth, which is notably more accurate than 2-stream codes where much of the diffuse light is reported as direct beam. Thus, 2-stream methods have large errors in the diffuse:direct ratio of PAR under clouds or aerosols. We estimate that PAR errors are level 2, but a more thorough analysis would need to couple the direct and diffuse PAR to a land biosphere model to evaluate the errors in primary productivity (equivalent of W m^{-2}).

Data Availability Statement

Dataset, Solar-J source code, and scripts for generating figures and tables are concurrently available to DRY-AD University of California, Irvine with DOI <https://doi.org/10.7280/D1PQ3W>.

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References

Baek, S., & Bae, S. Y. (2018). New optical properties of ice crystals for multiclass cloud microphysics. *Journal of Advances in Modeling Earth Systems*, 10, 2971–2982. <https://doi.org/10.1029/2018MS001398>

Bailey, M., & Hallett, J. (2009). A comprehensive habit diagram for atmospheric ice crystals: Confirmation from the laboratory, AIRS II, and other field studies. *Journal of Atmospheric Science*, 66(9), 2888–2899. <https://doi.org/10.1175/2009JAS2883.1>

Bankert, R. L., & J. E. SolBrig (2015). Cluster Analysis of A-train data: Approximating the vertical cloud structure of oceanic cloud regimes. *Journal Applied Meteorology and Climatology*, 54, 996–1008. <https://doi.org/10.1175/JAMC-D-14-0227.1>

Barker, H. W. (2008). Representing cloud overlap with an effective decorrelation length: An assessment using CloudSat and CALIPSO data. *Journal of Geophysical Research: Atmospheres*, 113, D24205. <https://doi.org/10.1029/2008JD010391>

Barker, H. W., Cole, J. N. S., Li, J., Yi, B., & Yang, P. (2015). Estimation of errors in two-stream approximations of the solar radiative transfer equation for cloudy-sky conditions. *Journal of the Atmospheric Sciences*, 72, 4053–4074. <https://doi.org/10.1175/JAS-D-15-0033.1>

Barker, H. W., Cole, J. N. S., Morcrette, J.-J., Pincus, R., Raisanen, P., von Salzen, K., et al. (2008). The Monte Carlo Independent Column Approximation: An assessment using several global atmospheric models. *Quarterly Journal of the Royal Meteorological Society*, 134, 1463–1478.

Barker, H. W., Kato, S., & Wehr, T. (2012). Computation of solar radiative fluxes by 1D and 3D methods using cloudy atmospheres inferred from A-train satellite data. *Surveys in Geophysics*, 33, 657–676. <https://doi.org/10.1007/s10712-011-9164-9>

Barker, H. W. & Li, J. (2019). Accelerating radiative transfer calculations for high-resolution atmospheric models. *Quarterly Journal of the Royal Meteorological Society*, 145, 2046–2069.

Baum, B. A., Yang, P., Heymsfield, A. J., Platnick, S., King, M. D., Hu, Y. X., et al. (2005). Bulk scattering properties for the remote sensing of ice clouds. 2: Narrowband models. *Journal of Applied Meteorology*, 44(12), 1896–1911.

Baum, B. A., Yang, P., Heymsfield, A. J., Schmitt, C., Xie, Y., Bansemer, A., et al. (2011). Improvements to shortwave bulk scattering and absorption models for the remote sensing of ice clouds. *Journal of Applied Meteorology and Climatology*, 50, 1037–1056.

Bergman, J. W., & Rasch, P. J. (2002). Parameterizing vertically coherent cloud distributions. *Journal of the Atmospheric Sciences*, 59, 2165–2182.

Bian, H. S., & Prather, M. J. (2002). Fast-J2: Accurate simulation of stratospheric photolysis in global chemical models. *Journal of Atmospheric Chemistry*, 41(3), 281–296.

Bi, L., & Yang, P. (2017). Improved ice particle optical property simulations in the ultraviolet to far-infrared regime. *Journal of Quantitative Spectroscopy and Radiative Transfer*, 189, 228–237.

Boucher, O. (1998). On aerosol direct shortwave forcing and the Henyey-Greenstein phase function. *Journal of the Atmospheric Sciences*, 55(1), 128–134.

Briegleb, B. P. (1992). Delta-Eddington approximation for solar-radiation in the NCAR community climate model. *Journal of Geophysical Research*, 97(D7), 7603–7612.

Chandrasekhar, S. (1950). *Radiative transfer* (pp. 393). Oxford University Press.

Chou, M.-D., & Suarez, M. J. (1999). A solar radiation parameterization (CLIRAD-SW) developed at Goddard climate and radiation branch for atmospheric studies. *NASA Technical Memorandum NASA/TM-1999-104606*, 15, 48.

Clough, S. A., Shephard, M. W., Mlawer, E., Delamere, J. S., Iacono, M., Cady-Pereira, K., et al. (2005). Atmospheric radiative transfer modeling: A summary of the AER codes. *Journal of Quantitative Spectroscopy and Radiative Transfer*, 91(2), 233–244.

Coakley, J. A., & P. Chylek (1975) The two-stream approximation in radiative transfer: Including the angle of the incident radiation. *Journal of the Atmospheric Sciences*, 32(2), 409–418.

Di Giuseppe, F., & Tompkins, A. M. (2015). Generalizing cloud overlap treatment to include the effect of wind shear. *Journal of the Atmospheric Sciences*, 72(8), 2865–2876.

Edwards, J. M., & Slingo, A. (1996). Studies with a flexible new radiation code. I: Choosing a configuration for a large-scale model. *Quarterly Journal of the Royal Meteorological Society*, 122, 689–719. <https://doi.org/10.1002/qj.49712253107>

Etmann, M., Myhre, G., Highwood, E. J., & Shine, K. P. (2016). Radiative forcing of carbon dioxide, methane, and nitrous oxide: A significant revision of the methane radiative forcing. *Geophysical Research Letters*, 43(24), 12614–12623. <https://doi.org/10.1002/2016GL071930>

Fang, T. M., Wofsy, S. C., & Dalgarno, A. (1974). Opacity distribution functions and absorption in Schumann-Runge bands of molecular-oxygen. *Planetary and Space Science*, 22(3), 413–425.

Franklin, C. N., Sun, Z., Bi, D., Dix, M., Yan, H., & Bodas-Salcedo, A. (2013). Evaluation of clouds in ACCESS using the satellite simulator package COSP: Global, seasonal, and regional cloud properties. *Journal of Geophysical Research: Atmospheres*, 118, 732–748. <https://doi.org/10.1029/2012JD018469>

Fu, Q. (1996). An accurate parameterization of the solar radiative properties of cirrus clouds for climate models. *Journal of Climate*, 9, 2058–2082.

Goody, R., West, R., Chen, L., & Crisp, D. (1989). The correlated-k method for radiation calculations in nonhomogeneous atmospheres. *Journal of Quantitative Spectroscopy Radiative Transfer*, 42(6), 539–550.

Grant, K. E., & Grossman, A. S. (1998). *Description of a solar radiative transfer model for use in LLNL. Climate and atmospheric chemistry studies* (pp. 17). UCRL-ID (129949).

Hall, S. R., Ullmann, K., Prather, M. J., Flynn, C. M., Murray, L. T., Fiore, A. M., et al. (2018). Cloud impacts on photochemistry: Building a climatology of photolysis rates from the atmospheric tomography mission. *Atmospheric Chemistry and Physics*, 18(22), 16809–16828. <https://doi.org/10.5194/acp-18-16809-2018>

Ham, S.-H., Kato, S., Barker, H. W., Rose, F. G., & Sun-Mack, S. (2015). Improving the modelling of short-wave radiation through the use of a 3D scene construction algorithm. *Quarterly Journal of the Royal Meteorological Society*, 141, 1870–1883. <https://doi.org/10.1002/qj.2491>

Harshvardhan, & King, M. D. (1993). Comparative accuracy of diffuse radiative properties computed using selected multiple scattering approximations. *Journal of the Atmospheric Sciences*, 50, 247–259. [https://doi.org/10.1175/1520-0469\(1993\)050<0247:CAODRP>2.0.CO;2](https://doi.org/10.1175/1520-0469(1993)050<0247:CAODRP>2.0.CO;2)

Heymsfield, A. J., Kraemer, M., Luebke, A., Brown, P., Cziczo, D. J., Franklin, H., et al. (2017). Cirrus clouds. Ice formation and evolution in clouds and precipitation: Measurement and modeling challenges. *Meteorological Monographs*, 58, Boston, MA: American Meteorological Society. <https://doi.org/10.1175/AMSMONOGRAPHIS-D-16-0010.1>

Heymsfield, A. J., Matrosov, S., & Baum, B. (2003). Ice water path-optical depth relationships for cirrus and deep stratiform ice cloud layers. *Journal of Applied Meteorology*, 42, 1369–1390.

Heymsfield, A. J., Schmitt, C., & Bansemer, A. (2013). Ice cloud particle size distributions and pressure-dependent terminal velocities from in situ observations at temperatures from 0 degrees to -86 degrees C. *Journal of the Atmospheric Sciences*, 70, 4123–4154.

- Hogan, R. J., & Bozzo, A. (2018). A flexible and efficient radiation scheme for the ECMWF model. *Journal of Advances in Modeling Earth Systems*, 10(8), 1990–2008.
- Hogan, R. J., & Illingworth, A. J. (2000). Deriving cloud overlap statistics from radar. *Quarterly Journal of the Royal Meteorological Society*, 126(569), 2903–2909.
- Hogan, R. J., Schäfer, S. A. K., Klinger, C., Chiu, J. C., & Mayer, B. (2016). Representing 3-D cloud radiation effects in two-stream schemes: 2. Matrix formulation and broadband evaluation. *Journal of Geophysical Research: Atmospheres*, 121, 8583–8599. <https://doi.org/10.1002/2016JD02487>
- Holz, R. E., Platnick, S., Meyer, K., Vaughan, M., Heidinger, A., Yang, P., et al. (2016). Resolving ice cloud optical thickness biases between CALIOP and MODIS using infrared retrievals. *Atmospheric Chemistry and Physics*, 16, 5075–5090. <https://doi.org/10.5194/acp-16-5075-2016>
- Hsu, J., Prather, M. J., Cameron-Smith, P., Veidenbaum, A., & Nicolau, A. (2017). A radiative transfer module for calculating photolysis rates and solar heating in climate models: Solar-J v7.5. *Geoscientific Model Development*, 10(7), 2525–2545.
- Jin, Z., Charlock, T. P., Smith, W. L., & Rutledge, K. (2004). A parameterization of ocean surface albedo. *Geophysical Research Letters*, 31, L22301. <https://doi.org/10.1029/2004GL021180>
- Jin, Z. H., Qiao, Y. L., Wang, Y. J., Fang, Y. H., & Yi, W. N. (2011). A new parameterization of spectral and broadband ocean surface albedo. *Optics Express*, 19(27), 26429–26443.
- Joseph, J. H., Wiscombe, W. J., & Weinman, J. A. (1976). Delta-eddington approximation for radiative flux-transfer. *Journal of the Atmospheric Sciences*, 33(12), 2452–2459.
- KärcherDörnbrack, B. A. & Sölch, I. (2014). Supersaturation variability and cirrus ice crystal size distributions. *Journal of Atmospheric Sciences*, 71, 2905–2926.
- Kato S., Sun-Mack S., Miller W. F., Rose F. G., Chen Y., Minnis P., & Wielicki, B. A. (2010). Relationships among cloud occurrence frequency, overlap, and effective thickness derived from CALIPSO and CloudSat merged cloud vertical profiles. *Journal of Geophysical Research*, 115, <http://dx.doi.org/10.1029/2009jd012277>
- King, M. D., & Harshvardhan (1986) Comparative accuracy of selected multiple scattering approximations, *Journal of the Atmospheric Sciences* 43 (8), 784–801.
- Lacis, A. A., & Oinas, V. (1991). A description of the correlated Kappa-distribution method for modeling Nongray gaseous absorption, thermal emission, and multiple-scattering in vertically inhomogeneous atmospheres. *Journal of Geophysical Research*, 96(D5), 9027–9063.
- Li, J., Barker, H., Yang, P., & Yi, B. (2015). On the aerosol and cloud phase function expansion moments for radiative transfer simulations. *Journal of Geophysical Research: Atmospheres*, 120, 12128–12142. <https://doi.org/10.1002/2015JD023632>
- Lin, Z. Y., Chen, N., Fan, Y. Z., Li, W., Stamnes, K., & Stamnes, S. (2018). New treatment of strongly Anisotropic scattering phase functions: The delta-M plus method. *Journal of the Atmospheric Sciences*, 75(1), 327–336.
- Liou, K.-N. (1974). Analytic two-stream and four-stream solutions for radiative transfer. *Journal of the Atmospheric Sciences*, 31, 1473–1475.
- Li, J., & Ramaswamy, V. (1996). Four-stream spherical harmonic expansion approximation for solar radiative transfer. *Journal of the Atmospheric Sciences*, 53(8), 1174–1186.
- Li, J., Scinocca, J., Lazare, M., McFarlane, N., von Salzen, K., & Solheim, L. (2006). Ocean surface albedo and its impact on radiation balance in climate models. *Journal of Climate*, 19(24), 6314–6333.
- Logan, J. A., Prather, M. J., Wofsy, S. C., & Mcelroy, M. B. (1978). Atmospheric chemistry - response to human influence. *Philosophical Transactions of the Royal Society A-Mathematical Physical and Engineering Sciences*, 290(1367), 187–234.
- Lu, P., Zhang, H., & Li, J. (2011). Correlated k-distribution treatment of cloud optical properties and related radiative impact. *Journal of the Atmospheric Sciences*, 68, 2671–2688. <https://doi.org/10.1175/JAS-D-10-05001.1>
- Meador, W. E., & W. R. Weaver (1980) Two-stream approximations to radiative transfer in planetary atmospheres: A unified description of existing methods and a new improvement, *Journal of the Atmospheric Sciences* 37(3), 630–643.
- Mishchenko, M. I., Videen, G., Babenko, V. A., Khlebtsov, N. G., & Wriedt, T. (2004). T-matrix theory of electromagnetic scattering by particles and its applications: A comprehensive reference database. *Journal of Quantitative Spectroscopy Radiative Transfer*, 88, 357–406.
- Mishchenko, M. I., Zakharova, N. T., Khlebtsov, N. G., Videen, G., & Wriedt, T. (2016). Comprehensive thematic T-matrix reference database: A 2014–2015 update. *Journal of Quantitative Spectroscopy & Radiative Transfer*, 178, 276–283.
- Mlawer, E. J., Payne, V. H., Moncet, J. L., Delamere, J. S., Alvarado, M. J., & Tobin, D. C. (2012). Development and recent evaluation of the MT_CKD model of continuum absorption. *Philosophical Transactions of the Royal Society A-Mathematical Physical and Engineering Sciences*, 370(1968), 2520–2556.
- Mlawer, E. J., Taubman, S. J., Brown, P. D., Iacono, M. J., & Clough, S. A. (1997). Radiative transfer for inhomogeneous atmospheres: RRTM, a validated correlated-k model for the longwave. *Journal of Geophysical Research*, 102(D14), 16663–16682.
- Myhre, G., Shindell, D., Bréon, F.-M., Collins, W., Fuglestvedt, J., Huang, J., et al. (2013). Anthropogenic and natural radiative forcing. In T. F. Stocker et al. (Eds.), *Climate change 2013: The physical science basis. Contribution of working group I to the fifth assessment Report of the intergovernmental Panel on climate change*. Cambridge, UK: Cambridge University Press.
- Naud, C. M., Del Genio, A., Mace, G. G., Neu, S., J. L., & Prather, M. J. (2012). Toward a more physical representation of precipitation scavenging in global chemistry models: Cloud overlap and ice physics and their impact on tropospheric ozone. *Atmospheric Chemistry and Physics*, 12, 3289–3310. <https://doi.org/10.5194/acp-12-3289-2012>
- Neu, J. L., & Prather, M. J., & (2012). Toward a more physical representation of precipitation scavenging in global chemistry models: cloud overlap and ice physics and their impact on tropospheric ozone. *Atmospheric Chemistry and Physics*, 12(7), 3289–3310. <https://acp.copernicus.org/articles/12/3289/2012/>
- Neu, J. L., Prather, M. J., & Penner, J. E. (2007). Global atmospheric chemistry: Integrating over fractional cloud cover. *Journal of Geophysical Research*, 112, D11306. <https://doi.org/10.1029/2006JD008007>
- Olson, J., Prather, M., Bernsten, T., Carmichael, G., Chatfield, R., Connell, P., et al. (1997). Results from the IPCC photochemical model intercomparison (PhotoComp). *Journal of Geophysical Research*, 102, 5979–5991.
- Oreopoulos, L., Lee, D., Sud, Y. C., & Suarez, M. J. (2012). Radiative impacts of cloud heterogeneity and overlap in an atmospheric General Circulation Model. *Atmospheric Chemistry and Physics*, 12(19), 9097–9111.
- Paynter, D., & Ramaswamy, V. (2014). Investigating the impact of the shortwave water vapor continuum upon climate simulations using GFDL global models. *Journal of Geophysical Research: Atmospheres*, 119(18), 10720–10737. <https://doi.org/10.1002/2014JD021881>
- PhotoComp (2010). In V. Eyring, T. G. Shepherd, & D. W. Waugh (Eds.), *Chapter 6 in SPARC CCMVal report on evaluation of chemistry-climate models*. SPARC Report No. 5, WCRP-30/2010, WMO/TD – No. 40. <https://www.sparc-climate.org/publications/sparc-reports/>
- Pincus, R., Barker, H. W., & Morcrette, J. J. (2003). A fast, flexible, approximate technique for computing radiative transfer in inhomogeneous cloud fields. *Journal of Geophysical Research*, 108(D13), -. <https://doi.org/10.1029/2002JD003322>

- Pincus, R., Hannay, C., Klein, S. A., Xu, K. M., & Hemler, R. (2005). Overlap assumptions for assumed probability distribution function cloud schemes in large-scale models. *Journal of Geophysical Research*, *110*(D15). <https://doi.org/10.1029/2004JD005100>
- Pincus, R., Mlawer, E. J., & Delamere, J. S. (2019). Balancing Accuracy, Efficiency, and Flexibility in Radiation Calculations for Dynamical Models. *Journal of Advances in Modeling Earth Systems*, *11*, (10), 3074–3089. <http://dx.doi.org/10.1029/2019ms001621>
- Pincus, R., Mlawer, E. J., Oreopoulos, L., Ackerman, A. S., Baek, S., Brath, M., et al. (2015). Radiative flux and forcing parameterization error in aerosol-free clear skies. *Geophysical Research Letters*, *42*(13), 5485–5492. <https://doi.org/10.1002/2015GL064291>
- Pincus, R., & Stevens, B. (2013). Paths to accuracy for radiation parameterizations in atmospheric models. *Journal of Advances in Modeling Earth Systems*, *5*, 225–233. <https://doi.org/10.1002/jame.20027>
- Platnick, S., Meyer, K., King, M. D., Wind, G., Amarasinghe, N., Marchant, B., et al. (2017). The MODIS cloud optical and microphysical products: Collection 6 updates and examples from Terra and Aqua. *IEEE Transactions on Geoscience and Remote Sensing*, *55*(1), 502–525. <https://doi.org/10.1109/TGRS.2016.2610522>
- Platt, C., & Martin R. (1997). A Parameterization of the Visible Extinction Coefficient of Ice Clouds in Terms of the Ice/Water Content. *Journal of the Atmospheric Sciences*, *54*(16), 2083–2098. https://journals.ametsoc.org/view/journals/atsc/54/16/1520-0469_1997_054_2083_apotve_2.0.co_2.xml
- Prather, M. J. (1974). Solution of the Inhomogeneous Rayleigh Scattering Atmosphere. *The Astrophysical Journal*, *192*, 787–792.
- Prather, M. J. (2015). Photolysis rates in correlated overlapping cloud fields: Cloud-J 7.3c. *Geoscientific Model Development*, *8*(8), 2587–2595.
- Prather, M. J., & Hsu, J. (2010). Coupling of nitrous oxide and methane by global atmospheric chemistry. *Science*, *330*, 952–954.
- Prather, M. J., & Hsu, J. C. (2019). A round Earth for climate models. *PNAS*, *116*(39), 19330–19335. <https://doi.org/10.1073/pnas.1908198116>
- Prather, M. J., Zhu, X., Flynn, C. M., Strode, S. A., Rodriguez, J. M., Steenrod, S. D., et al. (2017). Global atmospheric chemistry – which air matters. *Atmospheric Chemistry and Physics*, *17*(14), 9081–9102. <https://doi.org/10.5194/acp-17-9081-2017>
- Radel, G., K. P. Shine, & I. V. Ptashnik (2015) Global radiative and climate effect of the water vapour continuum at visible and near-infrared wavelengths. *Quarterly Journal of the Royal Meteorological Society*, *141*, 727–738, <https://doi.org/10.1002/qj.2385>
- Räisänen, P. (2002). Two-stream approximations revisited: A new improvement and tests with GCM data. *Quarterly Journal of the Royal Meteorological Society*, *128*, 2397–2416. <https://doi.org/10.1256/qj.01.161>
- Räisänen, P., Barker, H. W., Khairoutdinov, M. F., Li, J., & Randall, D. A. (2004). Stochastic generation of subgrid-scale cloudy columns for large-scale models. *Quarterly Journal of the Royal Meteorological Society*, *130*, 2047–2067.
- Séférian, R., Baek, S., Boucher, O., Dufresne, J.-L., Decharme, B., Saint-Martin, D., et al. (2018). An interactive ocean surface albedo scheme (OSAv1.0): Formulation and evaluation in ARPEGE-Climat (V6.1) and LMDZ (V5A). *Geoscientific Model Development*, *11*, 321–338. <https://doi.org/10.5194/gmd-11-321-2018>
- Shonk, J. K. P., & Hogan, R. J. (2008). Tripleclouds: An efficient method for representing horizontal cloud inhomogeneity in 1D radiation schemes by using three regions at each height. *Journal of Climate*, *21*, 2352–2370.
- Shonk, J. K. P., & Hogan, R. J. (2010). Effect of improving representation of horizontal and vertical cloud structure on the Earth's global radiation budget. Part II: The global effects. *Quarterly Journal of the Royal Meteorological Society*, *136*, 1205–1215. <https://doi.org/10.1002/qj.646>
- Shonk, J. K. P., Hogan, R. J., Edwards, J. M., & Mace, G. G. (2010). Effect of improving representation of horizontal and vertical cloud structure on the Earth's global radiation budget. Part I: Review and parameterization. *Quarterly Journal of the Royal Meteorological Society*, *136*, 1191–1204. <https://doi.org/10.1002/qj.647>
- Søvde, O. A., Prather, M. J., Isaksen, I. S. A., Bernsten, T. K., Stordal, F., Zhu, X., et al. (2012). The chemical transport model Oslo CTM3. *Geoscientific Model Development*, *5*, 1441–1469. <https://doi.org/10.5194/gmd-5-1441-2012>
- Spurr, R., & Natraj, V. (2011). A linearized two-stream radiative transfer code for fast approximation of multiple-scatter fields. *Journal of Quantitative Spectroscopy*, *112*(2011), 2630–2637.
- Szepszó, G. V., & Carver, S. G. (2019). Using the ECMWF OpenIFS model and state-of-the art training techniques in meteorological education. *Advances in Science and Research*, *16*, 39–47.
- Taylor, J. P., Edwards, J. M., Glew, M. D., Hignett, P., & Slingo, A. (1996). Studies with a flexible new radiation code. II: Comparisons with aircraft short-wave observations. *Quarterly Journal of the Royal Meteorological Society*, *122*, 839–861. <https://doi.org/10.1002/qj.49712253204>
- Thornberry, T. D., Rollins, A. W., Avery, M. A., Woods, S., Lawson, R. P., Bui, T. V., et al. (2017). Ice water content-extinction relationships and effective diameter for TTL cirrus derived from in situ measurements during ATTREX 2014. *Journal of geophysical Research: Atmospheres*, *122*, 4494–4507. <https://doi.org/10.1002/2016JD025948>
- Tompkins, A. M., & Di Giuseppe, F. (2007). Generalizing cloud overlap treatment to include solar zenith angle effects on cloud geometry. *Journal of Atmospheric Sciences*, *64*, 2116–2125. <https://doi.org/10.1175/JAS3925.1>
- Tompkins, A. M., & Di Giuseppe, F. (2015). An interpretation of cloud overlap statistics. *Journal of the Atmospheric Sciences*, *72*(8), 2877–2889.
- Wild, O., Zhu, X., & Prather, M. J. (2000). Fast-J: Accurate simulation of in- and below-cloud photolysis in tropospheric chemical models. *Journal of Atmospheric Chemistry*, *37*(3), 245–282.
- Wiscombe, W. J. (1977). Delta-M method - rapid yet accurate radiative flux calculations for strongly asymmetric phase functions. *Journal of the Atmospheric Sciences*, *34*(9), 1408–1422.
- Wiscombe, W. J., & Grams, G. W. (1976). The backscattered fraction in two-stream approximations. *Journal of the Atmospheric Sciences*, *33*, 2440–2451.
- Wood, R. (2012). Stratocumulus clouds. *Monthly Weather Review*, *140*, 2373–2423.
- Yang, P., Bi, L., Baum, B. A., Liou, K.-N., Kattawar, G., Mishchenko, M., et al. (2013). Spectrally consistent scattering, absorption, and polarization properties of atmospheric ice crystals at wavelengths from 0.2 to 100 μm . *Journal of the Atmospheric Sciences*, *70*(1), 330–347. <https://doi.org/10.1175/JAS-D-12-039.1>
- Yang P., Hioki S., Saito M., Kuo C.-P., Baum B., & Liou, K.-N. (2018). A Review of Ice Cloud Optical Property Models for Passive Satellite Remote Sensing. *Atmosphere*, *9*, (12), 499. <http://dx.doi.org/10.3390/atmos9120499>
- Zhao, W., Peng, Y., Wang, B., Yi, B., Lin, Y., et al. (2018). Comparison of three ice cloud optical schemes in climate simulations with community atmospheric model version 5. *Atmospheric Research*, *204*. <https://doi.org/10.1016/j.atmosres.2018.01.004>

Assessing Uncertainties and Approximations in Solar Heating of the Climate System

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Methods

Solar-J is based on Cloud-J (Prather, 2015) and has a long history of modeling and testing photolysis rates) for wavelengths 180-800 nm, which includes photolysis of O₃ and NO₃ in the 400-800 nm range (e.g., Olson et al, 1997; Photo Comp, 2010). For Solar-J, we adjusted the visible wavelength bands and adopted the RRTMG-SW spectral code for the infrared (see H2017). For wavelengths > 778 nm SJ simply takes the RR model. Since v7.5, SJ has shifted a wavelength boundary, 345-412-778 nm to 345-485-778 nm, to better separate the Rayleigh scattering region (345-485 nm) from the Chappuis ozone absorption (485-778 nm). We use the g-point sub-bins as specified for the 9 IR bands and as adapted to our last visible band (485-778 nm). In some test cases using alternative spectral models LLNL and CLIRAD, we have further reduced this latter band to 485-700 nm to match the IR bands of these two models.

The solar heating codes Solar-J and RRTMG-SW are included as modules within the UC Irvine chemistry-transport model (UCI CTM, Prather et al., 2017). The UCI CTM is coupled with meteorological fields from the European Centre Integrated Forecast System, open IFS cycle 38r1 run at T159N80L60 using the native Gaussian grid for atmospheric physics (about 1.1° horizontal with 60 layers). We take the archived 3-hour averages of the atmospheric column data: pressure on the layer edges; temperature, water vapor, cloud fraction, cloud liquid water content, cloud ice water content in each layer. A

standard ozone climatology is used. Cloud effective radii (liquid and ice) and scattering phase function are as specified in the CTM photolysis code Cloud-J (Neu et al., 2007; Prather, 2015). The heating rates reported here are calculated for the IFS fields of January 2015. Cloud fields change every 3 hours, and the solar zenith changes every hour, giving 744 hourly data for January. The OSA code generously provided by Séférian was modified slightly: the diffuse albedo was not used as Solar-J calculates albedos specifically for each scattered stream; the albedo goes to a constant for SZA > 90° in spherical atmospheres; and the parameter table for the white-cap variable 'XRWC' was reset from 0.0 to 0.2 for wavelengths <400 nm.

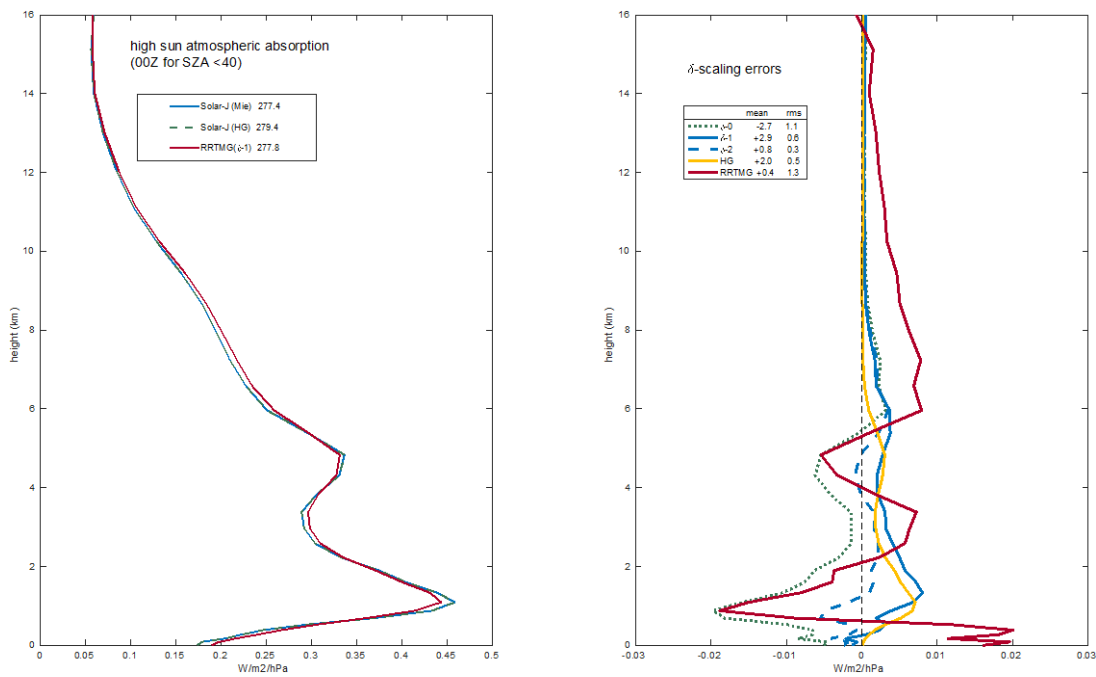


Figure S1. (left) January atmospheric absorption profiles ($\text{W m}^{-2} \text{hPa}^{-1}$) over the Pacific Ocean at high sun ($\text{SZA} > 40^\circ$ at 00Z, area within green oval in Figure 4). Liquid-water, but not ice-water clouds are included. The total heating rates (W m^{-2}) for Solar-J with standard Mie phase function, Solar-J using a Henyey-Greenstein (HG) phase function, and RRTMG (standard δ -1 scaling) are shown in the legend. These values are averaged only over high sun. **(right)** Profiles of the errors in atmospheric absorption caused by δ -scaling (0, 1, 2) and HG phase function for liquid-water clouds. Same conditions as for left panel. Total heating-rate and rms differences (both W m^{-2}) are given in the legend. Also shown is the difference, RRTMG (δ -1) minus Solar-J (Mie).

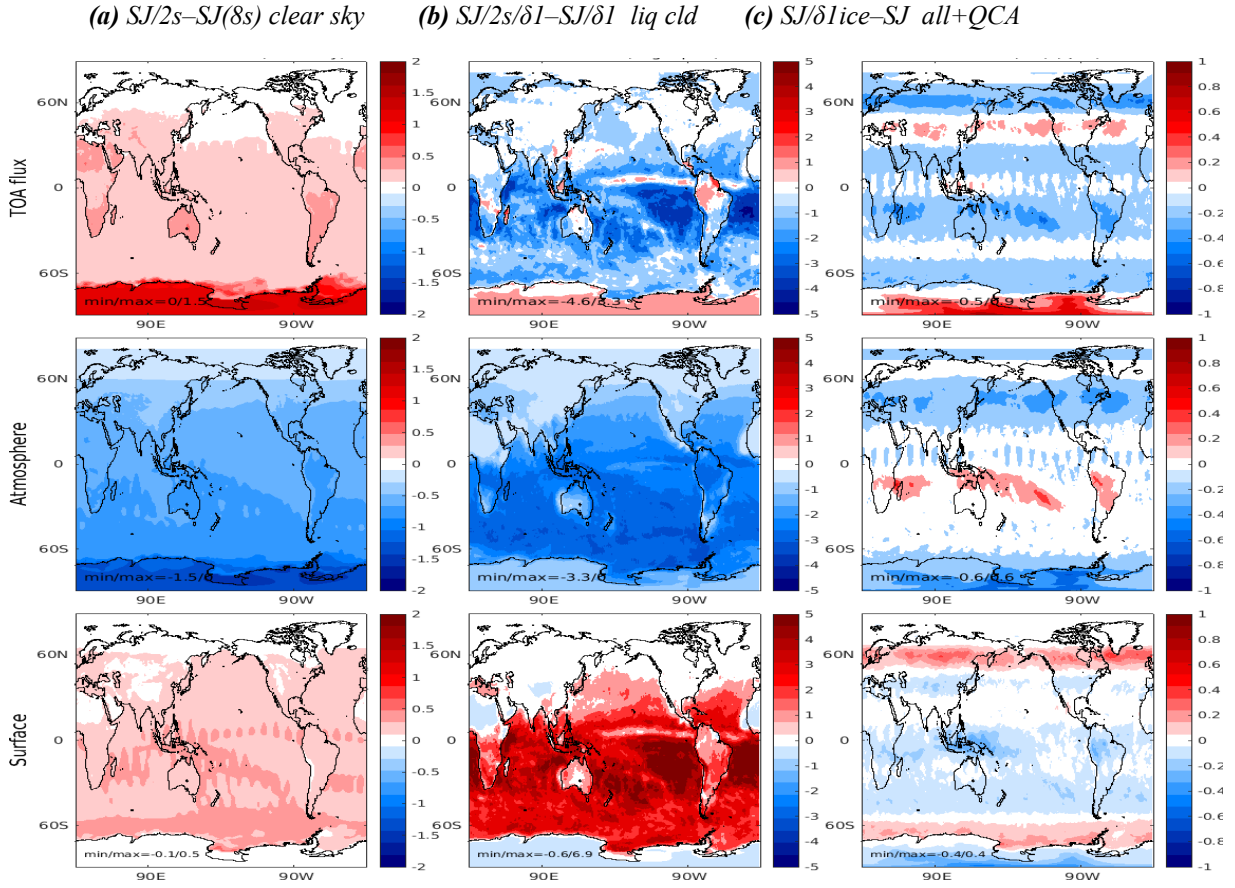


Figure S2. Geographic map of model differences in solar radiative budget (W m^{-2}) averaged over January, columns show (a) Solar-J’s 2-stream minus 8-stream under clear sky, (b) Solar-J’s 2-stream minus 8-stream with averaged liquid-only clouds and $\delta-1$ scaling for both, and (c) Solar-J with $\delta-1$ scaling of ice clouds minus the standard SJ where both calculations use ICAs sorted by cloud quadrature (QCA) and MAX-COR cloud overlap scheme. All calculations in this paper are aerosol-free. The small wave-24 noise seen in many of these panels results from calling the radiation code every hour.

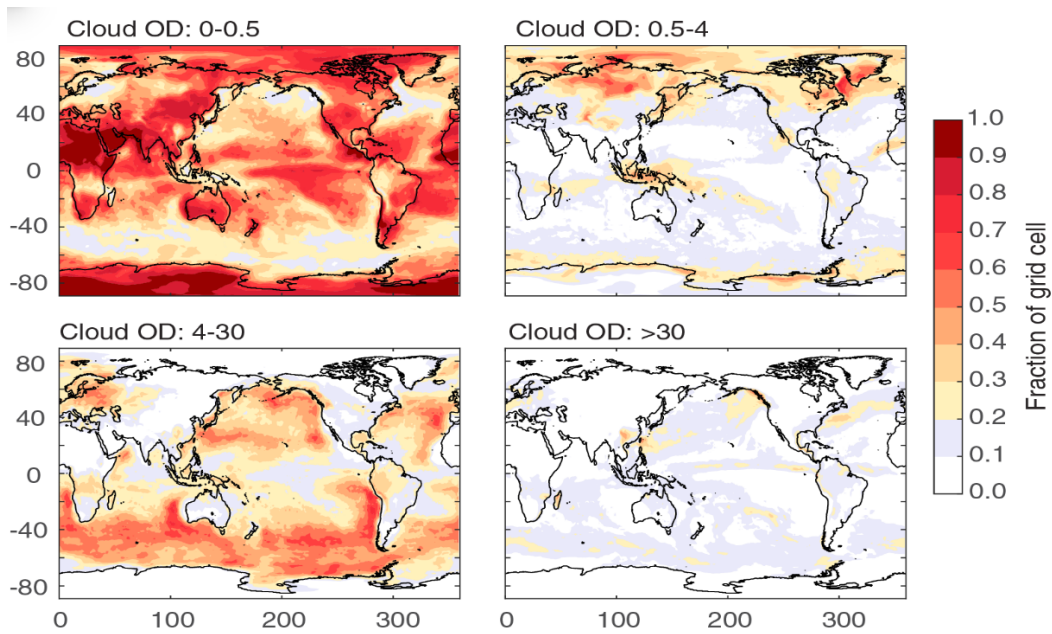


Figure S3. Frequency of occurrence for each of the four QCA bins using MAX-COR overlap for the January 2015 case study here. The QCAs are binned by 600 nm total cloud optical depth (liquid+ ice water: $0 - \frac{1}{2}$; $\frac{1}{2} - 4$; $4 - 30$; >30). For each time step, the fractional area assigned to each QCA is calculated and then averaged over the month to give a frequency that sums to 1.00 over the 4 QCAs.

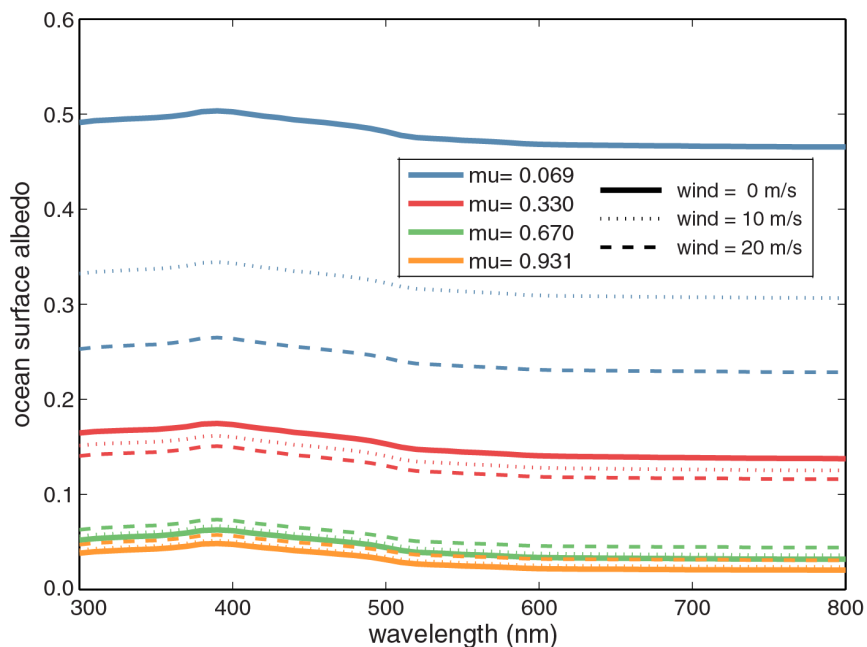


Figure S4. Ocean surface albedo (OSA, dimensionless) as a function of wavelength, shown for different incident angles (colors) and different wind speeds (solid, dashed or dotted). Four different incident zenith angles used in Solar-J's 8-stream scattering code are identified by their cosine values, μ : 0.931 (orange), 0.670 (green), 0.330 (red), 0.069 (blue). Three different wind speeds are shown: 0 m/s (solid), 10 m/s (dotted), 20 m/s (dashed).

Table S1. Versions of Solar-J and RRTMG-SW codes used here.	
notation	code description
SJ	The standard version of Solar-J version 7.6d as published here. It is a minor update of 7.6c published in Prather and Hsu (2019, doi.org/10.7280/D1096P) to make the MAX-RAN consistent with 7.6c changes. SJ uses a standard 8-stream Feautrier RT solver. Solar-J uses Cloud-J data tables for heating by O ₂ (bins 1:11, <291 nm) & O ₃ (bins 1:18, <778 nm) and RRTMG-SW tables for other gas-phase absorption (H ₂ O, CO ₂ , CH ₄ , O ₂) in IR bands 18:27 (83 sub-bin g-points with 5 in visible and 78 in IR). Full cloud treatment includes vertical decorrelation length for cloud overlap (MAX-COR) to generate independent column atmospheres (ICAs) and then 4 cloud quadrature atmospheres (QCA) to average over the ICAs. SJ can also be run with clear-sky or averaged cloud (full cloud cover in each cell as the average of cloudy and clear fractions), which does not invoke cloud-overlap and ICAs. SJ can be run in flat, spherical, refractive, and geometric options. SJ by default uses a constant ocean surface albedo (OSA = 0.06) but can invoke OSA to be of function of wavelength, incident angle (including scattered light) and surface wind.
SJ/RAN	SJ run with MAX-RAN cloud overlap.
SJ/RRX	The 78 IR sub-bins are replaced with the RRTM-SW benchmark code's 144 sub-bins.
SJ/CLIRAD	Solar-J with IR bands replaced with CLIRAD model: 0.70-1.22, 1.22-2.27, and 2.27-10.0 μm , each with 10 absorption sub-bins for each band (Chou and Suarez, 1996). The edge of the IR transition is shifted from 778 to 700 nm and cross section in bin 18 are adjusted. Only water vapor is included in the IR bands.
SJ/LLNL	The IR bands are replaced by the 3 large LLNL bands: 0.69-0.86, 0.86-2.27, and 2.27-3.85 μm , which include a total of 21 sub-bins. The edge of the IR transition is shifted from 778 to 700 nm and cross section in bin 18 are adjusted. Only water vapor is included in the IR bands. (Chou, 1992; Grant & Grossman, 1998).
SJ/hrv	SJ with high-resolution-visible version (<i>SJ/hrv</i>), putting 18 bands in the VIS region.
SJ/2S	SJ with 2-stream RT solver.
SJ/66b	SJ with a very high wavelength resolution in the IR used to resolve ice- and liquid-water cloud absorption. It is constructed using 0.05 to 0.10 μm wide bands, yielding 66 IR bands instead of the 9 in SJ. Without sub-bins, it cannot calculate any IR gas absorption.
SJ/Ray	SJ with the standard Rayleigh gas scattering phase function ($1 + \cos^2(\Theta)$) changed to isotropic (1).
SJ/ δ_0	SJ with all cloud optical depths and phase functions changed to δ_0 scaling, see text and Table 2.
SJ/ δ_1	SJ with all cloud optical depths and phase functions changed to δ_1 scaling, see text and Table 2.
SJ/ δ_1 /2S	SJ/ δ_1 with 2-stream RT solver.
SJ/ δ_1 ice	SJ standard, but with δ_1 scaling of ice clouds only.
SJ/ δ_2	SJ with all cloud optical depths and phase functions changed to δ_2 scaling, see text and Table 2.
SJ/HG	SJ with all cloud phase functions changed to Henyey-Greenstein, see text and Table 2.
SJ/OSA	SJ with OSA a function of wavelength, incident angle (including scattered light) and surface wind.
AER4.0	The standard RRTMG-SW version 4.0 code. If there are fractional clouds, this code uses MAX-RAN cloud overlap and McICA sampling of the ICAs.

Table S2. Global monthly mean SW radiation budget for labeled experiments in Table 1. See Table S1 for the model version.							
name	Table 1 row	code version	conditions.	Incident	Reflected	Atmos.	Surface
B0	1-6, 17	SJ (std)	flat-atmosphere, clear sky	351.37	49.92	70.71	230.74
B1	1	SJ/CLIRAD	flat-atmosphere, clear sky	351.37	50.86	65.03	235.47
B2	2	SJ/LLNL	flat-atmosphere, clear sky	351.37	50.46	63.05	237.87
B3	3	SJ/RRX	flat-atmosphere, clear sky	351.37	49.95	70.47	230.95
B4	5	SJ/hrv	flat-atmosphere, clear sky	351.37	50.25	70.25	230.86
B5	6	SJ/2S	flat-atmosphere, clear sky	351.37	50.23	70.18	230.96
MR	17	SJ/Ray	flat-atmosphere, clear sky isotropic Rayleigh scattering	351.37	49.91	70.72	230.74
C0	7-8	SJ/66b	sphere, QCA/MAX-COR overlap, no IR gases	352.85	127.37	28.96	196.53
C1	7	SJ/CLIRAD	sphere, QCA/MAX-COR overlap, no IR gases	352.86	125.23	32.80	194.83
C2	8	SJ	sphere, QCA/MAX-COR overlap, no IR gases	352.89	126.23	30.64	196.02
Mh	9	SJ/HG	flat-atmosphere, averaged clouds, no ice clouds	351.45	127.26	78.31	145.88
M0	9-14	SJ	flat-atmosphere, averaged clouds, no ice clouds	351.45	127.31	78.17	145.97
M1	10	SJ/δ0	flat-atmosphere, averaged clouds, no ice clouds	351.41	127.04	77.95	146.41
M2	11,15	SJ/δ1	flat-atmosphere, averaged clouds, no ice clouds	351.42	126.98	78.40	146.04
M3	12	SJ/δ2	flat-atmosphere, averaged clouds, no ice clouds	351.42	126.99	78.22	146.20
M4	13,16	SJ/δ1/2S	flat-atmosphere, averaged clouds, no ice clouds	351.42	126.30	76.66	148.46
M5	20	SJ	flat-atmosphere, averaged clouds (all)	351.46	131.19	78.49	141.77
MI	18	SJ/δ1ice	flat-atmosphere, QCA/MAX- COR overlap	351.43	111.57	76.88	162.98
D0	18,19	SJ	flat-atmosphere, QCA/MAX- COR cloud overlap	351.43	111.59	76.87	162.97
D1	19,21	SJ/RAN	flat-atmosphere, QCA/MAX- RAN overlap	351.43	110.24	76.81	164.38
O0	22	SJ/OSA	sphere, varied OSA	352.98	111.72	76.77	164.49
O1	22	SJ	sphere, fixed OSA at 0.06	352.98	112.40	76.93	163.65
RR0	4	AER4.0	clear sky	351.38	51.22	70.24	229.92
RR1	20	AER4.0	averaged clouds	351.38	130.82	75.82	144.74
RR2	14-16	AER4.0	averaged clouds, liquid only	351.38	126.17	77.35	147.86
RR3	21	AER4.0	McICA	351.38	108.14	74.45	168.79

Table S3. Cross sections for ozone absorption and Rayleigh scattering in the 3 visible bands with different weighting functions				
	weighting	320-345 nm	345-485 nm	485-778 nm
$\sigma_{\text{Rayleigh}} \text{ (cm}^2\text{)}$	Photons	3.644e-26	1.387e-26	3.136e-27
	Watts	3.645e-26 (+0%)	1.436e-26 (+4%)	3.367e-27 (+7%)
τ_{Rayleigh}		0.73	0.29	0.07
$\sigma_{\text{O}_3} \text{ (cm}^2\text{, 298 K)}$	Photons	6.520e-21	2.125e-22	2.325e-21
	Watts	6.522e-21 (+0%)	1.984e-22	2.427e-21 (+4%)
$\tau_{\text{O}_3} \text{ (troposphere)}$		0.005	0.0002	0.002