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1 **Modeled connectivity between Walleye Pollock (*Gadus chalcogrammus*) spawning**
2 **and age-0 nursery areas in warm and cold years with implications for juvenile**
3 **survival**

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23

24 ABSTRACT

25 Adult and early life stage distributions of the commercially important demersal
26 fish Walleye Pollock (*Gadus chalcogrammus*) have varied in relation to the warm and
27 cold environmental conditions on the eastern Bering Sea shelf. Previous modeling studies
28 indicate that transport alone does not account for the disparate juvenile distributions in
29 warm and cold years, but that spawning locations are important. Our objective was to
30 determine the connectivity of eastern Bering Sea pollock spawning areas with juvenile
31 nursery areas between warm and cold years. We calculated the connectivity between
32 sources and juvenile positions that were produced by a coupled biological-physical
33 individual-based model that simulated transport, growth, and vertical behavior of pollock
34 from the egg until the juvenile stage. Three connectivity patterns were seen in most
35 simulations: along-isobaths to the northwest, self-retention, and transport around the
36 Pribilof Islands. The major differences in connectivity between warm and cold years,
37 more northwards in warm years and more off-shelf in cold years, mimicked wind-driven
38 flow characteristics of those years that were related to the winter mean zonal position of
39 the Aleutian Low. Connectivity relationships were more sensitive to spatial alterations in
40 the spawning areas in cold years, while they were more responsive to spawn timing shifts
41 in warm years. Juvenile pollock prey and predator studies favor the Pribilof Islands and
42 Outer Shelf regions as advantageous juvenile habitats. The strongest connectivity to these
43 regions originated in the well-known spawning areas around Unimak Island and the
44 Pribilof Islands, and the less well-studied North Outer Shelf. The North Outer Shelf
45 emerged as a very large sink of pollock reaching the juvenile transition from all spawning

46 sources, suggesting more thorough sampling across multiple trophic levels of this
47 potentially important juvenile pollock nursery is needed.

48

49 KEY WORDS: Bering Sea, connectivity, fish early life history, pollock

50

51 INTRODUCTION

52 One of the most commercially important semi-demersal fishes in the eastern
53 Bering Sea (EBS) is Walleye Pollock (*Gadus chalcogrammus*), hereafter referred to as
54 “pollock.” Over the EBS shelf, pollock spawn consistently at at least two spawning sites,
55 one north of Unimak Island (Figure 1, region 1), and the other near the Pribilof Islands
56 (Figure 1, region 5), beginning near Unimak and spreading towards the Pribilof Islands
57 (Bacheler *et al.*, 2010). Eggs can be found in the water column from December to
58 August, but the annual peak occurs in either April or May (Bacheler *et al.*, 2010). Eggs
59 are predominantly found within the mixed layer (Smart *et al.*, 2013) and incubation time
60 depends on temperature (Blood, 2002). After hatching, larvae develop through a series of
61 stages known as the yolk sac, preflexion, and postflexion (or late), classified by
62 anatomical differences (Matarese *et al.*, 1989). Pollock transition from larvae to pelagic
63 juveniles between 30 and 40 mm (Matarese *et al.*, 1989), become more bottom-oriented
64 with age (Brodeur & Wilson 1996; Duffy-Anderson *et al.*, 2015), and finally recruit to
65 the fishery at age-3 or age-4 (Ianelli *et al.*, 2012b).

66 The EBS is a very productive region, supporting large marine bird and mammal
67 populations and several commercial fisheries. The EBS shelf extends approximately 500
68 km southwestward from the Alaskan coast to the shelfbreak and 1000 km northwestward

69 from the Alaska Peninsula to the Siberian coast (Figure 1). The shelf can be divided into
70 three regions based on bathymetry: Inner Shelf (<50 m), Middle Shelf (50-100 m), and
71 Outer Shelf (>100 m) (Coachman 1986). The EBS has been further divided into regions
72 for reference and comparisons across projects under the Bering Ecosystem Study –
73 Bering Sea Integrated Ecosystem Research Program (BEST-BSIERP). These 16 regions
74 (Figure 1) cover the EBS shelf and slope within the U.S. Exclusive Economic Zone and
75 were delineated by bathymetry, oceanography, animal distributions, ecological domains,
76 and established survey areas (Ortiz *et al.*, 2012). In numerical order, these regions are: 1
77 – AK Peninsula, 2 – South Inner Shelf, 3 – South Middle Shelf, 4 – South Outer Shelf, 5
78 – Pribilofs, 6 – Midnorth Middle Shelf, 7 – Midnorth Inner Shelf, 8 – North Outer Shelf,
79 9 – St. Matthews, 10 – North Middle Shelf, 11 – North Inner Shelf, 12 – St. Lawrence, 13
80 – South Bering Strait, 14 – Norton Sound, 15 – Off-shelf North, 16 – Off-shelf Southeast
81 (Ortiz *et al.*, 2012). Baker and Hollowed (2014) found that distinct biological
82 communities in the EBS are best delineated by depth, bottom temperature, frontal
83 boundaries, and position north or south of 60°N; these ecoregions are well-represented by
84 the BEST-BSIERP regions.

85 Recently the EBS experienced alternating periods of temperatures that were either
86 warmer (2001-2005) or colder (2007-2012) than average (Stabeno *et al.*, 2012). In cold
87 years, winter ice extends farther offshore creating a larger cold pool that could influence
88 the spawning and movements of demersal fishes (Mueter & Litzow, 2008). The water
89 circulation is also influenced by atmospheric and hydrographic conditions, which could
90 affect the dispersal of pelagic early life stages. For example, the predominantly
91 northwesterly winds in cold years that extend sea ice and the cold pool also promote off-

92 shelf Ekman transport (Danielson *et al.*, 2011b; Stabeno *et al.*, 2012; Danielson *et al.*,
93 2012). On the other hand, the more southeasterly winds in warm years resulted in
94 northward or weak on-shelf flows over the Middle Shelf (Stabeno *et al.*, 2012; Danielson
95 *et al.*, 2012). Moreover, the intensity and timing of setup of the Inner and Middle Fronts,
96 which separate the three shelves, varies with temperature, also potentially affecting the
97 transport of larval fishes on the shelf. Concomitant with the atmospheric and
98 oceanographic differences, the distributions of early life stages diverged in cold and
99 warm years (Smart *et al.*, 2012), likely due to thermally-mediated variability in spawning
100 locations of adults (Petrik *et al.*, 2014).

101 Early life stage transport contributes greatly to population connectivity and
102 persistence. Studies of connectivity between spawning regions and nursery habitats
103 advance our understanding of genetic exchange, processes during the early life period,
104 and recruitment (reviewed in Cowen & Sponaugle, 2009). Alternate transport pathways
105 between spawning grounds and nursery areas can influence early life survival because of
106 spatial differences in growth and mortality (Cowen & Sponaugle, 2009). Survival to
107 recruitment of individuals is further influenced by the distribution of individuals at the
108 end of the larval period in regard to spatial differences in prey and predators. For
109 example, the overlap between juvenile and cannibalistic adult pollock explains 30-50% of
110 interannual recruitment variability (Mueter *et al.*, 2011). Additionally, match-mismatch
111 between the spatial distribution of juvenile pollock and prey energy density may explain
112 the low and high year classes of 2005 and 2010, respectively (Siddon *et al.*, 2013b). Thus
113 it is important to identify the source regions and oceanographic conditions influencing the
114 cold and warm year patterns of juvenile pollock distribution and survival.

115 Modeling results suggested that spawning locations and their proximity to
116 different physical oceanographic features drove differences in the distributions of pollock
117 early life stages (Petrik *et al.*, 2014). Following these results, it was our primary
118 objective to determine the connectivity between spawning and nursery regions of walleye
119 pollock in the EBS (1995-2012). Our second objective was to resolve how connectivity
120 diverged in warm and cold years. This examination provides a projection of the potential
121 effects of predicted future warming (Wang *et al.*, 2012). To accomplish these objectives
122 we used an individual-based model (IBM) of pollock biology and behavior coupled to a
123 hydrodynamics model of the EBS to simulate transport during the period from spawning
124 to the juvenile transition. Model output was compared to observed distributions of
125 juvenile pollock (1996-2012) to groundtruth model results.

126

127 METHODS

128 **Coupled individual-based model**

129 A detailed description of the coupled physical-biological IBM can be found in
130 Petrik *et al.* (2014). To summarize, we used the Regional Ocean Modeling System
131 (ROMS; Shchepetkin & McWilliams 2009) developed for the Northeast Pacific (NEP6)
132 to represent the hydrodynamics of the EBS. ROMS is a free-surface, hydrostatic
133 primitive equation ocean circulation model. It is a terrain-following, finite volume
134 (Arakawa C-grid) model. The ROMS NEP6 model domain extends over a broader region
135 than the EBS Shelf, from 20°N to 71°N and reaching approximately 2250 km offshore
136 from the North American west coast. The spatial resolution is a 10 km horizontal grid in
137 a Lambert Conical projection that is rotated relative to lines of constant longitude and has

138 50 terrain-following depth levels stretched towards the surface boundary in each grid cell.
139 ROMS includes a fully parallel coupled sea-ice model (Budgell, 2005). An older version
140 of ROMS for the Northeast Pacific (NEP5) has been thoroughly validated against
141 observations and had significant skill reproducing aspects of the Bering Sea physical
142 oceanography, such as the horizontal and vertical structure of tidal currents and the
143 frequency of kinetic energy (Curchitser *et al.*, 2010; Danielson *et al.*, 2011a). The NEP6
144 has further improved upon this model through more comprehensive evaluations with the
145 newly available data collected under the BEST-BSIERP program, which had much more
146 complete spatial and temporal coverage than earlier data sets. The NEP6 coupled ocean-
147 sea ice model was integrated in hindcast mode for the period from 1994-2012 deriving
148 surface forcing from the Modern Era Retrospective-Analysis for Research and
149 Applications (MERRA; Rienecker *et al.*, 2011), air-sea fluxes computed using bulk
150 formulae (Large & Yeager, 2009), the Dai and Trenberth (2002) method as a surface
151 fresh water flux for the riverine inputs, initial and boundary conditions for this domain
152 from the Simple Ocean Data Assimilation (SODA) ocean reanalysis (Carton & Giese
153 2008) for early years, and boundary conditions from the global HYCOM assimilative
154 product (HYCOM Ocean Prediction website) for the later ones. Daily averages of
155 velocity, temperature, and mixed layer depth were saved and used as offline inputs to the
156 particle-tracking model.

157 We utilized the particle-tracking tool TRACMASS to simulate transport of
158 pollock early life stages. TRACMASS computes Lagrangian trajectories from Eulerian
159 velocity fields from general circulation model simulations through an offline coupling.
160 TRACMASS interpolates any general circulation model three-dimensional grid to its own

161 grid and solves the trajectory path through each grid cell with an analytical solution of a
162 differential equation, which depends on the velocities at the grid cell walls (Döös 1995,
163 Blanke & Raynaud, 1997; de Vries & Döös, 2001). The TRACMASS turbulence
164 subroutine was implemented to incorporate sub-grid scale motion. Physiological and
165 behavioral information was added to TRACMASS to create an IBM of EBS pollock
166 (Petrik *et al.*, 2014). Individuals transitioned from eggs to yolk sac larvae, preflexion
167 larvae, and late larvae, ultimately reaching the beginning of the juvenile stage. Stages
168 were defined by length (yolk sac: <6 mm SL; preflexion: 6-10 mm SL; late: 10-40 mm
169 SL) that was attained via stage-specific temperature-dependent growth (Petrik *et al.*,
170 2014). The nonfeeding eggs and yolk sac larvae were modeled as neutrally buoyant, while
171 the feeding stages, preflexion and late larvae, directed their vertical swimming towards
172 the middle of the mixed layer (Petrik *et al.*, 2014).

173

174 **Model initialization and simulations**

175 The model simulated spawning by initializing individuals at the egg stage.
176 Spawning locations and times were the same as those used in Petrik *et al.* (2014). The
177 locations concentrated on the major spawning regions identified by Hinckley (1987) and
178 Bachelier *et al.* (2010), with individual spawning polygons created from aggregated
179 locations of mature spawning females from observer data (Figure 1). The Bogoslof Island
180 region, defined by bottom depth > 250 m, was disregarded as a spawning ground as this
181 likely represents a distinct population (Ianelli *et al.*, 2012a). Following the literature and
182 the observer data, spawning started in mid-Jan near Unimak Island, then around the
183 Pribilof Islands in Feb. Spawning expanded to the northwest as the season progressed,

184 peaking in mid-Mar, and then contracting into Apr. The result was seven different
185 spawning times (15 Jan, 1 Feb, 15 Feb, 1 Mar, 15 Mar, 1 Apr, 15 Apr) with spawning
186 areas of varying sizes. Eggs were released at these times at the center of each ROMS
187 NEP6 grid cell within the spawning polygons. Ten eggs were released per 10 m depth
188 increment in each spawning grid cell in all cases with the exception of the “Contracted”
189 simulation (see description below). Because spawning occurred in fewer grid cells
190 compared to the other cases, 15 eggs per 10 m depth were spawned to achieve the total
191 number of particles needed for stable results.

192 Five distinct simulations were run (Petrik *et al.*, 2014) to test the effects of
193 advection, spawning location, and spawn timing on connectivity. In the “Transport Only”
194 case, all years 1995-2012 shared the same initial spawning areas and times and differed
195 only by their physical environments simulated with ROMS NEP6. This was considered
196 the “Base” simulation and was divided into “Base Cold” (1995, 1997, 1999, 2000, 2006-
197 2012) and “Base Warm” (1996, 2002, 2003, 2005) years using the temperature anomaly
198 criteria of Smart *et al.* (2012). To simulate hypothesized changes in spawning locations in
199 reference to the cold pool extent, spawning polygons were contracted off-shelf to the
200 southwest in cold years (“Contracted”) and expanded on-shelf to the northeast in warm
201 years (“Expanded”). Finally, to represent potential shifts in peak spawn timing (Smart *et*
202 *al.*, 2012), spawning times were delayed 40 d in cold years (“Late”) and advanced 40 d in
203 warm years (“Early”). In each simulation, individuals were followed from spawn until
204 they reached 40 mm, the size at the juvenile transition (Matarese *et al.*, 1989).

205

206 **Observational data**

207 *T/S Oshoro Maru*

208 A time series of age-0 pollock juvenile abundances from mid-summer collections
209 (1996-2005, 2007) was available from cooperative cruises conducted by the Alaska
210 Fisheries Science Center's (AFSC) Eco-FOCI program and the Graduate School of
211 Fisheries, Hokkaido University, Japan aboard the *T/S Oshoro Maru* (Busby *et al.*, 2014).
212 Sampling occurred at a grid of stations over the Outer and Middle Shelves, though not all
213 stations were sampled each year. Cumulative over the time series, sampled stations
214 resided within a minimum of two BEST-BSIERP regions (regions 3, 4) and a maximum
215 of seven BEST-BSIERP regions (1-7). A modified beam trawl with 5-m² mouth opening
216 fitted with a 3x2 mm oval mesh net and a 1-mm mesh cod end was towed obliquely
217 through the water column from 200 m depth (or 10 m off-bottom, whichever was
218 shallowest) to the surface. Use of a flowmeter permitted quantitative estimates of catch.
219 Samples were preserved in formalin, taxa were sorted, identified, and enumerated, and
220 pollock catch was expressed as catch 10 m⁻² sea surface area.

221

222 *BASIS surface trawls*

223 Surface trawl surveys (2008-2013) were conducted as part of the Bering Arctic-
224 Subarctic Integrated Survey (BASIS) aboard the NOAA Ship *Oscar Dyson*. The surveys
225 were typically conducted mid-Aug through late Sep, and occupied a systematic grid of
226 stations over the southeast Bering Sea shelf. Near-surface fish, inclusive of age-0 walleye
227 pollock, were collected using a 198-m long midwater rope trawl (CanTrawl) composed of
228 hexagonal mesh wings and fitted with a 1.2-cm mesh codend liner (Farley *et al.*, 2015).
229 The CanTrawl had a mouth opening of 55-m horizontal x 20-m vertical, with typical

230 sampling occurring between 0 m and 20 m depth. Near-surface orientation was
231 maintained by buoys attached to the headrope. The net was towed at speeds ranging from
232 3.5-5.0 knots for 30 min during daylight hours. Gridded stations typically spanned 14-16
233 of the BEST-BSIERP regions, but only six regions were covered in 2008.

234

235 *BASIS acoustic surveys*

236 Acoustic backscatter measurements were collected in conjunction with the BASIS
237 surface trawl surveys (above) to determine subsurface age-0 pollock biomass (2008-
238 2013). Measurements were collected aboard the NOAA Ship *Oscar Dyson* with Simrad
239 EK60 echosounders from five split-beam transducers (18, 38, 70, 120, and 200 kHz)
240 mounted on the bottom of the vessel's retractable centerboard. Data presented herein are
241 the results from 38 kHz-collected data. Verification of midwater and near-bottom
242 acoustic sign was accomplished through targeted trawl sampling conducted
243 opportunistically during daylight hours. Age-0 pollock abundance was estimated by
244 combining echo integration data with species and size-composition information derived
245 from targeted trawl sampling. Post-processing was conducted in accordance with other
246 AFSC surveys (Honkalehto *et al.*, 2011). Comprehensive descriptions of acoustic
247 methodologies are presented in Parker-Stetter *et al.* (2013) and De Robertis *et al.* (2014).
248

249 **Analyses**

250 Spawning release locations and positions of individuals upon reaching the
251 juvenile transition were both mapped to the BEST-BSIERP regions (Figure 1).
252 Connectivity from spawning region "A" to juvenile region "B" for a given time was

253 defined as the number of individuals released in region “A” that were in region “B” at the
254 time they reached juvenile size divided by the total individuals spawned in “A” at that
255 time. As connectivity is defined as a fraction of the total number of eggs spawned in that
256 region, it is independent of the number of eggs spawned. Connectivity for each spawning
257 region-juvenile region pair was calculated by spawning time, year, and grouped cold or
258 warm years. Connectivity is presented with connectivity matrices, grids with source
259 regions on one axis and sink regions on the other axis that indicate the strength of the
260 connection from source to sink as the value at their intersection. Retention was defined as
261 the special case of connectivity from a region to itself.

262 Connectivity was analyzed a number of ways. The local index of collocation
263 (LIC), a method for making cell-by-cell comparisons (Kotwicki & Lauth, 2013), was
264 used to assess similarities and differences between connectivity matrices of spawning
265 time, year, and simulation. Also, connectivity of each pair was classified as strong
266 (>0.10), moderate ($0.05-0.10$), or weak (<0.05) to facilitate comparisons. Additionally,
267 cross-shelf transport was evaluated by calculating the longitudinal difference between
268 spawn and juvenile locations and binning results as either $>10^{\circ}\text{W}$, $5-10^{\circ}\text{W}$, $<5^{\circ}\text{W}$, $<5^{\circ}\text{E}$,
269 $5-10^{\circ}\text{E}$, or $>10^{\circ}\text{E}$.

270 Eastward transport was correlated to different climate indices including the mean
271 winter cross-shelf wind velocity, the mean winter along-shelf wind velocity, the May sea
272 surface temperature (SST) anomaly, and the North Pacific Index in the winter (NPIw).
273 The mean winter cross- and along-shelf wind velocity anomalies represent Oct-Apr at the
274 NARR (NARR website) grid point of 60°N , 169.94°W (BEST-BSIERP region 9). Wind
275 velocity anomalies were NE+/SW- for cross-shelf velocities and NW+/SE- for along-

276 shelf velocities (Danielson unpub. data). The May SST index was calculated as mean
277 monthly SST averaged over the area 54.3°N-60.0°N, 161.2°W-172.5°W using the
278 NCEP/NCAR Reanalysis project data (Kalnay *et al.*, 1996). The index values are
279 deviations from the mean value normalized by the standard deviation for the period
280 1961-2000 (Kalnay *et al.*, 1996). The North Pacific Index is the area-weighted sea level
281 pressure over the region 30°N-65°N, 160°E-140°W that measures the intensity of the
282 Aleutian Low (AL) (Trenberth & Hurrell, 1994). A positive NPI corresponds to a weak
283 AL, while a negative NPI indicates a strong AL, although it does not differentiate
284 between variations in AL zonal displacements. The winter index is the average NPI from
285 Nov-Mar and the anomalies were normalized by the mean and standard deviation for
286 1961-2000 (Trenberth & Hurrell, 1994). The NPI data were provided by the Climate
287 Analysis Section, NCAR, Boulder, USA.

288 Comparisons between modeled distributions of individuals at the time of reaching
289 the juvenile transition and observations of age-0 pollock were also made with the local
290 index of collocation (LIC) as this method of cell-by-cell comparisons has been used to
291 compare distributions of populations (Kotwicki & Lauth, 2013; Petrik *et al.*, 2014) and is
292 similar to the Overlap Coefficient, also used in fisheries studies of distribution overlap
293 (e.g. Hinrichsen *et al.*, 2005). Observations of depth-integrated abundance (# surface
294 area⁻¹) were first binned by BEST-BSIERP region to calculate a mean abundance in each
295 region of each year. Mean depth-integrated abundances (# surface area⁻¹) were then
296 converted into mean total areal abundances (#) by multiplying the areal extent of each
297 region (surface area). Mean numbers of individuals were used to calculate the fraction of
298 individuals observed in each region out of the total number of individuals in all regions

299 observed that year. Since not all 16 BEST-BSIERP regions were sampled by all
300 observations in all years, model-observation comparisons were only made using the
301 regions sampled. Thus, modeled fractions of individuals were calculated as the number of
302 individuals that reached juvenile size in each region sampled by that observation data set
303 that year out of the total number of individuals that reached juvenile size in all regions
304 sampled by that observation data set that year. The LIC was then calculated year-by-year
305 for each observation data set by comparing each specific pair of observed and modeled
306 fractions.

307

308 RESULTS

309 Modeled spawning only occurred in the BEST-BSIERP regions (Figure 1) AK
310 Peninsula (1), South Middle Shelf (3), South Outer Shelf (4), Pribilofs (5), Midnorth
311 Middle Shelf (6), North Outer Shelf (8), and Off-shelf Southeast (16), yet individuals
312 were found in all 16 regions at the time each reached the juvenile transition. Connectivity
313 between spawning and juvenile regions differed by spawn date, year, and simulation.

314

315 **Base simulation**

316 The greatest mean connectivity over all spawn times and years from the Base
317 (“Transport Only”) simulation was retention within the North Outer Shelf (8) with 0.374
318 of all individuals spawned residing in the area as both eggs and juveniles (Table 1). The
319 South Middle Shelf (3) and Midnorth Middle Shelf (6) were other regions with strong
320 retention >0.10 (Table 1). Retention within the AK Peninsula (1) was moderate (0.05–
321 0.10), while it was weak (<0.05) in the other spawning areas (4, 5, 16) (Table 1).

322 Consistently high connections existed from the AK Peninsula (1), historically described
323 as a key spawning area, to the southern Middle Shelf regions and both Outer Shelf
324 regions (3, 4, 6, 8) (Table 1). Robust connections originating from the other historic
325 spawning ground, the Pribilofs (5), were to more northern Middle, Outer, and Off-shelf
326 regions (6, 8, 9, 15) (Table 1). Spawning in the South Middle Shelf (3) and Midnorth
327 Middle Shelf (6) bridged between the historic areas around Unimak (region 1) and the
328 Pribilof Islands (region 5) and shared their connectivity patterns, with the South Middle
329 Shelf (3) more like the AK Peninsula (1) and the Midnorth Middle Shelf (6) more like the
330 Pribilofs (5) (Table 1). The Outer and Off-shelf spawning regions (4, 8, 15) all connected
331 strongly to the North Outer Shelf (8) and Off-shelf North (15) (Table 1). Connectivity
332 between all spawning areas and the Inner Shelf regions (2, 7) and northernmost regions
333 (11-14) was very low (Table 1).

334

335 **Seasonal connectivity patterns**

336 The 15 Jan release date represented early in the season when spawning starts near
337 Unimak Island, only occurring in the AK Peninsula (1), South Middle Shelf (3), and
338 South Outer Shelf (4) (Figures 1, 2). As spawning spread to the northwest with time,
339 there was additional connectivity originating in the Pribilofs (5), Midnorth Middle Shelf
340 (6), North Outer Shelf (8), and Off-shelf Southeast (16) regions (Figures 1, 2). Seasonal
341 connectivity in the Base simulations generally mirrored the mean connectivity with a few
342 variations by spawn date. Retention within the North Outer Shelf (8) was the greatest
343 connectivity for all spawning times except Jan 1 (Figure 2). Connectivity to the southern
344 Middle Shelf (1, 3) was highest from 15 Jan – 1 Mar spawning dates, while connectivity

345 to Off-shelf North (15) was higher from the 1 Mar – 15 Apr spawning dates (Figure 2).
346 Connectivity to the Inner Shelf (2, 7) and northernmost regions (11-14) decreased from
347 all spawning regions as time progressed (Figure 2). The AK Peninsula (1) and South
348 Middle Shelf (3) tended to have the highest retention in the middle of the spawning
349 period, while retention in the Pribilofs (5), Midnorth Middle Shelf (6), North Outer Shelf
350 (8), and Off-shelf Southeast (16) regions increased over the spawning season (Figure 2).
351 Similarity between connectivity patterns from one spawning time to another decreased as
352 the time between them increased, with the exception of 15 Jan, which had connectivity
353 patterns least like those of 15 Mar spawning and most like those of 15 Apr (Table 2). Due
354 to its isolated spawning (Figure 1), the 15 Jan spawning simulation was the least like the
355 other connectivity patterns (Table 2). Conversely, the most widespread spawning in 15
356 Feb had the greatest similarity in connectivity to all other spawning times (Table 2).

357

358 **Interannual connectivity patterns**

359 Connectivity patterns varied from year to year, but were generally more similar to
360 each other than random, with mean annual LIC scores ≥ 0.55 (Table 3). Between-year
361 variability in the connectivity patterns ($\sigma^2=0.0094$) was greater than the variability
362 between spawning times ($\sigma^2=0.0064$). Seasonal variability in connectivity was largely
363 driven by the small spawning area in Jan that had a mean LIC of 0.58, while the other
364 months had mean scores ≥ 0.83 (Table 2).

365 The year with connectivity patterns most similar to all others was 1999 (cold),
366 with a mean of 0.82, though no individual comparisons were exceptionally high (Table 3)
367 and may be the result of more cold years simulated than warm ones. On the other hand,

368 the years 1997 (an El Niño with a warm summer) and 2002 and 2004 (both warm) were
369 very similar to 5 or 6 individual years (Table 3), typically to each other and warm years
370 but also to cold years 2000, 2009, and 2011 (Table 3). Generally, connectivity patterns of
371 individual warm or cold years did not have exclusively high LIC scores with either year
372 type (Table 3). The most striking result was the years with the lowest similarity in
373 connectivity to the other years: 2005 and 2007 (Table 3). Both of these years tended to
374 have lower connectivity from all spawning areas to all juveniles areas compared to the
375 mean Base simulation. For 11 of the 17 years, connectivity in 2007 was the most
376 different (Table 3). Connectivity in 2007 was much like that of 2008 (LIC=0.90), but
377 2005 was dramatically different from 2007 (LIC=0.29; Table 3). There were much higher
378 connections to the North Outer (8) and Off-Shelf North (15) regions from all spawning
379 areas in 2007 than 2005. In contrast, 2005 had greater connectivity from Middle Shelf
380 regions to regions just north of them (1, 3 to 3; 5, 6 to 9, 10). Retention in the North
381 Outer Shelf (8) was consistently above 16% in all years, which had the greatest retention
382 of all times and regions in 1997, whereas virtually no retention occurred Off-shelf
383 Southeast (16) that year (Table 4). The South Outer Shelf (4), Pribilofs (5), Midnorth
384 Middle Shelf (6), and Off-shelf Southeast (16) shared the years with maximum rates
385 (1998), but minimum retention rates varied (Table 4). The high retention year of 1998
386 produced above average rates in the other regions as well (Table 4).

387

388 **Connectivity patterns in different simulations**

389 Combining all spawning times and years within each simulation likewise
390 produced different connectivity patterns. Comparing between only the Cold and Warm

391 years of the Base simulation highlights differences based on temperature regime, and
392 better elucidates connections between spawning areas and different regions of the EBS
393 (Figure 4). In these comparisons, connectivity was classified as strong (>0.10), moderate
394 ($0.05\text{--}0.10$), or weak (<0.05 ; not pictured in Figure 3). Strong connectivity existed from
395 the AK Peninsula (1) to the South Middle Shelf (3) and North Outer Shelf (8) in both
396 Cold and Warm years (Figure 3A). Connections from most spawning areas (1, 3, 5, 6, 8)
397 to the northern Middle Shelf regions (6, 9, 10) and the southern Outer and Off-shelf
398 regions (4, 16) were stronger in Warm years, while to Off-Shelf North (15) was greater in
399 Cold years (Figures 3A-B,D-F). Connectivity patterns from both the South Outer Shelf
400 (4) and Off-Shelf Southeast (16) were identical in both temperature conditions: strong to
401 northern Outer and Off-shelf regions (8,15) and moderate to the Midnorth Middle Shelf
402 (6) (Figures 3C,G). With the exception of the North Outer Shelf (8) spawning region,
403 most moderate and strong connections were to regions north and/or west of spawning
404 regions, with higher off-shelf connectivity in Cold years and greater northward
405 connectivity in Warm years (Figure 3). Retention of juveniles (not visible in Figure 3)
406 decreased in all regions in Cold years, and increased in the South Middle Shelf (3),
407 Pribilofs (5), and North Outer Shelf (8) in Warm years (Table 4).

408 When only comparing the cold year simulations (Figure 4 top), the largest
409 differences existed for the Contract simulation because there was no spawning in the
410 Midnorth Middle Shelf (6) (Figure 4 top). Connectivity from southern Middle Shelf
411 regions (1, 3) to northern Middle, Outer, and Off-shelf regions (6, 8, 9, 15) increased and
412 lessened to southern regions (1, 2, 3, 16) when spawning was contracted in cold years
413 (Figure 4 top). Heightened connectivity was seen between the Pribilofs (5) and southern

414 Middle, Outer, and Off-shelf regions (1, 3, 4, 16), while links to the Inner Shelf regions
415 (2, 7) and north to St. Matthews (9) dropped (Figure 4 top). Contraction of spawning
416 areas in cold years reduced retention in the North Outer Shelf (8), but increased it in the
417 Pribilofs (5) and Off-shelf Southeast (16) (Table 4). Contraction of the spawning area had
418 the largest effect on connectivity from southern Middle Shelf spawning regions (1, 3),
419 and to the South Middle Shelf (3), North Outer Shelf (8), and St. Matthews (9) juvenile
420 regions.

421 Shifting spawning 40 d later in the Late simulation also produced differences in
422 connectivity from the Cold Base simulation. Connections to mid-north Middle Shelf
423 regions (5, 6) and southern Outer and Off-shelf regions (4, 16) were greater than or equal
424 to those of the Cold simulation (Figure 4 top). Late spawning weakened connectivity
425 from most spawning regions to the Inner Shelf (2, 7), and to the northern regions (8-11)
426 (Figure 4 top). Late spawning reduced retention in the South Middle Shelf (3) compared
427 to Base Cold, but was positive for all other regions (Table 4).

428 The warm year simulations also varied among each other, often as either an
429 increase or decrease to a juvenile region from all source regions. Expanding spawning
430 regions increased delivery to the southern inner and Middle regions (1, 2, 3) (Figure 4
431 bottom). In comparison to the Warm Base simulation, retention in the Pribilofs (5)
432 increased as well as its connectivity to the South Outer Shelf (4) and North Middle Shelf
433 (10) (Figure 4 bottom). Expanded spawning lowered connectivity from the Midnorth
434 Middle Shelf (6) to many Middle Shelf regions (3, 4, 5, 6, 9, 10) (Figure 4 bottom). With
435 the exception of this area, retention was enhanced in all regions when compared to
436 retention in the Warm Base simulation (Table 4).

437 Advancing spawning times by 40 d in the Early simulation created the greatest
438 differences between the warm simulations. Connectivity from the majority of the
439 spawning regions to the Inner Shelf and northern Middle Shelf regions (2, 7, 10, 11, 12)
440 was stronger, as well as to the South Middle Shelf (3) (Figure 4 bottom). Conversely,
441 links from nearly all source regions to the Outer and Off-shelf regions and mid-north
442 Middle Shelf regions (4, 5, 6, 8, 15, 16) fell in the Early simulations, with the exception
443 of retention in the Midnorth Middle Shelf (6) (Figure 4 bottom). In addition, connectivity
444 to St. Matthews (9) on the Middle Shelf increased from Outer and Off-shelf areas (4, 8,
445 16), but decreased from the mid-north Middle Shelf regions (5, 6) (Figure 4 bottom).

446 When comparing the Warm year simulations to the Cold year simulations, the
447 Cold years tended to have much stronger connectivity to the Off-Shelf North (15) and
448 slightly increased connectivity to the Inner Shelf (1, 2, 7), while connectivity to the
449 Middle Shelf (3, 6, 9, 10) was greater in the Warm year simulations (Figure 4). The
450 Contract simulation produced connectivity patterns that were the most dissimilar from
451 other simulations (Table 5). Contracting the spawning areas resulted in greater
452 connectivity differences from the Base Cold simulation than delaying spawning 40 d with
453 the Late simulation, almost as large as the difference of the Base Cold from the Base
454 Warm simulation (Table 5). In contrast, connectivity patterns with Expanded spawning
455 were more similar to the Base Warm simulation than when spawning was 40 d Early, and
456 the Base Warm simulation was much more like both the Expanded and Early simulations
457 than the Base Cold one (Table 5).

458

459 **East-West transport**

460 Connectivity to more on-shelf regions was further exemplified by the fraction of
461 juveniles transported 0° – 5° , 5° – 10° , or $>10^{\circ}$ longitude to the east of where they were
462 spawned (Figure 5). The total fraction of all juveniles found east of their spawning
463 locations never exceeded 0.4 in any year (Figure 5). Especially strong peaks in the
464 fractions transported 0° – 5° E and 5° – 10° E occurred in 1998, 2000, and 2009 (Figure 5 top
465 and middle). Dispersal $>10^{\circ}$ E of the spawning origin was much less common and showed
466 a different interannual pattern from transport $<10^{\circ}$ E (Figure 5). While 2009 was also a
467 year with the fraction greater than the mean; transport was very low in 1998 and 2000,
468 and above average in 2007 (Figure 5). None of the eastward transport metrics were
469 significantly correlated with temperature or wind indices (Table 6), but dispersal $>10^{\circ}$ E
470 was significantly positively correlated with the winter North Pacific Index (NPIw;
471 $r=0.55$, $p=0.02$, Table 6).

472 When the years were grouped by Cold and Warm in the Base simulation, there
473 were differences in the fractions of individuals found various distances east and west of
474 their spawning origins (Table 7). The Cold year simulations had greater fractions
475 transported $>5^{\circ}$ W compared to the Warm simulations, but also greater fractions $>5^{\circ}$ E
476 (Table 7). When comparing across simulations, roughly 80% or more of all juveniles
477 were found to the west of where they were spawned, with the plurality found a distance
478 of $<5^{\circ}$ W in all simulations (Table 7). Contracting the spawning grounds off-shelf led to
479 greater westward transport, whereas advancing spawning time by 40 d with the Early
480 simulations increased eastward transport more than Expanding the spawning areas on-
481 shelf (Table 7).

482

483 **Comparisons with juvenile survey observations**

484 Agreement between the model and different sets of observations varied by year
485 and observational method (Table 8). The BASIS surface observations had the lowest LIC
486 values in all years and simulations indicating less than 50% mean overlap (Table 8). On
487 the other hand, model agreement with the *Oshoro Maru* and BASIS acoustics
488 observations were much better, with 76% and 67% mean collocation, respectively (Table
489 8). The model showed the highest agreement with the *Oshoro Maru* observations in all
490 simulations except when spawning was delayed 40 d in the Late simulation, which had a
491 greater LIC with the BASIS acoustics (Table 8). There was only one year (2007) that had
492 all three types of observations collected. In this year, the model overlapped greatly with
493 the *Oshoro Maru* data and poorly with both of the BASIS data sets (Table 8, Figure 6).
494 This year was not the best representation of how the model and observations differed. In
495 most years the model overestimated juveniles in the AK Peninsula (1), South Outer Shelf
496 (4), and Midnorth Middle Shelf (6) and underestimated juveniles in the south Inner and
497 Middle Shelves (2, 3) compared to the *Oshoro Maru* observations (Figure 7), whereas in
498 2007 there are greater fractions of modeled juveniles in all regions except for the South
499 Outer Shelf (4) (Figure 6 top). The over-prediction of modeled juveniles in the North
500 Outer Shelf (8) compared to the BASIS surface trawl observations was seen in 2007
501 (Figure 6 middle), but this year lacked the general trend of more northward displaced
502 modeled juveniles (regions 6, 9, 10, 15) (Figure 8). Compared to the BASIS acoustics,
503 the model simulated fewer juveniles on the south Inner and Middle Shelves (2, 3), and
504 more juveniles elsewhere (Figure 9), of which the decreased fractions in the South
505 Middle Shelf (3) were seen in 2007 (Figure 6 bottom).

506

507 DISCUSSION

508 **Relationship with oceanography and climate**

509 The strong and moderate connectivity between spawning and juvenile regions
510 exhibited three patterns: along-isobath flow to the northwest, self-retention, and in the
511 vicinity of the Pribilof Islands. Most of the consistently high connections existed from
512 spawning regions to the southeast to juvenile regions to the northwest, suggesting
513 transport via flows like the BSC (Figure 1) and the baroclinic flows along the 100-m and
514 200-m isobaths. These along-isobath flows carried individuals to regions directly
515 northwest of the spawning region, or to regions located either off- or on-shelf from the
516 spawning region and sharing a common border with the 100-m and 200-m isobaths. The
517 highest connectivity for all spawning times (except 15 Jan), years, and simulations was
518 retention within the North Outer Shelf (8). Though technically retention, transport within
519 the North Outer Shelf could also be considered northwesterly. This long, narrow region
520 that spans a wide range of latitudes could easily retain individuals that were spawned in
521 the south of the region near the Pribilofs (5) and were transported along-shelf to more
522 northwestern parts of the same region. Unlike the North Outer Shelf, the AK Peninsula
523 (1), South Middle Shelf (3), and Midnorth Middle Shelf (6) are not as long in the
524 northwest direction. The strong (3, 6) and moderate (1) retention in these regions must be
525 related to their locations on the Middle Shelf. Some of the retention in region 1 could be
526 the result of on-shelf advection by the ACC (Figure 1) that flows northeastward along the
527 Peninsula, while other retention in regions 1, 3, and 6 is likely related to the Inner and
528 Middle Fronts inhibiting cross-shelf flow (Coachman, 1986; P. Stabeno, pers. comm.).

529 The third connectivity feature that appeared was transport in and around the
530 Pribilof region (5). The Pribilof Islands are a major spawning ground (Hinckley, 1987;
531 Bachelier *et al.*, 2010) and are known for anti-cyclonic flow that can retain drifters for 20-
532 200 d (Kowalik & Stabeno, 1999; Stabeno *et al.*, 2008). The weak modeled retention of
533 this region could be explained by its small size that would account for anticyclonic
534 transport into the adjacent regions (3, 6, 8) with strong and moderate connectivity.
535 Transport from the Outer Shelves (4,8) to the Midnorth Middle Shelf (6) is evidence of
536 the anti-cyclonic flow extending into other regions. Off- and on-shelf transport across
537 two major fronts rarely occurred, with a few exceptions. Movement from the Middle
538 Shelf into the slope region of Off-shelf North (15) persisted with strong connectivity
539 from the Pribilofs (5) and moderate connectivity from the South (3) and Midnorth Middle
540 (6) Shelves. The greatest on-shelf connection was moderate from Off-shelf Southeast
541 (16) to the Midnorth Middle Shelf (6). All of these on- and off-shelf pathways required
542 crossing both the Middle Front and the BSC, likely via the anti-cyclonic Pribilof flow.
543 The inadequate size of the BEST-BSIERP Pribilof region is further supported by a study
544 that characterized the spatial extent of the Pribilof Islands habitat using ecosystem
545 energetics and foraging theory with data spanning 1990-2000 (Ciannelli *et al.*, 2004).
546 Both methods estimated a radius of approximately 100 nautical mi (185.2 km), which is
547 consistent with the size of the clockwise gyre during summer (Stabeno *et al.*, 1999). This
548 area is much larger than the BEST-BSIERP region 5 (<50 nmi) and would contain parts
549 of regions 3, 4, 6, 8, 15, and 16.

550

551 *Temporal variability within a season*

552 General connectivity patterns did not vary much across spawning times, though
553 some connections changed with time. The strong link between the AK Peninsula (1) and
554 South Middle Shelf (3) was greatest earlier in the spawning season. Similarly, the weak
555 connections to northernmost regions 10-14 decreased as the spawning season progressed.
556 The fluctuations in connectivity and retention with the Early spawning simulations arose
557 from these relationships. Connectivity strengthened to Inner Shelf and northern regions
558 and weakened to Outer and Off-shelf regions with 40 d Early spawning.

559 The AK Peninsula (1) and South Middle Shelf (3) tended to have the highest
560 retention in the middle of the spawning period, while retention in the Pribilofs (5),
561 Midnorth Middle Shelf (6), North Outer Shelf (8), and Off-shelf Southeast (16) regions
562 increased over the spawning season. Modeled connectivity and retention with 40 d Late
563 spawning corresponded to these seasonal patterns and inverted the results of the Early
564 simulation. Connections to the Inner and north Middle Shelves and retention within the
565 South Middle Shelf (3) were reduced when spawning times were shifted later. In contrast,
566 Late spawning improved retention in all other regions, most of which had higher
567 retention rates later in the Base spawning season.

568 Taken together, these results suggest that northward and on-shelf transport was
569 greater in the late winter and early spring compared to the rest of the spawning season. In
570 fact, advancing spawning time by 40 d with the Early simulations increased eastward
571 transport more than Expanding the spawning areas on-shelf. Ekman surface currents are
572 strongest in the winter and decrease through the spring to minima in the summer
573 (Danielson *et al.*, 2011b; Stabeno *et al.*, 2012; Danielson *et al.*, 2014), accounting for the
574 greater overall transport in the Early simulations and greater retention in Late

575 simulations. In addition to strength, current direction varied by month. For example, at
576 station M2 on the Middle Shelf (region 3) the mean monthly surface currents during
577 1995-2010 were to the northwest in Jan and Feb, and then shifted to the west/southwest
578 as they weakened through the spring (Stabeno *et al.*, 2012). The warm years showed a
579 different pattern with strong northward flow in Jan and Feb and weak eastward flow in
580 May and Jun (Stabeno *et al.*, 2012). Similarly, the BSC is strongest Jan-Mar and weakens
581 as the year progresses (Ladd, 2014). Additionally, seasonal stratification establishes the
582 Inner Front and Middle Front in Apr and May (Kachel *et al.*, 2002; Gibson *et al.*, 2013).
583 The absence of these fronts in the early spring substantiates the strong connectivity from
584 region 1 to region 3 during early spawning times by allowing transport across the 50-m
585 isobath, while their presence in the late spring and early summer supports the higher
586 retention rates in most regions with Apr, May, and Jun spawning times by preventing
587 cross-shelf transport of individuals. Additionally, the warmer temperatures of later
588 spawning times would raise growth rates, thus reducing the time to reach the juvenile
589 stage and potentially transport distance thereby increasing retention.

590

591 *Spatial variability within a region*

592 The alterations in connectivity, retention, and E-W transport in simulations with
593 Contracted and Expanded spawning areas revealed spatial variability of currents within
594 regions. The Contract simulation had the largest differences among the cold year
595 simulations and was the most dissimilar from all simulations. These divergences included
596 no spawning in the Midnorth Middle Shelf (6) and increased off-shelf transport.
597 Connectivity from the Middle Shelf regions (1, 3) rose to the north Middle, Outer, and

598 Off-shelf regions (6, 8, 9, 15). Greater off-shelf advection was also demonstrated with
599 increased fractions of individuals transported west of their spawning locations.
600 Contracting spawning areas to the southwest increased south and southwestern transport
601 from the Pribilofs (5) as well as retention within this region. On the other hand,
602 contracted spawning reduced retention in the AK Peninsula (1), South Middle Shelf (3),
603 and North Outer Shelf (8) and led to greater transport to the northwest, most likely by
604 concentrating the individuals spawned in each region closer to northwestern baroclinic
605 flows along the 100-m and 200-m isobaths and the BSC.

606 When spawning areas were Expanded on-shelf, delivery from most source regions
607 increased to the south Inner and Middle Shelves and subsequently decreased to the Outer
608 Shelves and Off-shelf North. In contrast to the Contracted simulation, Expansion
609 increased the fractions of individuals transported to the east. Furthermore, retention was
610 amplified in all regions except for the Midnorth Middle Shelf (6). The Expanded on-shelf
611 spawning locations resulted in more eastward movement that facilitated improved
612 retention in all other regions by placing eggs in regions with weaker flows, that were
613 further from northwestward currents, and/or near eastward currents such as the ACC.
614 Alternatively, Expanded spawning lowered connectivity from the Midnorth Middle Shelf
615 (6) to south and west regions (3, 4, and 5), suggesting that more on-shelf individuals in
616 region 6 were transported northwestward, while more off-shelf individuals in the region
617 were influenced by the anti-cyclonic circulation near the Pribilofs. This is further
618 supported by the increased Pribilofs retention and southwestward connectivity when
619 spawning sites were contracted off-shelf.
620

621 *Interannual variability and cold/warm years*

622 The years with the lowest similarity in connectivity to the other years, 2005 and
623 2007, were anomalous for different reasons, thus accounting for their low similarity to
624 each other. 2005 was one of the warmest years on record in the EBS, leading to strongly
625 stratified conditions in the southern shelf during the summer (Stabeno *et al.*, 2010) and
626 very high on-shore Ekman transport (Danielson *et al.*, 2011b). Stratification was also
627 strong in 2007, but wind velocities and thus Ekman transport were very low (Danielson *et*
628 *al.*, 2014). This strong stratification and low wind mixing reduced nutrient flux to surface
629 waters and may have ultimately led to poor pollock recruitment success in 2007 through
630 bottom-up processes, an unanticipated outcome of cold year conditions (Gann *et al.*, *in*
631 *press*). Additionally, 2007 was a year where ichthyoplankton species and stations groups
632 characterized by Busby *et al.* (2014) were unique from all other years (1996-2007), with
633 pollock contributing most to the differences.

634 Especially high fractions of individuals were transported 0°–5°E and 5°–10°E in
635 1998, 2000, and 2009. Many physical and biological characteristics of the EBS were
636 anomalous in 1998 (Stabeno *et al.*, 2001). This year had strong winds from Jan into Jun
637 and enhanced on-shelf transport (Stabeno *et al.*, 2001). This on-shelf transport was
638 beneficial for retention, which was above average in all regions in 1998. Both 1998 and
639 2000 had negative winter North Pacific Index values, indicative of strong Aleutian Lows,
640 yet the fractions of individuals transported 0°–5°E and 5°–10°E were not significantly
641 correlated to the winter NPI, nor any climate indices. In contrast, dispersal >10°E of the
642 spawning origin was significantly positively correlated with the winter NPI, denoting
643 more eastward transport during weak ALs, often linked with cold temperatures

644 (Rodionov *et al.*, 2007), which is consistent with greater fractions transported $>10^{\circ}\text{E}$ in
645 the cold simulations. However, the fraction of juveniles found $>10^{\circ}\text{E}$ from their
646 spawning locations was very small in all simulations, and neither the fractions $0^{\circ}\text{--}5^{\circ}\text{E}$
647 nor the summed fractions $>0^{\circ}\text{E}$ were higher in cold years. This discrepancy between
648 eastward transport and AL strength is rectified by the fact that it is the mean position of
649 the AL, not the strength, that affects winds and cross-isobath fluxes on the Bering shelf
650 (Danielson *et al.*, 2014), and that the position is more closely related to winter air
651 temperatures in the Bering Sea than the strength of the AL (Rodionov *et al.*, 2007).

652 When the mean winter position of the AL is more westward, the EBS shelf
653 experiences southeasterly winds that drive coastal convergence and northward flow along
654 the Alaskan coast (Danielson *et al.*, 2014). Over the Middle Shelf flow is on-shelf in the
655 south (regions 1, 3) and is northwards along isobaths in the midnorth (regions 6, 9), while
656 the north Outer Shelf (region 8) currents cross isobaths to the northeast (Danielson *et al.*,
657 2014). Additionally, western storm paths drive on-shelf Ekman transport (Danielson *et*
658 *al.*, 2014). The warm period 2000-2005 coincided with westward-displaced ALs, thus
659 supporting the modeled warm year patterns of more eastward transport over the shelf,
660 stronger connectivity to northern regions, and greater connectivity from the North Outer
661 Shelf (8) northeast to the Midnorth Middle Shelf (6) and St. Matthews (9). When Ekman
662 flows in warm years were not to the north, they were very weak in the cross-shelf
663 direction (Danielson *et al.*, 2011b; Stabeno *et al.*, 2012). The high retention rates and
664 smaller fractions of individuals transported $>5^{\circ}$ to the west and east of their spawning
665 locations in warm years reflected these weak cross-shelf flows.

666 In contrast, the AL was shifted eastward into the Gulf of Alaska during the cold
667 period of 2006-2011 (Danielson *et al.*, 2014). With an eastward-displaced AL, winds
668 over the EBS shelf are more northwesterly, northerly, and northeasterly forcing coastal
669 divergence and advection along isobaths to the southwest over the southern Middle Shelf
670 (regions 1, 3) (Danielson *et al.*, 2014). These characteristics account for the greater
671 westward transport and stronger off-shelf connectivity to regions such as the North Outer
672 Shelf (8), Off-shelf North (15), and Off-shelf Southeast (16) in the cold year simulations.
673 Additionally, greater fractions of individuals traveled further distances, in both eastward
674 and westward directions, in cold years compared to warm years. Correspondingly, the
675 stronger off-shelf currents in cold years (Danielson *et al.*, 2011b; Stabeno *et al.*, 2012;
676 Danielson *et al.*, 2014) support the greater travel and decreased retention in the Cold Base
677 simulation.

678

679 **Comparisons with observations**

680 There are important considerations as to why modeled and observed juvenile
681 distributions were not identical. For one, there is a difference in time between the two
682 distributions. Modeled distributions were the locations of individuals when they reached
683 40 mm; this could be any time between Mar and the following Jan, but were typically
684 Jun-Jul. On the other hand, most of the observations were taken later in the year (late Jul-
685 early Oct) where age-0 pollock could include fish up to 130 mm (Moss *et al.*, 2009), and
686 exclude fishes ≤ 40 mm by gear selectivity (e.g. BASIS surface trawl). Over the course of
687 this time, simulated juveniles would continue to be transported, possibly to different
688 nursery areas. The reason that simulations were not run until this time is the second

689 potential difference between the model and observations: the fact that swimming ability
690 increases as pollock juveniles get older and larger (Brodeur & Wilson, 1996; Duffy-
691 Anderson *et al.*, 2015) and they cannot be considered passive drifters in the horizontal
692 direction. Thirdly, in addition to horizontal swimming, vertical migration increases with
693 age resulting in different depth distributions of juveniles, with midwater abundances
694 often exceeding those at the surface (Parker-Stetter *et al.*, 2013). The spatial distributions
695 of age-0 pollock at the surface and deeper in the water column were similar in warm
696 years (2006, 2007), yet diverged in cold years (2009, 2010) when densities were low in
697 the surface and high in the midwaters (Parker-Stetter *et al.*, 2013). Thus, surface net tow
698 observations, like those from the BASIS survey may be representative of the total water
699 column distribution of juvenile pollock in warm years, but not in cold ones where they
700 would diverge from model results and the BASIS whole water column acoustic surveys.
701 These differences in depth distributions may account for why the agreement between the
702 model and BASIS surface tow data was higher in the warm years 2004 and 2005 and
703 lower in the cold years 2007, 2009, and 2010, and why full water column observations
704 such as the *Oshoro Maru* and BASIS acoustics data sets were always a better match with
705 simulated results.

706 In contrast to observations, modeled juveniles were found on the Inner Shelf
707 (regions 2, 7, 11), especially in cold years. The simulations are consistent with the
708 observed extension of a Middle Shelf zooplankton community into the Inner Shelf in cold
709 years, which may be the result of a weaker Inner Front in cold years (Eisner *et al.*, 2014)
710 or one that develops in August rather than June as was observed in 1998 and 1999
711 (Kachel *et al.*, 2002). Observed transport of zooplankton to the Inner Shelf implies that

712 either behavior and/or mortality are responsible for the lack of observed juveniles on the
713 Inner Shelf. High mortality on the Inner Shelf may be related to its low prey energy
714 density (Siddon *et al.*, 2013b), the lack of stratification that prevents energy conservation
715 by vertical migration, and/or the shallower depths that inhibit vertical migration to avoid
716 visual predators. Analogously, mortality in all regions, which is not represented in the
717 simulations, is yet another factor that could account for differences between modeled and
718 observed juvenile distributions.

719

720 **Consequences of connectivity**

721 The connectivity presented here is representative of potential connectivity.
722 Ultimate population connectivity would be determined from survival processes both
723 during and after dispersal until recruitment into the fishery. Some information about
724 survival of age-0 juveniles can be gleaned from the spatial distributions of prey and
725 predators in reference to the modeled regions of juvenile metamorphosis.

726

727 *Nursery areas*

728 Historically good juvenile habitat is near the Pribilof Islands (region 5; Duffy-
729 Anderson *et al.*, 2015). Tides and winds around the Pribilof Islands create a structural
730 front that enhances nutrient vertical fluxes, and primary and secondary productivity
731 (Brodeur *et al.*, 1997; Stabeno *et al.*, 1999; Stabeno *et al.*, 2008). This high productivity
732 in combination with a retentive clockwise circulation pattern around the islands create an
733 ideal rearing location (Swartzman *et al.*, 2005). Studies of the Pribilof Islands (Brodeur *et*
734 *al.*, 1997; Ciannelli *et al.*, 2002; Swartzman *et al.*, 2002; Swartzman *et al.*, 2005)

735 examined nursery suitability of all mixed, frontal, and stratified zones of the southeastern
736 Bering Sea by using the Pribilof Islands, which contained all three, as a representative
737 mesocosm. Their findings suggest that the Pribilof Islands and other stratified (like the
738 Outer Shelf regions 4 and 8) and frontal habitats of the EBS would be favorable feeding
739 and rearing habitats for juvenile pollock.

740 More recent studies during the warm and cold time periods acknowledge the
741 Outer Shelf (regions 4 and 8) as beneficial feeding areas for age-0 pollock juveniles. EBS
742 zooplankton communities were significantly different in warm and cold year surveys
743 from mid Aug to early Sep for 2003-2009 (Eisner *et al.*, 2014). These species
744 compositions are consequential, as zooplankton differ in their size and lipid content,
745 which affects their suitability and quality as juvenile pollock prey. For example, the
746 equivalent biomasses and distributions of copepods *Oithona* spp. in warm and
747 *Pseudocalanus* spp. and *Acartia* spp. in cold years do not necessarily denote equal
748 juvenile feeding, condition, and survival. *Oithona* spp. is a less preferred prey in age-0
749 diet analyses (Strasburger *et al.*, 2014) that could be determined either passively if its
750 smaller size reduces encounter rates (Petrik *et al.*, 2009), or actively if individuals choose
751 to target other taxa when available. Juvenile pollock gut fullness was not significantly
752 different in warm years (Heintz *et al.*, 2013), which does not suggest that feeding success
753 suffered. Despite high abundance and statistically similar average weight and percent gut
754 fullness, age-0 pollock energy density was lower in warm years compared to cold (Heintz
755 *et al.*, 2013). The large zooplankton *Thysanoessa* spp. and *Calanus marshallae* and the
756 small zooplankton *Pseudocalanus acuspes* and *Pseudocalanus minutus* store sizeable
757 amounts of lipids (Coyle *et al.*, 2011). Due to these zooplankton, the Outer Shelf (4, 8) in

758 all years and the Middle Shelf (1, 3, 5, 6, 9, 10) in cold years have higher energy densities
759 of prey, while the small zooplankton of the Inner Shelf (2, 7, 11) give it the lowest energy
760 density (Coyle *et al.*, 2011; Siddon *et al.*, 2013b). There was greater coincidence of age-0
761 pollock with regions of high prey energy density in cold years (e.g. 2010; Siddon *et al.*,
762 2013b), accounting for the higher proportion of euphausiids in their diet and the higher
763 energy density of pollock in cold years compared to warm years (Moss *et al.*, 2009;
764 Coyle *et al.*, 2011; Heintz *et al.*, 2013).

765 During warm years, the reduced availability of large lipid-rich prey items led to
766 age-0 pollock becoming a large portion of the diet of many fishes, such as juvenile
767 pollock and salmon (*Oncorhynchus* spp.; Moss *et al.*, 2009; Coyle *et al.*, 2011), Pacific
768 herring (*Clupea pallasii*; Andrews *et al.*, *accepted*), and age-0 Pacific cod (*Gadus*
769 *macrocephalus*; Farley *et al.*, 2015). In addition to the absence of preferred high lipid
770 prey, the lower per capita prey availability produced by greater juvenile densities and
771 decreased prey biomass (Heintz *et al.*, 2013) could have been accountable for the higher
772 proportion of cannibalism in warm years (Moss *et al.*, 2009; Coyle *et al.*, 2011; Heintz *et*
773 *al.*, 2013).

774 The distribution of predatory fishes also varied across the shelf by region and
775 warm or cold period (L. Eisner, pers. comm.) producing different predation rates based
776 on the amount of overlap with each species. During warm years juvenile pollock
777 distributions often overlapped with herring and sockeye salmon (*O. nerka*) on the north
778 and south Middle Shelf (1, 3, 5, 6, 9, 10; L. Eisner, pers. comm.) Juvenile pollock co-
779 occurred with capelin (*Mallotus villosus*) and herring on the northern parts of shelf (6, 9,
780 10, 11) and with Pacific cod and yellowfin sole (*Limanda aspera*) along the Inner Front

781 (and in regions 1 and 3 in 2010) (L. Eisner, pers. comm.). Only the Outer Shelf regions
782 (4, 8) lacked strong associations between juvenile pollock and predatory fishes (L.
783 Eisner, pers. comm.). These Outer Shelf regions are home to Arrowtooth flounder, a
784 voracious predator of pollock (Lang *et al.*, 2000), that has increased in abundance and
785 biomass over the past 30 years (Zador *et al.*, 2011; Hunsicker *et al.*, 2013). The cluster of
786 arrowtooth flounder in the northern half of the North Outer Shelf (8) would pose the
787 greatest threat to age-0 pollock since these arrowtooth flounder were the largest, had the
788 lowest proportion of non-empty stomachs, and consumed age-0s more frequently than
789 clusters in other areas in the 2000s (Zador *et al.*, 2011).

790

791 *Juvenile survival*

792 Swartzman *et al.* (2005) found a significant difference in the 1995-2001 T/S
793 *Oshoro Maru* net tow abundance of age-0 pollock between the Pribilof Islands and non-
794 Pribilof shelf areas (1-4, 6-9), with greater abundances within 100 nmi of the islands. The
795 increased contribution of the Pribilofs to recruitment in 1997-1999 (Swartzman *et al.*,
796 2005) led to use of the abundance of age-0 pollock near the Pribilof Islands as an index of
797 year class strength (Macklin & Hunt, 2004). Year-class strength at recruitment (age-3)
798 was correlated with age-0 density, the spatial relationship between juveniles and
799 euphausiids, and the centricity of age-0 distributions around the Pribilof Islands for the
800 years 1994-1999, having a positive relationship with euphausiid overlap (Winter &
801 Swartzman, 2006). Consistent with later diet and condition studies (Moss *et al.*, 2009;
802 Heintz *et al.*, 2013; Siddon *et al.*, 2013a), Winter and Swartzman (2006) hypothesized
803 that age-0 juveniles must transition their diet from copepods to euphausiids by Sep to

804 maximize survival potential. The Pribilof Island studies strengthen the idea that transport
805 to the Islands alone are a good predictor of recruitment success that would be enhanced
806 when this area had high concentrations of euphausiids and low concentrations of
807 predators.

808 Later studies of recruitment demonstrate that recruitment at age-2 is related to
809 predation mortality on age-1 juveniles (Mueter *et al.*, 2006). The variability in survival to
810 age-1 is best described by a model that includes temperature variability and predation,
811 such that survival is higher when temperatures and predation are lower to average
812 (Mueter *et al.*, 2011). Survival from age-0 to age-1 was poor in warm years as indicated
813 by high abundances of age-0s but below average recruitment year-classes (Moss *et al.*,
814 2009; Hunt *et al.*, 2011). The higher energy density and smaller lengths of age-0s in cold
815 years vs. warm years suggest survival to age-1 via an energy allocation strategy towards
816 lipid storage rather than growth (Moss *et al.*, 2009; Siddon 2013a). However, growth to a
817 minimum size would be advantageous because comparisons of length frequencies of age-
818 0 pollock in stomachs and trawls imply that both age-1 pollock and small arrowtooth
819 flounder (<20cm) have age-0 pollock upper size limits of 62 mm and 55 mm respectively
820 (Lang *et al.*, 2000). The number of age-1 recruits per spawner during the warm and cold
821 periods (2001-2010) was significantly positively related to age-0 weight, energy density,
822 and total energy (weight \times energy density), with total energy density explaining the most
823 variability (Heintz *et al.*, 2013). This relationship with total energy density indicates that
824 survival to age-1 requires both high energy condition going into winter to prevent
825 starvation and a large size to store and use lipids more efficiently (Heintz *et al.*, 2013)
826 and to avoid predation. High age-0 survival and ultimately recruitment would occur from

827 transport to regions with lower temperatures or stratified water columns and high
828 concentrations of energy-rich prey, to promote gaining mass and lipids, as well as low
829 concentrations of predators to prevent predation mortality.

830 Through synthesis of the information above, the Pribilofs (5), South Outer Shelf
831 (4), and North Outer Shelf (8) emerge as the most favorable age-0 nursery habitats,
832 regardless of temperature regime, based on high-energy prey communities and lower
833 predation risk. Though arrowtooth flounder, salmon, and older pollock are present in
834 these regions, arrowtooth flounder consume more pollock of ages 1-3 than age-0 (Zador
835 *et al.*, 2011; Hunsicker *et al.*, 2013), and the deeper depths of the Outer Shelf allow
836 pollock to vertically avoid salmon near the surface and arrowtooth flounder and older
837 pollock near the bottom (Hollowed *et al.*, 2012). The Base simulation mean connectivity
838 to the North Outer Shelf (8) was strong from all spawning regions in all years except
839 from the Midnorth Middle Shelf (6) in warm years, with the greatest connectivity from
840 retention. Mean connectivity to the South Outer Shelf (4) was strong or moderate from
841 the AK Peninsula (1) and South Middle Shelf (3) spawning regions in all years. Mean
842 connectivity to the Pribilofs (5) was always weak, most likely due to its small areal size.
843 The greatest proportions of juveniles that were transported to the Pribilofs came from the
844 AK Peninsula (1) and the South Middle Shelf (3) followed by retention within the region
845 and was higher in warm years compared to cold. If the Pribilof habitat is best represented
846 by a radius of 100 nmi, then it would contain parts of regions 3, 4, 6, 8, 15, and 16.
847 Accordingly, some connectivity to these regions would be just as beneficial as to region
848 5, though it would depend on the final transport location being within 100 nmi of the
849 center of the Pribilof habitat.

850 In cold years, the Middle Shelf (1, 3, 6) also hosted large populations of high
851 energy prey, however none of these regions offer particularly favorable predation risk
852 with herring, salmon, Pacific cod, and older pollock present. Though large zooplankton
853 were less abundant and had reduced lipid stores in warm years, northern parts of the shelf
854 (regions 9, 10, 11) had moderate amounts of prey energy density (Siddon *et al.*, 2013b)
855 and euphausiids were prevalent in the diet of other fishes (Andrews *et al.*, *accepted*). The
856 North Inner Shelf (11) is too shallow to avoid the abundant capelin, which may be a
857 predator on or prey competitor with age-0 pollock, and connectivity to this region was
858 very weak. If there are refuges in regions 9 and 10 from herring, salmon, and older
859 pollock, then the simulated increased connectivity to these regions in warm conditions
860 may also be beneficial.

861 Unsurprisingly, the greatest connectivity to favorable juvenile nurseries stemmed
862 from the historic spawning areas off of Unimak Island (1, 3) and the Pribilof Islands (5).
863 Spawning adults in these vicinities could be protected if management goals are for
864 increased production of juveniles to advantageous habitats. As climate projections
865 forecast increasing temperatures (Wang *et al.*, 2012), the rather invariant Outer Shelf (4,
866 8; Siddon *et al.*, 2011; Eisner *et al.*, 2014) will become increasingly important for its high
867 abundances of energy rich prey that are advected onto the shelf from the slope waters
868 (Gibson *et al.*, 2013) and for its greater depths that allow vertical energy conservation and
869 predator avoidance. Transport patterns and connectivity in warm years increased
870 connectivity to the South Outer Shelf (4) from its dominant source regions the AK
871 Peninsula (1) and South Middle Shelf (3) in these years. However, these connections
872 were reduced if either spawning areas were expanded or spawning times were advanced

873 under warm conditions. The reduced off-shelf transport in warm years decreased mean
874 connectivity to the North Outer Shelf (8) from the historic spawning regions (1, 3, 5) in
875 all warm simulations (Base, Expand, Early), which would place increased emphasis on
876 spawning in all other regions with warmer water temperatures regardless of whether
877 spawning locations and times remain the same, extend on-shelf, or advance.

878 In addition to the regional differences in feeding and predation conditions,
879 juvenile survival rates may be the consequence of their origin or the specific transport
880 route taken such that all juveniles that reach a particular habitat do not fare equally well.
881 For example, individual condition and survival could be impacted by maternal condition
882 at spawning (Solemdal, 1997; Berkeley *et al.*, 2004) or sublethal effects of temperature at
883 the time and location of spawning (e.g. abnormal development; Blood, 2002).
884 Additionally, juvenile size and condition will be affected by the temperatures and prey
885 types and abundances encountered during transport.

886

887 **Perspectives**

888 There was strong connectivity to Off-Shelf North (15) from 2 of the 7 of the
889 spawning areas in warm years and 6 of the 7 of the spawning areas in cold years of the
890 Base simulations. Juvenile pollock surveys do not regularly include stations in this region
891 and Off-shelf Southeast (16), and only a fraction of the North Outer Shelf (8) is sampled
892 (south of 60°N). The overall retention rate in all 16 BEST-BSIERP regions ranged from
893 0.65 to 0.83 across the different simulations, with a retention rate of 0.73 of the total
894 spawning in the Base simulation (Table 9). Juveniles in regions 8, 15, and 16 represented
895 0.34-0.51 of all the individuals spawned, which was greater than the proportion in the

896 other 13 regions in all simulations except for Early spawning (Table 9). Acoustic surveys
897 by the AFSC Midwater Assessment and Conservation Engineering Program have found a
898 strong pollock acoustic signal all over the northern Outer Shelf (Honkalehto *et al.*, 2011).
899 These observations in conjunction with the simulations suggest expansion of the early life
900 and juvenile surveys into Outer Shelf and Slope regions of the EBS. Simulations such as
901 the ones run in this study could be used to aid sampling plan design to include times and
902 locations with large predicted patches of juveniles and other early life stages.

903 The EBS is not isolated from other sources of pollock, but is connected to the
904 Gulf of Alaska by passes through the Aleutian Islands. A similar biophysical individual-
905 based model of pollock in the Gulf of Alaska was used to assess connectivity of that
906 stock. Spawning in the Gulf of Alaska resulted in high modeled densities of surviving
907 age-0 pollock juveniles in potential nursery areas of the Bering Sea (C. Parada, pers.
908 comm.). Mean connectivity from the eastern, central, and western Gulf of Alaska varied
909 from 0.004-0.183 to the Inner and Middle Shelves combined and from 0.031-0.628 to the
910 Outer Shelf and Basin combined (C. Parada, pers. comm.). Individuals were advected
911 into the Bering Sea through Unimak pass and dominated by source regions in the western
912 Gulf of Alaska (C. Parada, pers. comm.). Though this high potential connectivity may
913 represent a substantial loss to the Gulf of Alaska, it is doubtful that transport of Gulf of
914 Alaska pollock subsidizes the EBS stock if these juveniles survive and remain, given the
915 fact that the Gulf of Alaska stock is an order of magnitude smaller than the EBS stock
916 (Ianelli *et al.*, 2012b). Some support for this idea is lent by one study that found that
917 pollock early life stages were not a significant contribution to the larval fish species
918 assemblage associated with Unimak Island during 2002-2008 (Siddon *et al.*, 2011). Flow

919 through Unimak Pass was unidirectional in cold years and bidirectional in warm years of
920 the study, suggesting that the increased density of early life stages of pollock near
921 Unimak Island in warm years was the result of reverse direction flow from the EBS shelf
922 population (Siddon *et al.*, 2011). Low mixing between the EBS and Gulf of Alaska stocks
923 is additionally substantiated by morphological, demographic, and genetic differences
924 (O'Reilly *et al.*, 2004; Grant *et al.*, 2010) that separate the two stocks. Yet genetics would
925 not resolve any increased input of juveniles into the EBS from shifts in spawning into the
926 western Gulf of Alaska post 1989 (Ciannelli *et al.*, 2007). As the realized connectivity
927 from the Gulf of Alaska to the adult EBS population is unknown, it would be beneficial
928 to validate the simulated potential connectivity, possibly with natural tags such as otolith
929 microchemistry.

930

931 **Conclusions**

932 The three connectivity patterns of along-isobath flow to the northwest, self-retention,
933 and transport around the Pribilof Islands were seen in most simulations. The major
934 differences in connectivity between warm and cold years, more northwards in warm
935 years and more off-shelf in cold years, mimicked wind-driven flow characteristics of
936 those years (Danielson *et al.*, 2011b; Stabeno *et al.*, 2012) that were related to the winter
937 mean zonal position of the Aleutian Low (Danielson *et al.*, 2014). Connectivity
938 relationships were more sensitive to spatial alterations in the spawning areas in cold
939 years, as seen with the Contracted case, while they were more responsive to spawn timing
940 shifts in warm years, demonstrated with the Early simulation. Early spawning favored
941 transport to the Inner Shelf and northern regions, whereas Late spawning increased

942 retention in most regions and delivery to the South Outer Shelf (4), Pribilofs (5),
943 Midnorth Middle Shelf (6) and Off-shelf Southeast (16). Contracted spawning increased
944 connectivity in the off-shelf direction, while Expanded spawning facilitated retention
945 within regions. Synthesis of juvenile pollock prey and predator studies promote the
946 Pribilofs (5), South Outer Shelf (4), and North Outer Shelf (8) regions as very favorable
947 juvenile habitats due to their high densities of energy rich prey and deeper water columns
948 that allow for vertical separation from predators. The strongest connectivity to these
949 regions originated in the well-known spawning areas within the AK Peninsula (1), South
950 Middle Shelf (3), and Pribilofs (5), as well as in the North Outer Shelf (8). Surprisingly,
951 the North Outer Shelf (8) emerged as a very large sink of pollock reaching the juvenile
952 transition from all spawning sources, and requires more thorough sampling across
953 multiple trophic levels since it may be a potentially important juvenile pollock nursery.

954

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962

963 REFERENCES

964

965 Andrews III, A.G., Strasburger, W.W., Farley Jr., E.V., Murphy, J.M., and Coyle, K.O.
966 *accepted*. Effects of warm and cold climate conditions on capelin (*Mallotus villosus*) and
967 Pacific herring (*Clupea pallasii*) in the eastern Bering Sea. Deep-Sea Res. II.
968

969 Bachelier, N.M., Ciannelli, L., Bailey, K., and Duffy-Anderson, J.T. 2010. Spatial and
970 temporal patterns of walleye pollock (*Theragra chalcogramma*) spawning in the eastern
971 Bering Sea inferred from egg and larval distributions. Fisheries Oceanography, 19: 107-
972 120.
973

974 Baker, M.R., and Hollowed, A.B. 2014. Delineating ecological regions in marine
975 systems: Integrating physical structure and community composition to inform spatial
976 management in the eastern Bering Sea. Deep-Sea Research, 109: 215-240.
977

978 Berkeley, S.A., Chapman, C., and Sogard, S.M. 2004. Maternal age as a determinant of
979 larval growth and survival in a marine fish, *Sebastes melanops*. Ecology, 85: 1258-1264.
980

981 Blanke, B., and Raynaud, S. 1997. Kinematics of the Pacific Equatorial Undercurrent: a
982 Eulerian and Lagrangian approach from GCM results. Journal of Physical Oceanography,
983 27: 1038-1053.
984

985 Blood, D.M. 2002. Low-temperature incubation of walleye pollock (*Theragra*
986 *chalcogramma*) eggs from the southeast Bering Sea shelf and Shelikof Strait, Gulf of
987 Alaska. Deep-Sea Research II, 49: 6095-6108.

988

989 Brodeur, R.D., and Wilson, M.T. 1996. A review of the distribution, ecology and
990 population dynamics of age-0 walleye pollock in the Gulf of Alaska. *Fisheries*
991 *Oceanography*, 5: 144-166.

992

993 Brodeur, R.D., Wilson, M.T., Napp, J.M., Stabeno, P.J., and Salo, S., 1997. Distribution
994 of juvenile pollock relative to frontal structure near the Pribilof Islands, Bering Sea. *In*
995 *Proceedings of the International Symposium on the Role of Forage Fishes in Marine*
996 *Ecosystems*, pp. 573–589. Alaska Sea Grant AK-SG-97-01, Fairbanks. 816 pp.

997

998 Budgell, W. 2005. Numerical simulation of ice-ocean variability in the Barents Sea
999 region. *Ocean Dynamics*, 55: 370–387.

1000

1001 Busby, M.S., Duffy-Anderson, J.T., Mier, K.L., and De Forest, L.G. 2014. Spatial and
1002 temporal patterns in summer ichthyoplankton assemblages on the eastern Bering Sea
1003 shelf 1996-2007. *Fisheries Oceanography*, 23: 270-287, doi:10.1111/fog.12062.

1004

1005 Carton, J.A., and Giese, B.S. 2008. A reanalysis of ocean climate using Simple Ocean
1006 Data Assimilation (SODA). *Monthly Weather Review*, 136: 2999-3017.

1007

1008 Ciannelli, L., Brodeur, R.D., Swartzman, G.L., and Salo, S. 2002. Physical and biological
1009 factors influencing the spatial distribution of age-0 walleye pollock (*Theragra*

1010 *chalcogramma*) around the Pribilof Islands, Bering Sea. Deep-Sea Research II, 49: 6109-
1011 6126.
1012
1013 Ciannelli, L., Robson, B.W., Francis, R.C., Aydin, K., and Brodeur, R.D. 2004.
1014 Boundaries of open marine ecosystems: an application to the Pribilof Archipelago,
1015 southeast Bering Sea. Ecological Applications, 14: 942-953.
1016
1017 Ciannelli, L., Bailey, K.M., Chan, K.-S., and Christian Stenseth, N. 2007. Phenological
1018 and geographical patterns of walleye pollock (*Theragra chalcogramma*) spawning in the
1019 western Gulf of Alaska. Canadian Journal of Fisheries and Aquatic Sciences, 64: 713-
1020 722.
1021
1022 Coachman, L.K. 1986. Circulation, water masses, and fluxes on the southeastern Bering
1023 Sea shelf. Continental Shelf Research, 5: 23-108.
1024
1025 Cowen, R.K., and Sponaugle, S. 2009. Larval dispersal and marine population
1026 connectivity. Annual Review of Marine Science, 1: 443-466.
1027
1028 Coyle, K., Eisner, L., Mueter, F., Pinchuk, A., Janout, M., Cieciel, K., Farley, E., *et al.*
1029 2011. Climate change in the southeastern Bering Sea: impacts on pollock stocks and
1030 implications for the Oscillating Control Hypothesis. Fisheries Oceanography, 20: 139-
1031 156.
1032

1033 Curchitser, E.N., Hedstrom, K., Danielson, S., and Weingartner, T.J. 2010. Modeling the
1034 circulation in the North Aleutian Basin. Department of Interior, OCS Study BOEMRE
1035 2010-028.
1036

1037 Dai, A., and Trenberth, K.E. 2002. Estimates of freshwater discharge from continents:
1038 Latitudinal and seasonal variations. *Journal of Hydrometeorology*, 3: 660-687.
1039

1040 Danielson, S., Curchitser, E., Hedstrom, K., Weingartner, T., and Stabeno, P. 2011a. On
1041 ocean and sea ice modes of variability in the Bering Sea. *Journal of Geophysical*
1042 *Research*, 116: C12034, doi:10.1029/2011JC007389.
1043

1044 Danielson, S., Eisner, L., Weingartner, T., and Aagaard, K. 2011b. Thermal and haline
1045 variability over the central Bering Sea shelf: seasonal and interannual perspectives.
1046 *Continental Shelf Research*, 31: 539-554.
1047

1048 Danielson, S., Hedstrom, K., Aagaard, K., Weingartner, T., and Curchitser, E., 2012.
1049 Wind-induced reorganization of the Bering shelf circulation. *Geophysical Research*
1050 *Letters*, 39: L08601-L08606.
1051

1052 Danielson, S.L., Weingartner, T.J., Hedstrom, K.S., Aagaard, K., Woodgate, R.,
1053 Curchitser, E., and Stabeno, P.J. 2014. Coupled wind-forced controls of the Bering-
1054 Chukchi shelf circulation and the Bering Strait throughflow: Ekman transport, continental

1055 shelf waves, and variations of the Pacific-Arctic sea surface height gradient. Progress in
1056 Oceanography, 125: 40-61.

1057

1058 De Robertis, A., McKelvey, D., Taylor, K., and Honkalehto, T. 2014. Development of
1059 acoustic trawl survey methods to estimate the abundance of age-0 walleye pollock in the
1060 eastern Bering Sea shelf during the Bering Arctic Subarctic Integrated Survey. U.S.
1061 Department of Commerce, NOAA Technical Memorandum, NMFS-AFSC-272, 46 pp.

1062

1063 de Vries, P., and Döös, K. 2001. Calculating Lagrangian trajectories using time-
1064 dependent velocity fields. Journal of Atmospheric and Oceanic Technology, 18: 1092-
1065 1101.

1066

1067 Döös, K. 1995. Inter-ocean exchange of water masses. Journal of Geophysical Research,
1068 100: 13499-13514.

1069

1070 Duffy-Anderson, J.T., Barbeaux, S., Farley, E., Heintz, R., Horne, J., Parker-Stetter, S.,
1071 Petrik, C., *et al.* 2015. State of knowledge review and synthesis of the first year of life of
1072 walleye pollock (*Gadus chalcogrammus*) in the eastern Bering Sea with comments on
1073 implication for recruitment. Deep-Sea Research II,
1074 <http://dx.doi.org/10.1016/j.dsr2.2015.02.001>.

1075

1076 Eisner, L., Napp, J., Mier, K., Pinchuk, A., and Andrews, A. 2014. Climate-mediated
1077 changes in zooplankton community structure for the eastern Bering Sea. Deep-Sea
1078 Research II, doi:10.1016/j.dsr2.2014.03.004.
1079

1080 Farley, Jr., E.V., Heinz, R.A., Andrews, A.G., and Hurst, T.P. 2015. Size, diet, and
1081 condition of age-0 Pacific cod (*Gadus macrocephalus*) during warm and cool climate
1082 states in the eastern Bering Sea. Deep-Sea Research II,
1083 <http://dx.doi.org/10.1016/j.dsr2.2014.12.011>.
1084

1085 Gann, J.C., Eisner, L.B., Porter, S., Watson, J.T., Cieciel, K.D., Mordy, C.W.,
1086 Yasumiishi, E.M., *et al. in press*. Possible mechanism linking ocean conditions to low
1087 body weight and poor recruitment of age-0 walleye pollock (*Gadus chalcogrammus*) in
1088 the southeast Bering Sea during 2007. Deep-Sea Research II.
1089

1090 Gibson, G.A., Coyle, K.O., Hedstrom, K., and Curchitser, E.N. 2013. A modeling study
1091 to explore on-shelf transport of oceanic zooplankton in the Eastern Bering Sea. Journal of
1092 Marine Systems, 121-122: 47-64.
1093

1094 Grant, W.S., Spies, I., and Canino, M.F. 2010. Shifting-balance structure in North Pacific
1095 walleye pollock (*Gadus chalcogrammus*). ICES Journal of Marine Science, 67: 1687-
1096 1696.
1097

1098 Heintz, R.A., Siddon, E.C., Farley Jr., E.V., and Napp, J.M. 2013. Correlation between
1099 recruitment and fall condition of age-0 walleye pollock (*Theragra chalcogramma*) from
1100 the eastern Bering Sea under varying climate conditions. Deep-Sea Research II, 94: 150-
1101 156.
1102
1103 Hinckley, S. 1987. The reproductive biology of walleye pollock, *Theragra*
1104 *chalcogramma*, in the Bering Sea, with reference to spawning stock structure. Fisheries
1105 Bulletin U.S., 85: 481-498.
1106
1107 Hinrichsen, H.-H., Schmidt, J.O., Petereit, C., and Möllmann, C. 2005. Survival
1108 probability of Baltic larval cod in relation to spatial overlap patterns with its prey
1109 obtained from drift model studies. ICES Journal of Marine Science, 62: 878-885.
1110
1111 Hollowed, A.B., Barbeaux, S.J., Cokelet, E.D., Farley, E., Kotwicki, S., Ressler, P.H.,
1112 Spital, C., *et al.* 2012. Effects of climate variation on pelagic ocean habitats and their role
1113 in structuring forage fish distributions in the Bering Sea. Deep-Sea Research II, 65-70:
1114 230-250.
1115
1116 Honkalehto, T., Ressler, P.H., Towler, R., and Wilson, C. 2011. Using acoustic data to
1117 estimate walleye pollock abundance in the eastern Bering Sea. Canadian Journal of
1118 Fisheries and Aquatic Sciences, 68: 1231-1242.
1119

1120 Hunsicker, M.E., Ciannelli, L., Bailey, K.M., Zador, S., and Stige, L.C. 2013. Climate
1121 and demography dictate the strength of predator-prey overlap in a subarctic marine
1122 ecosystem. PLoS ONE, 8: e66025. doi:10.1371/journal.pone.0066025.
1123

1124 Hunt, G.L., Coyle, K.O., Eisner, L.B., Farley, E., Heinz, R., Mueter, F., Napp, J.M., *et al.*
1125 2011. Climate impacts on eastern Bering Sea food webs: a synthesis of new data and an
1126 assessment of the Oscillating Control Hypothesis. ICES Journal of Marine Science, 68:
1127 1230-1243.
1128

1129 HYCOM Ocean Prediction website. <http://hycom.org/ocean-prediction>. Last accessed 25
1130 Nov 2013.
1131

1132 Ianelli, J.N., Barbeaux, S.J., McKelvey, D., and Honkalehto, T. 2012a. Assessment of
1133 walleye pollock in the Bogoslof Island Region. *In* Stock assessment and fishery
1134 evaluation report for the groundfish resources of the Bering Sea/Aleutian Islands regions,
1135 pp. 235-244. North Pacific Fishery Management Council, Anchorage. 1298?? pp.
1136

1137 Ianelli, J.N., Honkalehto, T., Barbeaux, S., Kotwicki, S., Aydin, K., Williamson, N.
1138 2012b. Assessment of the Walleye Pollock stock in the Eastern Bering Sea. *In* Stock
1139 assessment and fishery evaluation report for the groundfish resources of the Bering
1140 Sea/Aleutian Islands regions, pp. 51-156. North Pacific Fishery Management Council,
1141 Anchorage. 1298?? pp.
1142

1143 Kachel, N.B., Hunt Jr., G.L., Salo, S.A., Schumacher, J.D., Stabeno, P.J., and Whitley,
1144 T.E. 2002. Characteristics and variability of the inner front of the southeastern Bering
1145 Sea. *Deep-Sea Research II*, 49: 5889-5909.
1146
1147 Kalnay, E., Kanamitsu, M., Kistler, R., Collins, W., Deaven, D., Gandin, L., Iredell, M.,
1148 *et al.* 1996. The NCEP/NCAR 40-year reanalysis project. *Bulletin of the American*
1149 *Meteorological Society*, 77: 437-471.
1150
1151 Kotwicki, S., and Lauth, R.R. 2013. Detecting temporal trends and environmentally-
1152 driven changes in the spatial distribution of bottom fishes and crabs on the eastern Bering
1153 Sea shelf. *Deep-Sea Research II*, 94: 231-243.
1154
1155 Kowalik, Z., and Stabeno, P. 1999. Trapped motion around the Pribilof Islands in the
1156 Bering Sea. *Journal of Geophysical Research*, 104: 25,667–25,684.
1157 doi:10.1029/1999JC900209.
1158
1159 Ladd, C. 2014. Seasonal and interannual variability of the Bering Slope Current. *Deep-*
1160 *Sea Research II*, doi/10.1016/j.dsr2.2013.12.005
1161
1162 Lang, G. M., Brodeur, R.D., Napp, J.M., and Schabetsberger, R. 2000. Variation in
1163 groundfish predation on juvenile walleye pollock relative to hydrodynamic structure near
1164 the Pribilof Islands, Alaska. *ICES Journal of Marine Science*, 57: 265-271.
1165

1166 Large, W.G., and Yeager, S.G. 2009. The global climatology of an interannually varying
1167 air-sea flux data set. *Climate Dynamics*, 33: 341-364.
1168

1169 Macklin, S.A., and Hunt, Jr., G.L. (Eds.) 2004. The southeast Bering Sea ecosystem:
1170 implications for marine resource management (Final Report: Southeast Bering Sea
1171 Carrying Capacity). NOAA Coastal Ocean Decision Analysis Series No. 24, Silver
1172 Spring, 192 pp.
1173

1174 Matarese, A.C., Kendall, Jr., A.W., Blood, D.M., and Vinter, B.M. 1989. Laboratory
1175 guide to early life history stages of northeast Pacific fishes. NOAA/National Marine
1176 Fisheries Service, NOAA Technical Report NMFS 80. 652 pp.
1177

1178 Moss, J.H., Farley, E.V., Feldman, A.M., and Ianelli, J.N. 2009. Spatial distribution,
1179 energetic status, and food habits of eastern Bering Sea age-0 walleye pollock.
1180 *Transactions of the American Fisheries Society*, 138: 497-505.
1181

1182 Mueter, F.J., Ladd, C., Palmer, M.C., and Norcross, B.L. 2006. Bottom-up and top-down
1183 controls of walleye pollock (*Theragra chalcogramma*) on the Eastern Bering Sea shelf.
1184 *Progress in Oceanography*, 68: 152-183.
1185

1186 Mueter, F.J., and Litzow, M.A. 2008. Sea ice retreat alters the biogeography of the
1187 Bering Sea continental shelf. *Ecological Applications*, 18: 309-320.
1188

1189 Mueter, F.J., Bond, N.A., Ianelli, J.N., and Hollowed, A.B. 2011. Expected declines in
1190 recruitment of walleye pollock (*Theragra chalcogramma*) in the eastern Bering Sea
1191 under future climate change. ICES Journal of Marine Science, 68: 1284-1296.
1192
1193 NARR website. http://nomads.ncdc.noaa.gov/dods/NCEP_NARR_DAILY/narr-
1194 [a_221_uvscf.subset.info](http://nomads.ncdc.noaa.gov/dods/NCEP_NARR_DAILY/narr-a_221_uvscf.subset.info). Last accessed 6 May 2015.
1195
1196 O'Reilly, P.T., Canino, M.F., Bailey, K.M., and Bentzen, P. 2004. Inverse relationship
1197 between F_{ST} and microsatellite polymorphism in the marine fish, walleye pollock
1198 (*Theragra chalcogramma*): implications for resolving weak population structure.
1199 Molecular Ecology, 13: 1799-1814.
1200
1201 Ortiz, I., Wiese, F.K., and Grieg, A. 2012. Marine regions boundary data for the Bering
1202 Sea shelf and slope. UCAR/NCAR-Earth Observing Laboratory/Computing, Data, and
1203 Software Facility. Dataset. <http://dx.doi.org/10.5065/D6DF6P6C>.
1204
1205 Parker-Stetter, S.L., Horne, J.K., Farley, E.V., Barbee, D.H., Andrews III, A.G., Eisner,
1206 L.B., and Nomura, J.M. 2013. Summer distributions of forage fish in the eastern Bering
1207 Sea. Deep-Sea Research II, <http://dx.doi.org/10.1016/j.dsr2.2013.04.022>
1208
1209 Petrik, C.M., Kristiansen, T., Lough, R.G., and Davis, C.S. 2009. Prey selection by larval
1210 haddock and cod on copepods with species-specific behavior: an individual-based model
1211 analysis. Marine Ecology Progress Series, 396: 123-143.

1212

1213 Petrik, C.M., Duffy-Anderson, J.T., Mueter, F.J., Hedstrom, K., and Curchitser, E.N.
1214 2014. Biophysical transport model suggest climate variability determines distribution of
1215 Walleye Pollock early life stages in the eastern Bering Sea through effects on spawning.
1216 Progress in Oceanography, <http://dx.doi.org/10.1016/j.pocean.2014.06.004>.

1217

1218 Renner, H.M., Mueter, F., Drummond, B.A., Warzybok, J.A., and Sinclair, E.H. 2012.
1219 Patterns of change in diets of two piscivorous seabird species during 35 years in the
1220 Pribilof Islands. Deep-Sea Research II, 65-70: 273-291.

1221

1222 Rienecker, M.M., Suarez, M.J., Gelaro, R., Todling, R., Bacmeister, J., Liu, E.,
1223 Bosilovich, M.G., *et al.* 2011. MERRA: NASA's modern-era retrospective analysis for
1224 research and applications. Journal of Climate, 24: 3624-3648.

1225

1226 Rodionov, S.N., Bond, N.A., and Overland, J.E. 2007. The Aleutian Low, storm tracks,
1227 and winter climate variability in the Bering Sea. Deep-Sea Research II, 54: 2560-2577.

1228

1229 Shchepetkin, A.F., and McWilliams, J.C. 2009. Ocean forecasting in terrain-following
1230 coordinates: Formulation and skill assessment of the regional ocean modeling system.
1231 Journal of Computational Physics, 228: 8985-9000. doi :10.1016/j.jcp.2009.09.002.

1232

1233 Siddon, E.C., Duffy-Anderson, J.T., and Mueter, F.J. 2011. Community-level response of
1234 fish larvae to environmental variability in the southeastern Bering Sea. Marine Ecology
1235 Progress Series, 426: 225-239.
1236

1237 Siddon, E.C., Heinz, R.A., and Mueter, F.J. 2013a. Conceptual model of energy
1238 allocation in walleye pollock (*Theragra chalcogramma*) from age-0 to age-1 in the
1239 southeastern Bering Sea. Deep-Sea Research II, 94: 140-149.
1240

1241 Siddon, E.C., Kristiansen, T., Mueter, F.J., Holsman, K., Heinz, R.A., and Farley, E.V.
1242 2013b. Spatial match-mismatch between juvenile fish and prey explains recruitment
1243 variability across contrasting climate conditions in the eastern Bering Sea. PLoS ONE, 8:
1244 e84526, doi:10.1371/journal.pone.0084526.
1245

1246 Smart, T.I., Duffy-Anderson, J.T, and Horne, J.K. 2012. Alternating temperature states
1247 influence walleye pollock early life stages in the southeastern Bering Sea. Marine
1248 Ecology Progress Series, 455: 257-267.
1249

1250 Smart, T.I., Siddon, E.C., and Duffy-Anderson, J.T. 2013. Vertical distributions of the
1251 early life stages of walleye pollock (*Theragra chalcogramma*) in the Southeastern Bering
1252 Sea. Deep-Sea Research II, 94: 201-210.
1253

1254 Solemdal, P. 1997. Maternal effects – a link between past and future. Journal of Sea
1255 Research, 37: 213-227.

1256

1257 Stabeno, P.J., Schumacher, J.D., Salo, S.A., Hunt, G.L., and Flint, M. 1999. Physical
1258 environment around the Pribilof Islands. *In* Dynamics of the Bering Sea, pp. 193-215. Ed.
1259 by T.R. Loughlin, and K. Ohtani. Alaska Sea Grant Press AK-SG-99-03, Fairbanks, 838
1260 pp.

1261

1262 Stabeno, P.J., Bond, N.A., Kachel, N.B., Salo, S.A., and Schumacher, J.D. 2001. On the
1263 temporal variability of the physical environment over the south-eastern Bering Sea.
1264 Fisheries Oceanography, 10: 81-89.

1265

1266 Stabeno, P.J., Kachel, N., Mordy, C., Righi, D., and Salo, S. 2008. An examination of the
1267 physical variability around the Pribilof Islands in 2004. Deep-Sea Research II, 55: 1701-
1268 1716.

1269

1270 Stabeno, P., Napp, J., Mordy, C., and Whitley, T. 2010. Factors influencing physical
1271 structure and lower trophic levels of the eastern Bering Sea shelf in 2005: sea ice, tides
1272 and winds. Progress in Oceanography, 85: 180-196.

1273

1274 Stabeno, P.J., Kachel, N.B., Moore, S.E., Napp, J.M., Sigler, M., Yamaguchi, A., and
1275 Zerbini, A.N. 2012. Comparison of warm and cold years on the southeastern Bering Sea
1276 shelf and some implications for the ecosystem. Deep-Sea Research II, 65-70: 31-45.

1277

1278 Strasburger, W.W., Hillgruber, N., Pinchuk, A.I., and Mueter, F.J. 2014. Feeding ecology
1279 of age-0 walleye pollock (*Gadus chalcogrammus*) and Pacific cod (*Gadus*
1280 *macrocephalus*) in the southeastern Bering Sea. *Deep-Sea Research II*, 109: 172-180.
1281

1282 Swartzman, G., Napp, J., Brodeur, R., Winter, A., and Ciannelli, L. 2002. Spatial patterns
1283 of pollock and zooplankton distribution in the Pribilof Islands, Alaska nursery area and
1284 their relationship to pollock recruitment. *ICES Journal of Marine Science*, 59: 1167-
1285 1186.
1286

1287 Swartzman, G., Winter, A., Coyle, K., Brodeur, R., Buckley, T., Ciannelli, L., Hunt, G.,
1288 *et al.* 2005. Relationship of age-0 pollock abundance and distribution around the Pribilof
1289 Islands, to other shelf regions of the eastern Bering Sea. *Fisheries Research*, 74: 273-287.
1290

1291 Trenberth, K.E., and Hurrell, J.W. 1994. Decadal atmosphere-ocean variations in the
1292 Pacific. *Climate Dynamics*, 9: 303-319.
1293

1294 Wang, M., Overland, J.E., and Stabeno, P.J. 2012. Future climate of the Bering and
1295 Chukchi Seas projected by global climate models. *Deep-Sea Research II*, 65-70: 46-57.
1296

1297 Winter, A. G., and Swartzman, G.L. 2006. Interannual changes in distribution of age-0
1298 walleye pollock near the Pribilof Islands, Alaska, with reference to the prediction of
1299 pollock year-class strength. *ICES Journal of Marine Science*, 63: 1118-1135.
1300

1301 Zador, S., Aydin, K., and Cope, J. 2011. Fine-scale analysis of arrowtooth flounder
1302 *Atheresthes stomias* catch rates reveals spatial trends in abundance. Marine Ecology
1303 Progress Series, 438: 229-239.
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1324 TABLES

1325 **Table 1.** Mean percent connectivity across all spawning times and years in the Base

1326 (“Transport Only”) simulation.

Sink	Source						
	1	3	4	5	6	8	16
1	8.66	7.69	0.51	0.72	0.25	0.10	0.53
2	0.74	0.94	0.10	1.16	1.13	0.08	0.06
3	20.50	20.08	1.74	6.65	4.65	0.82	1.72
4	10.79	9.89	1.95	2.68	2.02	0.63	2.13
5	2.47	2.44	0.73	1.69	1.18	0.47	0.86
6	10.96	9.95	6.09	20.73	24.12	7.67	5.60
7	0.38	0.46	0.10	1.30	1.99	0.10	0.05
8	16.60	18.71	30.97	27.73	23.32	37.45	31.87
9	3.19	3.64	3.05	12.89	26.06	5.97	2.42
10	0.40	0.48	0.53	1.28	3.01	1.18	0.51
11	0.10	0.16	0.04	0.57	1.08	0.07	0.02
12	0.03	0.10	0.38	0.13	0.14	0.49	0.38
13	0.02	0.07	0.28	0.07	0.06	0.34	0.30
14	0.00	0.00	0.00	0.00	0.01	0.01	0.00
15	9.38	9.12	16.73	10.59	5.67	13.48	15.90
16	5.08	4.55	2.74	1.04	0.18	1.12	3.64

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1328 **Table 2.** Local Index of Collocation comparisons of connectivity patterns between

1329 spawning times in the Base (“Transport Only”) simulation.

	15-Jan	1-Feb	15-Feb	1-Mar	15-Mar	1-Apr	15-Apr
15-Jan	1.00						
1-Feb	0.58	1.00					
15-Feb	0.58	0.98	1.00				
1-Mar	0.55	0.96	0.98	1.00			
15-Mar	0.54	0.95	0.97	0.99	1.00		
1-Apr	0.59	0.86	0.85	0.85	0.87	1.00	
15-Apr	0.61	0.84	0.84	0.84	0.86	0.99	1.00
Mean	0.58	0.86	0.87	0.86	0.87	0.84	0.83

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1332 **Table 3.** Local Index of Collocation comparisons between years in the Base (“Transport
 1333 Only”) simulation.

	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	
1995	1.00																		
1996	0.70	1.00																	
1997	0.74	0.92	1.00																
1998	0.67	0.67	0.66	1.00															
1999	0.85	0.85	0.85	0.78	1.00														
2000	0.77	0.91	0.92	0.79	0.87	1.00													
2001	0.78	0.70	0.65	0.72	0.89	0.65	1.00												
2002	0.78	0.89	0.92	0.71	0.90	0.94	0.66	1.00											
2003	0.68	0.96	0.91	0.58	0.79	0.86	0.65	0.84	1.00										
2004	0.73	0.89	0.97	0.71	0.87	0.95	0.66	0.94	0.87	1.00									
2005	0.62	0.82	0.86	0.50	0.68	0.86	0.45	0.84	0.83	0.87	1.00								
2006	0.77	0.69	0.63	0.64	0.90	0.66	0.95	0.69	0.64	0.65	0.48	1.00							
2007	0.66	0.42	0.35	0.43	0.66	0.43	0.76	0.46	0.38	0.37	0.29	0.87	1.00						
2008	0.74	0.53	0.49	0.65	0.78	0.60	0.83	0.59	0.49	0.53	0.42	0.88	0.90	1.00					
2009	0.64	0.81	0.91	0.70	0.80	0.84	0.62	0.83	0.82	0.90	0.77	0.58	0.29	0.47	1.00				
2010	0.85	0.57	0.55	0.61	0.74	0.59	0.85	0.57	0.60	0.56	0.50	0.82	0.75	0.82	0.51	1.00			
2011	0.67	0.90	0.88	0.77	0.88	0.92	0.73	0.90	0.85	0.92	0.81	0.73	0.48	0.69	0.84	0.61	1.00		
2012	0.87	0.65	0.59	0.72	0.83	0.64	0.93	0.62	0.61	0.58	0.43	0.90	0.83	0.87	0.52	0.90	0.65	1.00	
Mean	0.74	0.76	0.75	0.67	0.82	0.78	0.73	0.77	0.73	0.76	0.65	0.73	0.55	0.66	0.70	0.67	0.78	0.71	

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1348 **Table 4.** Mean percent retention rate of each spawning region grouped by year (all
 1349 spawning times) in the Base simulation and by simulation (all spawning times and years).

Simulation	Time period	Region						
		1	3	4	5	6	8	16
Base all	1995	2.7	21.8	1.5	1.9	6.5	53.0	7.3
Base all	1996	11.3	14.9	1.9	2.9	24.7	46.5	4.3
Base all	1997	3.7	32.0	0.7	1.7	34.3	57.1	0.0
Base all	1998	18.5	23.7	12.7	4.2	50.8	41.1	16.7
Base all	1999	7.0	15.8	1.8	2.5	33.4	42.3	1.2
Base all	2000	12.0	36.3	8.0	2.9	23.5	55.2	11.5
Base all	2001	7.1	2.8	0.2	1.0	47.3	27.9	4.8
Base all	2002	7.0	40.6	2.1	2.2	19.6	52.9	0.8
Base all	2003	9.0	14.4	1.4	0.6	13.8	38.1	1.3
Base all	2004	5.4	41.0	0.5	2.5	34.9	44.2	0.2
Base all	2005	4.4	35.7	0.2	1.5	2.9	37.1	0.0
Base all	2006	4.9	4.4	0.3	1.1	31.3	30.9	0.1
Base all	2007	4.4	2.9	0.3	0.1	0.1	19.9	0.1
Base all	2008	17.6	12.1	0.4	0.7	10.0	17.6	0.7
Base all	2009	9.9	31.0	0.6	2.3	37.3	28.1	1.4
Base all	2010	6.1	3.9	0.2	0.1	11.0	16.6	0.0
Base all	2011	20.7	26.5	1.4	1.0	31.3	29.9	3.7
Base all	2012	4.0	1.4	1.0	1.3	21.5	35.6	11.3
Cold	Cold yrs	8.5	17.1	1.5	1.4	21.8	35.1	3.4
Contract	Cold yrs	4.5	7.9	1.5	1.7	--	32.4	4.4
Late	Cold yrs	13.3	16.4	2.3	3.0	34.5	36.0	7.2
Warm	Warm yrs	7.9	26.4	1.4	1.8	15.3	43.7	1.6
Expand	Warm yrs	10.3	32.6	1.7	2.7	14.2	44.2	1.6
Early	Warm yrs	8.8	31.9	1.1	1.2	17.5	33.8	0.9

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1352 **Table 5.** Local Index of Collocation comparisons of connectivity between simulations.

	Cold	Contract	Late	Warm	Expand	Early
Cold	1.00					
Contract	0.88	1.00				
Late	0.97	0.83	1.00			
Warm	0.86	0.76	0.81	1.00		
Expand	0.86	0.74	0.82	0.99	1.00	
Early	0.84	0.74	0.79	0.95	0.95	1.00

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1354 **Table 6.** Correlations between climate indices and the fraction of juveniles transported
 1355 varying distances east from their spawning locations in the Base simulation. Correlations
 1356 significant at the 0.05 level are in bold. NPIw: winter North Pacific Index, May SST:
 1357 May Sea Surface Temperature, Cross: cross-shelf wind velocities, Along: along-shelf
 1358 wind velocities.

	0°-5°E	5°-10°E	>10°E	>0°E
NPIw	-0.19	0.02	0.55	-0.17
MaySST	0.33	0.02	-0.11	0.30
Cross	-0.12	0.04	0.30	-0.10
Along	0.04	-0.16	-0.08	0.02

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1360 **Table 7.** Longitudinal transport of juveniles from spawning location presented as the
 1361 fraction of the total in each simulation.

	>0°W	>10°W	5-10°W	<5°W	<5°E	5-10°E	>10°E	>0°E
Base all	0.87	0.11	0.34	0.41	0.12	0.007	2.7E-05	0.13
Cold	0.89	0.13	0.35	0.40	0.10	0.007	3.1E-05	0.11
Contract	0.92	0.15	0.38	0.38	0.07	0.005	3.0E-05	0.07
Late	0.85	0.08	0.30	0.47	0.13	0.007	2.1E-06	0.14
Warm	0.86	0.06	0.33	0.47	0.13	0.004	1.6E-05	0.14
Expand	0.82	0.05	0.29	0.48	0.17	0.006	1.4E-05	0.17
Early	0.79	0.06	0.28	0.45	0.19	0.010	5.6E-05	0.20

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1369 **Table 8.** Local Index of Collocation values of model-observation comparisons for
 1370 individual years in the Base simulations and for all years in the alternate simulations.
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	<i>Oshoro Maru</i>	BASIS surface trawl	BASIS acoustics
Base 1995	--	--	--
Base 1996	0.72	--	--
Base 1997	0.68	--	--
Base 1998	0.77	--	--
Base 1999	0.58	--	--
Base 2000	0.99	--	--
Base 2001	0.53	--	--
Base 2002	0.89	--	--
Base 2003	0.47	0.28	--
Base 2004	0.98	0.65	--
Base 2005	0.80	0.76	--
Base 2006	--	0.58	0.64
Base 2007	0.97	0.02	0.43
Base 2008	--	0.67	0.74
Base 2009	--	0.31	0.44
Base 2010	--	0.29	0.91
Base 2011	--	0.48	0.89
Base 2012	--	0.21	0.61
Base Mean	0.76	0.42	0.67
Cold	0.81	0.37	0.67
Contract	0.70	0.31	0.60
Late	0.65	0.37	0.68
Warm	0.72	0.52	--
Expand	0.78	0.56	--
Early	0.82	0.58	--

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1377 **Table 9.** Mean percent retention of all individuals spawned in the eastern Bering Sea by
 1378 BEST-BSIERP juvenile region and simulation.

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	All regions	Regions 1-7, 9-14	Regions 8, 15, 16
Base all	73.3	26.1	47.2
Cold	70.3	22.8	47.5
Contract	64.8	16.8	48.0
Late	77.4	26.0	51.4
Warm	80.5	34.2	46.3
Expand	83.3	40.8	42.5
Early	74.3	40.4	33.9

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1382 **FIGURE CAPTIONS**

1383

1384 **Figure 1.** (Large panel) Map of the Eastern Bering Sea with bathymetry (50, 100, and
 1385 200-m isobaths) and major currents shown in black and the BEST-BSIERP regions in
 1386 blue. ACC: Alaska Coastal Current, BSC: Bering Slope Current, 1: AK Peninsula, 2:
 1387 South Inner Shelf, 3: South Middle Shelf, 4: South Outer Shelf, 5: Pribilofs, 6: Midnorth
 1388 Middle Shelf, 7: Midnorth Inner Shelf, 8: North Outer Shelf, 9: St. Matthews, 10: North
 1389 Middle Shelf, 11: North Inner Shelf, 12: St. Lawrence, 13: South Bering Strait, 14:
 1390 Norton Sound, 15: Off-shelf North, 16: Off-shelf Southeast. (Small subpanels) Spawning
 1391 initial locations released on the dates shown. Light grey: Contracted simulation, dark
 1392 grey: Base, Late, Early simulations, black: Expanded simulation. Polygons are
 1393 overlapping and all share the same western and southern boundaries.

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1395 **Figure 2.** Mean connectivity matrices by spawning time in the Base simulation.
1396 Connectivity is presented as the fraction of individuals in a given region when reaching
1397 40 mm length (columns) from a given spawning region (rows).

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1399 **Figure 3.** Mean connectivity in the Base simulation from the spawning regions AK
1400 Peninsula (A), South Middle Shelf (B), South Outer Shelf (C), Pribilofs (D), Midnorth
1401 Middle Shelf (E), North Outer Shelf (F), and Off-shelf Southeast (G) with Cold (blue)
1402 and Warm (red) years shown separately. Line thickness indicates the strength of
1403 connection; thick: strong connection >0.10 , thin: moderate connection $0.05-0.10$. Not
1404 pictured: weak connection <0.05 .

1405

1406 **Figure 4.** Mean connectivity matrices of the different simulations. The Base simulation is
1407 grouped by warm and cold years. Connectivity is presented as the fraction of individuals
1408 in a given region when reaching 40 mm length (columns) from a given spawning region
1409 (rows).

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1411 **Figure 5.** Fraction of individuals advected $<5^\circ$, $5^\circ-10^\circ$, or $>10^\circ$ longitude to the east by
1412 year in the Base simulation. Note differences in y-axis scales.

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1414 **Figure 6.** Fractions of 2007 observed T/S *Oshoro Maru* (top), BASIS surface trawl
1415 (middle), and BASIS water column acoustics (bottom) age-0 abundance and the fraction
1416 of modeled juveniles in the observed regions of each data type. Unsampled regions are
1417 white and zero abundance is marked with an X.

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1419 **Figure 7.** Model-*Oshoro Maru* differences for the Base simulation, represented as the
1420 fraction simulated in the observed regions only minus the fraction observed. Unsampled
1421 regions are in grey.

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1423 **Figure 8.** Model-BASIS surface trawl differences for the Base simulation, represented as
1424 the fraction simulated in the observed regions only minus the fraction observed.
1425 Unsampled regions are in grey.

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1427 **Figure 9.** Model-BASIS acoustics differences for the Base simulation, represented as the
1428 fraction simulated in the observed regions only minus the fraction observed. Unsampled
1429 regions are in grey.