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Six centuries of geomagnetic intensity variations recorded by royal Judean stamped jar handles

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The Earth's magnetic field, one of the most enigmatic physical phenomena of the planet, is constantly changing on various time scales from decades to millennia and longer. The reconstruction of geomagnetic field behavior in periods predating direct observations with modern instrumentation is based on geological and archaeological materials and has the twin challenges of 1) the accuracy of ancient paleomagnetic estimates and 2) the dating of the archaeological material. Here we address the latter by using a set of storage jar handles (fired clay) stamped by royal seals as part of the ancient administrative system in Judah (Jerusalem and its vicinity). The typology of the stamp impressions, which corresponds to changes in the political entities ruling this area, provides excellent age constraints for the firing event of these artifacts. Together with rigorous paleomagnetic experimental procedures, this study yielded an unparalleled record of the geomagnetic field intensity during the 8th – 2nd centuries BCE. The new record constitutes a substantial advance in our knowledge of past geomagnetic field variations in the southern Levant. While it demonstrates a relatively stable and gradually declining field during the 6th – 2nd centuries BCE, the new record provides further support for a short interval of extreme high values during the late 8th century BCE. The rate of change during this "geomagnetic spike" (defined as > 160 VADM ZAm²) is further constrained by the new data, which indicate an extremely rapid weakening of the field (losing ~27% of its strength over ca. 30 years).

archaeomagnetism | archaeointensity | levantine archaeomagnetism curve | paleosecular variation | archaeomagnetism spikes

1 Introduction

Reconstruction of geomagnetic secular variation during the Holocene has implications for various fields of research, from geophysics and other planetary sciences to biology and archaeology. Such reconstructions are based predominantly on heat-impacted geological and archaeological materials, whose thermal remanent magnetization (TRM) holds information on the geomagnetic field vector at the time of their last cooling. As evidence for fluctuating field behavior, including short (decadal) periods of rapid changes, is constantly growing (1-5), using records with excellent time resolution has become increasingly of interest.

In order to improve the accuracy and precision of age constraints associated with estimates of ancient geomagnetic field strength, the current study exploits a set of archaeological artifacts whose ages are exceptionally well constrained. This set is composed of well-studied ceramic jars from Judah/Yehud/Judea (Jerusalem and its vicinity), which bear royal stamp impressions on their handles (6-10). The stamped jars were part of the ancient administration of this region for about six hundred years, between the late 8th and late 2nd centuries BCE. As the types of stamp impressions changed with time according to the political situation, the jar handles provide an excellent record for geomagnetic intensity in the Levant during this time.

The geomagnetic intensity record of the Levant has recently improved with new data from Israel, Jordan, Syria and Cyprus (4, and references therein). These data indicate two very short episodes of extremely high field values (Virtual Axial Dipole

Moments [VADMs] in excess of 160 ZAm²) during the 10th and 8th centuries BCE respectively, which are referred to as the "Iron Age spikes" (2-4). However, as the unusually high field values, accompanied by apparently rapid changes in field strength, raise difficulties in core-flow models, the existence of the spikes have been questioned (11), and a scholarly debate has emerged (5, 12). Thus, an additional aim of the current study is to further investigate this phenomenon, using jar handles bearing successive seal types from the 8th c. BCE, the time of the later Iron Age spike.

2 Materials and Methods

2.1 Sampling

The focus of the current research is on royal Judean stamped jar handles that were found in surveys and excavations in Jerusalem and the hill country of Judah. As the archaeological context of these artifacts has no direct relation to the place of their firing (i.e., the location where magnetization was acquired), the entire assemblage is treated here as though coming from one central location in Judah. This location was chosen to be the archaeological site of Tel Sochoh (31.682°N, 34.975 °E), which several studies suggest was the production place of one of the major jar groups (the lmlk stamp type, 6, 7, 13). That said, as all of the stamped jars investigated in this study were produced within the boundaries of the political formations ruling the Judean region throughout the first millennium BCE (~31.2°N-32.2°N), the maximum expected uncertainty in estimated VADM is less than 1 ZAm².

Age estimates of the jar handles (Fig. 1, Table 1) are based on the typology of the stamp impressions found on them, which, except for one general type (the incised concentric circles), were

Significance

Understanding the geomagnetic field behavior in the past, and in particular its intensity component, has implications for various (and disparate) fields of research, including the physics of the Earth's interior, atmospheric and cosmologic sciences, biology and archaeology. This study provides substantial new data on variations in geomagnetic field intensity during the 8th – 2nd centuries BCE Levant, thus significantly improving the existing record for this region. In addition, it provides further evidence of extremely strong field in the late 8th century BCE ("geomagnetic spike"), and of rapid rates of change (> 20% over three decades). The improved Levantine record is an important basis for geophysical models (core-mantle interactions, cosmogenic processes and more) as well as a reference for archaeomagnetic dating.

Reserved for Publication Footnotes

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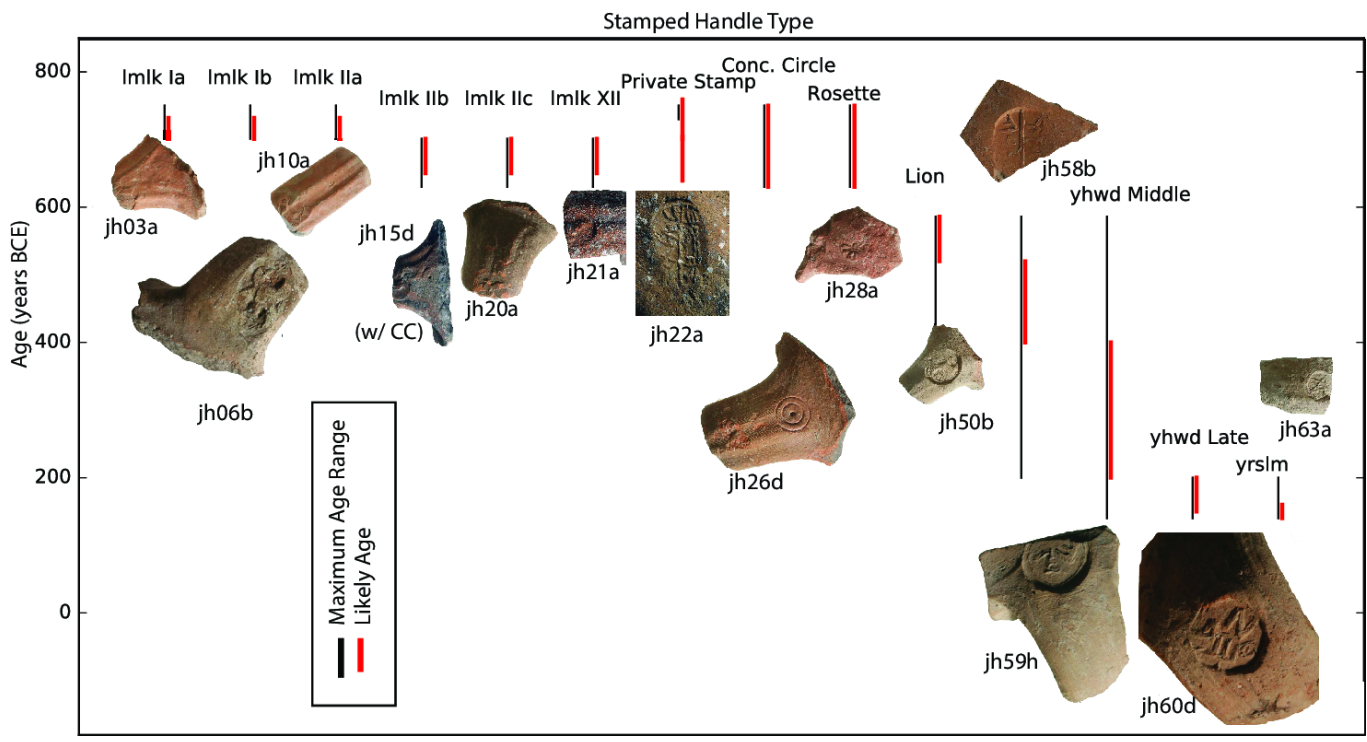


Fig. 1. Six centuries of royal Judean stamped handles: basic typology of the seal impressions and their ages (see Table 1 for references).

Table 1. Age ranges of the Judean stamped handled.

Stamp Type	Max Age Range (BCE)	Max Age References	Likely Age Range (BCE)	Likely Age References
Lmlk Ia	750-701	(37-45)	732-701	(6, 7, 34, the latter argues for a likely start date at ca. 715 BCE, 46)
Lmlk Ib	750-701		732-701	
Lmlk IIa	750-701		732-701	
Lmlk IIb	701-630	(38, 46)	701-650	
Lmlk IIc	701-630		701-650	
Lmlk XII	701-630		701-650	
Private Stamps	750-630		704-701	
*Conc. Circle Incisions	750-630	The dates refer to the firing of the jars (the incision was done after firing)	750-630	(40, 44, 47-49)
Rosette	630-586	(7, 8, 40, 50-53)	630-586	(40, 44, 47-49)
Lion	586-320	Limited stratigraphic evidence that the Lions do not persist to the end of the Persian Period	586-520	(54, 55)
Yhwd Early	586-200	(15, 56)	520-400	(57, 58)
Yhwd Middle	586-140	These types are found excavated with the previous type and both later types	400-200	(57, 58)
Yhwd Late	200-140	(59-61)	200-150	(57, 58)
Yrslm	200-140	(59, 60, 62, 63)	160-140	(64)

done by stamping a seal onto the wet clay just before firing. More than a century of research of these artifacts has resulted in good to excellent chronological constraints. These are based on their stratigraphic context (sharply confined by destruction layers at 701 and 586 BCE), stylistic considerations, the study of the script (Hebrew or Aramaic) and relevant historical events (e.g., 6, 7, 14, 15). While there is relatively broad scholarly agreement on the age ranges labeled “likely” in Table 1 (and used as a reference for our results), the maximum possible time intervals are also provided, with the references for the relevant literature.

Extensive detail about the artifacts used in this study is provided in the supplementary material (#1), including the context of their discovery, stamp impression typology, photographs and

references. Most of the handles used in this study were retrieved from the collections of the Ramat Rahel Expedition (16) and the Tel Sochoh survey (17). Each artifact, referred to here as a “sample”, is identified by a 5-character label that include the name of the study (JH = Judean Handles), the type/sub-type of the stamp impression (e.g., 50 = the lion type), and the sample running number (in letters). For the paleomagnetic experiments, five to six small (~2 mm) pieces were chipped from each sample. These chips are referred to here as “specimens” and are indicated by running numbers; for example, specimen JH50b3 is the third specimen from the second lion-type sample in this study.

2.2 Paleomagnetic experiments

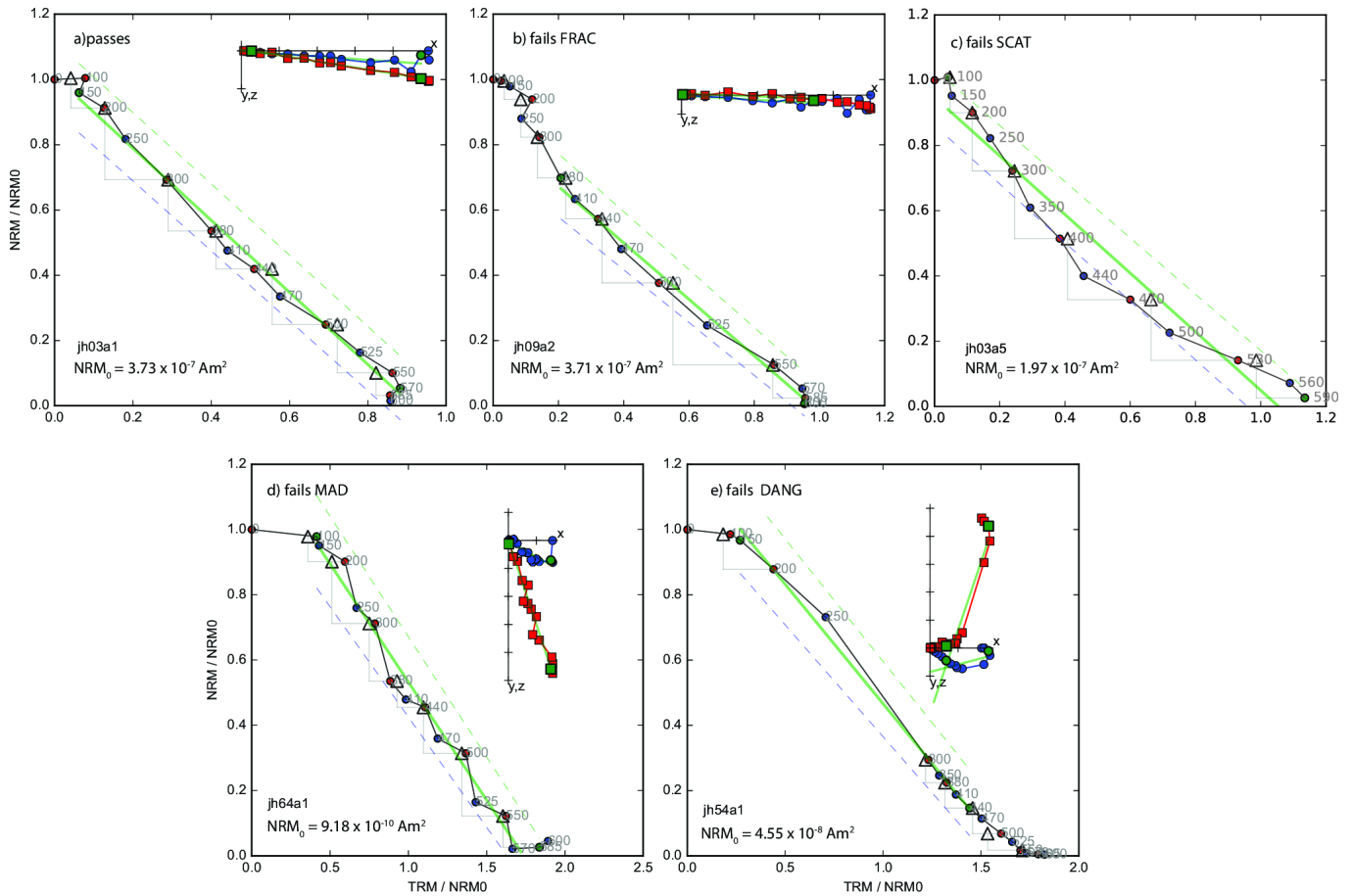


Fig. 2. Examples of behavior of specimens during the paleointensity experiment. Arai plots (Nagata et al., 1963) show NRM lost (NRM/NRM₀) versus pTRM gained (TRM/NRM₀). Blue symbols are from in-field cooling followed by zero-field cooling (IZ steps) and red symbols are from zero-field cooling followed by in-field cooling (ZI steps). Triangles are the pTRM check steps. Green line is the best-fit line through the data. The (absolute value of the) slope of this line multiplied by the laboratory field gives the ancient field value. The dashed lines are the 'SCAT' box. The insets are Zijderveld (1967) diagrams whereby the remanences measured after zero-field cooling are plotted as X,Y (blue circles) and X,Z (red squares). a) experiment passed all selection criteria. b) failed the FRAC criterion. c) failed the SCAT criterion. d) field the MAD criterion. e) failed the DANG criterion. (see Supplement for detailed description of the criteria used).

Paleointensity experiments were carried out in the Paleomagnetic Laboratory of Scripps Institution of Oceanography (SIO), University of California San Diego, using laboratory built computer-controlled paleomagnetic ovens and a 2G-SRM-760 3-axis superconducting magnetometer. Laboratory procedures and data analyses were done in the same manner as described in Shaar et al. (4). The procedure followed the IZZI protocol of Tauxe and Staudigel (18) with routine partial Thermal Remanent Magnetization (pTRM) checks at every second temperature step (19). A remanence tensor for anisotropy corrections was calculated from thermoremanent magnetizations (TRMs) acquired in six orthogonal positions, or with anhysteretic magnetizations (ARMs) acquired in nine position. Corrections for cooling rate effects were done assuming a logarithmic relationship between TRM overestimation from ratios of laboratory versus original cooling rates (20), and cooling time from 500°C to 200°C approximations of 0.1 hours, 3.7 hours, and 6 hours for the lab-fast, lab-slow, and ancient cooling times. In all experiments the field during 'in-field cooling' in the oven was 60 μT . Data analysis was done with the Thellier GUI program (21), which is part of PmagPy software (22), using the automatic interpretation technique described in detail in Shaar et al. (4, 23). The acceptance criteria follow Shaar et al. (4) and are described with references in the Supplementary Material.

3. Results

All data from our paleomagnetic experiments are provided in the MagIC online database (<https://earthref.org/MagIC/>). [Note to reviewers: the data will be available upon acceptance of the manuscript.] Out of 211 specimens, 158 passed the threshold values of the criteria used to establish paleomagnetic reliability (Supplementary Material #2), a success rate of 74%. This relatively high success rate for ceramic material (cf., 24), together with the strictness of the threshold values used in this study (cf., 25), demonstrates the high quality of the Judean jars as a paleomagnetic recorder.

Fig. 2 illustrates typical behavior of specimens during the paleomagnetic experiments. Most specimens have a single component magnetization and a blocking temperature compatible with magnetite. In addition, the original (or "natural") remanent magnetization (NRM) of the fired clay is relatively strong, in the range of $10^{-5} \text{ Am}^2/\text{kg}$, allowing the use of very small fragments ($\sim 20 \text{ mg}$) in the (destructive) archaeomagnetic experiments, which is important especially when working on rare archaeological materials such as inscribed clay.

Applying a minimum of 3 successful specimens and a maximum standard deviation of 3 μT or 8%, 27 out of the 67 samples measured yielded reliable paleomagnetic results (Table 2). These new data add to previously published geomagnetic intensity values for the Levant during the first millennium BCE (Fig. 3).

Table 2. Geomagnetic intensity results of samples with N≥3 and standard deviation ≤ 3 μT or 8%. N = number of successful specimens; Int. = intensity; VADM = Virtual Axial Dipole Moment.

Stamp Type	Sample	Specimens	N	Int. (μT)	Int. σ	VADM (ZAm ²)	VADM σ
lmlk Ia	jh03a	jh03a6;jh03a1;jh03a3	3	61.9	4.92	118	9.41
lmlk Ib	jh06b	jh06b1;jh06b2;jh06b3	3	84.1	2.98	161	5.7
lmlk IIa	jh12a	jh12a4;jh12a5;jh12a3	3	78	2.61	149	4.99
lmlk IIa	jh10a	jh10a3;jh10a4;jh10a5	3	71.4	2.38	137	4.55
lmlk IIb	jh15d	jh15d4;jh15d3;jh15d1	3	64.1	0.0595	123	0.114
lmlk IIc	jh20a	jh20a5;jh20a4;jh20a1;jh20a3	4	71.6	1.64	137	3.14
lmlk XII	jh21a	jh21a1;jh21a2;jh21a3;jh21a4;jh21a5	5	78.6	0.787	150	1.51
Private Stamp	jh24a	jh24a1;jh24a3;jh24a2	3	76.7	1.03	147	1.97
Private Stamp	jh24d	jh24d4;jh24d5;jh24d2;jh24d1	4	73	2.83	140	5.41
Private Stamp	jh24c	jh24c3;jh24c2;jh24c1;jh24c5	4	68.2	3.29	130	6.29
Conc. Circle	jh25b	jh25b3;jh25b5;jh25b4	3	65.9	1.44	126	2.75
Rosette	jh27a	jh27a2;jh27a3;jh27a1;jh27a4	4	72.3	0.0793	138	0.152
Rosette	jh28a	jh28a1;jh28a3;jh28a2	3	71.4	0.0779	137	0.149
Lion	jh55a	jh55a4;jh55a1;jh55a2	3	68.2	1.27	130	2.43
Lion	jh56a	jh56a4;jh56a2;jh56a3;jh56a1	4	64.7	0.119	124	0.228
Lion	jh57b	jh57b2;jh57b3;jh57b1;jh57b4	4	64.4	1.04	123	1.99
yhwd Early	jh58b	jh58b1;jh58b3;jh58b2;jh58b4	4	73.6	1.18	141	2.26
yhwd Early	jh58a	jh58a4;jh58a1;jh58a2;jh58a3	4	72.9	1.82	139	3.48
yhwd Early	jh58h	jh58h3;jh58h2;jh58h1;jh58h4	4	70.2	1.21	134	2.31
yhwd Early	jh58j	jh58j1;jh58j3;jh58j2;jh58j4	4	65.7	2.51	126	4.8
yhwd Middle	jh59l	jh59l4;jh59l2;jh59l3;jh59l1	4	70.3	0.0718	134	0.137
yhwd Middle	jh59e	jh59e4;jh59e1;jh59e3;jh59e2	4	66.7	0.0728	128	0.139
yhwd Middle	jh59h	jh59h2;jh59h3;jh59h1;jh59h4	4	59.9	4.7	115	8.99
yrslm	jh62a	jh62a4;jh62a3;jh62a2;jh62a1	4	56.1	0.0955	107	0.183
yrslm	jh65a	jh65a1;jh65a3;jh65a4	3	55.8	0.0533	107	0.102
yrslm	jh63a	jh63a2;jh63a3;jh63a1;jh63a4	4	50.9	2.89	97.4	5.53

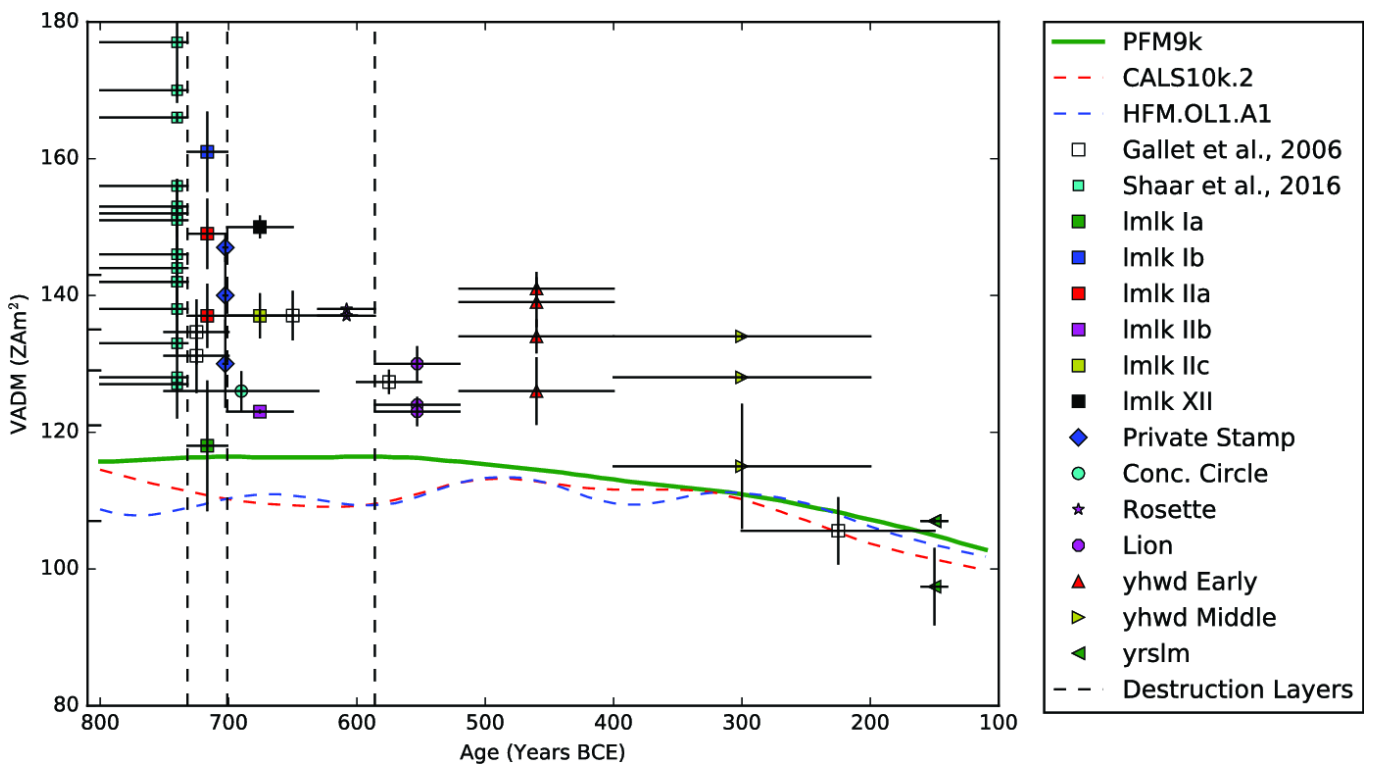


Fig. 3. Six centuries of geomagnetic intensity in the Levant (this study [Table 1], Shaar et al. (4) and Gallet et al. (28)). The reference curves (solid green, dashed red and blue lines respectively) are from PFM9K model of Nilsson et al. (26), and HOL.OL1.A1 and CAL510k.2 of Constable et al. (27). The vertical lines represent key chronological markers: the Assyrian campaign to the southern Levant in 734-732 BCE, the destruction of Judean cities by Assyria in 701 BCE and the destruction of Jerusalem by Babylon in 586 BCE. All data, including results of the current study, are available in the MagIC database (earthref.org)

4 Discussion

Our paleomagnetic experiments yielded excellent geomagnetic intensity values for all of the stamp impression types and sub-

types defined in Table 1 and shown in Fig. 1, except for one ("Late

545 yhw^d). The new data cover a period of ca. six hundred years, from
546 the late-8th to the late-2nd centuries BCE. In general, the results
547 indicate a gradual decrease in the field's intensity during the 7th
548 – 2nd centuries BCE, in agreement with the trends of the recent
549 paleosecular variation models PFM9K of Nilsson et al. (26) and
550 CALS10K.2 and HOL.OL1.A1 of Constable et al. (27), previously
551 published data of Gallet et al. (28). Following the peak, there
552 is a trough around 0 CE identified by Ben-Yosef et al. (29, ~77
553 ZAm² VADM). In general, however, it is evident that the secular
554 variation models predict significantly weaker fields and a much
555 smoother behavior than our data suggest.

556 Discrepancies between models and experimental data have
557 been observed in other recent publications of studies from the
558 southern Levant (e.g., 4), Cyprus (23) and other regions (e.g.,
559 30, 31), and they are most notable in the early Iron Age of
560 the Eastern Mediterranean (ca. 1200-700 BCE) when the field
561 fluctuated rapidly with intensity peaks reaching more than 150%
562 of the model-predicted values (cf. Fig. 3 for the 8th c. BCE). As the
563 models are based on the extensive data published over decades
564 of research, it is evident that they are smoothed by “noise” in
565 the data. These source of noise includes both faulty intensity es-
566 timations (inappropriate experimental protocol and/or selection
567 criteria) and erroneous dating. The latter issue has been under-
568 appreciated until recently, when more collaborative projects were
569 introduced and effort began in tackling the intricate problem
570 of dating archaeological contexts and artifacts. Thus, the next
571 generation of models need to take into account regional datasets
572 that were scrutinized for quality of their individual samples. The
573 Levantine curve presented here (Fig. 3a) includes only such data,
574 and our research on the Judean stamped jar handles underscores
575 the advantages of working with inscribed clay materials to tackle
576 the dating issue.

577 In addition to the “noise” in the database, rapid secular varia-
578 tions are not represented in the geomagnetic field models because
579 of their extremely short durations. To detect rapid changes such
580 as those observed for the 8th c. BCE southern Levant (Fig. 3) it
581 takes an extensive quantity of data obtained from materials that
582 represent a time-sequence of only several decades. Not only are
583 such efforts rare in common archaeomagnetic research, but the
584 archaeological record itself is often not continuous and biased
585 towards major events of destructions or abandonment. Several
586 ways to overcome this issue have been suggested in previous
587 research, including working with materials from waste piles and
588 industrial debris (2, 29).

589 Our new data support the existence of an interval of extremely
590 high field intensity during the late 8th century BCE. These high
591 values are in agreement with recently published data by Shaar et
592 al. (4) and represent one of the Levantine Iron Age “geomagnetic
593 spikes”. These anomalies, first reported by Ben-Yosef et al. (2),

594 were defined by Cai et al. (32) as “a sharp increase in the field
595 intensity to more than twice the present value (~160 ZAm²
596 VADM) in less than 500 years”. Following this definition and the
597 current data available for the Levant (4), there is evidence for
598 at least two such spikes, one during the 10th century BCE (cf., 2,
599 3; note that evidence of a 9th century BCE spike failed the more
600 rigorous selection criteria applied in the current study, 30) and the
601 other during the 8th century BCE. Both the 10th century and 8th
602 century BCE spikes occurred during a time span of generally high
603 field values world-wide (33), which appears to promote rapidly
604 fluctuating and unstable fields (see discussion in 3). The data of
605 the current study add new information on the 8th century BCE
606 spike, as it provides strong evidence of the rapidly decreasing
607 intensity over the interval after 732 BCE, an interval not covered
608 by previous studies (Fig. 3). Age constraints from archaeological
609 contexts and stamped jar handles during the second half of the
610 8th BCE southern Levant are exceptionally tight, as the region
611 was influenced by Assyrian interventions that resulted in excellent
612 chronological markers in the archaeological record (10). These
613 include military campaigns that left destruction layers of the
614 major Israelite and Judahite cities (in 734-732/722-720 BCE and
615 701 BCE respectively, Fig. 3). Moreover, the interaction with
616 Assyria and preparation for possible conflicts had direct bearing
617 on the administration of Judah, which is reflected in changes in
618 the stamp impressions on the jar handles (Table 1, and refer-
619 ences therein). Thus, the data indicate a sharp drop of ~27% in
620 field intensity over 31 years (732 – 701 BCE), or – if accepting
621 Na’aman’s (34) chronology – over 14 years (715-701 BCE). This
622 well-constrained time interval of the decaying 8th century BCE
623 spike is important new evidence that should be taken into account
624 as part of the ongoing discussion on this phenomenon, its sources
625 and its effects (e.g., 11, 12, note that the rates here are around
626 ~0.75/1.5 μT/year, within the limits of the suggested models).

627 Recently, more evidence of extremely high field values
628 around the time of the Levantine Iron Age spikes (~3000 BP)
629 was found in nearby regions, including Turkey (35) and Georgia
630 (36). Altogether, the available data suggest that this is a re-
631 gional phenomenon, similar in scale to the current South Atlantic
632 Anomaly (cf., 4); however, the exact geographic expanse of this
633 phenomenon has yet to be investigated, and the fact that these
634 are very short lived features that can be easily missed suggest that
635 there is much more to discover. As demonstrated here, special
636 archaeological materials such as inscribed clay are one of the keys
637 for increasing time resolution in future research.

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