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ABSTRACT

The California Current Ecosystem Long Term Ecological Research (CCE-LTER) site has been in existence **02** since 2004. One of its primary objectives is to understand the response of the southern California Current ecosystem to climate forcing. The CCE-LTER site cooperates with the California Cooperative Oceanic Fisheries Investigations (CalCOFI) program and complements CalCOFI's work through more extensive observations, process studies, and a modeling program. This special issue is focused on the long-term observations made by the CCE-LTER and CalCOFI programs, describing and understanding long-term changes in the physical, chemical, and biotic environment in the region. The papers in this issue highlight the climatological conditions during recent years and employ modeling to diagnose the principal forcing of meridional currents and eddy transport, both of which affect biotic responses. Changes in source waters in the region, and altered flushing of the Santa Barbara Basin, are considered. Temporal variations in inherent optical properties and in higher trophic levels, including seabirds and marine mammals, are presented. Key methodological developments presented include the incorporation of subsurface phytoplankton and light distributions in order to improve remotely sensed measures of primary production, and the validation of multi-frequency acoustic estimates of mesopelagic fish biomass. Results also highlight significant spatial differences across the CCE-LTER region, including cross-shore trends in microbial assemblages, and glider-resolved frontal features and zones of mixing associated with abrupt topography. Alterations to the spatial structure of the pelagic ecosystem must also be considered when evaluating future climate-related changes.

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A central question environmental scientists face today is: how will natural ecosystems respond to climate variability and climate change? One of the important resources available to address this question is the network of Long-Term Ecological Research (LTER) sites supported by the U.S. National Science Foundation. The LTER network includes 25 sites spanning biomes from the Arctic to the Antarctic and from the pelagic ocean to the Rocky Mountains. At all sites long-term measurements are carried out in five core areas: disturbance patterns, primary production, movement of inorganic nutrients, organic matter formation and decomposition, and studies of key populations. In addition, each site conducts intensive hypothesis-based research designed to understand processes that structure populations or mediate the fluxes of energy and matter between trophic levels. These observations and experiments are tied together with modeling at levels, ranging from organisms to the ecosystem (Robertson et al., 2012).

The California Current Ecosystem LTER (CCE-LTER) site was established in 2004. The site builds on the foundation of the California Cooperative Oceanic Fisheries Investigations (CalCOFI) program (Ohman and Venrick, 2003), as well as a suite of independent measurements. CalCOFI has made observations in the California Current System (CCS)

since 1949. These observations were initially focused on small pelagic fishes in the CCS, Pacific sardine and northern anchovy, and their physical and biological environment. The 65-plus year time series has been invaluable for characterizing the response of the CCS to changing ocean climate. However, CalCOFI is primarily an ocean observing program, while CCE-LTER seeks to understand the underlying mechanisms that drive the response of the system to changing ocean climate. The central motivating questions of the CCE-LTER program are: What are the mechanisms leading to different ecosystem states in a coastal pelagic ecosystem? What is the interplay between changing ocean climate, community structure and ecosystem dynamics? Nested within these overarching questions, CCE-LTER is currently focused on the specific mechanisms leading to abrupt transitions in ecosystem state **03**81 (Ohman and Venrick, 2003).

CCE-LTER uses three basic approaches to address these questions: time series observations, process studies, and modeling. One foundation of the time series observations is the CalCOFI program that currently surveys 70-85 stations off southern and central California four times a year, supplemented by several types of autonomous measurements. CCE-LTER augments CalCOFI observations to further characterize the biogeochemistry and the population structure of the pelagic ecosystem. Related programs make observations of marine mammals and seabirds. The results of these efforts

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are time series of differing length. The time series of core CalCOFI data, salinity, temperature, oxygen, zooplankton biomass and abundance, and ichthyoplankton abundance is, at the time of this writing, 66 years long. Since 1984 inorganic nutrients, chlorophyll *a* and primary production have been measured consistently. Seabirds have been enumerated since 1987 and marine mammals since 2002. The CCE-LTER program started making its observations in the fall of 2004. The main variables observed by the CCE program are concentrations of dissolved and particulate organic matter, bacterial biomass, and the biomass and population structure of phyto- and zooplankton using a variety of approaches.

CCE time series observations also include a series of autonomous measurements from *Spray* ocean gliders that operate continuously along two lines off Southern California, two interdisciplinary moorings, one in the core area of the CC and a second one off Pt. Conception, and remotely sensed data (Ohman et al., 2013b).

2. CCE-LTER special issue

The present issue of Deep-Sea Research II addresses some of the new insights that have been gained in the time series component of the CCE-LTER program since the advent of the CCE site. We do not attempt to summarize the results of other elements of the CCE program (Franks et al., 2013; Ohman et al., 2013a), but refer the reader to the CCE site bibliography for this purpose (http://cce.lternet.edu/publications).

28 A multidisciplinary approach is at the core of the study of 29 ecosystems over long periods of time by members of the LTER 30 network. This is illustrated by the diverse nature of the articles in 31 this volume. Miller et al. (2014) provide the physical framework 32 for the papers in this issue. They first summarize the oceano-33 graphic and climatological conditions in the CCS, describe changes 34 in the perspectives in CCS physical processes and lastly, using the 35 Regional Ocean Modeling System (ROMS), hindcast the physical 36 oceanographic evolution during the CCE Process cruises. Over the 37 last 10 years physical conditions in the CCE varied widely. The El 38 Nino/Southern Oscillation (ENSO) varied between warm, neutral, 39 and cool states over this time period, never reaching extremes. The 40 North Pacific Gyre Oscillation (NPGO) was strongly positive over 41 most of the time period, which, in our region, is indicative of 42 enhanced upwelling. Remarkably cool conditions were observed in 43 the CCS from late 2007 to early 2009 and from mid-2010 through 44 2012, which is consistent with the negative state of the PDO and 45 the positive state of the NPGO during those intervals. Miller et al. 46 (2014) note that the most important conceptual breakthroughs in 47 the understanding of the dynamics of the CCS were: (a) the 48 identification of the NPGO, which appears to control many aspects 49 of the CCS such as salinity, chlorophyll, nutrients and oxygen; (b) 50 the identification of a class of energetic small-scale variations in 51 the upper ocean that are now referred to as sub-mesoscale 52 variations that may enhance lateral and vertical mixing processes; 53 (c) the realization that offshore Ekman pumping by wind-stress 54 curl is equally important as coastal Ekman upwelling in supplying 55 nutrients to the CCE; and (d) that sub-thermocline concentrations 56 of oxygen have been declining over the last decades, driven likely 57 by increasing stratification in the CCS and/or by changes in the 58 oxygen content in the source of these deep waters that are 59 advected into the region. The ROMS hindcasts of physical condi-60 tions during the CCE Process cruises showed that these cruises 61 encountered widely varying climate states and physical condi-62 tions. Surface current fields for the 2011 and 2012 cruises will be 63 used as a physical framework for the interpretation of the in situ 64 biological measurements made during the process cruises.

Our ship-based observations are too coarse in space and time to resolve the physical processes that drive the CCE, in particular at the

mesoscale. Davis and Di Lorenzo (2014a) utilize an eddy-resolving 67 Regional Ocean Modeling System (ROMS) primitive equation model 68 69 with terrain following coordinates to diagnose the primary pro-70 cesses forcing meridional currents and eddies in the California Current System (CCS) over the period from 1950 to 2008. First, the 71 authors analyze alongshore currents in the California Current 72 73 System. They consider the relative importance of local wind-stress curl (WSC) as contrasted with remote coastally-trapped waves 74 (CTW) originating in the equatorial region, as dominant sources of 75 the variability of mesoscale circulation and meridional transport at 76 interannual and longer time scales. This is a problem to which 77 modeling is well-suited because even the CalCOFI data set does not 78 79 have sufficiently closely spaced stations to resolve the meandering finer-scale currents that would be linked to the wind-stress curl. 80 Davis and Di Lorenzo (2014a) address this issue using two model 81 configurations, both with and without coastal trapped waves of 82 equatorial origins, to diagnose the equatorward meridional flow of 83 the California Current. They focus on the region offshore of the 84 narrow continental shelf. The authors find that while CTW explain 85 much of the variation in sea-surface height in the California Current, 86 it does not explain the variation in alongshore currents and 87 88 associated transports, which instead are more closely related to wind-stress curl. This result is important for understanding the 89 transport of nutrients, plankton, and other particles throughout the 90 California Current System. 91

Davis and Di Lorenzo (2014b) illustrate another advantage of 92 93 models in extending the observational record. Approximately 20 94 years of sea-surface height (AVISO) altimetry data were available to describe the mesoscale eddy field of the California Current, but 95 nearly 60 years of modeled sea-surface height. Using the longer 96 modeling record, they address the extent to which mesoscale eddy 97 variability in the southern and northern CCS is controlled by 98 variations in the wind-stress curl. The wind-stress curl gradient 99 forces much of the mesoscale eddy activity in the northern sector, 100 but relatively little of the eddy variability in the southern sector. 101 They suggest that the eddies in the southern California (CCE-LTER) 102 region are more related to intrinsic variability in the system. They 103 hypothesize that southern California eddies are less directly linked 104 to the winds because of the longer residence time of the eddies, 105 hence greater opportunity for the eddies themselves to influence the 106 background mean conditions, generating stronger feedbacks and 107 lower predictability. The inference that there is greater intrinsic 108 variability in eddy activity in the CCE-LTER region suggests that 109 predicting future changes in eddy dynamics will be challenging. 110

Changes in hydrographic and chemical properties at depth in the 111 CCS have been reported previously (Bograd et al., 2008; McClatchie 112 et al., 2010). Bograd et al. (2014) analyze 28-year trends in dissolved 113 nutrients and dissolved oxygen in source waters entering the CCE-114 LTER region to identify drivers of observed changes. Dissolved O₂ has 115 been decreasing and nitrate, phosphate, and silicic acid have gen-116 erally been increasing throughout the region at a density surface 117 σ_{θ} =25.8, corresponding to the upper pychocline region. At a deeper 118 density surface (σ_{θ} =26.5) in the lower pycnocline, O₂ has also been 119 decreasing and nitrate and phosphate increasing, but changes in 120 121 silicic acid are more complex. Silicic acid has been increasing in the 122 offshore region but decreasing in the nearshore region influenced by waters of the California Undercurrent coming from the south. This 123 altered nutrient content of waters in the California Undercurrent is 124 also expressed as a decline in the Si:N ratio throughout most of the 125 region. There has also been a decline in the N:P ratio in deeper 126 waters but an increase in N:P ratio in shallower waters. The 127 agreement of the sign and magnitude of the shallower water trends 128 with previously documented trends in the western and eastern 129 130 subarctic Pacific (Whitney et al., 2013) suggests that these trends are linked to basin-scale climate forcing. The authors also use global 131 132 ocean reanalysis products to illustrate the time lags necessary

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for circulation to propagate anomalies into the CCE-LTER region.
They call attention to the possible role of altered Si:N in altering
phytoplankton community composition.

Changes in oxygen and nutrients at depth have the potential to 4 5 affect the biogeochemistry of the basins of the California Border-6 lands. Goericke et al. (2014) study the unusual biogeochemical 7 conditions at the bottom of the Santa Barbara Basin that have been 8 observed intermittently since 2006. During this time extremely 9 low concentrations of nitrate and extremely high concentrations of 10 nitrite were observed. 10 times higher than those observed 11 previously. Such conditions are unprecedented relative to the 12 1984 to 2005 time period, suggesting that fundamental changes 13 in the forcing of the bottom waters of the basin have occurred. The 14 loss of nitrate from the system is driven by denitrification, which 15 requires low concentrations of oxygen. Oxygen is replenished in 16 the basin when it is flushed by dense waters that are raised above 17 the sill of the basin by upwelling, and it declines during other 18 times due to the remineralization of organic matter in the water 19 column and in sediments. Indeed, the frequency of low-oxygen 20 events in the bottom waters has increased. Increased rates of 21 primary production and thus increased rates of sedimentation of 22 organic matter to the bottom of the base are unlikely causes. 23 Rather these changes are due to the decreased frequency and/or 24 extent of flushing of the basin combined with a decrease of oxygen 25 concentrations in these waters that enter the basin. The inter-26 mittent, dramatic increases of nitrite at the bottom of the basin 27 likely represent a tipping point in the biogeochemical system 28 driven by decreasing concentrations of oxygen in the bottom 29 waters. This study thus illustrates that gradual changes in the 30 drivers of a system can lead to a highly non-linear response of its 31 properties, in this case concentrations of nitrite.

32 Near the beginning of the CCE-LTER program, a regular series of 33 cross-shore sampling lines was established using the Sprav ocean 34 gliders (Sherman et al., 2001), which now includes CalCOFI lines 35 80 and 90 in the CCE region and line 66.7 off Monterey Bay. 36 Johnston and Rudnick (2014) analyze an extensive series of six 37 years of measurements of the cross-shore patterns in mixing, as 38 inferred from Acoustic Doppler Profiler-measured shear and CTD-39 derived strain. They uncover a striking contrast between the 40 patterns of mixing along the two more northern lines in compar-41 ison with the southernmost line that transits the southern 42 California Bight and region offshore. Mixing is elevated close to 43 shore along the two more northern lines and decays with distance 44 offshore, while mixing along line 90 is elevated in a region located 45 approximately 200 km from shore, near the Santa Rosa-Cortes 46 Ridge. This zone of offshore elevated mixing, associated with 47 rough topography, is linked to dissipation of both diurnal (D1) 48 and semidiurnal (D2) internal tidal energy. The biological and 49 biogeochemical consequences of such topographically-linked mix-50 ing have not been fully explored.

51 The same glider series, from lines 80 and 90, was indepen-52 dently analyzed by Powell and Ohman (2014) in order to under-53 stand the extent of co-variability of biological and physical fronts. 54 While fronts tend to be located closer to shore along line 80, they 55 tend to be further offshore along line 90, near the region of altered 56 mixing identified by Johnston and Rudnick (2014). Horizontal 57 density fronts are regions of marked changes in zooplankton 58 acoustic backscatter and phytoplankton fluorescence, a biomass 59 proxy. Frontal zones also tend to be regions of enhanced biomass 60 of phytoplankton and mesozooplankton, as inferred from the 61 glider-based sensors. These authors also illustrate the enhanced 62 foraging success that a mobile predator could experience by 63 searching along a frontal gradient. Because satellite imagery 64 suggests that the incidence of Chl-a and sea-surface temperature 65 fronts in the Southern California region have increased over the 66 past 14-29 years (Kahru et al., 2012), the processes occurring at fronts are expected to become more important in the future. Such glider-derived results have contributed to making frontal processes a central focus of the current CCE-LTER experimental process cruise studies (Landry et al., 2012; Ohman et al., 2013a)

The CCE-LTER program relies primarily on observations made by ships, gliders or moorings, which are either physically located in one place or move slowly through the ocean, not allowing for a synoptic view of the system. Observations made by satellites do not suffer from this shortcoming. However, merging observations made by different satellites and relating these to in situ ocean properties is often difficult. Kahru et al. (2013) address these issues. They estimate inherent optical properties in the California Current from observations made by four different ocean color sensors (OCTS, SeaWiFS, MODISA and MERIS). Using inversion models they derive a 16-year time series for the phytoplankton absorption coefficient at 440 nm, absorption by dissolved and detrital organic matter at 440 nm, and the particle backscattering coefficient at 490 nm. The uncertainty of these derived properties is still high due to a limited number of in situ match-ups, their uneven distribution in space, and uncertainty associated with the raw sensor data. However, the time series show very interesting patterns. Along a wide band along the coast the phytoplankton absorption coefficient, a proxy for phytoplankton biomass, has been increasing over the last 16 years (1996-2012). In contrast, a significant decrease has been observed in the oligotrophic North Pacific gyre. The trend along the coast is positively correlated with increasing wind speeds, related to upwelling, suggesting likely drivers for these trends.

Over the last decade remote ocean color observations have incr-95 easingly being used not only as a proxy for chlorophyll a (Chl a) but 96 97 phytoplankton carbon biomass, depth-integrated primary production (PP) and export production. For the estimation of the latter 98 99 two variables the subsurface distribution of phytoplankton biomass and the physiology of the resident populations are potentially 100 important. Jacox et al. (2013) ask how sensitive models of PP based 101 on remotely sensed data are to the parameterization of phytoplankton 102 physiology and the subsurface Chl *a* and light fields. The authors base 103 their study on the 1985–2011 CalCOFI record of Chl a and primary 104 production. The authors only find modest improvement to model-105 based estimates of PP when measured surface Chl a is substituted for 106 107 remotely sensed surface Chl a. However, dramatic improvements are 108 observed when in situ depth-resolved data for Chl a and light are used. The authors point out that today such in situ data are easily 109 110 obtained from gliders and potentially from Bio-Argo floats. The aut-111 hors show that the inclusion of glider data in the PP models substantially improves estimates of PP and suggest that this be done on a 112 routine basis. 113

Taylor et al. study microbial community patterns in the CCE-114 115 LTER region using a set of samples collected between 2004 and 2010 on 25 CalCOFI cruises as part of the CCE augmented 116 sampling. They use flow cytometry and digital epifluorescence 117 microscopy to understand the spatial and seasonal variations of 118 different components of the autotrophic and heterotrophic plank-119 120 ton $< 200 \,\mu$ m. Although diatom biomass was elevated nearshore 121 along line 80 and in the Santa Barbara Basin, a surprising finding is the relatively large contribution of autotrophic dinoflagellate 122 biomass to the microplankton assemblage at all times of year. In 123 some years (e.g., 2005-2006) dinoflagellate biomass exceeded 124 diatom biomass, even in the upwelling regions. Autotrophic 125 picoplankton (principally the cyanobacteria Synechococcus and 126 Prochlorococcus) were ubiquitous, but their abundance less vari-127 able across the region. The size composition of the assemblage was 128 consistently dominated by larger microplankton ($> 20 \mu m$) in the 129 130 vicinity of Pt Conception, while the nanoplankton $(2-20 \,\mu m)$ dominated the autotrophic assemblage everywhere else, including 131 132 along line 90. The cross-shore patterns were markedly different

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along line 90, where the microplankton did not dominate biomass nearshore; peak biomass was instead found approximately 190 km from shore. This offshore peak in autotrophic biomass coincides spatially with the zone of enhanced mixing documented by Johnston and Rudnick (2014). On average, living cells (phytoplankton, heterotrophic protists, and heterotrophic bacteria) are found to comprise about half of the total particulate organic carbon present. Taylor et al. also find that the ratio of autotrophic C: Chl-*a* varies in a systematic manner with depth in the euphotic zone and, in addition, with the depth of the nitracline. These relationships will be especially valuable for CCE modeling efforts.

12 The last three papers of this volume study some of the higher 13 trophic levels of the CCE: mesopelagic fishes, marine mammals, 14 and seabirds. The importance of these trophic levels for our 15 understanding of the ecosystem can be twofold. Such organisms 16 may be important mediators of fluxes of energy and matter, as has 17 recently been shown for mesopelagic fishes (Davison et al., 2013). 18 These organisms may also be sensitive indicators of ecosystem 19 state, because they tend to integrate changes in the ocean 20 environment over larger space and time scales. However, estimat-21 ing the biomass and abundance of these groups poses special 22 challenges as the organisms are often sparsely and patchily 23 distributed in space and time. Davison et al. (2014) use both 24 trawling and active acoustics to estimate the biomass of mesope-25 lagic fish in the CCE. The uncertainty associated with trawls is the 26 trawl capture-efficiency of the fish which can vary with size, 27 species, and time of day. Trawls are required for the validation of 28 acoustic estimates of biomass, thus impacting the accuracy of 29 acoustic estimates as well. The authors show that variations in 30 trawl capture efficiency can affect acoustic estimates of biomass 31 estimates 5-fold. In addition differences in acoustic backscattering 32 among fishes (e.g., the presence and size of a swimbladder) and 33 unknown mixtures of populations with dramatically differing 34 backscattering characteristics also add uncertainty to such acous-35 tic estimates. The authors carefully correct their estimates of 36 mesopelagic fish biomass for such biases and quantify the uncer-37 tainty associated with these estimates. They conclude, consistent 38 with their earlier studies (Davison et al., 2013), that the biomass of 39 mesopelagic fishes of southern California is $25-37 \text{ g m}^{-2}$. Such 40 high levels of biomass suggest that these fish have a significant 41 role in the ecosystem. A future question is how this role will 42 change as their biomass has been decreasing over the last decades 43 (Koslow et al., 2011).

Cetaceans are among the charismatic species of the marine 44 45 environment and the CCS provides a habitat for numerous taxa. 46 They are potentially impacted by human activities such as fishing, 47 shipping, and naval operations. Campbell et al. (2014) use visual 48 surveys carried out on CalCOFI cruises to access distributions and 49 trends in the abundance of the six most commonly encountered 50 cetacean species off southern California over the last decade. The 51 unique aspect of this study is the long-term continuous observations 52 that allow the analysis of trends with time. Blue whales, fin whales, 53 and humpback whales were the most abundant baleen whales. 54 Short-beaked common dolphins, Pacific white-sided dolphins, and 55 Dall's porpoise were the most frequently encountered small ceta-56 ceans. Significant spatial and seasonal variability is observed for most 57 species. A generalized additive model is used to estimate long-term 58 trends accurately. No significant long-term trends are observed for 59 blue whale, fin whale, humpback whale, short-beaked common 60 dolphin or Dall's porpoise abundance. Pacific white-sided dolphins 61 showed a significant decrease over the 10-year study period. 62 Environmental parameters that correlated with this decrease are 63 not identified; complicating the interpretation is the presence of 64 northern and southern subpopulations of this species. The data that 65 were collected as part of this study will be used to model and predict the habitat of these species. This model will be used by the US Navy 66

off Southern California to minimize adverse impacts of their operations on these species.

69 Sydeman et al. (2014) analyze trends and interannual varia-70 bility in seabirds, enumerated for the last 25 years at sea on CalCOFI cruises. They ask whether variability in seabird densities 71 72 covaries with euphausiids (krill) or forage fish abundance, and whether seabirds and their prey are affected by changes in ocean 73 stratification and upwelling. The authors use the CalCOFI ichthyo-74 plankton data as an index for the abundance of the adult fish that 75 spawned them, as others have before them. They find an overall 76 decline in seabirds (combined abundance of 23 species), averaging 77 -2.2% vr⁻¹ over the past 25 years. The decline was more pro-78 nounced along the northern line 80 than on the more southern 79 line 90. The decline in seabirds occurred despite a temporal 80 increase in abundance of euphausiids (Thysanoessa spinifera and Euphausia pacifica), and is instead related to the decrease in abundance of nearshore fishes, especially northern anchovy. However, in contrast to the longer term trends, Sydeman et al. (2014) find that detrended interannual variability in springtime seabird density is best explained by the density of euphausiids (T. spinifera and E. pacifica) measured in springtime plus the density of larval fish measured in the preceding winter. They also suggest that seabird density may have become more closely linked to euphausiid density after 1999. Overall, their results suggest that the effects of ocean climate on seabird abundance are mediated through the availability of their prey.

93 In summary, the papers in this special issue provide an introduction to some of the diverse responses of the ecosystem, 94 studied by the CCE-LTER, to changing ocean climate. Many of the 95 studies are based on CalCOFI data and highlight the distinct 96 advantages to linkage of the LTER site with the deep historical 97 perspective provided by the 66-year long CalCOFI time series. The 98 studies illustrate the importance of conducting not merely single 99 point measurements over time, but rather space-resolving time 100 101 series that resolve spatial patterns on a regional scale. Especially in dynamic Eastern boundary current ecosystems, there are impor-102 tant cross-shore and regional variations that must be considered in 103 understanding susceptibility to climate forcing. This volume also 104 illustrates the importance of autonomous measurements, multi-105 frequency acoustics, and satellite remote sensing, in uncovering 106 key elements of frontal dynamics, ocean mixing, and differential 107 responses of the nearshore and offshore ocean to climate forcing. 108 This collection of papers also illustrates the vital role of integration 109 of different measurement approaches. Each method, whether 110 shipboard, robotic vehicles, or remote sensing, has advantages 111 and disadvantages, and the greatest understanding comes from 112 combinations of diverse measurements spanning different time-113 space domains. Multiple measurement approaches also permit 114 independent validatation. 115

In some cases, the new time series developed here are too short to discern and interpret temporal trends. However, as these observational programs continue into the future and continue to be assimilated into models, we have a framework for understanding both natural climate variability and the anthropogenically induced forcing that will alter the pelagic ecosystem in the future.

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