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Modeling the Effects of Dimerization and Bulk Diffusion on the Evaporative Behavior of Secondary Organic Aerosol Formed from α -Pinene and 1,3,5-Trimethylbenzene

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- ¹ Modelling the Effects of Dimerization and Bulk
- ² Diffusion on the Evaporative Behavior of Secondary
- ³ Organic Aerosol Formed from α -Pinene and 1,3,5-

4 Trimethylbenzene

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16 ABSTRACT: Volatility determines the gas-particle partitioning of organic compounds. Volatility 17 is thus a key property needed to understand the behavior of organic aerosol (OA) in the atmosphere. 18 Various studies have been conducted to experimentally measure and numerically simulate 19 distributions of OA volatility. The observed OA evaporation rates have generally been slower than 20 the rates assuming instantaneous gas-particle equilibrium and volatility estimated from secondary 21 organic aerosol (SOA) formation experiments. Particle phase diffusion and/or low volatility 22 compounds, such as oligomers and highly oxygenated molecules, could limit the evaporation of 23 OA, though the relative contributions of these factors are still uncertain. In this study, we 24 conducted model simulations using a volatility basis set framework with consideration of kinetic gas-particle partitioning, formation and dissociation of dimers, and particle-phase diffusion to 25 26 reproduce observed evaporative behaviors of SOA formed from α-pinene ozonolysis and 1,3,5-27 trimethylbenzene (TMB)/NO_x photooxidation. Based on simulations constrained by various 28 volatility distributions derived from chemical analysis or heating experiments, we found that both 29 dimerization and slow particle-phase diffusion contributed to the observed slow evaporation under 30 dry conditions. In contrast, particle-phase diffusion did not practically inhibit SOA evaporation 31 under humid conditions. The similarity of the fitted parameters, including dimer 32 formation/dissociation rates and bulk diffusivity, for SOA from α -pinene and 1,3,5-TMB under 33 dry conditions suggested that these processes are important for both monoterpene and aromatic 34 SOA. Evaporation rates of SOA from α -pinene in this study were slower than the rates reported in 35 previous experimental studies. This difference could be partly explained by differences in the 36 experimental setups, including the treatment of organic vapors.

37 KEYWORDS: secondary organic aerosol, dilution-induced evaporation, dimer formation, bulk
38 diffusion, kinetic model, aerosol particle viscosity, vapor wall loss.

39 1. Introduction

Organic aerosol (OA) represents a large fraction (20–90%) of submicron aerosol,¹ and thus, 40 41 understanding the behavior and sources of OA is necessary. For accurate representation of OA by 42 numerical simulation models, we need to understand both chemical and physical properties of OA. 43 For chemical properties, there is a large diversity in the composition of OA but not all constituents of OA have been identified.^{2, 3} Although there are several approaches to explicitly simulate OA 44 45 composition,⁴⁻⁷ the computational burden of these explicit simulations is large and OA is 46 represented by grouping or lumping compounds. To simulate secondary organic aerosol (SOA), a 47 wide variety of organic compounds have been classified based on their physicochemical properties, 48 such as volatility (or vapor pressure), oxidation state, carbon number, or number of functional groups.⁸⁻¹¹ In addition to semi-volatile organic compounds (SVOC), highly oxygenated organic 49 molecules (HOMs)¹²⁻¹⁴ and oligomers,^{15, 16} both of which can be extremely low volatility or low 50 51 volatility organic compounds (ELVOC or LVOC, respectively), make significant contributions to SOA. The formation of these compounds has thus been introduced in some SOA models.¹⁷⁻²⁰ 52 53 Physical properties, particularly volatility and bulk diffusion coefficients, are also important in 54 accurately representing OA, because these processes strongly influence the gas/particle 55 partitioning process. In many chemical transport models, rapid equilibrium between the gas and 56 particulate-phases is assumed for organic compounds,^{21, 22} though studies during the last decade 57 have suggested that this assumption may not hold for all conditions and systems. For example, 58 evaporation of laboratory-generated or ambient SOA after dilution or scavenging of vapors are associated with a timescale longer than several hours.^{23, 24} Trump and Donahue ¹⁷ and Kolesar, et 59 al.¹⁸ have indicated that oligomer formation and its slow dissociation could be one explanation for 60 the slow evaporation of SOA, while Yli-Juuti, et al.²⁵ have shown that particulate-phase diffusion 61

62 is a limiting factor of slow evaporation under dry conditions. In contrast, Saleh, et al. ²⁶ and Saha 63 and Grieshop ²⁷ have indicated that the timescale for gas-particle equilibrium is less than 1 h and 64 that the equilibrium assumption can be applied in chemical transport models given typical model 65 timesteps. Applicability of the equilibrium assumption is an undecided issue, and there is presently 66 no consensus about its use.

67 Understanding of both physical and chemical properties is necessary to reveal the factors that 68 control SOA evaporation. It has been pointed out that the SOA volatility distribution estimated 69 from formation experiments using a traditional volatility basis set (VBS) framework includes more 70 volatile components than that estimated from dilution experiments, heating experiments, or chemical analysis combined with parameterization techniques²⁷⁻³⁰. This suggests that the presence 71 72 of low-volatility compounds or kinetic inhibition affects SOA evaporation processes. To our 73 knowledge, few previous modeling studies have considered both dimerization and bulk diffusion ^{20, 31}, and the constraints needed to determine the controlling factors of SOA evaporation are still 74 75 lacking. Recently, the C^* distribution of α -pinene SOA has been evaluated from chemical analyses ^{32, 33} and heating experiments ^{27, 34}, and these data could potentially be used as constraints for the 76 77 estimation of the model parameters that determine SOA evaporation. For the simulation of bulk diffusion in viscous particles, Shiraiwa, et al. ³⁵ and DeRieux, et al. ³⁶ have developed a 78 79 parameterization process for estimation of the glass transition temperature (T_g) of individual SOA 80 compounds. Using their parameterization process, bulk diffusivity can be calculated from the 81 chemical composition of SOA. However, the consistency of this parameterization process for 82 predicting SOA evaporation is a concern and should be evaluated using independent estimates.

83 In addition, Sato, et al. ³⁷ conducted dilution and heating experiments in which they measured 84 the evaporative behaviors and chemical composition of SOA from 1,3,5-trimethylbenzene (TMB).³⁷ Most research on the evaporative behavior of SOA has concerned the SOA from biogenic volatile organic compounds (VOC). However, SOA from anthropogenic VOC can make important contributions to atmospheric OA in urban regions,^{38, 39} and the evaporative behavior of anthropogenic SOA should therefore be understood as well. SOA mass yields and SOA properties (volatility and O:C ratios) differ between anthropogenic and biogenic SOA, and experiments with anthropogenic SOA should therefore enhance understanding of the behavior of ambient SOA.

91 In this study, we conducted model simulations that took into consideration kinetic gas-particle 92 partitioning, formation and dissociation of dimers, and particle-phase diffusion, and we assessed 93 the effects of these processes on the evaporative behavior of SOA. We first conducted simulations 94 with single-component SOA to evaluate the sensitivity of its evaporative behavior to model 95 parameters related to volatility, dimer formation/dissociation, and particle-phase diffusion. We 96 then conducted simulations for multi-component SOA in which estimates of dimer 97 formation/dissociation rates and bulk diffusivity were constrained by SOA concentrations 98 observed during the formation and dilution experiments and by the volatility distributions 99 estimated from previous chemical analyses or heating experiments. Finally, we compared the 100 simulated evaporative behavior and estimated model parameters between monoterpene and 101 aromatic SOA. Based on these analyses, we assessed the influence of dimerization and bulk 102 diffusion processes on the formation and evaporation of SOA.

103

104 2. Methods

105 2.1. Chamber Experiments of SOA Formation and Dilution

106 Our modeling relies on chamber experiments performed by Sato, et al. $^{33, 37}$. Sato, et al. $^{33, 37}$ 107 conducted twelve experiments using SOA from α -pinene ozonolysis and seven using SOA from

108 1,3,5-TMB/NO_x photooxidation to estimate SOA volatility distributions. We conducted box model 109 simulations for ten of these previous experiments (Table 1), in which SOA behavior after 110 isothermal dilution was measured. Because the experimental setup for α -pinene ozonolysis and 1,3,5-TMB photooxidation have been described in Sato, et al. ³³ and Sato, et al. ³⁷, respectively, 111 we provide only a brief description of those experiments here. A 6-m³ Teflon-coated stainless-112 113 steel chamber was used for the experiments under dry conditions (relative humidity (RH) < 1%), 114 and a 6-m³ fluorinated ethylene polyethylene (FEP) film bag ($1.81 \times 1.81 \times 1.81 \text{ m}^3$, 50-µm film 115 thickness) was used for experiments under humid conditions (RH = 40%). A comparison of the 116 background concentrations of particulate matter and VOC in the chamber and film bag is shown 117 in Table S2 of the Supporting Information (SI). Considering these background concentrations and 118 the maximum SOA yields²¹, the contribution of these background concentrations to the produced 119 SOA was about 0.03% in the chamber used for the dry experiments and about 0.3% in the film 120 bag used for the humid experiments. Thus, we considered the contributions of the background to 121 be negligible in our analysis. The requisite amounts of precursors (i.e., α -pinene and 1,3,5-TMB), 122 ozone, or other gaseous compounds were injected into these chambers (Table 1). Particle size 123 distributions were measured with a scanning mobility particle sizer (SMPS) (Model 3034, TSI Inc., 124 USA). The effective particle density of the SOA from α -pinene ozonolysis was estimated to be 125 1.34 g cm⁻³ by using a custom-made differential mobility analyzer and aerosol particle mass 126 analyzer (APM) (Model 3600, Kanomax Inc., Japan).

127 Offline chemical analysis was performed using a positive electrospray ionization-liquid 128 chromatograph/time-of-flight-mass spectrometer (LC-TOF-MS, Agilent Technologies, UK). In 129 each experiment, an aerosol sample for LC-TOF-MS analysis was collected on a Teflon filter 130 (Sumitomo Electric Industries, Japan, 47 mm φ , pore size 1 μ m) at a rate of 16.7 L min⁻¹ for 30

min after the reaction finished. Sato, et al. ^{33, 37} have estimated volatility distributions of SOA from 131 132 the concentrations and molecular formulas of the organic species identified through LC-TOF-MS analysis. Various parameterizations were used to relate the molecular formulas to the saturation 133 concentrations (C*). In one method, for the SOA from both α -pinene and 1,3,5-TMB, the C^{*} 134 135 distribution was estimated with a two-dimensional function (C^* was parameterized as a function of the number of carbon and oxygen atoms, 2-D method).²⁸ To assess the uncertainty, Sato, et al. 136 ³³ used a one-dimensional function (C^* was parameterized as a function of molecular weight 137 (MW))²⁹ with a binary fit for products of α -pinene ozonolysis with low MW (m/z < 300) and high 138 MW (m/z > 300) (binary method). Sato, et al. ³⁷ also used the two-dimensional function to fit to 139 140 $\log_{10}(C^*)$ data calculated by the SPARC (SPARC Performs Automated Reasoning in Chemistry) calculator⁴⁰ for products from TMB photooxidation (SPARC method). Sato, et al. ^{33, 37} have shown 141 142 that only a small portion of HOMs decompose and that the root mean square error (RMSE) and 143 mean bias of the $\log_{10}(C^*)$ estimated by SPARC and the two-dimensional function parameterized in accord with Sato, et al. ³⁷ are 1.9 and 0.0, respectively. The RMSE and mean bias between the 144 145 $\log_{10}(C^*)$ estimated by the SPARC and the original parameterization of the two-dimensional function²⁸ were 2.2 and -0.4, respectively. Quantitative analysis of the uncertainty of the estimated 146 147 C^* distribution is difficult because the mass identified by the LC-MS analysis could only explain at most only 30% of the total SOA mass,³ however, these values still provide a measure of the 148 149 uncertainty.

Evaporative behaviors associated with isothermal dilution were measured by injecting air from the reaction chamber into a clean 6 m³ FEP film bag (external dilution chamber, EDC) using a dilution ejector (FPS-4000, Dekati Ltd., Finland). The clean dilution air used was from a clean-air generator (Model 1160, Thermo Fisher Scientific, USA) for dry experiments (RH < 1%) and from

154 a purified air generator (DAR-2000, Horiba Stec Ltd., Japan) for humid experiments (RH = 40%). 155 The dilution ratios were 20–77 (α -pinene) and 20–86 (TMB). The temperature of the laboratory 156 was controlled at 298 ± 1 K. The particle size distribution, particle density, and CO concentration 157 in the EDC were monitored after dilution of the SOA, and the CO concentration was used to estimate the dilution ratio. It has been shown by McFiggans, et al.⁴¹ that the composition of SOA 158 from α -pinene photooxidation was changed by the addition of CO. However, Sato, et al. ³³ have 159 160 shown that the composition of SOA from α -pinene ozonolysis did not markedly change even with 161 addition of a large amount of CO (0.83 torr = 1092 ppmv). In the dilution experiments of Sato, et al. ³³, CO was added as a dilution tracer at a much lower concentration (~ 20 ppmv), and therefore 162 163 we consider that the added CO did not greatly change the composition of the SOA products.

164

165 2.2. Framework of Organic Aerosol Model

166 For the modeling analyses of SOA formation and evaporation after isothermal dilution, we used a one-dimensional volatility basis set (1D-VBS) framework⁹ that accounted for formation and 167 dissociation of dimers^{17, 18} and kinetic limitations to diffusion inside the particle-phase.¹⁹ Products 168 169 of VOC oxidation (i.e., α -pinene + ozone or 1,3,5-TMB/NO_x + light) were represented as five surrogate compounds with saturation concentration (C_i^*) of 0.1, 1, 10, 100, and 1000 µg m⁻³. The 170 171 mass yield of compound i was represented by a_i . We should note that compounds with a C^* of 0.1 $\mu g m^{-3}$ were practically non-volatile under our experimental conditions. We therefore did not 172 consider monomeric compounds with a $C^* < 0.1 \ \mu g \ m^{-3}$ in this study. The formulation of the C^* 173 174 distribution is explained in Section 3.2.

For the calculation of kinetic gas-particle partitioning, we adopted the equations formulated by
Zaveri, et al. ¹⁹ These equations were used to calculate condensation/evaporation rates of organic

177 compounds while taking into consideration volatility, gas-phase diffusion, interfacial mass 178 accommodation, particle-phase diffusion, and particle-phase reactions (oligomerization in this 179 study). Section S1 of the SI provides details of these equations. We did not calculate coagulation, 180 given that mean volume diameter was larger than 150 nm immediately after the start of the 181 formation experiments (Figure S1 of the SI) and contribution of coagulation on the size distribution 182 is small.

An accommodation coefficient (α) and particle-phase diffusivity (D_b) were uncertain parameters in these simulations. Recent work has found that the mass accommodation coefficient for both laboratory and field aerosols is close to or equal to 1.⁴² Without an explicit treatment of the bulk diffusion coefficient, it is likely that the mass transfer limitations observed in earlier work may have been attributed to a low value of the mass accommodation coefficient (≤ 0.1).^{26, 27} As the bulk diffusion coefficient is explicitly modeled in this work, we assumed the mass accommodation coefficient to be equal to 1 for all our simulations.

190 The basic formulations of dimer formation and dissociation reactions were based on the formulations of Trump and Donahue¹⁷ and Kolesar, et al.¹⁸. We simulated reversible, condensed-191 phase reactions with dimer formation and dissociation rates as $R_f = k_f w_m^2$, $R_r = k_r w_d$, 192 193 respectively, where w_m and w_d represent the mass fractions of monomer and dimer in the particle 194 phase, and $k_f(s^{-1})$ and $k_r(s^{-1})$ are the formation and dissociation rate constants, respectively, for 195 dimers. Formation and dissociation rates of dimers should differ among dimer types, such as hemiacetals, acetals, esters, and ethers.^{43, 44} However, these differences were not considered in this 196 197 study because the differences in k_f and k_r among dimer types are not sufficiently constrained to 198 warrant separate treatment in the simulation. We therefore used average k_f and k_r values in this study. We assumed that dimers are non-volatile in our simulations because the C^* values of dimers 199

are less than 1 μ g m⁻³. In addition, we assumed that only homodimers (dimers formed from monomers with the same C^*) are formed, as in Kolesar, et al. ¹⁸, and that k_f and k_r are the same for all C^* species. Such simplification might not be realistic as already discussed in Trump and Donahue ¹⁷ and Kolesar, et al. ¹⁸. However, for the parameter estimation from the experimental data, number of model parameters should be reduced, so that we applied such simple module.

For the simulations with single C^* SOA (Section 3.2.1), size distributions were tracked using a sectional approach with 60 sections of particle diameters ranging from 10 to 1000 nm. Initial size distributions were based on observational data, and particle growth due to condensation or shrinkage by evaporation was considered. We assumed log-normal distributions of particle sizes for the simulations with multi-component SOA (Section 3.2.2) during both the formation and dilution experiments, because the observed size distributions were reasonably mono-modal as shown in Section 3.1.

212

213 2.3. Model simulation of chamber experiments

214 As noted in Section 2.1, we conducted simulations for experiments of both SOA formation and 215 dilution (Table 1). The particle wall loss (PWL) rate in the SOA reaction chamber was calculated 216 from experimental data. For the α -pinene ozonolysis experiments, after the α -pinene had reacted 217 away the change in the geometric mean diameter was small (-3 nm/hr), whereas the SOA mass concentrations decreased at a rate of 0.128 h⁻¹ because of particle deposition to the walls and the 218 219 small amount of dilution associated with instrument sampling. In the case of the SOA produced 220 by photooxidation of TMB, SOA-forming chemical reactions between SVOC and OH could 221 continue beyond the point at which all the TMB had reacted, and we were thus unable to estimate 222 the PWL rate of TMB-SOA directly from our experimental data. Instead, we assumed a PWL rate

of 0.22 h^{-1} , which is the PWL rate reported for ammonium sulfate particles in the same diameter 223 range.³⁷ PWL rates in the ejector were calculated from the changes of the size distribution and 224 225 mass concentration of the SOA. The geometric mean volume diameter of the SOA before and just 226 after passage through the ejector changed by 0.2–2% under dry conditions and by 4–7% under 227 humid conditions. Previous experiments have shown that the PWL rate in the ejector does not depend strongly on the distribution of particle diameters around the peak diameter.⁴⁵ Based on this 228 229 assumption, we estimated the evaporation rate of SOA in the ejector from the change of the 230 geometric mean volume diameter. Thus, the fraction of SOA evaporated during passage through the ejector was $[1 - (d_{vm}(\text{after ejector})/d_{vm}(\text{before ejector}))^3]$. Then, a further decrease of the 231 232 SOA mass was caused by size-independent PWL. In other words, the fraction of particles 233 deposited on the ejector wall to particles introduced into the ejector was assumed to equal [1 - $\frac{C_{SOA}^*(\text{after ejector})/C_{SOA}^*(\text{before ejector})}{(d_{vm}(\text{after ejector})/d_{vm}(\text{before ejector}))^3}]$, where C_{SOA}^* is the mass concentration of the SOA (corrected 234 by the dilution ratio), and d_{vm} is the volume-mean diameter. We compared the measured and 235 calculated volume fraction remaining (VFR) from the equation VFR(t) = $\left(\frac{d_{vm}(t)}{d_{vm}(before ejector)}\right)^3$ for 236 237 the analysis of the SOA evaporation rate; in this way uncertainties associated with the PWL in the 238 EDC were minimized. The volatility dependent vapor wall loss (VWL) in both the reaction chamber and the EDC were calculated by applying the formulation of Krechmer, et al. ⁴⁶ Details 239 240 are provided in Section S2 of the SI.

As shown in Sections 3.2.2 and 3.2.3, we estimated plausible sets of parameters by fitting the model to observational data. Parameter estimation was conducted by using a nonlinear programming solver, the Nelder-Mead simplex algorithm⁴⁷ of MATLAB (R2018a, The MathWorks, Inc., MA, USA).

246 3. Results and Discussion

247 3.1. Experimental Results

We first compared the observed time-dependent VFR from α -pinene³³ and 1,3,5-TMB³⁷ with 248 249 the rates reported in other studies (Figure 1). Details of the experimental setups of the other studies 250 are provided in the cited publications, and we briefly introduce the methodologies used in Table S3 of the SI. In Sato, et al. $^{33, 37}$, the time-dependent VFR of the SOA from α -pinene and 1,3,5-251 252 TMB were similar functions of time under dry conditions: the VFR of the SOA decreased by 14– 253 20% within 1 h after dilution, and there was continued slow evaporation even after 2-3 h. 254 Evaporation was slightly faster for the dry SOA from α -pinene versus the SOA from 1,3,5-TMB. To our knowledge, no studies other than Sato, et al. ³⁷ measured evaporative behaviors of aromatic 255 256 SOA by the isothermal dilution method, though this result was consistent with other heating experiments,^{48, 49} which have shown that dry SOA from aromatic VOC is less volatile than the 257 SOA from monoterpenes. Of course, factors other than precursors, including oxidants, ^{48, 50, 51} SOA 258 concentrations,⁴⁸ or reaction timescales^{52, 53} should also be important factors controlling SOA 259 260 volatility. Therefore, we cannot conclude that anthropogenic SOA is more (or less) volatile than 261 biogenic SOA.

The evaporation rate of SOA from α -pinene was significantly higher under humid conditions (RH = 40%) versus dry conditions. The VFR decreased by about 35% in 1 h after dilution and continued decreasing even after 3 h. This RH dependence is consistent with previous studies,^{24, 25, 54} who have measured isothermal SOA evaporation rates at different RH. By contrast, D'Ambro, et al. ⁵⁵ have shown that the evaporation rates of SOA from α -pinene ozonolysis did not indicate RH dependence within a range of 20%–80% RH. It should be noted that D'Ambro, et al. ⁵⁵ measured evaporation rates of SOA on filters, whereas evaporation rates of suspended SOA were

measured in other studies. In addition, the lowest RH in the experiments of D'Ambro, et al.⁵⁵ was 269 270 20%, which is higher than the RH in other dry experiments, and this difference could explain the different evaporation rates.⁵⁴ Consideration of these differences between D'Ambro, et al. ⁵⁵ and 271 272 other studies may account for the apparently different behavior of the SOA evaporation in these 273 studies. Another difference in the humid experiments is the RH during SOA formation. SOA produced under dry conditions was exposed to different RH by Yli-Juuti, et al.²⁵ and D'Ambro, et 274 275 al. ⁵⁵, whereas SOA were formed under humid conditions (RH = 40%) in the humid experiments performed by Sato, et al. ^{33, 37}. However, the SOA evaporation rate was not affected by the RH 276 (dry or 30%) during SOA formation in the experiments conducted by Yli-Juuti, et al.²⁵. Overall, 277 278 because the experimental setups of the individual studies differed, there were differences in the 279 evaporative behaviors of SOA from α -pinene under both dry and humid conditions (Figure 1). In 280 some previous studies, evaporation was accelerated because semi-volatile vapors were removed artificially^{23, 55} or they were likely removed within a short time in a stainless steel chamber.²⁵ 281 However, semi-volatile vapors were not intentionally removed in other chamber experiments.^{33, 37,} 282 ⁵⁶ In addition, concentrations, diameters, and the conditions of particles (e.g., on filters in D'Ambro, 283 et al. 55) differed between experiments. The implications of these differences are discussed in 284 285 Section 3.2.3. We used our experimental data for the modeling analysis presented in Section 3.2 286 and Section 3.3. We should note that during both the dry and humid experiments, evaporation still 287 occurred, even several hours after the start of dilution. This behavior is consistent with the findings 288 from other previous experiments (Figure 1). It is reasonable that evaporation still occurs in an open 289 system (with continuous removal of vapors), even at several hours after the perturbation. In a 290 closed system (with a finite amount of vapor, chamber experiments in this case), not only kinetic 291 limitation, but also VWL would contribute to this continuous decrease of VFR.

The derived C^* distributions of total (gas + particle) organic compounds ^{33, 37} were also compared 292 293 with those reported in the literature (Figure 2). We found that the fraction of total gas and particulate-phase oxidation products ($f_i = a_i / \Sigma(a_i)$) of C^* were in the range of $10^{-9} - 10^3 \,\mu\text{g m}^{-3}$. A 294 295 brief summary of the derivations in the previous studies is given in Table 2 and Section S3 of the 296 SI. Even though our estimate includes uncertainties (Section 2.1), the derived C^* distributions of total (gas + particle) organic compounds $^{33, 37}$ based on the measured particle composition of the 297 298 SOA from α -pinene and 1,3,5-TMB showed similar trends: the low-volatility (mostly dimers³³) 299 and semi-volatile ($C^* = 10-1000 \ \mu g \ m^{-3}$) organic compounds made marked contributions after 300 completion of the reaction in the reaction chamber (Figure 2). In these cases, the dimer fractions 301 (*f_{dimer}* = [dimer]/[SOA]) calculated using data from the LC-TOF-MS analysis were similar for both 302 SOA types, 0.59 for the SOA from α -pinene and 0.58 for the SOA from 1,3,5-TMB. Previous 303 experimental studies have also indicated the important contributions of oligomers, but their contributions significantly varied from a few percent to ~70% among studies^{2, 16, 33, 37, 57, 58}. The 304 305 f_{dimer} estimated by Sato, et al. ^{33, 37} is within a range of these previous estimates. The C^{*} distributions shown in Figure 2 were derived from chemical analysis ³⁴ or heating experiments ^{27, 32, 33}. We 306 307 should note that these estimates do not include the contributions of dimers to the distributions. Isaacman-VanWertz, et al. ³⁴ measured concentrations of gaseous compounds and estimated 308 309 aerosol contributions from their vapor pressures. Thus, it is presumed that the contributions of 310 LVOC and dimers were not properly calculated. In heating measurements, the oligomers could 311 evaporate or decompose, so understanding of the heat-induced evaporative behaviors of oligomers 312 is complicated ^{17, 59}. Considering this uncertainty, the contributions of dimers might not be 313 accurately estimated from heating experiments. However, even with these limitations, all the 314 estimates showed that SVOC made important contributions to the distributions (fi of compounds

with $C^* = 100-1000 \ \mu g \ m^{-3}$ was 0.36-0.69). For the practical model simulation, C^* distributions should be determined from the limited, currently available datasets, and thus, the C^* distributions and *f_dimer* values were used as constraints in the simulations, and the implications of the contributions of SVOC are discussed in Section 3.2.2.

The evolution of particle size during the formation and dilution experiments is shown in Figure S1. Size distributions during both experiments were reasonably unimodal and were well fit to lognormal functions throughout the experiments (geometric mean volume diameter around 200–400 nm). Therefore, we assumed log-normal distributions for the simulation of the multi-component SOA.

324

325 3.2. Sensitivity Simulations for SOA from α -Pinene

326 3.2.1. Simulations for single-component SOA

327 We simulated the evaporation behavior of SOA in the EDC assuming that the SOA was composed of compounds with the same C^* value, but with other parameters allowed to vary, for 328 329 the two objectives. First, this allowed for assessment of the sensitivity of simulated SOA concentrations obtained from dilution experiments to the assumed saturation concentration C^* , the 330 331 monomer-dimer conversion rates (dimer formation rate k_f and dissociation rate k_r), and the bulk 332 diffusivity D_b (Figure 3). Second, the simulation models often need to be simplified because of 333 limitations of computational resources, and therefore simulation results with a simpler framework 334 provide information that is useful for practical application. Here, we focus on understanding the 335 influence of RH and thus consider only SOA from a-pinene ozonolysis. As already noted, gas-336 particle equilibrium was not attained during the experiments under both dry and humid conditions. 337 However, under dry conditions, evaporation rates slowed even on a logarithmic timescale during the experiments of this study and those of Grieshop, et al. ⁵⁶ In contrast, the evaporation rates did not slow on a logarithmic timescale during experiments conducted by other groups.^{23, 25, 55} This point is also discussed in Section 3.2.3.

341 It was apparent from simulations without consideration of dimer formation ($k_f = 0$) and bulk 342 diffusion that time-dependent VFR was particularly sensitive to C^* (Figure 3). To quantitatively evaluate the model performance, the RMSE of measured and simulated VFR during the dilution 343 344 experiments was calculated in each case. In this study, derivation of logarithm of time t (i.e., 1/t) was used for weighting to reproduce the initial evaporative behavior (RMSE = 345 $\sqrt{\frac{\sum (VFR_{model,i} - VFR_{obs,i})^2/t_i}{\sum (1/t_i)}}$, where VFR_{model,i} and VFR_{obs,i} indicate simulated and observed VFR at 346 347 time t_i). Here, the observed SOA concentrations after dilution were best reproduced in the cases with $C^* = 1.3 - 2.5 \ \mu g \ m^{-3}$ for experiments under dry conditions and 3.2-12.6 $\mu g \ m^{-3}$ under humid 348 349 conditions. However, the shapes of the evaporation curves were not well reproduced over the whole period for any single C^* for either RH. Nonetheless, these estimated C^* values were 350 somewhat similar to an effective mass-weighted C^* (1-4 µg m⁻³) estimated from observed 351 evaporation rates of α -pinene ozonolysis SOA by D'Ambro, et al. ⁵⁵ These C^{*} values are 352 nonetheless greater than those estimated from chemical analysis (mass-weighted C^* was from 353 $0.0017 \ \mu g \ m^{-3}$ [the binary method] to $0.25 \ \mu g \ m^{-3}$ [the 2-D method]).³³ This gap indicates that the 354 355 evaporation and chemical properties of SOA cannot simply be represented by a single component calculation. In our experiments, OA concentrations decreased from $\sim 10^3 \ \mu g \ m^{-3}$ in the reaction 356 chamber to ~ $10^1 \,\mu\text{g} \text{ m}^{-3}$ in the EDC.^{33, 37} Because gas-particle partitioning at equilibrium is 1:1 for 357 a compound with C^* identical to the OA concentration $\left(\frac{C_{gas,i}}{C_{aerosol,i}}\right)$ for organic compound *i* is equal 358

to $\frac{C_i^*}{C_{OA}}$), it is reasonable that the mass loss resulted predominately from evaporation of compounds with $C^* = 10-1000 \ \mu g \ m^{-3}$ (Figure 3).

361 Results of the sensitivity simulations to C^* , k_f (or f_{dimer}), k_r , D_b are summarized in Section S4 of 362 the SI. For the parameter space considered here, the overall best results under the single C^* assumption were obtained when $D_{\rm b} \ge 10^{-18} \text{ m}^2 \text{ s}^{-1}$, $f_{\rm dimer} = 0.6$, $k_{\rm r} = -3 \times 10^{-6} \text{ s}^{-1}$, and $C^* = -10 \text{ }\mu\text{g}$ 363 364 m⁻³ for dry experiments (Figures S2 and S3 of the SI). In the humid experiments (Figures S4 and 365 S5 of the SI), generally similar results were obtained as in the dry experiments. However, there was a much clearer minimum in the RMSE that occurred at a comparably higher D_b (= 10^{-16} m² s⁻ 366 ¹), lower f_{dimer} (= 0.4), and larger k_r (~3 × 10⁻⁵ s⁻¹). Also the simulation-observation agreement at 367 low D_b ($\leq 10^{-18}$ m² s⁻¹) was notably worse for all combinations of the other parameters. We should 368 note that SOA evaporation was not delayed in the simulations with $D_b \ge 10^{-16.5} \text{ m}^2 \text{ s}^{-1}$ for all the 369 370 compounds, suggesting that SOA was not diffusion-limited in these simulations.

The changes of SOA sizes during dilution-induced evaporation are similar in scenarios with different k_f , k_r , and D_b values (Figure S6 of the SI). Namely, larger particles evaporated faster, and mean diameter decreased as evaporation proceeded. This particle-size evolution is consistent with the observed size changes. We therefore cannot constrain unique parameter sets of C^* distributions, k_f , k_r , and D_b from only SOA concentrations and size distributions during formation and dilution experiments.

Overall, the C^* and k_f (or f_{dimer}) determined the end-of-experiment concentrations, and sensitivities to these parameters were thus significant over the whole dilution experiment. In contrast, k_r and D_b determined equilibration timescales, and sensitivities to these parameters were thus apparent within the equilibration timescales calculated from these parameters. Dilutioninduced evaporative behaviors have been shown to be highly sensitive to C^* , and we can thus expect that this evaporative behaviors can be reproduced if we optimize the C^* distributions in multi-component systems, even with wide ranges of parameter values for k_f , k_r , and D_b^{25} . In the next section, we constrained C^* distributions from chemical analysis data³³.

- 385 3.2.2. Simulations for multi-component SOA
- 386 (a) Without constraint of C^* distributions

The C^* distribution of VOC oxidation products (i.e., mass yield (a_i) of oxidation product with 387 388 C_i^*) is one of the key factors that control the simulated SOA evaporation behavior, as discussed 389 above. Here, we assess the ability of the model to reproduce the observed evaporation profiles where it is assumed that the compounds comprising the SOA have a distribution of C^* values (C^* 390 = 0.1, 1, 10, 100, and 1000 μ g m⁻³). Simulations are performed for a variety of discrete k_f , k_r , and 391 D_b combinations; for each pair of k_f and k_r , the D_b are varied over the range 10^{-19} – 10^{-16} m² s⁻¹. The 392 393 C^* distributions are determined by minimizing the difference between observed and simulated 394 SOA concentrations in both the reaction chamber and the EDC. Here, an objective function for the 395 above parameter optimization was the sum of RMSE of SOA concentrations in the reaction 396 chamber (normalized by the observed SOA concentrations at the end of the formation experiments) 397 and VFR in the EDC. Both are normalized parameters, so the weighting factors for RMSE for 398 normalized SOA concentrations in the reaction chamber and VFR in the EDC were set equal. 399 Simulations with this multiple- C^* system were conducted both for the reaction chamber and EDC. 400 Thus, we explicitly calculated f_{dimer} after the formation experiment by inputting both k_f and k_r .

Figure 4 shows examples of fitting results for experiments under dry conditions. We find that, after optimization of the C^* distribution, any assumed combination of k_f and k_r used here allows for a reasonable match with the observed evaporation profile for the dry SOA when $D_b \ge 10^{-18} \text{ m}^2$ s^{-1} . For all cases, when $D_b = 10^{-19} \text{ m}^2 \text{ s}^{-1}$, evaporation proceeds too slowly initially, leading to

overall marginal agreement with the observations even after optimization. The optimized C^* 405 distributions obtained for the $D_b \ge 10^{-18} \text{ m}^2 \text{ s}^{-1}$ simulations depend on the assumed k_f and k_r . In 406 general, when the dimer fraction is assumed small ($k_f = 10^{-6} \text{ s}^{-1}$ and $k_r = 10^{-3} \text{ s}^{-1}$, so $f_{dimer} \sim 0.001$) 407 the fractional contribution of the lower volatility components ($C^* \le 10 \,\mu g \, m^{-3}$) is greater. However, 408 when substantial dimer formation is allowed ($k_f = 10^{-3} \text{ s}^{-1}$ and $k_r = 10^{-6} \text{ s}^{-1}$ or $k_r = 10^{-4} \text{ s}^{-1}$), there are 409 notable contributions from all C^* components and the dimers comprise the largest fraction. 410 411 Consequently, this suggests that evaporation for these cases is controlled largely by dimer 412 dissociation. The optimized C^* distributions from the high-dimer-fraction simulations are more consistent with the C^* distributions derived from the measured particle composition shown in 413 414 Figure 2(a).

415 Under humid conditions, SOA evaporation behavior was best reproduced in the cases with $D_b =$ 10^{-16} m² s⁻¹, while the evaporation rates immediately after the dilution were slightly 416 underestimated in the cases with $D_b = 10^{-17} \text{ m}^2 \text{ s}^{-1}$. The evaporation rate was significantly 417 underestimated in the cases with $D_b \le 10^{-18} \text{ m}^2 \text{ s}^{-1}$, suggesting that higher D_b (lower viscosity) was 418 419 necessary to reproduce the observed time-dependent VFR under humid conditions. The extent and 420 rate of dimer formation also influenced the simulation results. In particular, even in the lowviscosity case (i.e., $D_b = 10^{-16} \text{ m}^2 \text{ s}^{-1}$), the evaporation rate was underestimated in the simulations 421 with relatively fast dimer formation ($k_f = 10^{-3} \text{ s}^{-1}$) and slow dimer dissociation ($k_r = 10^{-6} \text{ s}^{-1}$) 422 (Figure 5). However, in the simulation with relatively fast dimer formation ($k_f = 10^{-3} \text{ s}^{-1}$) and fast 423 dimer dissociation ($k_r = 10^{-4} \text{ s}^{-1}$), the C^* distribution and SOA evaporative behavior were both 424 consistent with observations. In the simulation with little dimer formation ($k_{\rm f} = 10^{-6} \, {\rm s}^{-1}$ and $k_{\rm r} =$ 425 10^{-3} s⁻¹), good agreement with the evaporation profile was obtained when $D_b \ge 10^{-17}$ m² s⁻¹, but the 426 427 optimized C^* distributions differed notably from that estimated from the particle composition.

428 (b) With constraint of C^* distributions

429 The above solutions provide general insight into the factors controlling the SOA evaporation, 430 but are not unique owing to a lack of constraints. To better understand the factors controlling the 431 SOA evaporative behavior, we determined optimized values of D_b , k_f , and k_r using the observation-432 based C^* distributions and f_{dimer} (Table 2) as a constraint. Based on chemical analysis and heatinduced evaporation experiments, Sato, et al. ³³ have estimated that SOA components with 433 relatively high volatility ($C^* \ge 10^2 \,\mu \text{g m}^{-3}$) contribute 20–50% of the total SOA from α -pinene. In 434 435 addition, C^* distributions estimated in other previous studies also show that the SVOC had high 436 contributions (33–77%), suggesting that these compounds made important contributions to the 437 total products from α -pinene ozonolysis. These C^* distributions (f_i) were used as initial conditions, 438 and we tolerated the variation of f_i within $\pm 10\%$ from the initial values. As noted in Section 3.1, 439 the contributions of dimers have not been adequately estimated in the literature, except by Sato, et al. ³³. Thus, the following assumptions were made in the sensitivity simulations: (1) compounds 440 with $C^* \le 10^{-2} \ \mu \text{g m}^{-3}$ were considered as dimers, (2) compounds with $C^* \le 10^{-2} \ \mu \text{g m}^{-3}$ were 441 considered as low-volatility monomers ($C^* = 10^{-1} \ \mu g \ m^{-3}$ is practically non-volatile in this 442 calculation), and (3) in the case of Chhabra, et al. ³², the dimer fraction was set to be the same as 443 or half that in Sato, et al. ³³. These assumptions are summarized in Table 2. 444

We minimized the RMSE between the observed and simulated SOA concentrations for the formation experiments and between the observed and simulated VFR for the dilution experiments using the Nelder-Mead simplex algorithm, as noted in Section 2.3. For the formation experiments, an optimal set k_f and $\Sigma(a_i)$ was calculated by minimizing sum of the RMSE of the observed and simulated time dependent SOA concentrations in the reaction chamber (normalized by the observed SOA concentration at the end of the formation experiments) and the RMSE of *fdimer* at the end of the formation experiments with the same weight. For the dilution experiments, an optimal set k_r , D_b , and f_i was calculated by minimizing the RMSE of the observed and simulated VFR in the EDC. Fitting calculations for the reaction chamber and EDC were sequentially conducted. When the criteria of the RMSE for the reaction chamber and EDC were simultaneously met, the fitted parameters are considered as an optimal solution. Initial values of objective variables (i.e., a_i , k_f , k_r , and D_b) were pseudo-randomly chosen within plausible ranges, and these sets of fitting calculations were iterated 20 times.

458 The simulated VFR during dilution-induced evaporation is summarized in Figure 6(a). In the simulations with consideration of the observed fraction of dimers (CA-SA18-STD³³ and CA-459 460 CH15-HighD³² in Table 2), the observed VFR was reasonably reproduced with RMSE = 0.01. In 461 contrast, in the simulations without consideration of dimers, the observed VFR was not accurately 462 reproduced (RMSE = 0.05-0.08). As shown in Section 3.2.2 (a), the observed evaporation could 463 be reproduced by introducing large fractions of LVOC, even in the simulations unconstrained by 464 the C^* distribution. However, when constraining with these C^* distributions from the previous 465 studies, the observed VFR could not be well reproduced even after optimization of k_r and D_b . As already noted, SVOC ($C^* = 10^2 - 10^3 \ \mu g \ m^{-3}$) made important contributions in the observation-466 467 based C^* distributions (Figures 2(a)). When constraining with the C^* distributions, dimer formation 468 and decomposition appear to play important roles in SOA evaporation. This point is clearly 469 indicated in Figure 6(c). By introducing half of the observed dimer fraction (CA-CH15-HalfD), 470 the model performance (RMSE) was in between the case considering the observed dimers (59%) 471 and that not considering the dimers. The relationship between the model performance (RMSE) and 472 the estimated D_b is shown in Figure 6(d). With lower dimer fractions, the estimated D_b was lower, 473 although the observed VFR after the dilution was not adequately reproduced. In the simulations

474 with reasonable model performance, estimated D_b was around $10^{-17.8}$ m² s⁻¹. Hereafter, we mainly 475 show the results in the case of CA-SA18-STD.

The fitting parameters in the case of CA-SA18-STD were as follows: $\log_{10}(k_f [s^{-1}]) = -2.8, -5.9$ $\leq \log_{10}(k_r [s^{-1}]) \leq -5.1$, and $-18.0 \leq \log_{10}(D_b [m^2 s^{-1}]) \leq -17.8$ (Table 3). We should note that k_f and D_b were well constrained in relatively narrow ranges over the 20 iterations and over the three experiments under the dry conditions, while the estimated k_r indicated a variability by more than an order of magnitude. In this case, SOA evaporation was associated mostly with the evaporation of compounds with relatively high C^* ($\geq 10^2 \ \mu g \ m^{-3}$) and a fraction of dimer increased after the dilution (Figure 7).

483 In Sato, et al. ³³, filter sampling was not conducted during the experiments under the humid conditions. We therefore assume that the C^* distributions are the same for the dry and humid 484 485 experiments. This is reasonable given that variation in RH has little influence on the formation of low-volatility HOMs and dimers⁶⁰ or on the formation of SOA.^{57, 61} Under humid conditions, the 486 estimated fitting parameters were as follows: $-2.9 \le \log_{10}(k_f[s^{-1}]) \le -2.7, -4.0 \le \log_{10}(k_f[s^{-1}]) \le -2.7, -4.$ 487 $3.9, -15.4 \le \log_{10}(D_b[\text{m}^2 \text{ s}^{-1}]) \le -13.3$ (Table 3). These D_b values ($\le 10^{-16} \text{ m}^2 \text{ s}^{-1}$) indicate that 488 489 particle mixing is sufficiently fast (equilibrium timescale of particles with diameters of 0.1–0.4 µm was less than 1 min^{19, 62}) compared to the evaporation timescale (33-46 min³³), thus bulk 490 491 diffusion was not a factor that limited evaporation under humid conditions. In contrast, bulk 492 diffusion should inhibit evaporation under dry conditions.

Dimer dissociation rates estimated in this study were consistent with the previous estimate $(10^{-5}-$ 10⁻⁴ s⁻¹) based on isothermal evaporation by D'Ambro, et al. ⁵⁵ Chemical analysis has shown that dimers make important contributions to SOA from monoterpene and aromatic VOC ^{33, 37, 55}. Dimer formation should therefore be considered for the accurate reproduction of isothermal evaporation 497 rates. Kolesar, et al. ¹⁸ have estimated a faster k_r of 1.6×10^{-3} to 2.8×10^{-2} s⁻¹, probably because 498 they used the isothermal-evaporation data of Vaden, et al. ²³, which show faster evaporation than 499 those of Sato, et al. ³³ and D'Ambro, et al. ⁵⁵ (Figure 1(a)). There are obviously large uncertainties 500 in isothermal evaporation rates, as already discussed in Section 3.1. The cause of the differences 501 between experiments is further discussed in Section 3.2.3.

502 The D_b for dry particles estimated in this study was similar to that of Zaveri, et al. ³¹, who 503 estimated D_b (2×10⁻¹⁸ m² s⁻¹) from the experimental change of the sizes of SOA from α -pinene 504 under dry conditions (RH ~ 10%). Higher D_b values under humid conditions is also consistent with previous measurements of SOA viscosity^{36, 63, 64} or modeling analysis of SOA reactivity.^{65, 66} 505 Shiraiwa, et al. ³⁵ and DeRieux, et al. ³⁶ have developed a parameterization for the estimation of 506 507 T_g of individual SOA compounds, as detailed in Section S5 of the SI. With the estimated T_g , the 508 dynamic viscosity (η) and D_b can be calculated from the modified Vogel-Tammann-Fulcher equation and the Stokes–Einstein equation, respectively^{35, 36}. These equations and estimated η and 509 510 D_b are summarized in Section S5 and Table S4 of the SI. The value of D_b calculated from chemical composition data and an estimated $T_g (D_{b,chem})^{35, 36}$ was much lower than the value of D_b estimated 511 512 from parameter fitting of dilution-induced evaporation rates $(D_{b,dil})$ (Table 3) under both dry and 513 humid conditions. This difference is consistent with the results of Tikkanen, et al. ⁶⁷, who have 514 shown that the observed dilution-induced evaporation rate is greatly underestimated by simulations using a kinetic model with $D_{b,chem}$. Tikkanen, et al. ⁶⁷ have also shown that the observed dilution-515 induced evaporation rate can be reproduced if the T_g values of all compounds are reduced to 516 517 account for the uncertainty of T_g within a plausible range (30 K). As shown in Table S4 of the SI, the $D_{b,chem}$ derived from T_g but artificially reduced by 30 K was closer to $D_{b,dil}$ (Table 3). The 518 519 implication is that $D_{b,dil}$ and $D_{b,chem}$ were at least consistent within the uncertainty of T_g , though

these estimated $D_{b,chem}$ values include large uncertainty. As noted in the Introduction, DeRieux, et al. ³⁶ derived an equation to calculate $D_{b,chem}$ from the composition of SOA (Equations S11-17 of the SI). Thus, the relationship between D_b and SOA composition can be used to simultaneously optimize $D_{b,chem}$ and C^* . However, because of the large discrepancy between $D_{b,dil}$ and $D_{b,chem}$ as shown above, we could not conduct such an optimization analysis in this study.

525 3.2.3. Comparison with previous studies

526 As already noted in Section 3.1, the dilution-induced evaporation rates used here were lower 527 than the rates reported in most previous studies (Figure 1(a)). Differences in the treatment of VWL, 528 the range of SOA concentrations, and SOA composition that arise from different reaction 529 conditions are all possible explanations for the difference in SOA evaporation rates. Vaden, et al. ²³ removed gaseous compounds by activated charcoal denuders. Yli-Juuti, et al. ²⁵ used a stainless 530 531 steel chamber for an evaporation chamber and assumed in their simulations that gaseous semi-532 volatile compounds instantaneously deposited on the walls. To estimate the impact of VWL, we compared the VFR simulated with the VWL module of Krechmer, et al. ⁴⁶ and the VFR simulated 533 with the instantaneous VWL that was assumed in previous studies^{23, 25} (Figure 8). In this simulation, 534 535 we used the fitted parameters a_i , k_f , k_r , and D_b estimated in Section 3.2.2 (Table 3), and we changed 536 only the VFR calculation methods. Our simulations indicated that 3 h after the start of dilution the 537 VFR in our experimental setups was higher than the VFR in experiments with an instantaneous 538 VWL by about 0.06. This result indicates that treatment of the VWL range could partly explain 539 the differences in SOA evaporation rates among the experiments. This point has also been 540 discussed in a more general context by Li and Shiraiwa⁶⁸ based on comparison of SOA equilibrium 541 timescales in an open system (with a constant vapor concentration) and a closed system (with a 542 finite amount of vapor). By contrast to the above discussion, the difference of VFR after dilution

between the simulation without VWL and that with a module of Krechmer, et al.⁴⁶ was only a few 543 544 percent, suggesting that contributions of VWL to the SOA evaporation in our closed system was modest. D'Ambro, et al. 55 have measured evaporation rates of SOA on filters. Quantitative 545 546 assessment of the differences of the VFR between SOA suspended in the air and on filters is 547 difficult, though such differences could also change evaporation rates after isothermal dilution. In 548 addition, the VFR after evaporation was also sensitive to the total OA concentration. Simulations 549 used to evaluate the sensitivity of evaporative behavior to OA concentrations just before dilution 550 revealed that the VFR after dilution decreased by about 0.1 when the initial concentration of total 551 product (particles + gas) was reduced by one order of magnitude with the proportional C^* 552 distribution.

553

3.3. Simulations of Evaporative Behaviors of SOA from 1,3,5-TMB

555 During the formation experiment of 1,3,5-TMB, SOA formation did not start until after the decay of the precursor VOC. This behavior, which has already been reported⁶⁹, could be accounted for 556 557 by several explanations, including dependence of SOA yields on NO_x concentrations because of differences of chemical pathways⁷⁰ or differences of induction periods before SOA formation.⁷¹ 558 Ng, et al. ⁷⁰ have shown that the presence of NO initially suppresses the formation of SOA from 559 560 aromatic VOC. In this study, we simply set $a_i = 0$ when the NO concentration was higher than the 561 critical concentration (fitting parameter). This parameter setting might not represent the actual 562 mechanism of SOA formation, though this choice of a_i had little influence on the simulation of 563 SOA evaporation by isothermal dilution.

As already noted in Section 3.1, evaporation rates and C^* distributions were similar for dry SOA

565 from α -pinene and 1,3,5-TMB. The estimated parameters of k_f , k_r , and D_b were likewise similar to

566 each other as well (Table 3). The estimated values among several experiments were as follows: –

567 $18.6 \le \log_{10}(D_b \text{ [m}^2 \text{ s}^{-1}\text{]}) \le -18.1 \text{ and } -8.3 \le \log_{10}(k_r \text{ [s}^{-1}\text{]}) \le -3.7.$

568 Comparisons of evaporation rates or viscosities between SOA from monoterpene (mostly a-569 pinene) and aromatic VOC (toluene, xylene, or TMB) have recently been conducted. These rates continue to be a controversial issue. Based on the bead-mobility technique,⁷² it has been shown 570 571 that the viscosity of SOA particles is higher under dry condition than under humid condition for the SOA from α -pinene⁷³ and toluene.⁶³ The D_b values for the SOA from α -pinene and toluene 572 were estimated previously to be very low $(10^{-20} \text{ m}^2 \text{ s}^{-1})$ under dry conditions. However, other 573 574 groups have shown that diffusion of the SOA from monoterpene and aromatic VOC behaves differently when based on measurements of evaporation rates from organic films⁷⁴ or 575 measurements of the mixing state of isotopically labeled SOA.⁵⁴ In both cases,^{54, 74} evaporation of 576 577 the SOA from α -pinene was not diffusion limited under either dry nor humid conditions, whereas 578 the SOA from aromatic VOC (toluene or m-xylene) are diffusion limited under dry conditions but 579 not under humid conditions. The above-mentioned studies were consistent with respect to the 580 behavior of diffusion of the SOA from aromatic VOC, though they were inconsistent with respect 581 to the behavior of diffusion of the SOA from α -pinene under dry conditions. Our results indicated 582 that evaporation of the SOA from α -pinene was diffusion limited under dry conditions. This 583 behavior is consistent with the viscosity estimates based on the poke-flow technique⁷³ or particle formation experiments,³¹ but inconsistent when based on measurements of evaporation rates of 584 organic films⁷⁴ and SOA mixing experiments.⁵⁴ DeRieux, et al. ³⁶ have shown that the viscosity of 585 dry SOA from toluene (10⁹-10¹² Pa s) are higher than that of dry SOA from α -pinene (10⁷-10¹⁰ 586 587 Pa s). Thresholds of diffusion limitation might differ among methodologies, and SOA properties

could differ because of factors other than precursor types (e.g., concentrations or presence ofoxidants). These differences might account for the different conclusions.

The SOA yields and O:C ratios of mixtures of anthropogenic and biogenic SOA are seemingly explained as a linear combination of properties of purely anthropogenic and biogenic SOA.^{48, 75} However, other studies^{41, 76} have shown that SOA yields cannot be represented by linear combinations of different types of SOA, probably due to the nonlinear chemistry of the HOMs, dimer formation, and/or cycles of radicals. The evaporative behavior of these SOA mixtures should therefore be the focus of future studies.

596

597 4. Conclusions

598 We have previously measured the evaporative behavior of SOA from α -pinene ozonolysis and 599 1,3,5-TMB/NO_x photooxidation by conducting isothermal dilution experiments.^{33, 37} We have also 600 measured the volatility distributions of the SOA based on chemical analyses using LC-TOF-MS. 601 These observations, which have revealed that the distributions of SOA volatility estimated from 602 SOA yield curves during formation experiments, isothermal dilution experiments, and chemical 603 analysis all differ, suggest that it is important to model the multi-generational gas- and particle-604 phase reactions to produce ELVOC and LVOC along with the kinetics in gas-particle conversion 605 if we are to explain these differences. In this study, we conducted model simulations using the 606 VBS framework with consideration of kinetic gas-particle partitioning, formation and dissociation 607 of dimers, and limitation by particle-phase diffusion to understand the evaporative behaviors of 608 SOA from biogenic and aromatic VOC and their controlling factors.

609 SOA evaporation rates after dilution were particularly sensitive to C^* distributions, though k_f , k_r , 610 and D_b were also important factors that controlled the evaporation rates. Generally, C^* and k_f

611 determined equilibrium concentrations after evaporation, whereas k_r and D_b determined mainly 612 equilibration timescales. Fitting of the observed SOA formation and evaporation rates could not 613 produce unique solutions of all these parameters because the observed time-dependent VFR could 614 be reproduced by optimizing C^* distributions with different sets of k_f , k_r , and D_b . We then used the 615 various C^* distributions estimated from chemical analysis or heating experiments as a constraint, and in this way we estimated k_f , k_r , and D_b . Among the various C^* distributions, the observed 616 617 evaporation behavior could be reproduced in the experiment with significant dimer contributions (59%) and low D_b (10^{-17.8} m² s⁻¹). These results suggest that both dimerization and bulk diffusion 618 619 contributed to the observed slow evaporation rates under dry conditions. By contrast, particle-620 phase diffusion did not practically inhibit SOA evaporation under humid conditions. This 621 dependence on RH was consistent with some previous studies. However, several previous studies 622 have shown that evaporation of α -pinene SOA is not diffusion limited under dry conditions, but 623 that evaporation of aromatic SOA is diffusion limited. The reason for this difference of diffusion 624 limitation has not yet been resolved, and it should be examined in future studies. We also found 625 that model parameters estimated for the SOA from 1,3,5-TMB were similar to those estimated for 626 the SOA from α -pinene under dry conditions. The implication is that dimer formation and bulk 627 diffusivity are important in reproducing evaporative behaviors of both biogenic and aromatic SOA. 628 Evaporation rates of SOA from α -pinene varied among experiments. These variations could be 629 partly explained by differences in the treatment of the VWL. The experimental design is thus a 630 critical consideration in the simulation of SOA evaporative behavior.

631

632 ASSOCIATED CONTENT

633 Supporting Information

634 The Supporting Information is available free of charge on the ACS Publications website.

635

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Exp, #	Precursor	RH	[HC]0	[O ₃] ₀	[Et ₂ O]	[NO]0	[NO ₂]0	[CH ₃ ONO] ₀	Dilution Ratio
			ppb	ppb	ppm	ppb	ppb	ppb	
1	α-pinene	<1%	839	949	32				75
2	α-pinene	<1%	544	1023	23				20
3	α-pinene	<1%	510	1078	23				43
4	α-pinene	40%	455	1091	23				33
5	α-pinene	40%	455	1091	23				20
6	α-pinene	40%	455	1091	23				77
7	1,3,5-TMB	<1%	1500			1208	30	10	20
8	1,3,5-TMB	<1%	1526			1214	7	10	40
9	1,3,5-TMB	<1%	1515			1195	14	10	63
10	1,3,5-TMB	<1%	1488			1203	6	10	86

Table 1. Experimental setups. [X]₀ indicates initial concentrations of species X.

Abbreviations: RH: relative humidity, HC: hydrocarbon, Et₂O: diethyl ether, TMB:
trimethylbenzene.

926	Table 2 . List of C^*	[*] distributions estimated	from chemical	analysis and	heating experiments	in
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927 previous studies. Experiments examining secondary organic aerosol from ozonolysis of α -pinene

Distribution name	Method	System	Relative humidity	Dimer fraction	Reference
CA_SA18_STD	Chemical analysis (offline ESI-LC- TOF-MS & parameterization of C^*)	Smog chamber	$\leq 1\%$ and 40%	59%	Sato, et al. ³³
CA_SA18_halfD	''a	"	"	30%	"
HE_SA18_wD	Heating experiments (TD-AMS)	"	"	18%	Sato, et al. ³³
HE_SA18_woD		"	"	0%	"
CA_CH15_ wD	Chemical analysis (online-CIMS & parameterization of C^*)	PAM flow reactor	30%	11%	Chhabra, et al. 32
CA_CH15_HighD	"	"	"	59%	"
CA_CH15_HalfD	"	"	"	30%	"
CA_CH15_woD		"	"	0%	"
HE_VW18_wD	Heating experiments (TD-AMS)	Smog chamber	Dry	15%	Isaacman- VanWertz, et al.
HE_VW18_woD	"	"	"	0%	"
HE_SA16_wD	Heating experiments (TD-AMS)	Smog chamber	15–20%	8%	Saha and Grieshop ²⁷
HE_SA16_woD	"	"		0%	"

928 were selected, and details of these estimates are provided in Section S3 of the SI.

929 ^aSame as above. Abbreviations: PAM, potential aerosol mass; ESI, electrospray ionization; LC-

930 TOF-MS, liquid chromatograph/time-of-flight-mass spectrometer; TD, thermodenuder; AMS,

931 aerosol mass spectrometer; CIMS, chemical ionization mass spectrometer.

Exp, #	Precursor	RH	$\log_{10}(k_f [s^{-1}])$	$\log_{10}(k_r [s^{-1}])$	$\log_{10}(D_b \ [\text{m}^2 \ \text{s}^{-1}])$
1	α-pinene	<1%	-2.74 ± 0.07	-5.85 ± 1.45	-17.83 ± 0.05
2	α-pinene	<1%	-2.81 ± 0.09	-5.08 ± 0.18	-17.86 ± 0.01
3	α-pinene	<1%	-2.76 ± 0.07	-5.32 ± 1.25	-17.98 ± 0.06
4	α-pinene	40%	-2.88 ± 0.31	-3.97 ± 0.36	-15.41 ± 1.56
5	α-pinene	40%	-2.92 ± 0.06	-3.98 ± 0.08	-13.30 ± 0.83
6	α-pinene	40%	-2.74 ± 0.21	-3.90 ± 0.31	-15.14 ± 1.76
7	1,3,5-TMB	<1%	-2.81 ± 0.09	-8.30 ± 1.94	-18.13 ± 0.18
8	1,3,5-TMB	<1%	-2.72 ± 0.08	-6.26 ± 1.90	-18.06 ± 0.14
9	1,3,5-TMB	<1%	-2.73 ± 0.08	-6.44 ± 1.85	-18.13 ± 0.07
10	1,3,5-TMB	<1%	-2.52 ± 0.25	-3.70 ± 0.35	-18.62 ± 0.12

933 **Table 3.** Fitted parameters k_f , k_r , and D_b determined on the basis of formation and dilution 934 experiments for the case of the CA-SA18-STD (Table 2). Estimated average parameters over 20 935 iterations are shown with one standard deviation.

936 Abbreviations: RH: relative humidity, TMB: trimethylbenzene.



939Figure 1. Measured volume fraction remaining of SOA from (a) α-pinene under dry conditions,940(b) α-pinene under humid conditions, and (c) 1,3,5-TMB under dry conditions after dilution.941Experimental results of Sato, et al. $^{33, 37}$ and other previous studies $^{23, 25, 55, 56}$ are shown (Table S3942of the SI).

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944



946 Figure 2. C^* distributions of SOA from (a) α -pinene under dry conditions, (b) α -pinene under 947 humid conditions, and (c) 1,3,5-TMB under dry conditions. Data estimated with a molecular

948 corridor approach of the both 1-D function (binary method) and 2-D function (2-D and SPARC

methods) are shown, and the details of these methods are described in Section 2.1. C^*

950 distributions estimated in the previous studies (Table 2) are also shown.





953

954 Figure 3. Volume fraction remaining (VFR) during dilution experiments (a and c) and mass size

distributions $(dM/d\log D_p)$ after 2 h from dilution (b and d) under dry (a and b, experiment #1)

and humid (c and d, experiment #6) conditions. Black lines indicate observational results, and

957 colored lines indicate results of simulation with different C^* compounds. Simulations results with

the lowest RMSE are shown by thick lines. Simulations were conducted without consideration of

959 dimer formation and bulk diffusion.



964 Figure 4. (left) Volume fraction remaining (VFR) during a dilution experiment and (right) 965 simulated C* distributions after a formation experiment under dry conditions (experiment #1). (a and b) Dimer formation rates (k_f) were 10⁻⁶ s⁻¹ and dimer dissociation rates (k_r) were 10⁻⁶ s⁻¹, (c 966 and d) $k_f = 10^{-3} \text{ s}^{-1}$ and $k_r = 10^{-4} \text{ s}^{-1}$, (e and f) $k_f = 10^{-3} \text{ s}^{-1}$ and $k_r = 10^{-6} \text{ s}^{-1}$. Observational results 967 968 are shown in black circles or triangles (by the parameterization of the 2-D and binary methods, 969 respectively, as explained in Section 2.1) and simulation results are shown in each color with

- 970 different D_b . Average and standard deviation of 20 simulations in each setup are shown. For the
- 971 left figures, lines for $\log_{10}(D_b) = -16$ sometime overlap with the results for $\log_{10}(D_b) = -17$









Figure 5. Same with Figure 4, but for an experiment under humid conditions (experiment #6).



Figure 6. (a) Volume fraction remaining (VFR) during the dilution experiments for α -pinene SOA, as constrained by using C^* distributions obtained from the literature (Table 2). (b) C^* distributions of SOA (f_i) just before dilution. Relationship between root mean square error (RMSE) of the measured and simulated VFR and (c) the dimer fractions after the formation experiment or (d) diffusion coefficient (D_b) . Error bars indicate one standard deviation of 20 iterative calculations in individual cases.



Figure 7. Volume fraction remaining (VFR) during dilution experiments for α-pinene SOA under dry (a, experiment #1) and humid (b, experiment #6) conditions or 1,3,5-TMB SOA under dry conditions (c, experiment #9). Each color indicates the contribution of monomer or dimer compounds with a saturation concentration (C^*) range from 0.1 to 1000 µg m⁻³.



Figure 8. Volume fraction remaining (VFR) during dilution experiments under dry (a, experiment
#2) and humid (b, experiment #6) conditions. Vapor wall loss was calculated with the methodology
of Krechmer, et al. ⁴⁶ (open circles) and with an assumption of instantaneous deposition (crosses).



Dilution-induced SOA evaporation