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# Remote sensing-based energy balance for lettuce in an arid environment: influence of management scenarios on irrigation and evapotranspiration modeling

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## Abstract

Efficient irrigation is critical for managing scarce water resources where precipitation is minimal. Field-scale irrigation is largely unaccounted for in landscape evapotranspiration models, primarily due to the unavailability of data and the lack of water balance components in energy balance-based evapotranspiration models. To overcome these challenges, we implemented a remote sensing-based energy and water balance model BAITSSS (Backward-Averaged Iterative Two-Source Surface temperature Solution) to calculate evapotranspiration (ET) and irrigation requirements of winter lettuce in the arid environment of the Lower Colorado River Basin. Predicted evapotranspiration and irrigation were compared against data from twelve eddy covariance (EC) sites for wide range of soil hydraulic properties operating between 2016 and 2020 and the applied irrigation, respectively. BAITSSS estimated evapotranspiration and irrigation based on vegetative formation, weather demand, soil hydraulic characteristics, and predefined management allowed depletion (MAD) (0.4–0.6). Ground-based weather data, Sentinel-2-based vegetation indices, and SSURGO (NRCS soil survey database) soil moisture characteristics were model inputs. The results showed mean seasonal ET from BAITSSS and EC were comparable, differing on average by about 7% based on a constant rooting depth (500 mm) and MAD of 0.5 for entire crop growth stages. Variations in daily and seasonal ET were mainly due to differences in applied and model-simulated irrigation. Seasonal values of applied and simulated irrigation closely agreed (~6%) in most sites, though some sites applied irrigation more effectively than others. Overall, this study provided insight into consumptive water use and field-scale irrigation practices, as well as the capabilities and limitations of model-simulated ET and irrigation.

## Introduction

Remote sensing-based energy balance models are frequently utilized to quantify crop water use (evapotranspiration) for water management (Allen et al. 2005; Anderson et al. 2012). To manage agricultural water, water managers compare evapotranspiration (ET) to applied water (groundwater pumping and irrigation diversions) (Allen et al. 2005). Field-scale irrigation practices (applied water) depend on multiple factors including farmers' experience, behavior, adaptability, investment in infrastructure, availability of water resources, and many others (Leng et al. 2017; de Vito et al. 2017; Foster et al. 2019). Landscape irrigation mapping needs information regarding land management (where and when humans have provided water or supplemented rain-fed crops) (Ozdogan et al. 2010). Accounting for field-scale irrigation in remote sensing-based landscape evapotranspiration (ET) maps is challenging due to inadequate spatial and temporal

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resolution (Ha et al. 2013). Irrigation alters soil moisture, and high spatial and temporal-resolution soil moisture maps can capture this critical piece of information related to the applied irrigation timing and amount. However, currently available soil moisture maps are mostly coarser resolution ( $> 25$  km) (Brocca et al. 2017; Peng et al. 2021) and often provide information only about the upper soil moisture layer, for instance, microwave ( $\sim 50$  mm) (Kerr 2007; Collow et al. 2012; Ye et al. 2015; Akbar et al. 2018; Massari et al. 2021; Peng et al. 2021). Thus, these maps are generally useless where higher spatial resolution information is needed (Fang et al. 2019).

Thermal-based ET models often do not utilize irrigation directly in their surface energy balance algorithm, whereas the soil water balance-based ET models need information regarding irrigation and precipitation. However, the BAITSSS (Backward-Averaged Iterative Two-Source Surface temperature and energy balance Solution) (Dhungel et al. 2016, 2019) model accommodates water balance components in the energy balance algorithm and is used in operational settings (Dhungel et al. 2020). As field-scale irrigation information is not generally available, auto-irrigation in BAITSSS is based on weather demand, canopy formation (vegetation indices and physical properties), soil moisture characteristics, and predefined manageable allowed depletion (MAD). The estimated ET from BAITSSS represents the crop water requirements based on the irrigation management practice that should generally correspond to the applied irrigation in the field. However, discrepancies may arise due to the limitations in capturing the farmer's irrigation practices. Thus, a robust sensitivity analysis is needed to implement this model for regional irrigation scheduling and analysis effectively.

Previously, detailed evaluations of ET estimated with BAITSSS were conducted using eddy covariance (EC) data from 12 lettuce fields in the Lower Colorado River Basin (Dhungel et al. 2022), and using lysimeter ET data for corn and sorghum fields in Bushland, Texas (Dhungel et al. 2021). Both studies used the applied irrigation data as a model input, and the predicted ET closely agreed with both lysimeter and EC. In addition, the results showed the increased accuracy in ET estimation was due to the availability of applied irrigation in the field that was critical to capturing ET spikes right after the irrigation and tracking the continued decrease of ET after irrigation, especially during the partial cover period.

We extended the previous work focusing on lettuce in the Lower Colorado River Basin at the same sites and years (2016–2020). However, in this work, model calculations are made assuming the applied irrigation amounts, and timings are unknown. Irrigation scheduling in this region is commonly done based on the farmers'/producers' experience, rather than tracking soil moisture using field-based sensors

or using specific computer-based irrigation simulation models. Farmers tend to well water and not stress crops due to the severe yield consequences of depleting short-season vegetable crops with shallow root zones (Kuslu et al. 2008; Xu and Leskovar 2014). To avoid water-related stress in most crop models, a management allowable depletion (MAD) of 0.5 is generally adopted (Clark 2000; Callison 2012; Malik and Dechmi 2019). However, this rough guideline is largely based on major grain/cereal crops; the sensitivity of vegetable crops is less known and may have lower permissible MAD (Hartz 1996; Thompson et al. 2007). General characteristics of ET simulation in winter vegetables including lettuce are shallow root depth compared to grain crops, short growing seasons, and longer partial cover period (Thorup-Kristensen 2001; Escarabajal-Henarejos et al. 2015; Roux et al. 2016; Fisher et al. 2017).

The overall objective of this study was to compare ET from eddy covariance (EC) with ET estimated from BAITSSS using auto-irrigation based on multiple predefined MAD's (0.4–0.6) (Allen et al. 1998; Bartlett et al. 2015). Irrigation and subsequent ET were estimated from the widely recommended MAD procedure suggested by university extension (Aguilar et al. 2015; Peters et al. 2013) which had not been fully investigated in landscape-scale remote sensing energy balance algorithm previously because of the aforementioned limitations.

## Methodology

### Study area

The study area comprised about 600 km<sup>2</sup> in commercially managed lettuce fields around Yuma, Arizona in the Lower Colorado River Basin. The Colorado River serves as the primary water source as precipitation is minimal ( $\sim 80$  mm/year) (Arguez et al. 2012). Sprinkler and furrow irrigation were used within the same season, with sprinkler irrigation used to help establish a crop and furrow irrigation used for the remainder of the season. Evapotranspiration estimated from BAITSSS was compared to twelve EC sites between 2016 and 2020. These sites were NGIDD 19–20a, YID 18b, YCWUA 18a, YID 17a, YID 19–20a, YCWUA 19–20a, YID 17d, YID 17b, YID 17c, YCWUA 17–18b, YCWUA 17–18a, and YID 16. NGIDD represents North Gila Valley Irrigation and Drainage District, YID represents Yuma Irrigation District, and YCWUA represents Yuma County Water Users' Association. These lettuce fields were mostly rectangular shapes with a minimum and maximum area of 29,630 m<sup>2</sup> (YID19–20a) and 136,290 m<sup>2</sup> (YCWUA19–20a), respectively with an average of 105,699 m<sup>2</sup>. Planting dates varied between September and November.

System components included three-dimensional sonic anemometers (CSAT3 and CSAT3A, Campbell Scientific, Logan, UT), open path infrared gas analyzers (IRGASON—Campbell Scientific and Licor 7500—Licor Inc., Lincoln, NE), net radiometers (NR-Lite and CNR 4, Kipp and Zonen, Delft, Netherlands), soil heat flux plates (Hukseflux, Delft, Netherlands), and air temperature and humidity probes (Vaisala, Helsinki, Finland). Loggers and covariance sensors were calibrated by the manufacturer in 2016 and 2017. Zero and span of infrared gas analyzers (IRGA) were done in July 2017 and again in July 2018.

Micrometeorological observations (~ 108 variables per 30 min. time step), configured under Campbell Scientific's EasyFlux DL™ (Logan, UT) program,<sup>1</sup> were collected to allow continuous data measurements during the cropping cycle. High-frequency three-dimensional wind, sonic temperature, and CO<sub>2</sub> and H<sub>2</sub>O raw observations were collected at 20 Hz. A 30-min block-averaged fluxes, including the Webb–Pearman–Leuning (Webb et al. 1980) corrections were simultaneously stored and raw data were processed with EddyPro software in Express Mode (Fratini and Mauder 2014). We computed the changes in energy storage within the soil mass above the two heat flux plates. We removed the data spike followed the outlined methodology described by Vickers and Mahrt (1997). The online gap-filling tool (<http://www.bgc-jena.mpg.de/~MDIwork/eddyproc/method.php>); employing techniques as described in Falge et al. (2001) and Reichstein et al. (2005) were used to fill time gaps including those when the friction velocity ( $U^*$ ) was less than 0.15 m/s. We conducted energy balance closure at daily time steps by regressing turbulent energy (latent heat and sensible heat) against radiative energy (net radiation minus ground heat flux) and correcting for net energy storage via photosynthesis (Anderson and Wang 2014). We enforced energy balance closure at 30-min time steps by assigning residuals to latent heat (LE) fluxes (Rosa and Tanny 2015). We conducted two-dimensional flux footprint analyses using an R script provided by Kljun et al. (2015). Eighty percent of the flux footprints lay within plot boundaries with minor exceptions where daily fluxes were computed by summing the 30-min LE flux values. Further details of EC data and processing can be found in French et al. (in. prep) and Dhungel et al. (2022).

<sup>1</sup> Mention of trade names or commercial products in this publication is solely for the purpose of providing specific information and does not imply recommendation or endorsement by the U.S. Department of Agriculture.

## BAITSSS model

The detailed model description (BAITSSS configuration as shown in Fig. 1) and complete set of the equations utilized in the BAITSSS model can be found in Dhungel et al. (2016, 2019, 2020). This section summarizes the model and presents some primary equations. Flux gradient equations were used to compute latent heat flux (LE) and sensible heat flux ( $H$ ), whereas Jarvis-type canopy resistance ( $r_{sc}$ ) was used for estimating transpiration. BAITSSS estimates ET based on weather, soil moisture conditions, and canopy development (vegetation indices and physical properties), where the surface temperature required for the energy balance is computed iteratively for each time step. Irrigation ( $I_{rr}$ ) in BAITSSS is triggered when root zone soil moisture ( $\theta_{root}$ ) falls below the threshold moisture content ( $\theta_t$ ),

$$I_{rr} = \begin{cases} (\theta_{fc} - \theta_{root})d_{root} & \text{if } \theta_{root} < \theta_t \\ 0 & \text{if } \theta_{root} \geq \theta_t \end{cases} \quad (1)$$

RAW is computed based on the management allowable depletion (MAD) and total available water (TAW),

$$RAW = MAD(\theta_{fc} - \theta_{wp}) = MAD \cdot TAW \quad (2)$$

where  $TAW = \theta_{fc} - \theta_{wp}$  and  $\theta_{wp}$  is the wilting point soil moisture content and  $\theta_{fc}$  is field capacity. In this study, a constant MAD (0.4–0.6) for entire crop growth stages was used. The threshold  $\theta_t$  is computed from readily available water (RAW) and  $\theta_{fc}$  (Eq. 3),

$$\theta_t = \theta_{fc} - RAW \quad (3)$$

A reduction in MAD causes a corresponding reduction in RAW and consequently increases  $\theta_t$ . When MAD is equal to 1,  $\theta_t$  becomes  $\theta_{wp}$ , i.e., the model delays irrigating until  $\theta_{root}$  becomes  $\theta_{wp}$ . When  $MAD = 0$  ( $\theta_t = \theta_{fc}$ ), the model does not let the soil moisture go below  $\theta_{fc}$ . Each irrigation event restores soil moisture to  $\theta_{fc}$  for both the soil surface layer ( $d_{sur}$ ) and the root zone layer ( $d_{root}$ ) (as assumed sprinkler irrigation). Two concurrent accountings of soil moisture are made, one for the soil surface and one for the full root zone. Water content at the soil surface ( $\theta_{sur}$ ) (Eq. 4) is computed as

$$\theta_{sur} = \theta_{sur(i-1)} + \frac{(P + I_{rr} - S_{run}) - E_{ss} - T_e}{d_{sur}} + CR_e - DP_e \quad (4)$$

where  $\theta_{sur(i-1)}$  is the volumetric water content (VWC) of the surface layer at the previous time step ( $m^3 m^{-3}$ ),  $E_{ss}$  is soil surface evaporation (mm),  $P$  is precipitation (mm),  $S_{run}$  is surface runoff (mm),  $I_{rr}$  is irrigation,  $d_{sur}$  (= 150 mm) is soil surface depth,  $DP_e$  is deep percolated water from the upper soil layer to the root zone ( $m^3 m^{-3}$ ),  $CR_e$  is capillary rise



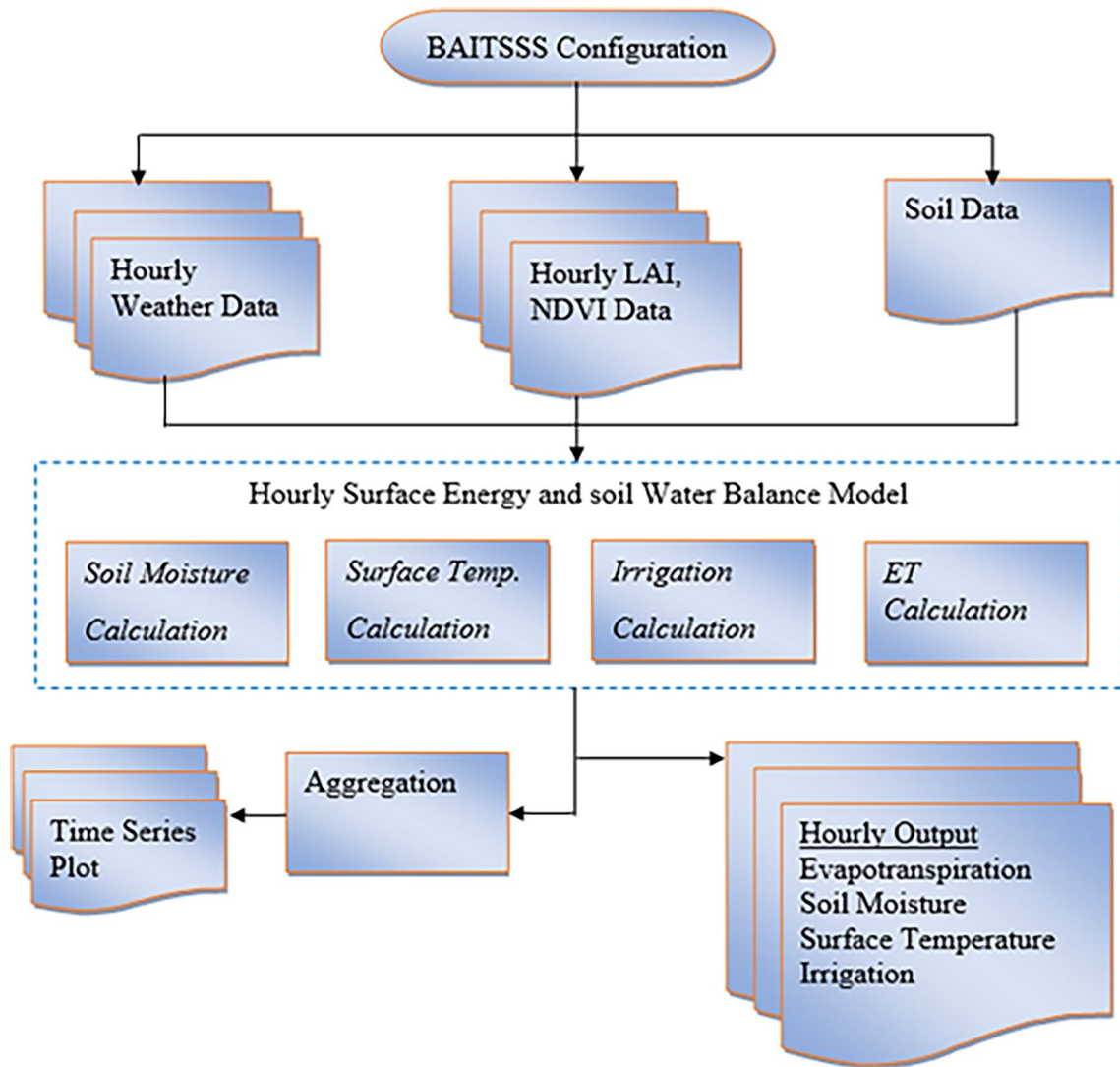


Fig. 1 BAITSS configuration with inputs (weather, vegetation indices, and soil data), intermediate, and output

from the root zone into the first soil surface layer ( $\text{m}^3 \text{m}^{-3}$ ), and  $T_e$  is transpiration from the soil surface layer (mm).

The water content of the root zone ( $\theta_{\text{root}}$ ) is similarly evaluated as Eq. (5), where root zone control volume consists of the soil surface layer.

$$\theta_{\text{root}} = \theta_{\text{root}(i-1)} + \frac{(P + I_{\text{rr}} - S_{\text{run}}) - E_{\text{ss}} - T}{d_{\text{root}}} + \text{CR} - \text{DP} \tag{5}$$

where  $\theta_{\text{root}(i-1)}$  is VWC of the root zone from previous time steps,  $T$  is transpiration from vegetation,  $d_{\text{root}}$  ( $= 500 \text{ mm}$ ) is rooting depth for entire growth stages, DP is deep percolation below the root zone, and CR is capillary rise from the third layer to the root zone.  $\text{CR}_e$ , CR, and  $T_e$  were neglected in the present analysis to simplify the soil water balance. Currently, the model permits soil to be dry at the surface

( $d_{\text{sur}}$ ) to an air-dry condition (Campbell and Norman 2000). The  $\theta_{\text{sur}}$  can elevate to saturated soil moisture ( $\theta_{\text{sat}}$ ), while  $\theta_{\text{root}}$  is limited to  $\theta_{\text{fc}}$ .

A Jarvis-type equation was used to compute canopy resistance ( $r_{\text{sc}}$ ) with weighting functions representing plant response to solar radiation ( $F_1$ ), air temperature ( $F_2$ ), vapor pressure deficit ( $F_3$ ), and soil moisture ( $F_4$ ) (Alfieri et al. 2008; Kumar et al. 2011),

$$r_{\text{sc}} = \frac{R_{\text{c\_min}}}{f_c \text{LAI} F_1 F_2 F_3 F_4} \tag{6}$$

where LAI is leaf area index and  $f_c$  is fraction of canopy cover. The fraction of canopy cover ( $f_c$ ) is calculated based on NDVI as suggested by Gutman and Ignatov (1998), Dhungel et al. (2016).

The minimum value of canopy resistance ( $R_{c\_min}$ ) is set at  $40 \text{ sm}^{-1}$  as per Kumar et al. (2011), Dhungel et al. (2019) for irrigated agricultural landscapes. When  $F_1 = F_2 = F_3 = F_4 \sim 1$ ,  $r_{sc}$  becomes the minimum.

A logistic growth-type equation ( $W$ ) is used to calculate the effects of the available water fraction (AWF) for  $F_4$  in Eq. (6) (Anderson et al. 2007).

$$AWF = \frac{\theta_{root} - \theta_{wp}}{\theta_{fc} - \theta_{wp}} \tag{7}$$

$$F_4 = \frac{\ln(W)}{\ln(W_f)} \tag{8}$$

$$W = \frac{W_0 W_f}{W_0 + (W_f - W_0) \exp(-\mu AWF)} \tag{9}$$

$W_0$ ,  $W_f$ , and  $\mu$  are logistic growth equation coefficients, where  $W_0 = 1$ ,  $W_f = 800$ , and  $\mu = 12$ .

BAITSSS adjusted energy and water balance components using  $f_c$  to partition soil and canopy components.

### Weather, soil, and vegetation data

The primary inputs of BAITSSS are weather, vegetation indices, and soil data (Fig. 1). Soil hydraulic characteristics data ( $\theta_{fc}$  and  $\theta_{wp}$ ) were acquired from the area and depth-averaged NRCS soil survey database SSURGO as described by Wiczorek (2014) and were added in Table 1. AZMET (Arizona Meteorological Network) and the EC sites provided the hourly weather data. Normalized difference vegetation index (NDVI) was derived from linearly interpolated Sentinel-2 and leaf area index (LAI) was estimated from the empirical relationship relating NDVI to LAI. Relatively dry initial surface ( $\theta = 0.05 \text{ m}^3 \text{ m}^{-3}$ ) and root zone ( $\theta = \theta_{wp}$ ) soil moisture contents were considered for all sites since minimal precipitation occurred 15 days before planting (i.e., start of the simulation).

## Results

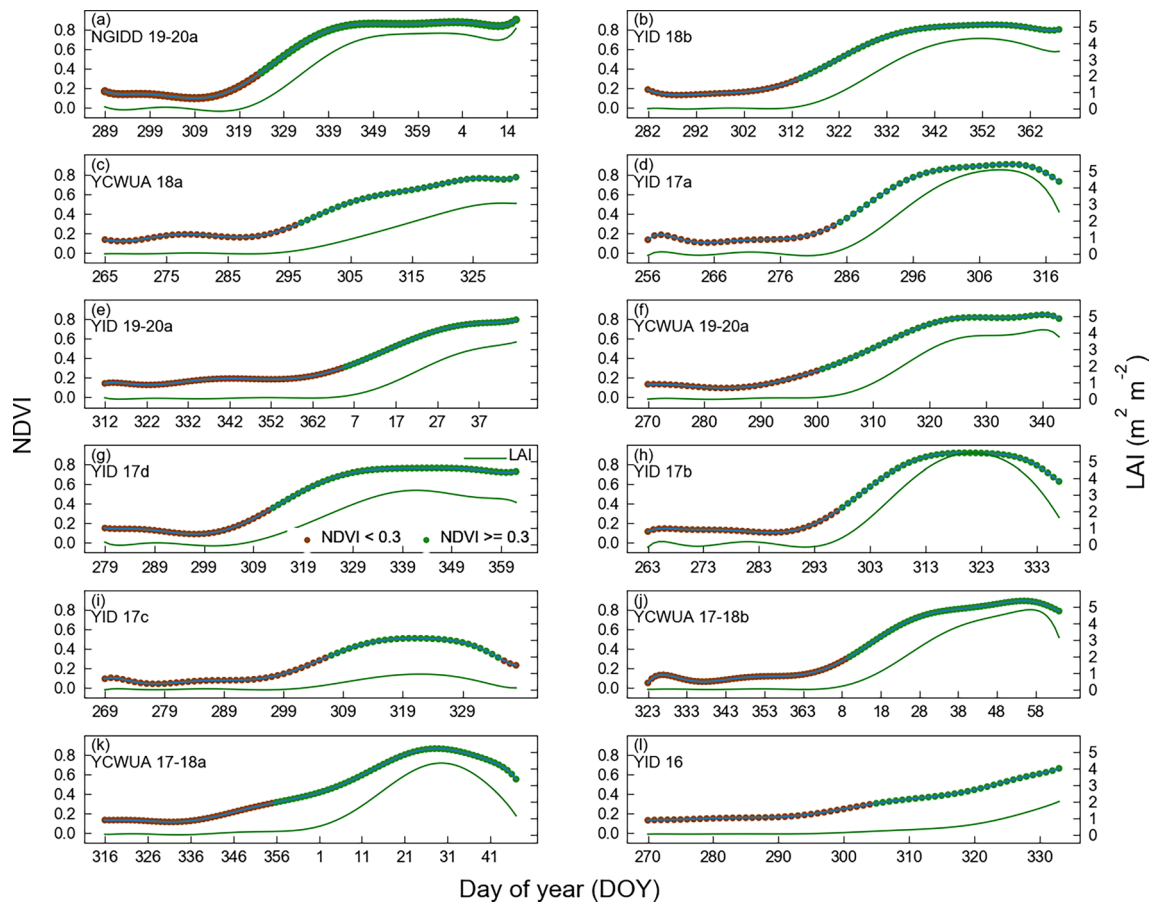
### Weather and vegetation indices

The average flux (diurnal cycle; 24 h) of solar radiation decreased from  $\sim 250$  to  $\sim 150 \text{ W m}^{-2}$  and air temperature decreased from  $\sim 30$  to  $\sim 10 \text{ }^\circ\text{C}$  toward the winter solstice. During this period, vegetation indices were mostly increasing creating phase differences between solar radiation and air temperature (Fig. 2). To reduce daily scatter, Polyfit and polyval functions of the polynomial module were used within the Python-based library NumPy to smooth the

**Table 1** Soil moisture characteristics and water balance components based on MAD of 0.5 of lettuce for multiple sites for various years around Yuma, AZ

S.no	Name	Irr <sub>applied</sub> (mm)	Irr <sub>BAITSSS</sub> (mm)	ET <sub>EC</sub> (mm)	ET <sub>BAITSSS</sub> (mm)	P (mm)	$\theta_{fc}$ ( $\text{m}^3 \text{ m}^{-3}$ )	$\theta_t$ ( $\text{m}^3 \text{ m}^{-3}$ )	Change ( $I_r$ ) %	Change (ET) %
a	NGIDD 19-20a	308.6	271.9	314.1	229.1	32.3	0.19	0.10	-11	-27
b	YID 18b	274.2	304.8	247.0	251.2	17.8	0.32	0.22	11	1
c	YCWUA 18a	442.9	362.0	263.2	303.2	24.1	0.42	0.34	-18	15
d	YID 17a	257.0	300.9	269.7	238.3	0.0	0.30	0.20	17	-11
e	YID 19-20a	435.1	227.1	297.5	227.4	58.1	0.19	0.10	-47	-23
f	YCWUA 19-20a	406.8	361.3	257.4	327.5	34.3	0.42	0.34	-11	27
g	YID 17d	281.2	401.5	249.0	328.2	1.3	0.30	0.20	42	31
h	YID 17b	328.6	401.3	306.9	306.9	1.3	0.30	0.20	22	0
i	YID 17c	243.1	305.1	267.6	249.9	1.3	0.32	0.22	25	-6
j	YCWUA 17-18b	411.5	458.5	302.8	374.5	2.0	0.32	0.22	11	23
k	YCWUA 17-18a	335.3	441.6	234.1	396.8	2.0	0.42	0.34	31	69
l	YID 16	310.3	305.4	245.6	221.2	1.3	0.32	0.22	-1	-9

Data are reported for each cropping season, which varied in length



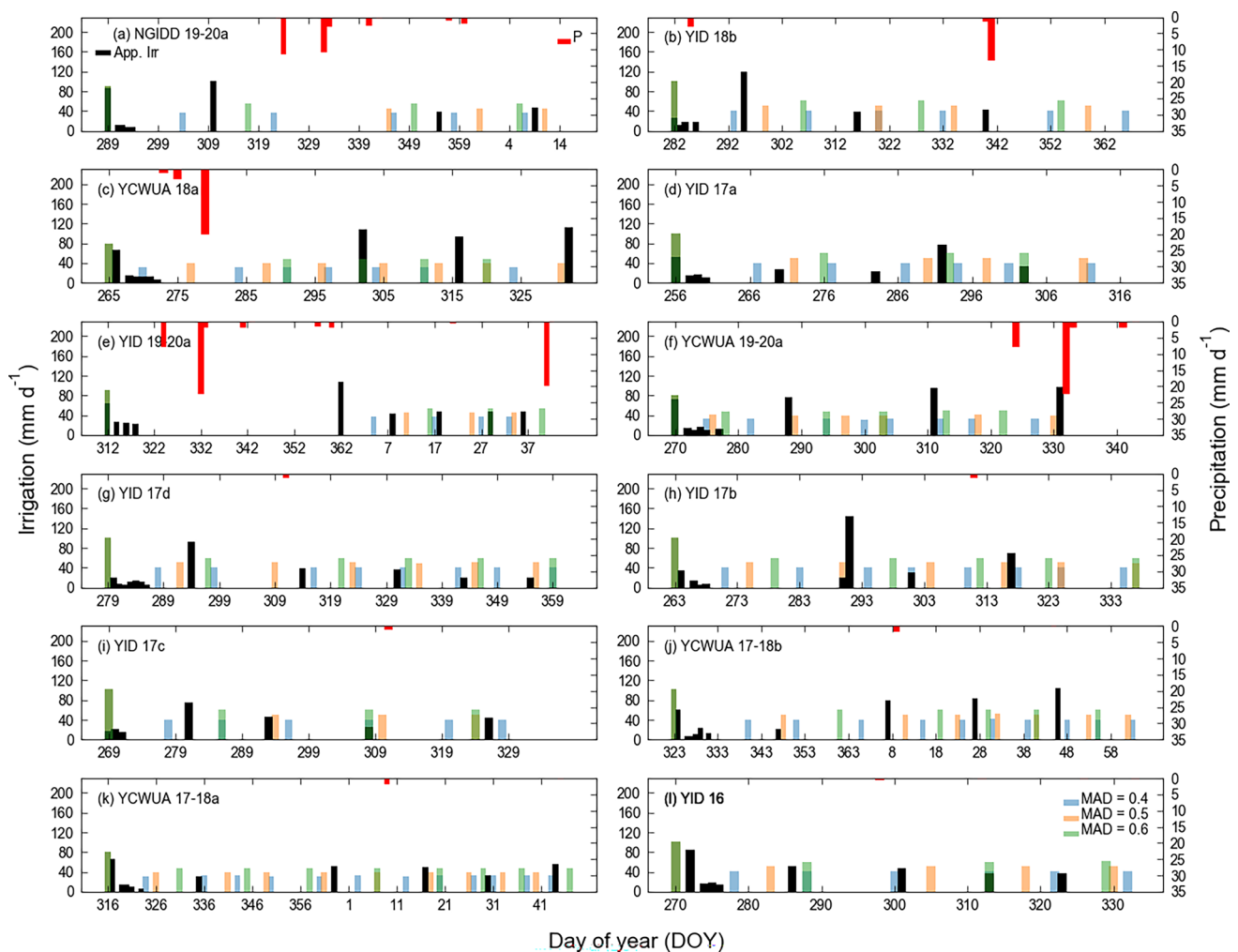
**Fig. 2** Daily values of vegetation indices (LAI and NDVI) from Sentinel-2 for various sites for multiple years around Yuma, AZ

NDVI and LAI values. NDVI period ( $< 0.3$  and  $\geq 0.3$ ) were separated for differentiating bare soil and the non-vegetated period from the vegetative period (based on color code). The maximum duration when  $NDVI < 0.3$  was 58 days for the site YID 19–20a (Fig. 2e) while the minimum was 29 days for the site YID 17a (Fig. 2d) with an average value of 38 days. Lettuce harvesting is done at peak or close to peak vegetative cover. Most of the sites showed cut-off in NDVI and LAI at the peak with some exceptions. The vegetation indices showed some variability among the sites, with the maximum values of NDVI and LAI being  $\sim 0.9$  and  $\sim 5.5 \text{ m}^2 \text{ m}^{-2}$ , respectively. The site YID 17c had the smallest vegetation index values (Fig. 2i, maximum NDVI  $\sim 0.4$  and LAI  $\sim 1 \text{ m}^2 \text{ m}^{-2}$ ) followed by YCWUA 17–18a (Fig. 2k, maximum NDVI  $\sim 0.6$  and LAI  $\sim 2 \text{ m}^2 \text{ m}^{-2}$ ).

### Irrigation frequency and depth

Figure 3 shows applied irrigation, precipitation, and the progression of simulated irrigation based on multiple MADs on a daily time scale. Field data showed that irrigation was generally applied on the first day of planting. Due to the

assumed low initial moisture content ( $= \theta_{wp}$ ) of the root zone, BAITSSS auto-irrigation also simulated on the first day ( $\theta_{wp} < \theta_t$  at the start of simulation) for all considered MADs. The irrigation depth (mm) from BAITSSS for the first application was the largest of the simulation because the model did not let the moisture go below the  $\theta_t$  thereafter (Fig. 3). For all sites, the second simulated irrigation was triggered earliest when  $MAD = 0.4$ , followed by  $MAD = 0.5$  and  $MAD = 0.6$ . Applied irrigation in the field showed a higher frequency of small irrigation events ( $< 20 \text{ mm}$ ) right after planting in all sites, which were mostly on daily basis or some days apart after the first irrigation. The frequent water applications were due to sprinkler irrigation being used during crop establishment to create a cool microclimate (Sanchez et al. 2009). However, the average difference between the first and second simulated irrigation events among all sites was 13, 18, and 24 days for MAD of 0.4, 0.5, and 0.6, respectively. The shortest difference was 5 days for YCWUA 18a (Fig. 3c) for MAD of 0.4 while the longest difference was 69 days for YID 19–20a (Fig. 3e) for MAD of 0.6, which was pushed back by multiple precipitation events. It was evident that precipitation delayed both applied and



**Fig. 3** Daily applied irrigation and precipitation, and progression of simulated irrigation from BAITSSS based on various management allow-depletions (MADs) around Yuma, AZ

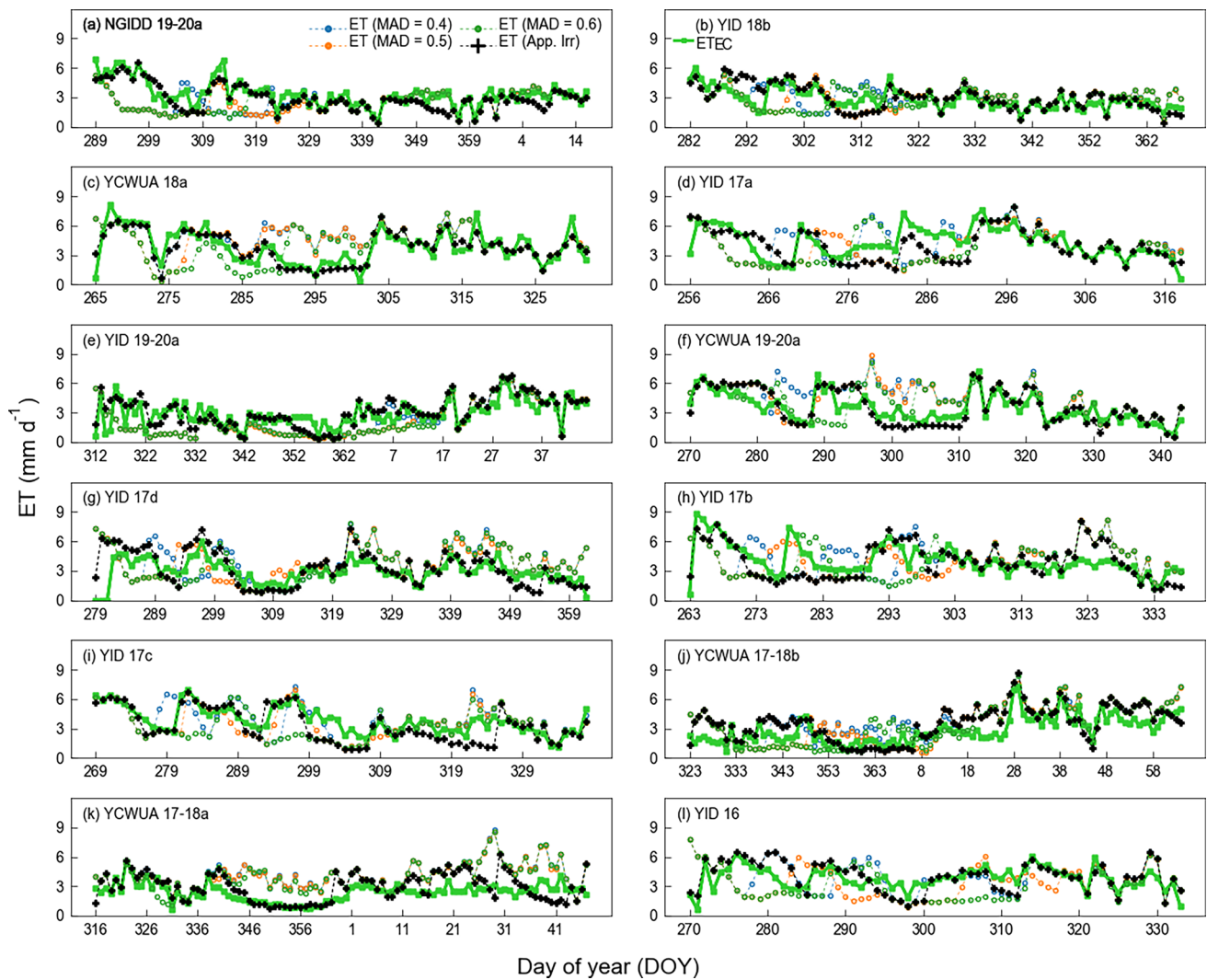
simulated irrigation (e.g., NGIDD 19–20a, Fig. 3a, YID 19–20a, Fig. 3e). The frequency and the application depth of simulated irrigation differed from applied irrigation later in the season as well. As indicated earlier, irrigation in BAITSSS (timing and depth) was triggered based on vegetative and weather demand with predefined management scenarios. Farmers normally rely on visual inspection of the crop, historical experience based on previous seasons, their knowledge about soil water retention, and energy cost. The applied irrigation depth varied from  $< 20$  to  $> 80$  mm.

### Evapotranspiration time series

Visual inspection of daily ET data (Fig. 4) showed that BAITSSS simulated ET closely followed EC measured values. However, differences were evident during the partial cover period mainly due to soil evaporation (YID 17a, Fig. 4d; YCWUA 19–20a, Fig. 4f; YID 17b, Fig. 4h). The

ET spikes from BAITSSS followed the simulated irrigation based on the respective MADs. The mismatch between EC and BAITSSS daily ET spikes was generally smaller during the full cover period, though it was not completely eliminated (YID 17d, Fig. 4g; YCWUA 17–18a, Fig. 4k). Obviously, the simulated ET when using applied irrigation [ET (App.Irr)] data as inputs was in closer agreement with EC measurements compared to simulated irrigation (Fig. 4). Eddy covariance measured evapotranspiration [ET (EC)] generally stayed higher right after planting because of the multiple small irrigation events, whereas ET predicted from BAITSSS showed a decline after the first irrigation event due to larger gaps in simulated irrigation events (NGIDD 19–20a, Fig. 4a; YID 17a, Fig. 4d). The maximum daily ET value was  $\sim 8$  mm from both EC and BAITSSS (YID 17b, Fig. 4h); however, ET was generally less than 6 mm most of the time. Evapotranspiration from EC for YCWUA 17–18a (Fig. 4k) was unrealistically stable ( $\sim 3$  mm) and





**Fig. 4** Daily evapotranspiration (ET) for irrigated lettuce at various sites around Yuma, AZ.  $ET_{EC}$  is ET measured with eddy covariance,  $ET_{BAITSSS}$  is ET simulated with BAITSSS using irrigation data

less responsive to irrigation events after the new year than BAITSSS ET.

### Soil moisture at root and $F_4$

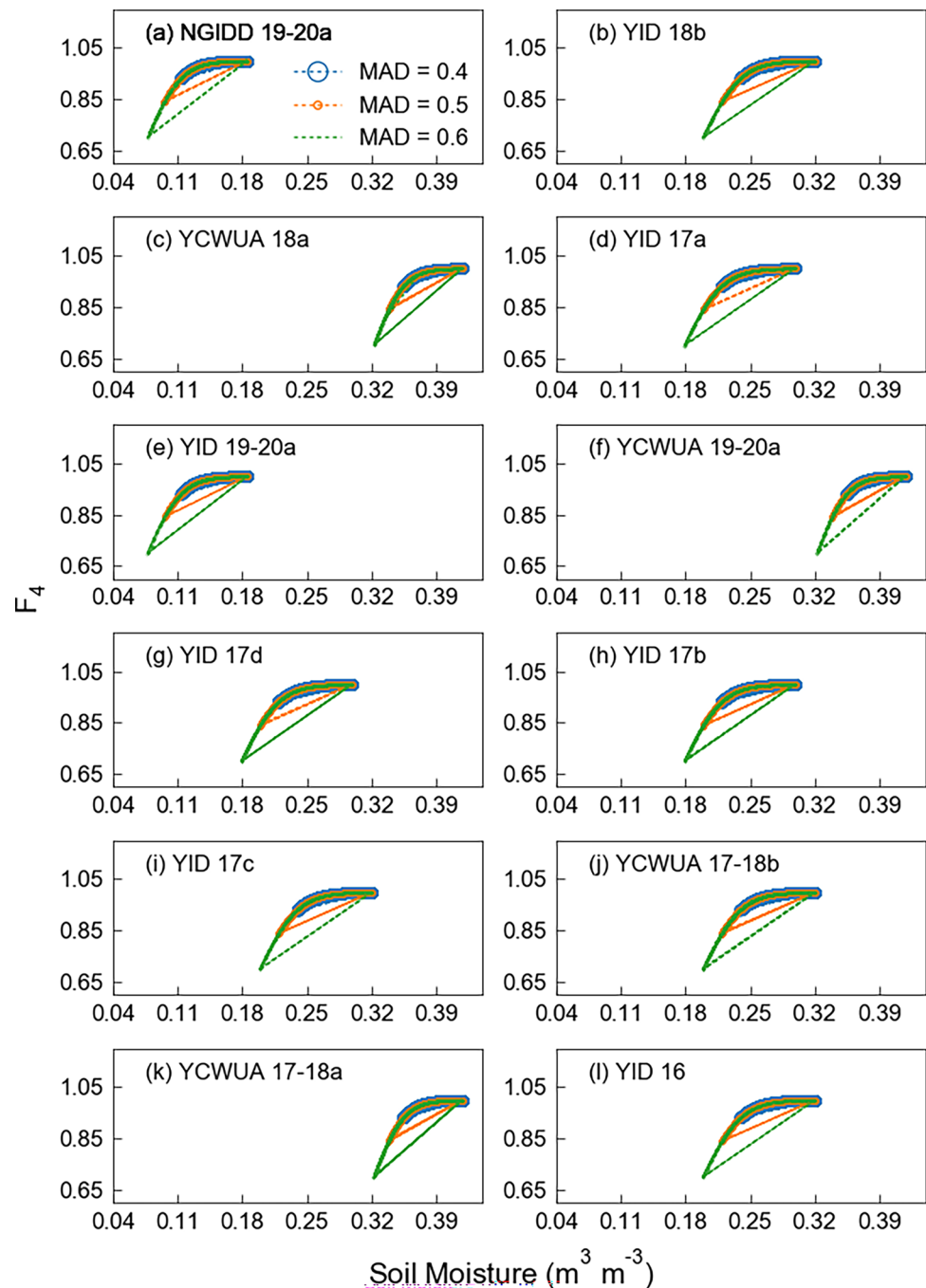
Figure 5 shows the relationship between  $F_4$  and  $\theta_{root}$  based on the logistic equation (Eqs. 7–9). The gradual change in shape and the transition points of these curves provide the basis for triggering irrigation. The upper value of soil moisture on the horizontal axis of each plot is field capacity, whereas the lower value was threshold moisture content (Fig. 5). NGIDD 19–20a (Fig. 5a) and YID 19–20a (Fig. 5e) had the lowest field capacity ( $\theta_{fc}$ ) of  $0.19 \text{ m}^3 \text{ m}^{-3}$ , while YCWUA 18a (Fig. 5c), YCWUA 19–20a (Fig. 5f), and YCWUA 17–18a (Fig. 5k) had the highest  $\theta_{fc}$  of  $0.42 \text{ m}^3 \text{ m}^{-3}$ . As per the irrigation rules, the model does not

as input, and the remaining curves are ET simulated with BAITSSS using various levels of management allowable depletion (MAD) to trigger simulated irrigation

permit the soil moisture to go below  $\theta_l$  to avoid moisture-related stress and each irrigation event brings moisture back to  $\theta_{fc}$  (Fig. 3). The  $F_4$ - $\theta_{root}$  curves for individual MADs overlap each other for individual sites because of the identical shape of AWF ( $\theta_{wp}$ ,  $\theta_{fc}$ ). The  $F_4$  reached the upper limit i.e.,  $\sim 1$  when  $\theta_{root} = \theta_{fc}$  for all MADs. The lowest value of  $F_4$  was  $\sim 0.9$ ,  $\sim 0.85 \sim 0.7$  for MADs of 0.4, 0.5, and 0.6, respectively. As a reminder, higher values of the weighing functions  $F_1$ ,  $F_2$ ,  $F_3$ , and  $F_4$  create lower  $r_{sc}$  and ultimately smaller resistance to transpiration.

The simulated root zone soil moisture ( $\theta_{root}$ ) is shown in Fig. 6. The soil moisture time course corresponded to the progression of MAD-based irrigation shown in Fig. 3 and ET spikes shown in Fig. 4. The first irrigation event at the start of the simulation restored  $\theta_{root}$  to  $\theta_{fc}$  in all sites. The subsequent, i.e., second irrigation (excluding the start

**Fig. 5** Hourly weighting functions relating plant response ( $F_4$ ) to root zone soil moisture  $\theta_{\text{root}}$  for various sites around Yuma, AZ. The straight line connecting the endpoints was added to help distinguish three curves in each plot



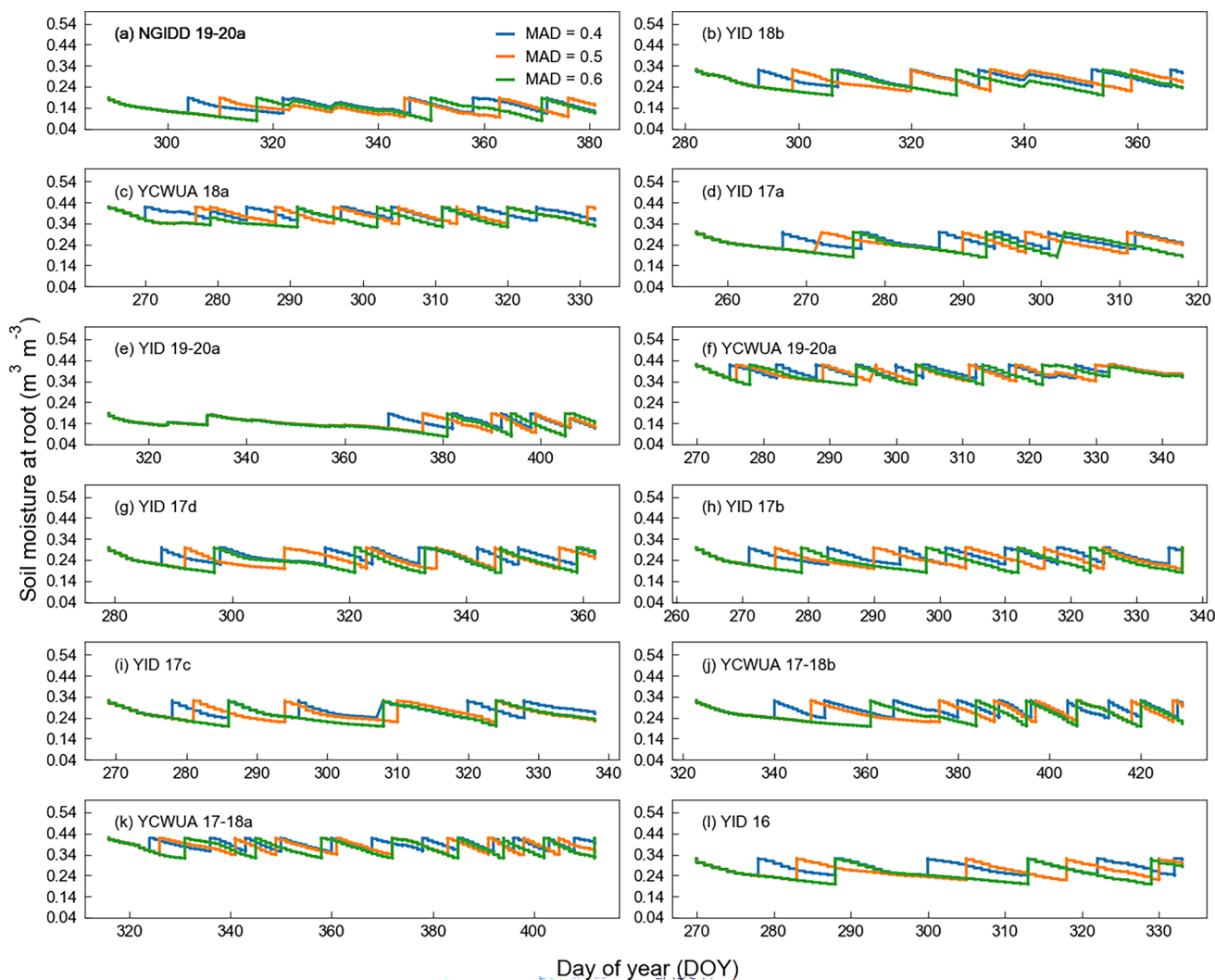
of simulation) was triggered first for MAD=0.4 and last for MAD=0.6. The  $\theta_{\text{sur}}$  followed a similar trend of  $\theta_{\text{root}}$  i.e., continuously decreased until the next irrigation or precipitation occurred (not shown).

### Seasonal cumulative evapotranspiration

A comparison of seasonal ET between EC and BAITSSS showed mixed results (Fig. 7). For instance, some sites showed good agreement (within  $\pm 10\%$ ) (YID 18b, Fig. 7b;

YID 17a, Fig. 7d; YID 17b, Fig. 7h; YID 17c, Fig. 7i; YID 16, Fig. 7l); some had positive bias (YCWUA 18a, Fig. 7c; YCWUA 19–20a, Fig. 7f; YID 17d, Fig. 7g; YCWUA 17–18b, Fig. 7j; YCWUA 17–18a, Fig. 7k), while others had negative bias as well (NGIDD 19–20a, Fig. 7a; YID 19–20a, Fig. 7e). The % values in the legend of each plot indicate the difference between simulated and measured cumulative ET from EC when BAITSSS was run with a MAD of 0.5 (Table 1). The largest positive bias was 69% for YCWUA 17–18a (Fig. 7k) while the largest negative





**Fig. 6** Estimated daily root zone soil moisture from BAITSSS for various sites around Yuma, AZ

bias was  $-27\%$  for NGIDD 19–20a (Fig. 7a). For YCWUA 17–18a, unrealistically stable ET from EC, (i.e., less responsive to irrigation events during the latter part of the simulation after the start of the new year) led to the large differences. For BAITSSS simulations with  $MAD = 0.5$ , the average difference between BAITSSS and EC cumulative ET was 7% over all sites and years. The average difference dropped to 2% when NGIDD 19–20a was excluded from the analysis. The MAD producing the best cumulative ET agreement varied by site. As expected, seasonal ET from BAITSSS was largest from  $MAD = 0.4$  and smallest from  $MAD = 0.6$  in all sites due to fewer irrigation events and less soil evaporation as MAD increases. The largest difference between the seasonal ET values between MAD of 0.4 and 0.6 was 19% for YID 16 (Fig. 7l), while the lowest difference was 2% for NGIDD 19–20a with average of  $\sim 10\%$  among the sites.

We further assessed cumulative ET by separating the period when the soil evaporation was dominant ( $NDVI < 0.3$ ) and afterward (Fig. 8). It was interesting to see that some sites showed better agreement during the period  $NDVI < 0.3$  (YID 18b, YCWUA 18a, YID 17d, YCWUA 17–18b) while other showed during  $NDVI > 0.3$  (NGIDD 19–20a, YID 17a, YID 19–20a, YID 17c, YID 16). The largest negative bias between EC and BAITSSS seasonal ET ( $-49\%$ ) occurred when  $NDVI < 0.3$  at the sites with the largest underestimation of seasonal ET (Figs. 7a, e and 8a, e, respectively). These sites had nominal differences in cumulative ET ( $-6\%$  to  $-1\%$ , respectively) during the period when  $NDVI > 0.3$ . Conversely, the largest positive bias during the period when  $NDVI < 0.3$  was 54%, and 79% during the period when  $NDVI > 0.3$  for YCWUA 17–18a. This site was the largest contributor to the seasonal ET difference between EC and BAITSSS (Fig. 7k).

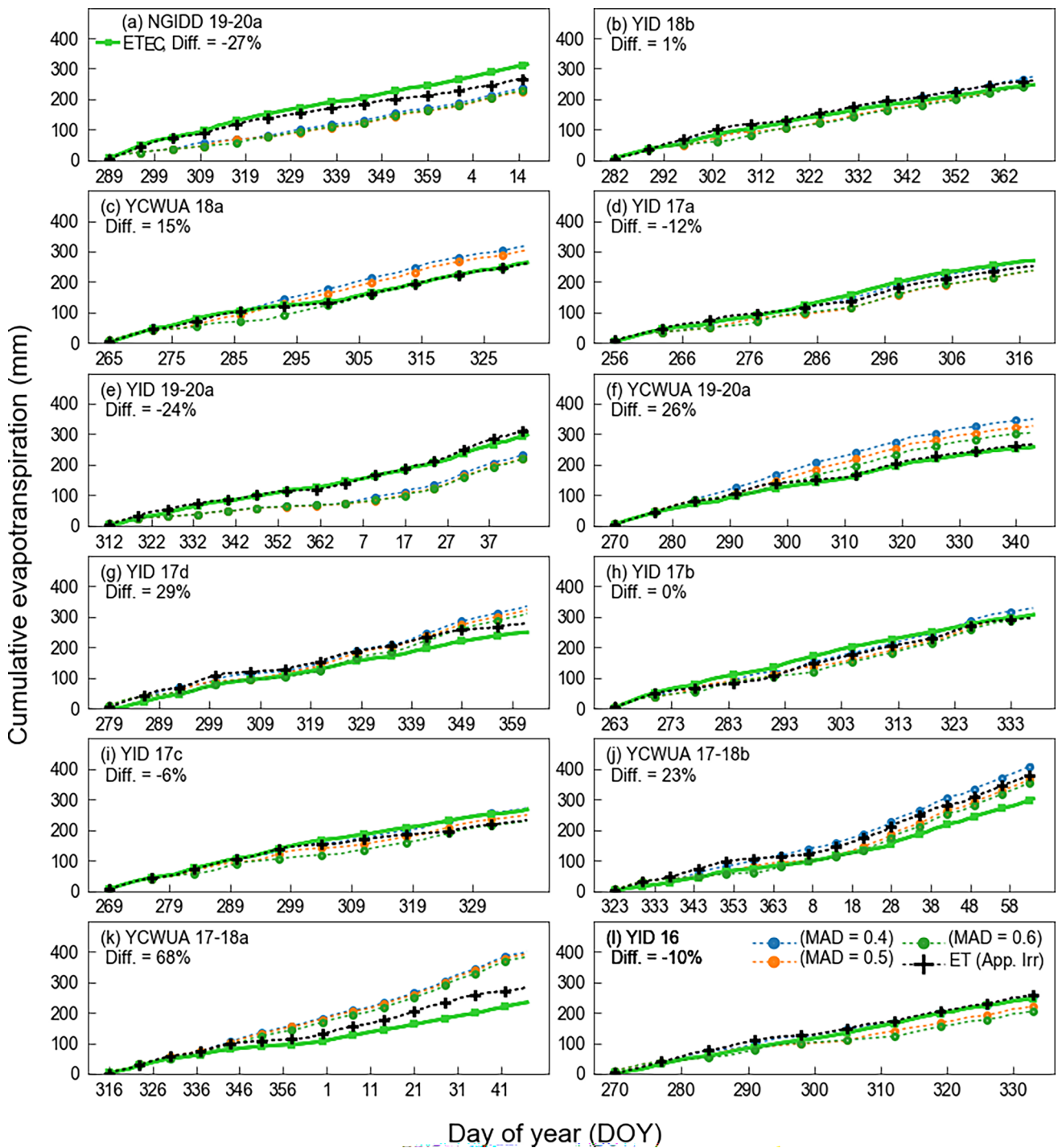
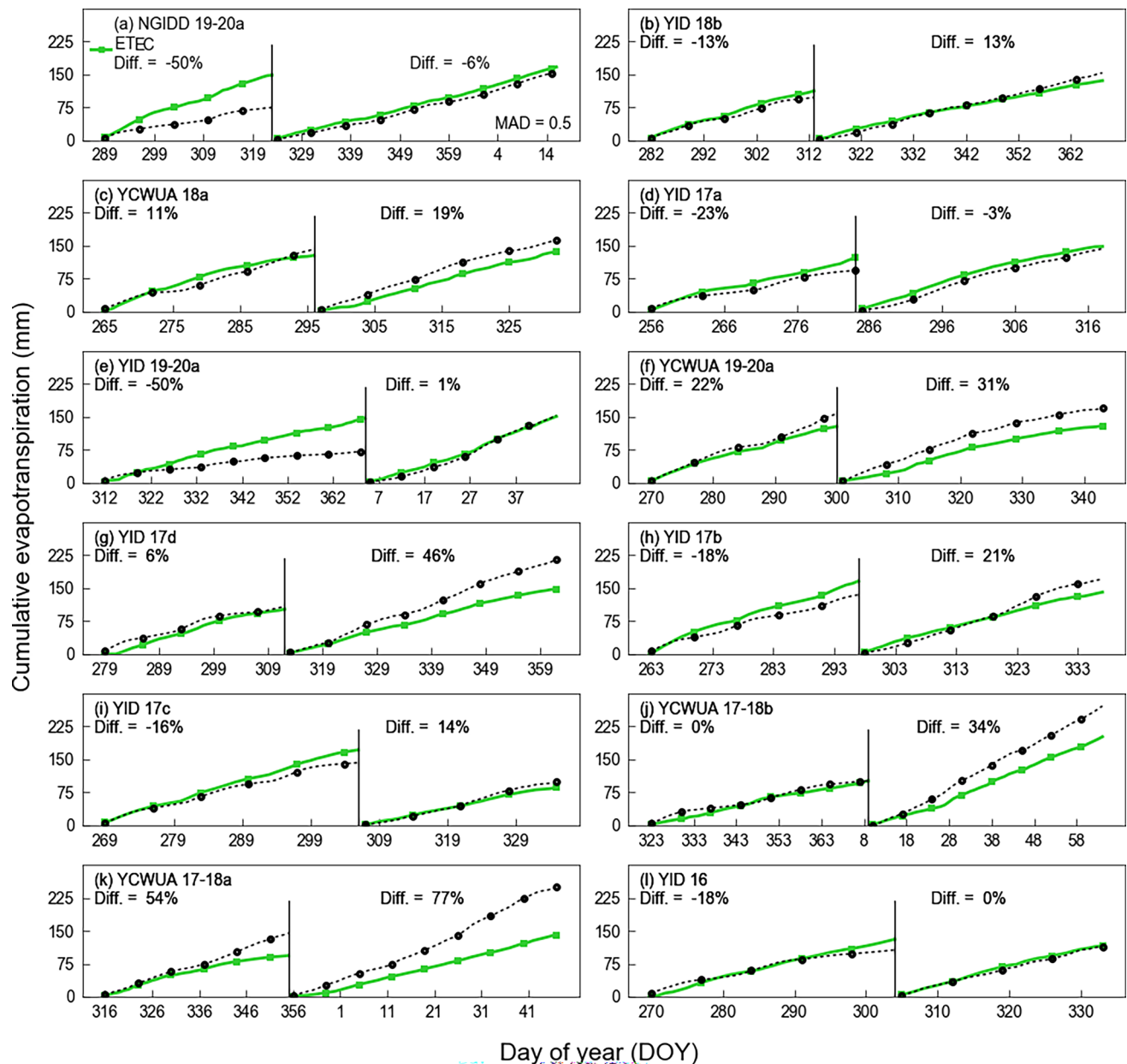


Fig. 7 Cumulative seasonal ET from EC and BAITSS of lettuce for various sites Yuma, AZ. Percent difference (Diff.) is based on MAD of 0.5

### Seasonal cumulative irrigation

Applied and simulated cumulative seasonal irrigation agreed well at some sites (NGIDD 19–20a, Fig. 9a; YID 18b, Fig. 9b; YCWUA 19–20a, Fig. 9f; YCWUA 17–18b, Fig. 9j; YID 16, Fig. 9l) and less well at others (Table 1). As an interesting fact, differences between simulated

and measured seasonal ET at sites NGIDDD 19–20a (Fig. 7a) and YID 19–20a (Fig. 7e) were of similar magnitude (both around –25%) with MAD=0.5 (Fig. 9). In contrast, the corresponding differences in seasonal irrigation differed significantly around –11% for NGIDDD 19–20a (Fig. 9a) and around –47% for YID 19–20a (Fig. 9e) although both showed negative bias. Seasonal



**Fig. 8** Cumulative seasonal ET from EC and BAITSS for lettuce at various sites around Yuma, AZ. The two periods in each plot are the period when soil evaporation was dominant ( $NDVI < 0.3$ ) and afterward. Percent difference (Diff.) is based on MAD of 0.5

precipitation was similar at both sites (32 mm for NGIDD 19–20a and 58 mm for YID 19–20a). The variability of seasonal irrigation among the MADs was also small for both of these sites (NGIDD 19–20a and YID 19–20a) and both had identical soil moisture characteristics with the lowest  $\theta_{fc}$  among all sites,  $0.19 \text{ m}^3 \text{ m}^{-3}$ . Applied irrigation similar to or smaller than simulated (i.e., a positive bias) generally indicates efficient water application (YID 18b, YID 17a, YID 17d, YID 17b, YID 17c, YCWUA 17–18b). The maximum positive bias (42%) in cumulative

irrigation was for YID 17d (Fig. 7g; Table 1). The mean seasonal differences between simulated and applied irrigation among the sites and years were 6%, which was close to seasonal cumulative ET (7%). Like cumulative ET, simulated cumulative irrigation was largest for MAD of 0.4. The largest difference between the seasonal irrigation values among the sites between MAD of 0.4 and 0.6 was 20% for YCWUA 17–18b (Fig. 9j), while the lowest difference was  $-7\%$  for YID 19–20a (Fig. 9e) with average of  $\sim 10\%$ .

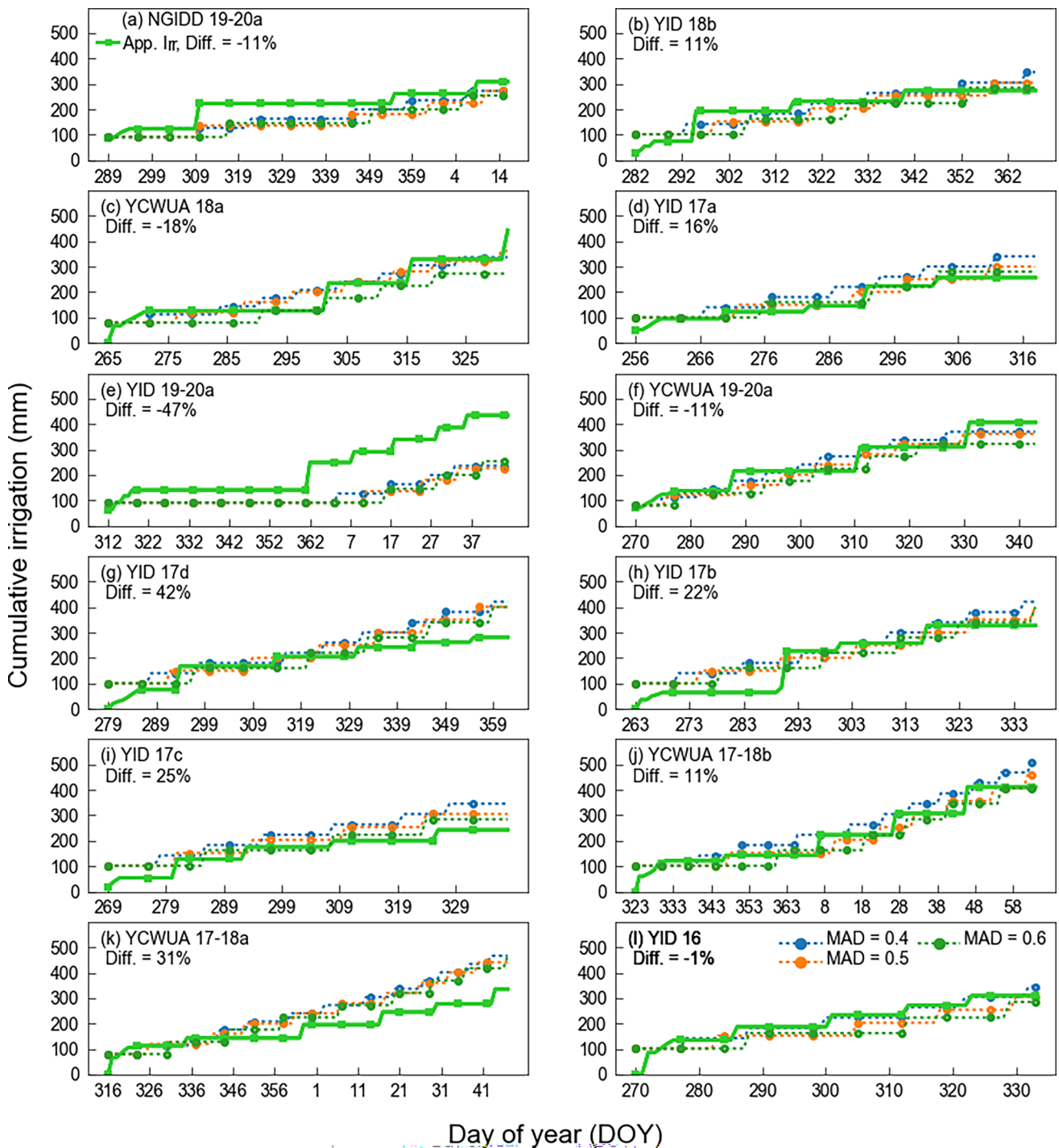


Fig. 9 Cumulative seasonal irrigation of lettuce for various sites around Yuma, AZ. Percent difference (Diff.) is based on MAD of 0.5

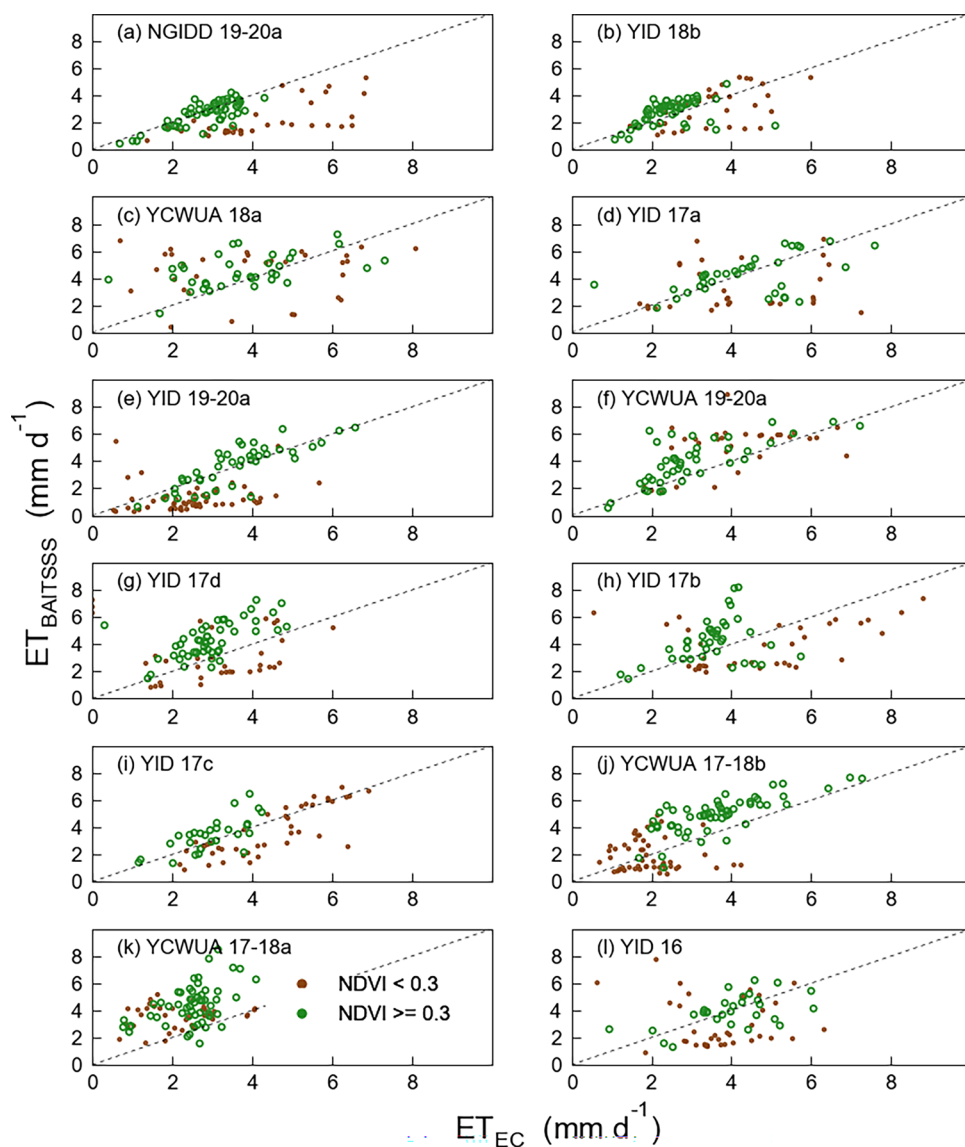
**Accuracy assessment**

Figure 10 presents scatterplots between EC and BAITSSS simulated daily ET where points lie on both sides of the 1:1 line. However, larger differences were observed during the partial cover period in some sites (NDVI < 0.3 indicated by a smaller brown color circle symbol and afterward a larger

green symbol). As observed in Fig. 4, the BAITSSS ET showed a continuous decline after the first irrigation event, i.e., right after planting and during the crop establishment period, whereas EC ET mostly stayed high due to frequent, small irrigation events. These different irrigation patterns were a primary cause of differences in ET during that period. The higher agreement (points close to 1:1) during



**Fig. 10** Scatterplot of daily evapotranspiration (ET) derived from eddy covariance ( $ET_{EC}$ ) versus ET computed with BAITSSS ( $ET_{BAITSSS}$ ). The BAITSSS model was run using a management allowable depletion (MAD) of 0.5. The results are for various lettuce fields around Yuma, AZ



the latter part (mostly during the full cover period) was perhaps because evaporation was less influential, and also because ET may have attained near reference conditions. BAITSSS didn't compute reference ET ( $ET_0$ ) based on the FAO-56 method (Allen et al. 1998), instead, the minimum canopy resistance ( $r_{sc}$ ) of  $40 \text{ s m}^{-1}$  limits the maximum value of transpiration.

## Discussion

A major advantage of remote sensing-based energy balance ET models is that there is no direct need for ground-based moisture conditions and does not require irrigation. Limited studies have been done utilizing remote sensing-based energy balance to simulate irrigation and compare it to applied irrigation applications. For instance, Droogers

et al. (2010) estimated applied irrigation application with reasonable accuracy (95%) by remotely sensed evapotranspiration observations (Landsat) provided that data are available at an interval of 15 days or shorter and the accuracy of the signal is 90% or higher. METRIC (Allen et al. 2007), a widely used remote sensing-based energy balance model, incorporates precipitation-based soil water balance to evaluate the upper soil moisture condition of hot pixel (Allen et al. 2013; Tasumi 2019). However, neither applied nor simulated irrigation is accounted for in these water balance as this information is not readily available. Chen et al. (2018) and Taghvaeian et al. (2020) documented hydrological and crop growth models that had the capability of irrigation simulation (manual and auto-irrigation) which were Soil and Water Assessment Tool (SWAT; watershed scale model) (Arnold et al. 1998), the Agricultural Policy/Environmental eXtender (APEX; small watershed/field-scale

model) (Williams 1995), the Decision Support System for Agrotechnology Transfer (DSSAT; crop model) (Jones et al. 2003), Root Zone Water Quality Model (RZWQM; field-scale model) (Ahuja et al. 2000), MODFLOW (groundwater model) (McDonald and Harbaugh 1988), and AquaCrop (Raes et al. 2009).

SWAT and APEX models had auto-irrigation which is either based on a water stress identifier of either plant water demand or soil water content. Chen et al. (2018) further developed auto-irrigation algorithms for the SWAT model based on uniform, single-season MAD as well as growth stage-specific MAD with options for seasonal growth stage partitioning based on the scheduled date and accumulated heat units. They indicated improved model performance for simulations of irrigation amount and frequency and actual evapotranspiration compared with observed data from an irrigated lysimeter field. Some other studies also indicated promising results for detection irrigation applications using microwave in some regions, however, because of spatial mismatch between model and satellite data, effects of topography, vegetation, frozen soils, and Radio-frequency Interference (RFI) led to substantial uncertainties in others (Kumar et al. 2015; Massari et al. 2021).

The irrigation depth in BAITSSS depends on the rooting depth, MAD, and soil moisture characteristics, along with vegetation cover and weather demand. A constant rooting depth of 500 mm (relatively shorter than for grain crops) and a constant MAD were adopted for the entire simulation period. Excluding the first irrigation event (at the start of the simulation), the rest of the irrigation depths from BAITSSS was nearly identical (because of identical  $\theta_t$ ) for a given MAD. During crop establishment and early growing periods, rooting depths are shallower than those at maturity, i.e., full canopy cover. In addition, applied irrigation showed that MAD right after planting can be significantly smaller compared to the later period. Implementing dynamic rooting depth (a gradual increase of rooting depth mimicking root growth) as implemented by Song et al. (2013), El Masri et al. (2015), Liu et al. (2020) in land surface models and growth stage-specific MAD may help to increase the frequency of irrigation right after planting, though it may increase computational challenges when using high spatial and temporal resolution data at the landscape scale.

Additionally, the planting dates and rooting depth may also vary based on crop types which may not be available. Because of these limitations, it is preferred to have a generalized solution from BAITSSS that can accommodate multiple crop types with fewer field data requirements. In future, with the help of high-resolution remote sensing products like Planet (Planet 2017) (daily scale, 3–5 m), soil maps like Polaris (30 m) (Chaney et al. 2016), shortwave-infrared-based water index (Yue et al. 2019), ingesting thermal band and microwave as well as combining multiple bands (for

instance near-infrared and blue bands) may assist in improving/identifying these field-scale moisture conditions and irrigation. In this study, we didn't intend to revise simulated irrigation to match the applied irrigation as one of the objectives of this study was to understand the irrigation practices and behavior of farmers and growers as well as MAD-based irrigation.

Earlier study discussed that one of the primary controls on ET was solar radiation and air temperature where ET decreased toward the winter solstice and increased afterward into the new calendar year (Dhungel et al. 2022). During that period, vegetation indices were generally increasing creating phase differences with solar and air temperature. We also found similar behavior in this study (Fig. 4). We did not observe a significant underestimation of ET based on these implemented MADs from moisture-related stress. A higher value of MAD delays the irrigation, thus pushing  $\theta_t$  near to  $\theta_{wp}$  ( $\theta_t = \theta_{wp}$  when MAD = 1). Even though ET estimated from BAITSSS appeared to be similar among the MADs during the full cover, delaying irrigation during germination and crop establishment cover period would not be realistic for healthy crop growth. Evaluating applied irrigation during the crop establishment period where a cooler micro-climate was created by frequent irrigation for the crop's physiological demand, the irrigation from BAITSSS may need to be adjusted if intended to match field conditions.

## Conclusion

We evaluated ET and auto-irrigation from BAITSSS for multiple sites and years for lettuce in an arid environment. The results showed the model was competent for estimating ET utilizing MAD based on auto-irrigation. A constant rooting depth of 500 mm and constant MAD (0.4–0.6) for all crop growth stages were evaluated. The model-simulated auto-irrigation and ET based on predefined irrigation rules (MADs), soil moisture characteristics, weather, and vegetation indices. It produced mixed results where obvious differences in ET spikes were observed in some sites during the partial cover period due to variations in irrigation patterns and frequencies. The differences in ET were generally reduced during the full cover period, which may be due to the reduced effect of soil evaporation. The agreement between applied and simulated irrigation showed the model's capability as well as effective irrigation application in the field. However, fields with significantly larger applied irrigation compared to simulated may guide farmers/producers for effective irrigation application. The behavior of applied irrigation, i.e., timing and depth widely varied for the same season (< 20 to > 80 mm) when compared simulated MAD-based irrigation. One of the common behaviors of applied irrigation in all sites was a larger irrigation



frequency of small irrigation events (< 20 mm) during the initial growing period for creating a cooler microclimate. However, due to constant rooting depth, constant MAD for entire growth stages, and a longer partial cover period (average of 38 days when NDVI < 0.3) the model was unable to mimic this behavior in some sites. Overall, there was not an obvious MAD value to be chosen to model the fields absent grower provided irrigation times and depths. However, shallower rooting depth or smaller MAD value, or a combination of both can help to increase the frequency of irrigation right after planting. The average difference between both seasonal ET and irrigation was ~ 10% between MAD of 0.4 and 0.6. The study highlighted the capabilities and limitations of model-simulated auto-irrigation and ET.

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**Author contributions** RD, RGA, AF, and TS designed and performed research, analyzed data, and originated manuscript; TS, MS and CAS contributed data; ES guided project development. All authors reviewed and revised the manuscript.

**Data availability** Weather data can be found (<https://ag.arizona.edu/azmet/>). We thank the United States Department of Agriculture (USDA) Agricultural Research Service (ARS) and the University of Arizona for providing the Eddy Covariance data.

## Declarations

**Conflict of interest** The authors declare no conflicts of interest relevant to this study.

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