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Full Title: Comparison of Methods for Analyzing Salmon Habitat Rehabilitation Designs For Regulated Rivers

Short title: Comparing Tools for River Rehabilitation

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Key words: river restoration, river modeling, river hydraulics, gravel-bed rivers, physical habitat quality, sediment transport regime

1 ABSTRACT

2

3 River restoration practices aiming to sustain wild salmonid populations have received
4 considerable attention in the United States and abroad, as cumulative anthropogenic impacts have
5 caused fish-population declines. An accurate representation of local depth and velocity in
6 designs of spatially complex riffle-pool units is paramount for evaluating such practices, because
7 these two variables constitute key instream habitat requirements and they can be used to predict
8 channel stability. In this study, three models for predicting channel hydraulics- 1D analytical, 1D
9 numerical, and 2D numerical- were compared for two theoretical spawning habitat rehabilitation
10 designs at two discharges to constrain the utility of these models for use in river-restoration
11 design evaluation. Hydraulic predictions from each method were used in the same physical
12 habitat quality and sediment transport regime equations to determine how deviations propagated
13 through those highly nonlinear functions to influence site assessments. The results showed that
14 riffle-pool hydraulics, sediment transport regime, and physical habitat quality were very poorly
15 estimated using the 1D analytical method. The 1D and 2D numerical models did capture
16 characteristic longitudinal profiles in cross-sectionally averaged variables. The deviation of both
17 1D approaches from the spatially distributed 2D model was found to be greatest at the low
18 discharge for an oblique riffle crest with converging cross-stream flow vectors. As decision
19 making for river rehabilitation is dependent on methods used to evaluate designs, this analysis
20 provides managers with an awareness of the limitations used in developing designs and
21 recommendations using the tested methods.

1 INTRODUCTION

2
3 Spawning habitat rehabilitation (SHR) is widely performed for regulated rivers in the
4 western United States and other semi-arid regions globally (Zeh and Donni, 1994; Wheaton et
5 al., 2004a; Gard, 2006; Elkins et al., 2007) to mitigate the decline in anadromous fish
6 populations associated with dam impacts and excessive fishing (Yoshiyama et al., 1998; Graf,
7 2001). A component of SHR involves adding washed gravel and cobble, 8-256 mm in diameter,
8 to a stream (aka gravel augmentation) to increase the quantity and quality of spawning habitat at
9 a placement site (Harper et al., 1998; Wheaton et al., 2004a) as well as to provide coarse
10 sediment to transport downstream where it may form diverse habitats (Trush et al., 2000). In
11 past decades, the design and construction of instream alluvial spawning habitat using augmented
12 gravels largely involved creating flat homogenous spawning beds supported by rock weirs
13 (Kondolf et al., 1996; Walker et al., 2004; Slaney et al., 1997; Newbury et al., 1997; CDFG,
14 1998; CDWR, 2000; Saldi-Caromile et al. 2004). Moreover, the analysis and evaluation of these
15 features also relies on the assumption of steady, uniform flow to estimate hydraulic variables
16 used in subsequent geomorphic and ecological predictions.

17 Based on observed deficiencies of past projects, there is a growing recognition of the
18 importance of design and implementation of complex alluvial features that utilize channel non-
19 uniformity to promote habitat heterogeneity (Pasternack et al., 2004; Wheaton et al. 2004b, c;
20 Elkins et al., 2007; Sawyer et al., 2008). There is also a growing need for accurate and cost-
21 effective means to represent habitat (Maddock, 1999; Moir and Pasternack, 2008).
22 Consequently, this study presents a comparison of three contemporary analytical and numerical
23 methods used for predicting channel hydraulics, spawning habitat quality, and sediment transport
24 regime (as defined by ranges in dimensionless Shields stress; see methods section) over riffles

1 and pools typical of regulated rivers where salmon spawn.

2 The overall goal of this study was to determine what quantitative differences arise from
3 the three different approaches for estimating channel hydraulics over riffle and pool sections and
4 to what extent, if any, the differences impact physical habitat quality and sediment transport
5 predictions. Specific questions were: 1) How do velocity and water surface profiles compare
6 between each approach for two different discharges and two different riffle-pool morphologies-
7 one with orthogonal and one with oblique gravel bars?, 2) What are the effects of 1D and 2D
8 hydraulic models on the prediction of sediment transport regime and spawning habitat quality?,
9 3) Can the hydraulics of even a spatially uniform broad flat riffle be determined within reason
10 using only a 1D analytical approach?, 4) How does a blanket-fill gravel placement design differ
11 in sediment transport regime and physical habitat quality from one that strongly accentuates
12 riffle-pool relief?, and 5) What would be the implications of each approach on decision making
13 for the test designs? The significance of this study lies in the practical aid it provides regulated-
14 river managers in deciding what level of detail they need in their design evaluation framework to
15 reasonably predict the outcome of their habitat rehabilitation projects in terms of ecological
16 success and geomorphic stability.

17 **Riffle-Pool Units**

18 Riffles and pools are respectively defined as topographic highs and lows along a channel
19 thalweg (Leopold et al., 1964; Richards, 1976; Wohl et al., 1993). They are a common
20 geomorphic unit for the aquatic community in gravel-bed rivers with slopes ranging from 0.001-
21 0.02, and are thus very important components in spawning habitat rehabilitation projects
22 involving direct gravel augmentation (NRC, 1992; Newbury et al., 1997; Kondolf, 2000;
23 Wheaton et al., 2004a,b; Saldi-Caromile et al., 2004; Elkins et al., 2007). Naturally sustained

1 riffles and pools are used by anadromous fish species through multiple life-stages (Bjornn and
2 Reiser, 1991) and represent contrasting hydraulic and geomorphic environments as indicated by
3 variations in water surface and energy gradients (Leopold et al., 1964; Keller, 1971;
4 Richards, 1976; Wohl et al., 1993), mean and local bed hydraulic parameters such as velocity and
5 shear stress (Keller, 1971; Carling, 1990; Clifford and Richard, 1992; MacWilliams et al., 2006),
6 grain size (Keller, 1971; Lisle, 1979; Milne, 1982), and channel area, which lead to further
7 variations in flow divergence and convergence (Thompson et al., 1999; MacWilliams et al.,
8 2006). Water-surface and energy gradients over riffles are typically steeper than over pools at
9 low flows and generally converge as stage increases and relative roughness decreases (Leopold
10 et al., 1964; Keller, 1971; Richards, 1976; Wohl et al., 1993). Similarly, mean velocity and shear
11 stress are greater over riffles than pools at low flows, but the reverse may occur at high flows
12 where channel conditions promote that phenomenon (Keller, 1971; Carling, 1991; Clifford et al.,
13 1993; Wilkinson et al., 2004; MacWilliams et al., 2006; Harrison et al., 2007). Variations in
14 depth, velocity, and water surface gradients between riffles and pools at low flows lead to distinct
15 sediment sorting patterns that may even persist at high flows and affect channel hydraulics
16 (Keller, 1971; Lisle, 1979). The convergence and/or “reversal” of riffle-pool water surface
17 elevation and velocity profiles can also cause sediment transport capacity to become greater
18 through pools than riffles (Sear, 1996).

19 The contrasting hydraulic and geomorphic environments provided by riffle and pool units
20 also yield a nested hierarchy in the ecological community. In the context of SHR, spawning
21 habitat, defined as suitable areas for salmon to deposit eggs and for those eggs to incubate, is
22 typically associated with riffle areas, while adult holding habitat is typically located in pools
23 (Hunter, 1991; Wheaton et al., 2004b; Gard 2006; Brown and Pasternack, 2008). The linkage

1 between channel hydraulics and what organisms do is often through what is termed “physical
2 habitat”. Physical habitat quality refers to the degree of suitability of local depth, velocity and
3 river-bed substrate size in a stream to support a particular ecological function. It is a common
4 metric used to evaluate existing channel conditions for instream flow needs (Smith, 1973,
5 Milhous et al., 1989, Bovee et al., 1998, Lacey and Millar, 2004, Brown and Pasternack, 2008)
6 as well as in design evaluation for SHR projects (Pasternack et al., 2004; Gard, 2006; Elkins et
7 al., 2007). Although other variables such as temperature, primary and secondary productivity,
8 and water quality affect the location where fish choose to spend time during a given lifestage
9 (Torgersen et al., 1999; Merz and Setka, 2004), it has been shown that physical habitat quality is
10 often a very strong predictor of some lifestages, especially spawning (Leclerc et al., 1995; Elkins
11 et al., 2007).

12 Predicting channel stability and physical habitat quality are important components of
13 evaluating SHR project designs. The manipulation of channel form that is typical of SHR is
14 related to these key goals through channel hydraulics, which can be represented at different
15 spatial scales. Engineering and design aspects of direct gravel augmentation for SHR often
16 require that target depths and velocities are present and that gravels are stable at low discharges
17 when spawning and embryo incubation are occurring (Merz et al., 2004). In contrast, channel
18 change and bed turn-over are desired at other times of the year, because the gravel bed needs to
19 be kept free of silt and fine sand that can cause subsurface oxygen deficits and embryo death
20 (Merz and Setka, 2004; Merz et al., 2004; Merz et al., 2006). This requirement of bed turnover
21 has stimulated research into design of “flushing flows” that partially mobile the bed (Wilcock et
22 al., 1996a). Gravel stability and bed-material transport are both related to channel conditions
23 through shear stress, which is a direct function of channel hydraulics (Yalin, 1977; Chang, 1998).

1 As physical habitat is also directly linked to channel hydraulics, capturing and distinguishing the
2 hydraulic attributes of riffles and pools is a key need in the successful design and evaluation for
3 SHR projects (Ghanem et al., 1996).

4

5 **METHODS**

6

7 To answer the study questions, a hydraulic analysis was performed using three methods
8 for two hypothetical SHR designs (Fig. 1) actually proposed for an individual pool-riffle-pool
9 sequence in the Lewiston Dam Reach (river mile 111.8-111.2; USFWS 1999) of the Trinity
10 River in Northern California. Since the test designs are hypothetical constructs, the numerical
11 models could not be validated for these specific scenarios. However, the models were validated
12 for the real Lewiston Dam Reach on the Trinity River (Brown and Pasternack, 2008), so the use
13 of these models for this investigation is appropriate. The hydraulic output for each test method
14 was propagated into standard sediment transport regime and spawning habitat quality algorithms
15 to evaluate the sensitivity of these metrics to choice of hydraulic estimation method. The
16 sediment transport regime was characterized by ranges of Shields stress reported in the literature
17 to produce various transport intensities in gravel-bed rivers. Spawning habitat quality was
18 characterized using a global habitat suitability index based on depth and velocity. Both of these
19 indices are described in more detail below. Although a specific river was used to obtain a
20 baseline topography and flow regime, the results should be applicable to channels worldwide
21 with similar non-dimensional geometric and Froude number scaling.

22 The comparison of three hydraulic estimation methods required establishing some values
23 for variables and parameters common to all three methods. To account for the flow dependence

1 of sediment transport regime and physical habitat quality, the three methods were evaluated at
2 two discharges- the prescribed spawning and embryo incubation discharge for the Trinity River
3 ($8.5 \text{ m}^3/\text{s}$) and the maximum flow release that occurred 1999-2004 ($170 \text{ m}^3/\text{s}$). For all
4 calculations of design bed material, the D_{50} (i.e. median size) and D_{90} (i.e. size that 90% of grains
5 are smaller than) were assumed to be 76.2 mm and 152.4 mm respectively, based on consultant-
6 recommended grain sizes for gravel placed into the Trinity River (McBain and Trush, 2003).
7 Similarly, for all parameterizations of channel roughness, Manning's n was used and set at 0.043,
8 which is typical of the channel roughness of well-mixed, double-washed placed gravels
9 (Pasternack et al., 2004). This value has been validated for use on the Trinity River in the
10 rehabilitation reach at Lewiston Dam over the range of flows evaluated (Brown and Pasternack,
11 2008). Effects of channel curvature were not evaluated in this study of straight-channel
12 mechanics, but would be important to consider for designs of meandering channels.

14 **Test Designs**

15 Each test design was situated within the LHR giving the test scenarios several physical
16 constraints. The average channel width is 35 m at low-flow ($8.5 \text{ m}^3/\text{s}$) and the reach is
17 approximately 600 m long with a slope of 0.0022 under current conditions. However, for this
18 study, the test scenarios were only for the upper 350 m of the reach. For more information on
19 the LHR the reader is referred to Brown and Pasternack (2008).

20 The morphological features of design 1 (Fig.1a) involved a "blanket-fill" of gravel with
21 all areas receiving a moderate amount of material. The riffle crests were oriented orthogonal to
22 the channel and included a broad flat riffle (BFR), an "L-shaped" central bar (LCB), and two
23 constricted pools (hereafter referred to as Pool One and Pool Two). The BFR is a commonly

1 used SHR structure when design is based on 1D methods (Kondolf et al., 1996; CDFG, 1998;
2 CDWR, 2000; Saldi-Caromile et al., 2004). The LCB is an analog to the central bar, an alluvial
3 morphology that has been found to provide habitat for multiple life stages of salmonids
4 (Wheaton et al., 2004b). The middle of the LCB has an asymmetrical cross-section where there
5 is a deeper chute on the river left. Constricted pools were used downstream of the BFR and the
6 LCB to potentially invoke hydraulic reversal and channel self-stability by nature of the riffles
7 being somewhat wider than pools (Carling, 1990; MacWilliams et al., 2006).

8 Design 2 (Fig. 1b) involved placing gravel everywhere, but strongly accentuating the
9 relief between the main riffle crest and the pools around it. The primary feature was a transverse
10 oblique riffle (TOR) crest in between two pools. The crest in this case was narrow and attached
11 to an alternate-type bar on river right. This type of bar has been seen in both natural and
12 artificial riffles. It is similar in function to chevron-shaped riffles, but converges flow along one
13 bank preferentially, which may enhance local scour there, promoting habitat heterogeneity.

14 Digital terrain models (DTMs) of both designs were based on the actual channel
15 topography of the Lewiston Dam reach, which was mapped in august 2003 with a resolution of
16 ~ 1.4 points per m^2 (Brown and Pasternack, 2008). For this study, that real surface was re-
17 contoured in AutoCAD Land Desktop 3 to obtain the desired channel features for each design.
18 The design DTMs (Fig. 1) were used to generate the topographic inputs for each test method.

20 **Hydraulic Prediction**

21 Hydraulic variables in regulated rivers serve as key metrics for a plethora of management
22 issues from habitat assessment, sediment management issues, and flood control. In the realm of
23 professional practice and applied science they aid decision making by facilitating the prediction

1 of other key parameters of interest to the management of regulated rivers. Since hydraulic
2 variables such as depth and velocity are now widely used as intermediate links connecting flow
3 processes (i.e. sediment transport and deposition) with ecological conditions in habitat evaluation
4 (Bovee et al., 1998; Maddock, 1999; Gard, 2006; Elkins et al., 2007), as well as geomorphic
5 processes associated with sediment transport and deposition, their accuracy is vital in attempts to
6 manage and restore regulated rivers (Pasternack et al., 2006). Currently in professional practice
7 there are three main approaches for predicting channel hydraulics that can be delineated by the
8 solution procedure and the physical dimensions they are capable of predicting, and they include a
9 1D analytical procedure, a 1D numerical model, and a 2D numerical model.

10

11 *1D Analytical*

12 The 1D analytical method involves predicting open channel processes by coupling some
13 combination of a mass-conservation equation, empirical hydraulic-geometry equations, empirical
14 flow-resistance equation, and an empirical or semi-empirical sediment-transport equation (Dunne
15 and Leopold, 1978; Yen, 1991; Rosgen, 1996; Chang, 1998). Flow resistance equations have
16 been typically derived from non-alluvial channel boundaries under steady, uniform flow and use
17 roughness coefficients that have no true theoretical basis (Yen, 1991). Because they assume
18 steady, uniform flow conditions, they are the easiest to perform and represent the lowest cost
19 approach to quantitative design evaluation (Shields et al., 2003).

20 The Manning-Gauckler equation, hereafter referred to as the Manning Equation, is a
21 frequently prescribed flow resistance equation used in the United States for evaluating channel
22 hydraulics for physical habitat and sediment transport (Rosgen, 1996; Bovee et al., 1998; Saldi-
23 Caromile et al. 2004), so it was used in this study to calculate cross-sectionally averaged velocity

1 (\bar{V}) for each specified water surface elevation (WSE):

$$2 \quad \bar{V} = \left(\frac{1}{n}\right)R^{2/3}S^{1/2} \text{ and } Q = A\bar{V} \quad (1,2)$$

3 where Q is discharge (m³/s), R is hydraulic radius (m), A is cross-sectional area (m²), n is
4 Manning's roughness coefficient, and S is typically the average bed slope. R and the
5 corresponding A were obtained iteratively in AutoCAD Land Desktop 3 until the continuity
6 equation was solved for the discharge of interest. The equation is very sensitive to S, and the
7 local slope can vary considerably in a gravel bed river, depending on the spatial scale being
8 examined. To avoid bias, it was assumed that S was related to geomorphic slope, defined as the
9 elevation difference between an upstream feature and the next downstream feature of the same
10 type (i.e. bar or pool) divided by the distance between them (Fig. 2).

11 As the 1D analytical method is based on the concept of steady-uniform flow, in which
12 temporal and spatial changes are neglected and driving and resisting forces balance each other
13 out (Yen, 1991), relevant hydraulic processes within and between cross-sections may not be
14 accounted for. In particular, backwater effects and other forms of channel-wide and local
15 convective accelerations are not accounted for using one-dimensional methods. These
16 phenomena have been shown to influence both pool-riffle hydraulics, sediment transport regime,
17 and physical habitat predictions (MacWilliams et al., 2006; Elkins et al., 2007).

18

19 *1D Numerical Model*

20 Numerical approaches to hydraulic estimation employ computers to approximate
21 solutions of 1D, 2D, or 3D equations of motion where the solution procedures are dependant
22 either on adjacent nodes or cross sections. 1D models such as HEC-RAS and MIKE11 can

1 consider unsteady conditions and to some extent non-uniform conditions. They cannot account
2 for transitional dynamics where no cross-sections are measured, nor can they account for
3 secondary flow processes that may occur at measured cross sections (Darby and Van de Wiel,
4 2003; Nelson et al., 2003). They use a standard step method to iteratively solve the energy
5 equation from one cross section to the next to calculate water surface profiles. These models
6 solve the energy equation for steady gradually varied flow and are also capable of calculating
7 subcritical, super critical, and mixed flow regime water surface profiles (Brunner, 1998). The
8 energy equation at two cross-sections denoted by subscript numbers 1 and 2 is described as

9

$$\bar{H}_2 + Z_2 + \frac{\alpha_2 \bar{V}_2^2}{2g} = \bar{H}_1 + Z_1 + \frac{\alpha_1 \bar{V}_1^2}{2g} + h_e \quad (3)$$

10 where \bar{H} is cross-sectionally averaged water depth (m), Z is the invert elevation (m), α is the
11 velocity weighting coefficient, g is gravitational acceleration (m/s^2), and h_e is energy head loss
12 (m). Eq. (1) is used implicitly in the solution of (Eq. 3).

13 HEC-RAS was used in this study and the boundary conditions required to run it were the
14 discharge entering the upstream boundary, the associated downstream water surface elevation,
15 and channel topography at designated cross-sections. Cross sections were sampled from the
16 DTM of each test design in 7.6-m increments. Some additional cross sections were augmented
17 to capture transitional dynamics and to match cross section locations used for the 1D analytical
18 method.

19 1D hydraulic models usually provide fairly accurate predictions of water surface
20 elevation for flood stages where relative roughness is low and/or there are gradual changes in
21 bed slope and channel width (Ghanem et al., 1996; Brunner, 1998; Lacey and Millar, 2004).
22 They have the advantage over 1D analytical approaches in that they can account for cross-

1 sectionally averaged convergence and divergence effects at cross-sections. They cannot account
2 for topographic variability between cross-sections nor even localized convective accelerations
3 within only a small part of a cross-section, such as in the vicinity of a boulder cluster. Previous
4 studies highlight that these models may not be appropriate for predicting channel hydraulics at
5 the spatial scales necessary to represent physical habitat and sediment transport (Ghanem et al.,
6 1996; MacWilliams et al., 2006; Lacey and Millar, 2004). Despite this, they are readily
7 available, require relatively sparse information, and are frequently used by professional
8 consultants.

9

10 *2D Numerical Model*

11 2D (depth-averaged) models such as FESWMS, RIVER2D, HIVEL2D, RMA2, MIKE21,
12 SRH-2D, TUFLOW, and TELEMAC further add the ability to consider full lateral and
13 longitudinal variability down to the sub-meter scale, including effects of alternate bars,
14 transverse bars, islands, and boulder complexes, but require highly detailed topographic maps of
15 channels and floodplains (French and Clifford, 2000). 2D models are also more realistically
16 linked to flow, sediment transport, and biological variables measured in the field at the same
17 spatial scale (Ghanem et al., 1996; Lacey and Millar, 2004; Pasternack et al., 2006). 2D models
18 have been used to study a variety of hydrogeomorphic processes (Bates et al. 1992; Leclerc et al.
19 1995; Miller and Cluer 1998; Cao et al., 2003). Recently, they have been evaluated for use in
20 regulated river rehabilitation emphasizing spawning habitat rehabilitation by gravel placement
21 (Pasternack et al., 2004, Wheaton et al., 2004b; Pasternack et al., 2006; Elkins et al., 2007;
22 Brown and Pasternack, 2008).

23 In this study, the 2D model known as Finite Element Surface Water Modeling System

1 3.1.5 (FESWMS) was used to simulate hydraulics (Froehlich, 1989). FESWMS solves the
 2 vertically integrated conservation of momentum and mass equations using a finite element
 3 method to acquire depth-averaged 2D velocity vectors and water depths at each node in a finite
 4 element mesh. The model is capable of simulating both steady and unsteady 2-D flow as well as
 5 subcritical and supercritical flows. The basic governing equations for vertically integrated
 6 momentum in the x- and y- directions under the hydrostatic assumption are given by

$$\begin{aligned}
 & \frac{\partial}{\partial t}(HU) + \frac{\partial}{\partial x}(\beta_{uu}HUU) + \frac{\partial}{\partial y}(\beta_{uv}HUV) + gH\frac{\partial z_b}{\partial x} + \frac{1}{2}g\frac{\partial H^2}{\partial x} \\
 & + \frac{1}{\rho}[\tau_x^b - \frac{\partial}{\partial x}(H\tau_{xx}) - \frac{\partial}{\partial y}(H\tau_{xy})] = 0
 \end{aligned}
 \tag{4a}$$

7 and

$$\begin{aligned}
 & \frac{\partial}{\partial t}(HV) + \frac{\partial}{\partial x}(\beta_{vu}HUV) + \frac{\partial}{\partial y}(\beta_{vv}HVV) + gH\frac{\partial z_b}{\partial y} + \frac{1}{2}g\frac{\partial H^2}{\partial y} \\
 & + \frac{1}{\rho}[\tau_y^b - \frac{\partial}{\partial x}(H\tau_{yx}) - \frac{\partial}{\partial y}(H\tau_{yy})] = 0,
 \end{aligned}
 \tag{4b}$$

8 where H is local water depth (m), U and V are local depth-averaged velocity components (m/s)
 9 in the horizontal x- and y- directions, respectively, z_b is the bed elevation (m), β_{uu} , β_{uv} , β_{vu} , and
 10 β_{vv} are the momentum correction coefficients that account for the variation of velocity in the
 11 vertical direction, τ_x^b and τ_y^b are the bottom shear stress components (Pa) acting in the x- and y-
 12 directions, respectively, and τ_{xx} , τ_{xy} , τ_{yx} , and τ_{yy} are the shear stress components (Pa) caused by
 13 fluid turbulence. Conservation of mass in two-dimensions is given by

$$\frac{\partial H}{\partial t} + \frac{\partial}{\partial x}(HU) + \frac{\partial}{\partial y}(HV) = 0.
 \tag{5}$$

14 FESWMS was implemented using Surface Water Modeling System v. 8.1 graphical user
 15 interface (EMS-I, South Jordan, UT). The boundary conditions required to run FESWMS were

1 the input hydrograph, the exit water surface elevation, and high-resolution channel topography.
2 In addition, model parameters are needed to describe channel roughness and provide turbulence
3 closure. Values for all boundary conditions and parameters were selected to be physically
4 realistic and were not numerically calibrated. The turbulence parameter eddy viscosity (E) was a
5 variable in the system of model equations, and it was computed as $E = c_0 + 0.6 \cdot H \cdot u^*$, where u^* is
6 shear velocity (m/s) and c_0 is a minimal constant added for numerical stability. This equation
7 was implemented in FESWMS to allow eddy viscosity to vary throughout the channel, which
8 yields more accurate transverse velocity gradients. However, a comparison of 2D and 3D
9 models for a shallow gravel-bed river demonstrated that even with this spatial variation, it is not
10 enough to yield as rapid lateral variations in velocity as occurs in natural channels, presenting a
11 fundamental limitation of 2D models like FESWMS (MacWilliams et al., 2006). Although eddy
12 viscosity is an oversimplifying turbulence closure parameter, observations of depth and velocity
13 may be used to compute values for it as a check on model performance (Fischer et al., 1979).
14 This model has previously been heavily validated for use in shallow gravel-bed rivers
15 (Pasternack et al., 2004; Wheaton et al., 2004b; Pasternack et al., 2006; Elkins et al., 2007;
16 Brown and Pasternack, 2008; Moir and Pasternack, 2008).

17 DEM $\{x,y,z\}$ contour and grid points for each test design were imported from AutoCAD
18 into the 2D model where they were used to interpolate the elevations of the nodes in a finite
19 element mesh consisting of triangular and quadrangular elements. Inter-nodal spacing ranged
20 from 0.2-0.6 m. To reduce model instability associated with mesh-element wetting and drying at
21 a threshold of 9-cm depth, meshes were iteratively trimmed to exclude dry areas, yielding
22 slightly different final meshes for each discharge simulation.

23 3D models such as UNTRIM and SSIIM go further by including vertical fluid fluxes, but

1 do not require much more data input than 2D models to get going (MacWilliams et al., 2006).
2 However, model validation of 3D processes is difficult and time-consuming. Also, few
3 ecological and geomorphic processes have been quantitatively linked to 3D flow dynamics yet,
4 limiting the interpretation of 3D model output for practical applications, such as SHR design.
5 Consequently, 3D models are not considered in this study.

6 There are obvious discrepancies between analytical calculations versus 1D and 2D
7 hydraulic models (Cao and Carling, 2002), however, the effect of these differences on sediment
8 transport regime and physical habitat predictions for riffle and pool units have not been explored.
9 Some previous studies have shown that 1D numerical models are insufficient to describe
10 hydraulics and sediment transport through pool-riffle sequences (Keller et al., 1993; Rathburn et
11 al., 2003). Moreover, studies have shown that 2D models provide a better description of channel
12 hydraulics and physical habitat than 1D models (Ghanem et al., 1996; Lacey and Millar, 2004;
13 MacWilliams et al., 2006). Such efforts provide a foundation for a thorough investigation of the
14 trade-offs when using models for river-rehabilitation design evaluation.

16 **Sediment Transport Regime Predictions**

17 In both analytical and numerical approaches Shields stress was used as the representative
18 variable that indicates channel stability via a sediment transport regime, defined below. For
19 purposes of comparison Shields stress was calculated as a section average for 1D analytical and
20 numerical approaches and on a node basis for the 2D numerical approach. Shields stress is
21 defined as

$$22 \quad \tau^* = \frac{\tau_b}{(\gamma_s - \gamma_f) D_{50}} \quad (6)$$

23 where, τ^* is Shields stress, τ_b is the predicted bed shear stress (Pa), γ_s is the specific weight of

1 sediment (N/m^3), γ_f is specific weight of water (N/m^3), and D_{50} is the median grain size of the
 2 bed surface (m). The sediment transport regime was characterized by the range of values that τ^*
 3 falls into, as defined by Lisle et al. (2000): Values of $0.00 < \tau^* < 0.01$ mean no transport is
 4 occurring, $0.01 < \tau^* < 0.03$ indicates intermittent, localized transport in response to infrequent
 5 turbulent bursts and/or bed vibrations, $0.03 < \tau^* < 0.06$ corresponds with Wilcock's (1996b)
 6 domain of "partial transport" in which grains move in proportion to their relative exposure on the
 7 bed surface, and $\tau^* > 0.1$ represents full mobility of a "carpet" of sediment $1-2 \cdot D_{90}$ thick. Higher
 8 thresholds for channel-altering conditions may exist, but have not been delineated in the
 9 literature. These threshold delineations may vary depending on how well compacted the bed is
 10 and other factors, but they provide a reasonable basis for characterizing sediment transport
 11 conditions. At present there is disagreement as to the breadth of applicability of these
 12 dimensionless ranges across all types of gravel- and sand-bed rivers. However, a review of
 13 sediment transport studies in gravel-bed rivers with the characteristics of the LHR found that
 14 they are applicable for the purposes of this study.

15 To obtain τ^* , predictions of τ_b were made using the hydraulic predictions from each
 16 method in the same Einstein's log-velocity equation for turbulent flows over rough beds

17

$$18 \quad \tau_b = \rho_w \left(\bar{u} / \left(5.75 \log \left(\frac{1.22H}{2D_{90}} \right) \right) \right)^2 \quad (7)$$

19

20 where \bar{u} is depth-average velocity (m/s), D_{90} = grain size in which 90% are finer than (m). When

21 used with the 1D analytical method, \bar{V} and R are substituted for \bar{u} and H , respectively. With

22 the 1D numerical method, \bar{V} and \bar{H} are substituted for \bar{u} and H , respectively. With the 2D

1 numerical method, \bar{u} is taken to be the magnitude of the velocity vector at a point.

2

3 **Physical Habitat Predictions**

4 A global habitat suitability index (GHSI) for Chinook salmon spawning habitat quality
 5 was calculated using depth (DHSI) and velocity (VHSI) habitat suitability curves that were
 6 previously derived for the Trinity River based on field observations of flow conditions at
 7 spawning sites (USFWS, 1997). These curves are typical across rivers in California. The best-
 8 fit numerical equations using H or \bar{H} to match reported DHSI were

9 For $H < 0$ and $H > 1.374$ m: $DHSI = 0$ (8a)

10 For $0 < H < 1.374$ m:

11
$$DHSI = 0.0351 - 0.7248 \cdot (H/0.3048) + 4.9307 \cdot (H/0.3048)^2 -$$

 12
$$5.2964 \cdot (H/0.3048)^3 + 2.4818 \cdot (H/0.3048)^4 - 0.6007 \cdot (H/0.3048)^5 + 0.007235 \cdot (H/0.3048)^6 -$$

 13
$$0.002514 \cdot (H/0.3048)^7 - 0.000274 \cdot (H/0.3048)^8 + 2.1277 \cdot (H/0.3048)^9$$
 (8b)

14 Similarly, using either U or \bar{V} those for VHSI were

15 For $U < 0.01$ and $U > 1.68$ m/s: $VHSI = 0$ (9a)

16 For $0.01 < U < 0.40$ m/s: $VHSI = -0.0131 + 0.5523 \cdot (U/0.3048) -$
 17
$$0.8246 \cdot (U/0.3048)^2 + 1.8208 \cdot (U/0.3048)^3 - 0.8013 \cdot (U/0.3048)^4$$
 (9b)

18 For $0.40 < U < 1.68$ m/s:

19
$$VHSI = -5.1205 + 13.1440 \cdot (U/0.3048) - 9.5769 \cdot (U/0.3048)^2 +$$

 20
$$2.2576 \cdot (U/0.3048)^3 + 0.4525 \cdot (U/0.3048)^4 - 0.3746 \cdot (U/0.3048)^5 + 0.08619 \cdot (U/0.3048)^6 -$$

 21
$$0.009046 \cdot (U/0.3048)^7 + 0.000369 \cdot (U/0.3048)^8$$
 (9c)

22 Similar analyses were done for steelhead and coho as well as for all three species in fry, juvenile,

23 and adult life stages, but only the Chinook spawning analysis is presented here for brevity to

1 exemplify the results and how differences in computing depth and velocity propagate through a
2 typical habitat suitability equation. An analysis of all species' lifestage habitats in the Lewiston
3 Dam reach of the Trinity river is available in Brown and Pasternack (2008). We assumed the test
4 scenarios in this study all had similar grain size distributions, making the use of a substrate
5 suitability index null in this particular study. DHSI and VHSI were combined using the standard
6

$$\text{GHSI} = \text{DHSI}^{0.5} \times \text{VHSI}^{0.5} \quad (10)$$

7 GHSI was classed as very poor (0-0.1), low (0.1-0.4), medium (0.4-0.7), and high (0.7-1.0)
8 quality habitat (Leclerc et al., 1995). One-dimensional analytical predictions were made using
9 \bar{V} and R , while 1D numerical predictions were made with \bar{V} and \bar{H} . The 2D model calculated
10 GHSI on a nodal basis to estimate physical habitat for the design. GHSI values are only reported
11 at the low flow, because at the high flow the depths and velocities were too high to yield much
12 habitat.

14 **Comparison Between Methods**

15 Although, the three approaches use somewhat different variables in their equations it was
16 possible to calculate comparable quantities to understand similarities and differences. Water
17 surface elevation is a common variable produced by all models, and thus is directly comparable.
18 Because the 2D model provides U-values and H-values at all computational nodes across the
19 channel, weighted nodal averages of each comparable to \bar{V} and \bar{H} were calculated by
20 accounting for the width along which each node's values were representative. To compare 1D
21 section-averaged values of τ^* among methods, two different averaging methods were used to
22 obtain section-averages from the 2D model output. First, equations 6-7 were applied at each
23 node, and then weighted nodal averages of τ^* across the channel were computed using the same

1 approach described above for U, H, and GHSI. Second, weighted nodal averages of U and H
2 were put into equations 6-7. This “smoothed” method ought to be less representative, but was
3 thought to be useful for evaluating the effects of different averaging methods on assessing
4 erosion potential.

5 Given a set of test metrics common to all three methods evaluated in this study, the key
6 analysis is the determination of the degree of deviation in the metrics that constitutes a
7 “significant” difference. Rather than using a statistical characterization of significance that
8 assumes values are probabilistically distributed, the appropriate approach for evaluating a
9 deterministic system is to identify the processes that ought to be operating and then evaluate the
10 different models relative to their ability to capture those processes. For the basic hydraulic
11 variables (WSE, \bar{H} , and \bar{V}), the key process evaluated was the degree to which each model
12 captured the longitudinal variations associated with backwater conditions upstream of riffle
13 crests and flow acceleration downstream of them. For the geomorphic and ecologic variables (τ^*
14 and GHSI), evaluation focused on whether deviations in values were large enough to cause the
15 method to predict the wrong sediment transport regime or habitat quality type, since sharp
16 thresholds delineated classes for these variables. Given that the differences in the predictions
17 between the methods were in fact large enough to cause such misidentification of classes, this
18 analytical approach provides an objective and useful basis for comparison. How badly a model
19 must perform before it is rejected is ultimately a philosophical question that each individual and
20 the river-management community at large must wrestle with.

21

22 **RESULTS**

23

1 **Hydraulic Prediction**

2

3 *Test Design 1*

4 Direct comparison of the three test methods for design 1 at the low flow, 8.5 m³/s, found
5 that the two numerical methods yielded very similar cross-sectionally averaged velocity and
6 water surface elevation values down the channel (Table 1), with the 2D velocity values showing
7 slightly greater longitudinal variation (Fig. 3a,b). In contrast, the 1D analytical method yielded a
8 highly uniform longitudinal velocity profile that was insensitive to the bed variation. Also, it
9 predicted a water surface profile that paralleled the bed profile. Consequently, it overpredicted
10 pool velocity and underpredicted riffle velocity, which was the expected outcome of neglecting
11 backwater effects. A useful attribute of the 2D numerical method was the provision of a range of
12 velocities across each section, not just an average (Fig. 3a). This at-a-section lateral variability
13 varied downstream in response to local flow convergence and divergence, as indicated by flow
14 vectors (Fig. 4b).

15 The hydraulic-variable profiles at 170 m³/s for numerical models in design 1 exhibited
16 similar relative predictions as found at low discharge (Table 2; Fig. 3c,d). In this case, the 1D
17 numerical model predicted slightly higher values than the 2D model. Meanwhile, the 1D
18 analytical method yielded greater variations than at low flow, with generally lower velocities and
19 higher water surface elevations than the numerical models. The lateral variation in velocity
20 predicted by the 2D model was much stronger at high flow than low flow, with peak velocities
21 focused in the channel center (Fig. 5b). Also, the longitudinal location of peak velocities shifted
22 at high flow from riffle exits to riffle entrances.

23

1 *Test Design 2*

2 Direct comparison of the three test methods for design 2 at 8.5 m³/s (Table 3) revealed
3 dramatic differences between analytical and numerical predictions. The high-relief riffle crest at
4 30 m downstream produced a strong backwater effect with higher water surface elevations and
5 lower velocities in both numerical models (Fig. 6a,b). In contrast, the analytical method
6 predicted that the water surface would step up with the bed surface, which was wrong. In terms
7 of velocity, the methods yielded similar values over the riffle crest, but the 1D analytical method
8 significantly overpredicted pool velocity (Table 3). The 2D model showed greater hydraulic
9 complexity and lateral variation in velocity in design 2 compared with design 1, which was
10 caused by the transverse oblique riffle (TOR). This is exemplified by the presence of two
11 distinct peak-velocity zones and two whirlpools (Figs. 7a,b). One peak-velocity zone occurred
12 where flow converged vertically over the oblique section of the riffle crest. The other occurred
13 in the laterally convergent, horseshoe-shaped (in plan view) riffle exit on river right. Both
14 whirlpools occurred on river left behind submerged lateral bars with very steep backslopes (Fig.
15 7b).

16 For the high discharge test, the 1D analytical method was even worse off than at low
17 discharge (Table 4; Figs. 6c,d). It predicted the peak velocity over the pool and minimum
18 velocity over the riffle crest, whereas the numerical methods yielded the opposite pattern,
19 correctly accounting for the pool backwater effect and convective acceleration leading into the
20 riffle crest. A stronger difference between 1D and 2D numerical velocity predictions was
21 evident in design 2 compared with design 1, driven by the stronger lateral channel non-
22 uniformity in design 2. Both methods did show that the location of peak velocity moved from
23 the riffle exit to the riffle entrance from low flow to high flow. Also, the 2D model flow vectors

1 did not show any whirlpools (Fig. 8a,b).

2

3 **Sediment Transport Regime Prediction**

4 In light of the strong differences between the 1D analytical method and the numerical
5 models, one would anticipate that propagation of the error into highly nonlinear functions, such
6 as those for sediment transport regime assessment, would accentuate the differences further.
7 Lumping all cross sections for all four tests, the mean and median differences of Shields stress
8 between the 1D analytical method and the 2D model were 175 and 41 %, respectively. Only one
9 out of 20 comparisons had less than a 10 % difference in Shields stress. Comparing the 1D and
10 2D models similarly, the mean and median differences of Shields stress were only 14 and 5 %, respectively.
11 In light of the poor predictive capability of the 1D analytical method, the use of
12 broadly defined classes of sediment transport regime did help diminish the impact, but 11 out
13 of 20 locations were misclassified by that method. In comparison, only one out of 20 locations
14 were classified differently between the 1D and 2D models, and that was the TOR exit at low
15 flow, where the 1D model could not handle strong local flow convergence properly.

16 Differences in predictions of Shields stress between methods varied sharply between the
17 two test designs. The mean difference in Shields stress between the 1D analytical method and
18 2D model was 4.6 times higher for design 2 than for design 1. Similarly, the mean difference in
19 Shields stress between the 1D and 2D models was 3.1 times higher. Overall, design 2 was
20 predicted to be much more at risk of riffle destruction at low and high flows than design 1.

21

22 *Test Design 1*

23 Among the six cross-sections analyzed for design 1 at 8.5 m³/s, most Shields stress

1 predictions were in the range of 0-0.01, corresponding with no transport (Table 1, Fig. 9a). For
2 most channel units, the 1D analytical predictions had a high range, but still remained in the
3 intermittent transport regime. The 1D analytical model underpredicted Shields stress on the two
4 riffle crests, likely yielding an incorrect classification of no transport, instead of intermittent
5 transport, as predicted by the more sophisticated numerical models. The spatial pattern predicted
6 by the 2D model showed longitudinal variation between riffle crests and pools, but little lateral
7 variation (Fig. 4c) given the simple channel geometry.

8 At $170 \text{ m}^3/\text{s}$, maximum Shields stresses corresponding to a state of partial transport were
9 predicted by all methods to occur over the broad flat riffle (BFR) (Table 2, Fig. 9b). In Pool 1,
10 the numerical models predicted continued partial transport due to the steep water surface slope
11 and associated acceleration, whereas the 1D analytical method predicted intermittent transport.
12 Over the middle of the LCB, the numerical models predicted intermittent transport due to the
13 backwater effect behind the next crest, whereas the 1D analytical model predicted partial
14 transport, neglecting any backwater effect. On the LCB riffle crest, the numerical models
15 predicted partial transport driven by vertical convergence, whereas the 1D analytical method
16 only predicted intermittent transport. Very little difference in quantitative or qualitative
17 prediction of cross-sectionally averaged shear stress was evident between the 1D and 2D
18 numerical models. The only place in design 1 where there was significant cross-channel
19 variability in the 2D model that could not be captured by the 1D model was in Pool 1. In that
20 case, the center of the channel was predicted to be in partial transport and the lateral margins in
21 intermittent transport (Fig. 5c).

22

23 *Test Design 2*

1 For test design 2, the 1D analytical method performed very poorly at $8.5 \text{ m}^3/\text{s}$, predicting
2 the wrong sediment transport regime at three of the four cross-sections, with the fourth one
3 barely correct (Table 3; Fig. 10a). The 1D and 2D numerical models yielded similar outcomes
4 everywhere except the TOR exit, where the 1D model predicted intermittent transport and the
5 latter predicted partial transport. The smoothed averaging using the 2D model, performed poorly
6 at this location relative to the nodal averaging method for estimating Shields stress. Looking at
7 the spatial pattern of Shields stress predicted by the 2D model (Fig. 7c), the 2D model predicts a
8 significant zone of full bed mobility on the riffle crest and along the left bank. Partial transport
9 is predicted in the horseshoe (in plan view) convergence zone along the right bank.

10 At $170 \text{ m}^3/\text{s}$, the numerical models continued to predict the maximum shields stress
11 corresponding to partial transport over the TOR crest, while the 1D analytical method predicted
12 the maximum over Pool One (Table 4; Fig. 10b). The 1D analytical method predicted the same
13 sediment transport regime as the numerical models in two of four cross-sections, but the
14 magnitude of Shields stress in the TOR exit was close to the lower threshold for partial transport.
15 The difference in cross-sectionally averaged conditions between 1D and 2D models was
16 diminished at the high discharge. The spatial pattern of Shields stress predicted by the 2D model
17 showed the same lateral variation as seen in design 1, but with full mobility in the channel center
18 and partial transport along the banks (Fig. 8c).

20 **Habitat Quality Prediction**

21 The governing habitat suitability equations (equations 8 and 9) are best-fit functions and
22 highly nonlinear, yielding dramatic differences between the hydraulic models in the assessment
23 of habitat quality for Chinook spawning. Lumping all cross-sections for both test designs, the

1 mean and median differences in predicted GHSI between the 1D analytical method and the 2D
2 model were 118 and 73 %, respectively. Only two of 10 comparisons had less than a 20 %
3 difference in GHSI. Comparing the 1D and 2D models similarly, the mean and median
4 differences of GHSI were 38 and 27 %, respectively. Only one of 10 comparisons had less than
5 a 5 % difference in GHSI. Even using broad quality classes did not help diminish the differences
6 between test methods, with seven out of ten locations misclassified by the 1D analytical method.
7 In comparison, three out of ten locations were classified differently between the 1D and 2D
8 models. Most importantly, the 1D analytical method and 1D model both consistently predicted
9 significantly higher quality of physical habitat than the 2D model.

10 Differences in GHSI predictions between methods varied about the same as those for
11 Shields stress between the two test designs. The mean difference between the 1D analytical
12 method and 2D model was 3.9 times higher for design 2 than for design 1. Similarly, that
13 between the 1D and 2D models was 2.8 times higher. Median differences showed a similar
14 effect. Overall, design 1 yielded significantly higher quality spawning habitat than design 2.

16 *Test Design 1*

17 For test design 1 at low discharge, the numerical models exhibited the expected
18 longitudinal variation of cross-sectionally averaged habitat quality (i.e. riffles are higher quality
19 and shallow pools are lower quality), while the 1D analytical and 2D smoothed averages did not
20 (Table 1; Fig. 11a). The 1D analytical method significantly overpredicted habitat quality,
21 because it predicted uniform, high velocities for this design. The 1D model also consistently
22 predicted higher quality habitat than the 2D model, with the largest difference occurring for the
23 BFR crest. The 2D model predicted the highest quality spawning habitat on the middle of the

1 BFR, the river right of Pool One and the LCB middle, the LCB crest, and the exit slope of the
2 LCB crest (Fig. 4d). Low quality habitat was predicted in deepest section of Pool 1, the river left
3 of the LCB chute, and the deepest section of Pool 2.

4

5 *Test Design 2*

6 A similar pattern of longitudinal variation in GHSI for design 2 at low flow as a function
7 of morphological unit type was observed for the test methods as reported for design 1, but with
8 significantly different GHSI values across methods (Table 3; Fig. 11b). The 1D analytical
9 method predicted high quality habitat at three of four cross-sections, the 1D model predicted
10 medium quality habitat at three of four, and the 2D model predicted medium quality habitat at
11 two and low quality habitat at two. The 2D model's spatial pattern of GHSI for design 2 showed
12 a paucity of high quality habitat, except for the channel margin and the entrance and exit slopes
13 of the TOR (Fig. 7d). High quality habitat was only predicted on riffle crests and the TOR riffle
14 exit leading to Pool 2, while medium quality habitat was predicted along the right side of the
15 TOR. Low quality habitat was predicted along the channel margins of both pools, while very
16 poor habitat was predicted in the center of both pools.

17

18 **DISCUSSION**

19 The focus of this paper was evaluating the hydraulic predictions associated with three test
20 methods using two different riffle-pool unit types at two different discharges to determine what
21 impact method choice has on sediment transport regime and physical habitat quality predictions
22 relevant for river rehabilitation design evaluation. The results illustrate not only the differences
23 between the test methods, but also how different riffle-pool morphologies function in a straight

1 channel with no floodplain connectivity.

2

3 **Evaluation of Hydraulic Predictions**

4 Considering known patterns of water surface and velocity profiles over riffle and pool
5 sequences, methodologies and tools to assess said features should at least account for general
6 trends found in the stream-wise direction. For example, based on common observation, when
7 discharge is at or below bankfull, flow over riffles should be shallow and fast, while over pools it
8 should be deep and slow. During rising floods, as the ratio of water depth to local bed variation
9 decreases, water surface slope, depth, and velocity should become more uniform. At the very
10 least, tools used to assess channel hydraulics over riffles and pools for any purpose should
11 capture these basic traits. In this study, the test designs represented two different types of
12 channel morphologies- one with features generally orthogonal to the channel banks and one with
13 a single large feature oblique to the channel banks. The topographical setting of these designs
14 had a direct impact on the performance of each approach to represent channel hydraulics and
15 thus, they are discussed individually.

16 The key result consistently observed in this study was that the 1D analytical method
17 performed poorly at predicting depth and velocity. This divergence in predictive capability
18 between analytical and numerical methods lay in the ability of each approach to properly account
19 for complex flow dynamics resulting from channel non-uniformities associated with riffle-pool
20 units common to gravel-bed rivers. Complex hydraulics involve convective accelerations and
21 decelerations at the scale of the entire cross-section (i.e. those due to channel-wide narrowing
22 and widening or channel deepening and shallowing between cross-sections) as well as those at
23 the sub-section scale (i.e. a slice of an oblique feature, a boulder cluster, a small gravel bar, etc).

1 Neither of these scales of variation is accounted for in the 1D analytical method. Only the
2 former is accounted for in a 1D numerical model. A 2D model can account for both of these
3 realities.

4 The comparison of 1D and 2D numerical models found that they both yielded the same
5 longitudinal patterns in cross-sectionally averaged hydraulic variables. The values for these
6 variables were close enough that there was no objective way to determine which was more
7 correct in the absence of observational data for a real stream. However, the fact that the
8 predictions for design 2, with its complex oblique morphology, had larger differences than those
9 for design 1, suggests that the 1D model was the one that was deficient in that case. For the low-
10 flow test, the velocity vectors predicted by the 2D model showed strong convergence and
11 divergence at the sub-section scale as well as two sizable whirlpools- all between the cross-
12 sections where the two models were compared. The whirlpools were large enough that sediment
13 in transport could deposit there and fish searching for desirable adult holding habitat could
14 choose to rest there. We have even observed fish building “backwards” redds on the Yuba
15 River, where whirlpools create adequate upstream velocities. In the high-flow test, the 2D model
16 showed very strong lateral velocity gradient for the TOR design, so it is likely that the 1D
17 model’s simple average was in error in that case.

18

19 **Sediment Transport Regime Prediction**

20 The ability of the different test methods to predict Shields stress and sediment transport
21 regime varied strongly and was different between the two test designs. Even though 1D
22 analytical methods are the dominant approach used to estimate Shields stress in a wide range of
23 fluvial geomorphic applications, including habitat assessments, flood impact assessments, and

1 landscape evolution models, this study found that more than half of the cross-sections were
2 misidentified for their sediment transport regime. The worst Shields stress value predicted using
3 the 1D analytical model was off by 1450 %, while the best was off by 8 %, creating large
4 uncertainty. Also, while the 1D numerical model performed well overall, it was unable to
5 correctly identify the special risk of erosion associated with convergent, horseshore-shaped (in
6 plan view) riffle exits where knickpoint migration can lead to systemic failure of gravels placed
7 in any curved morphology, such as the commonly used chevron, which has two downstream
8 converging horseshoes- one along either bank. Many riffles in natural are associated with
9 oblique bars that also have convergent, horseshoe-shaped locations along their riffle exits.

10

11 **Physical Habitat Prediction**

12 Physical habitat quality is a common metric for predicting fish utilization of instream
13 channel units. While this approach does not capture all relevant parameters (e.g. bioenergetics,
14 temperature, predation, and water quality), it does provide an objective basis for assessing
15 competing methods to predict channel hydraulics. Physical spawning habitat should exhibit
16 longitudinal variations that correspond with riffle and pool topography, with riffle sections
17 having a higher quality habitat than pools. Both numerical models showed that pattern, while the
18 1D analytical method did not. However, the 1D approaches consistently overpredicted habitat
19 quality relative to the 2D model. Also, relative to their ability to predict Shield stresses, both 1D
20 approaches performed worse at predicting GHSI values, likely due to the highly nonlinear nature
21 of the empirical functions describing habitat conditions. Since fish species in any freshwater
22 lifestage are capable of distinguishing and utilizing laterally isolated pockets of habitat, the
23 unique ability of the 2D model to resolve those features, as demonstrated for both design

1 morphologies tested (Figs. 4d, 7d), is notable and useful for design evaluation. Even though 3D
2 models can further resolve vertical velocity gradients that we have observed fish utilize, we are
3 not aware of any habitat suitability metric for velocity that is available in 3D at this time, nor
4 does it seem cost-effective or likely that biologists will pursue that capability in the near future.

6 **Broad Flat Riffles**

7 The broad flat riffle, as the name implies, is a uniformly graded riffle that has cross
8 sectional symmetry about its midpoint. Consequently, it is a frequently used feature in SHR for
9 its apparent ease in design, evaluation, and constructability. However, this study shows that
10 broad flat riffles do not exhibit the hydraulic uniformity assumed in many designs, especially
11 over the crest of the structure. Consequently, even using the 1D analytical method at BFR's is
12 highly problematic. This is illustrated by comparing predictions for the cross sections at the
13 middle and crest of the BFR in design 1 (Fig. 1, first two cross-sections). Both numerical models
14 show the flow accelerating toward the crest, which the 1D analytical method does not (Fig. 3a).
15 Consequently, the 1D analytical method underpredicted Shields stress on the BFR crest. In the
16 high-discharge test, the BFR crest was predicted by the numerical models to be in the upper
17 range of partial transport, whereas the 1D analytical method barely predicted partial transport at
18 all (Fig. 9b). Even at low flow, the numerical models predict moderate intermittent transport,
19 which for freshly placed, loose gravels could actually correspond with partial transport, not
20 intermittent transport. These considerations explain why commonly built BFR's with uniform
21 gravels fail rapidly, even though designers do not expect them to. To make matters worse, both
22 1D methods predicted high spawning habitat quality over the BFR crest, whereas the 2D model
23 predicted medium quality habitat there. This explains why designers are likely to think that

1 BFR's are more desirable than they end up being. We have heard several anecdotal stories by
2 SHR project managers that BFR projects failed to be used by fish after being built. Overall, use
3 of the 1D analytical method overpredicts habitat quality and channel stability, and thus leads to
4 an overvaluing of broad flat riffles in SHR design.

6 **Comparison of Design Features**

7 The two different designs tested in this study have been pointed out to differ in the
8 orientations of their features relative to the channel banks. Both the BFR and the LCF were
9 orthogonal to channel banks, while the TOR was oblique to them. This difference resulted in the
10 1D methods performing worse than the 2D model in characterizing the conditions at the TOR,
11 because such bed features invoke flow patterns that are predominantly controlled by local bar
12 topography (Whiting and Dietrich, 1991). Many natural and artificial riffles have oblique
13 features, and thus it is evident from this study that they need to be analyzed using a 2D model,
14 even to obtain accurate cross-sectionally averaged predictions, let alone to characterize sub-
15 section variability. Without a 2D model, it would not be possible to evaluate the risk posed by
16 convergent, horseshoe-shaped riffle exits to riffle stability.

17 Another important distinction between design 1 and design 2 is that design 1 involved a
18 "blanket" fill approach that entails placement of a moderate amount of gravel everywhere,
19 whereas design 2 was an accentuated fill involving most gravel being placed on the riffle crests.
20 The advantage of accentuating topography as in design 2 is that it creates much more complex
21 hydraulics and thus more diverse habitats. Each freshwater lifestage of a fish species has
22 different needs, and design 2 is more likely to have habitat for all lifestages. However, design 2
23 not only imposed prediction-based problems, but the results of this study suggest that design 2

1 would not be geomorphically stable. One reason for that is that the channel has uniform width,
2 and thus it is highly improbable that a stage-dependent reversal in cross-sectional area, velocity,
3 and Shields stress between riffles and pools can be established to maintain channel relief
4 (MacWilliams et al., 2006). No such reversal was evident in the test at $170 \text{ m}^3/\text{s}$, which was 20
5 times higher than the low-flow test. The other reason is that the steep drop-off at each riffle exit
6 inevitably creates excessive velocities and localized bed scour. This can cause knickpoint
7 migration that can quickly cut through the entire feature. Such an effect was observed for the
8 2001 project site on the lower Mokelumne River (Merz et al., 2006) when less gravel was
9 provided than called for by the design of Wheaton et al. (2004b), resulting in a steep, horse-shoe
10 shaped riffle exit at the downstream end of the project site. For the actual SHR project on the
11 Trinity River, a thorough analysis was done for three different accentuated-relief designs and
12 three different blanket-fill designs. In the end, the selected design was a blanket-fill morphology
13 using one broad flat riffle and two L-shaped bars. Details of all these designs and their
14 evaluations are available at <http://shira.lawr.ucdavis.edu/trinity.htm>.

16 **Implications to SHR Design and Decision Making**

17 Decision making for river management often relies on model predictions of physical
18 habitat to forecast potential impacts of flow regime changes as well as for evaluating potential
19 gravel augmentation designs. In this study a sediment transport regime metric and a global
20 habitat suitability index were used to transform quantitative hydraulic predictions into relevant
21 qualitative descriptors of geomorphic and ecologic conditions. Deviations in hydraulic
22 predictions led to contrasting results for section-average Shields stress and GHSI for each test
23 design, ultimately being related to whether the topography of the features initiated cross-stream

1 flow vectors and lateral variation in stream-wise vectors that deviated from cross-sectionally
2 averaged estimates.

3 Manning's equation and other analytical flow resistance equations are often considered
4 adequate for engineering tasks involving open channel flows in deep channels with a high depth
5 to bed-feature height ratio (>100 - 1000 as a conservative estimate). In this study, that ratio
6 ranged from ~ 1 - 10 . The results of this study show that for shallow gravel-bed rivers with
7 complex morphologies, including common river rehabilitation design structures of simple
8 geometry, the 1D analytical method is inappropriate for estimating depth and velocity. These
9 limitations can be further assumed to affect any non-uniform channel section with a width to
10 depth ratio in this range, and probably up to at least 50. Despite this inadequacy, there is a
11 practice of "fine-tuning" Manning's equation using detailed accounting of roughness in the
12 channel and velocity-profile effects. The benefits of these and other attempts to bend Manning
13 equation to fit the desire for a simple method are unlikely to be of use for SHR design
14 evaluation.

15 The results of this study suggest that a river manager, policy maker, scientist, or engineer
16 faced with evaluating the effect of a potential gravel augmentation design, a channel
17 modification, or a flow regime change on physical habitat would come to very different
18 conclusions depending on whether the hydraulics of the channel were evaluated using Manning's
19 equation or a 1D or 2D numerical model. The quality of spawning habitat was consistently over
20 predicted by 1D methods. Moreover, considering reach scale (10^2 - 10^3 channel widths) averages
21 of spawning habitat quality by averaging GHSI among sections for both designs, 1D methods
22 still grossly over-predict habitat such that medium and low quality habitat would be predicted as
23 high quality habitat. As 1D reach-scale approaches to physical habitat are the most prevalent,

1 this could lead decision makers to believe that existing and proposed physical habitat is of higher
2 quality and quantity, leading to false conclusions about river health. Considering the rapid
3 decline of salmonids despite the known linkages to habitat degradation, it is not implausible to
4 postulate these over-predictions of existing and proposed physical habitat have already
5 manifested themselves into key management plans of gravel bed rivers and dam re-licensing.

6 Predictions regarding channel stability and sediment transport would be impacted as well
7 from using analytical or numerically based tools. Under predicting Shields stress over riffle
8 crests could be catastrophic to SHR based gravel augmentation as knickpoint migration at the
9 riffle crest could lead to systemic failure, especially for a sequence of riffle crests, which could
10 experience a “reverse domino” toppling of bedforms (Pasternack et al., in press). Moreover,
11 small areas of high stress may be present that no section averaged approach can capture. Failure
12 to identify these areas could also lead to riffle failure.

13

14 **CONCLUSIONS**

15 Typical goals of most spawning habitat rehabilitation projects involve increasing quantity
16 and quality of one or more life stages of a species. Placed gravels must be stable during low
17 discharges to avoid embryo mortality and have some sediment transport at high discharges to
18 rejuvenate gravels by removing aquatic vegetation and interstitial fines. Generally speaking, a
19 project is considered successful if the number of redds and consequent fry production increases
20 and the project is stable during the spawning and incubation seasons, with some transport during
21 large flow events after incubation. Therefore, the evaluation of potential designs rests on two
22 key parameters- sediment transport regime and physical habitat quality. In this study, three
23 methods were tested against two different channel morphologies at two different discharges.

1 Numerical models captured the general stream-wise hydraulics associated with riffles and pools
2 that a 1D analytical method could not predict. Failure to account for convective accelerations
3 within and between riffles and pools explains why the 1D analytical method was wrong in most
4 of its predictions. Although 1D and 2D numerical predictions of cross-sectionally averaged
5 values had good agreement for both designs at both discharges, significant variations in between
6 cross-sections and at the sub-section scale were not resolved by the 1D model. For example, the
7 2D model properly predicted the presence of whirlpools and zones of convergent flow in
8 horseshoe-shaped riffle exits. For spatially distributed phenomena such as patterns of sediment
9 erosion and fish habitat quality, it is highly recommended that 2D models be used to resolve
10 lateral variations, without which 1D models appear to over-predict habitat quality, affecting
11 long-term river management policies and practices.

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21 22 **REFERENCES**

23
24 Bates PD, Anderson MG, Baird L, Walling DE, Simm D. 1992. Modeling floodplain flow with a
25 two-dimensional finite element scheme *Earth Surface Processes and Landforms* **17**: 575-
26 588.
27

- 1 Bjornn TC, Reiser DW. 1991. Habitat requirements of salmonids in streams. In *Influences of*
2 *forest and rangeland management on salmonid fishes and their habitats*, Meehan WR. (Ed).
3 *American Fisheries Society Special Publication*. **19**: 83-138.
4
- 5 Bovee K D, Lamb BL, Bartholow JM, Stalnaker C.B, Taylor J, and Henriksen, J. 1998. Stream
6 habitat analysis using the instream flow incremental methodology. *United States Geological*
7 *Survey, Information and Technology Report USGS/BRD/ITR-1998-0004*. Fort Collins, CO.
8
- 9 Brown RA, Pasternack, GB. 2008. Engineered Channel Controls Limiting Spawning Habitat
10 Rehabilitation. *Geomorphology* **97**: 631-654. .
11
- 12 Brunner GW. 1998. HEC-RAS: River Analysis System. *Unpublished Users Manual, Version*
13 *2.2*. U.S. Army Corps of Engineers: CA.
14
- 15 Carling P A. 1991. An appraisal of the velocity-reversal hypothesis for stable pool-riffle
16 sequences in the River Severn, England. *Earth Surface Processes and Landforms* **16**: 19-31.
17
- 18 California Department of Fish and Game. 1998. *California Salmonid Stream Habitat*
19 *Restoration Manual*. 3rd edition.
20
- 21 California Department of Water Resources. 2000. *Merced River Robinson-Gallo Project -*
22 *Ratzlaff Reach Engineering Report*.
23
- 24 Cao Z. and Carling PA. 2002. Mathematical modeling of alluvial rivers: reality and myth. Part
25 1: general overview. *Water and Maritime Engineering*, **154**: 207-219.
26
- 27 Cao Z, Carling P, Oakey R. 2003. Flow Reversal over a Natural Pool-Riffle Sequence: A
28 Computational Study. *Earth Surface Processes and Landforms* **28**: 689-705.
29
- 30 Chang HH. 1998. *Fluvial Processes in River Engineering*. John Wiley and Sons: New York,
31 NY,
32
- 33 Clifford NJ, Richard KS. 1992. The Reversal Hypothesis and the Maintenance of Riffle-Pool
34 Sequences: A Review and Field Appraisal. In *Lowland Floodplain Rivers:*
35 *Geomorphological Perspectives*, Carling, PA, Petts J. (Eds). Wiley and Sons: USA.
36
- 37 Darby SE, Van de Wiel MJ. 2003. Models in fluvial geomorphology. In *Tools in fluvial*
38 *geomorphology*, Kondolf, GM, Piégay, H (Eds). Wiley: Chichester.
39
- 40 Dunne T, Leopold LB. 1987. *Water in Environmental Planning*. W. H. Freeman: San Francisco,
41 CA.
42
- 43 Elkins EM, Pasternack GB, Merz, JE. 2007. The use of slope creation for rehabilitating incised,
44 regulated, gravel-bed rivers. *Water Resources Research* **43**. DOI:10.1029/2006WR005159.
45
- 46 Fischer HB, List JE, Koh CR, Imberger J., Brooks, NH. 1979. *Mixing in Inland and Coastal*

- 1 *Waters*. Academic Press, Inc: New York, NY.
2
- 3 French JR, and Clifford NJ. 2000. Hydrodynamic modeling as a basis for explaining estuarine
4 environmental dynamics: some computational and methodological issues. *Hydrological*
5 *Processes* **14**: 2089-2108.
6
- 7 Froehlich DC. 1989. HW031.D - Finite Element Surface-Water Modeling System: Two-
8 Dimensional Flow in a Horizontal Plane-Users Manual. *Federal Highway Administration.*
9 *Report FHWA-RD-88-17*. Federal Highway Administration: Washington DC.
10
- 11 Gard M. 2006. Modeling changes in salmon spawning and rearing habitat associated with river
12 channel restoration. *International Journal of River Basin Management*. **4**: 1-11.
13
- 14 Ghanem A, Steffler P, Hicks F. 1996. Two-dimensional hydraulic simulation of physical habitat
15 conditions in flowing streams. *Regulated Rivers Research and Management* **12**: 185-200.
16
- 17 Graf WL. 2001. Damage Control: Dams and the Physical Integrity of America's Rivers. *Annals*
18 *of the Association of American Geographers* **91**: 1-27.
19
- 20 Harper D, Ebrahimnezhad M, Cot FCI. 1998. Artificial riffles in river rehabilitation: setting the
21 goals and measuring the successes. *Aquatic Conservation-Marine and Freshwater*
22 *Ecosystems* **8**: 5-16.
23
- 24 Harrison, LR, Keller EA. Modeling forced pool--riffle hydraulics in a boulder-bed stream,
25 southern California. *Geomorphology* **83**: 232-248
26
- 27 Hunter CJ. 1991. *Better Trout Habitat: A Guide to Stream Restoration and Management*.
28 Montana Land Alliance. Island Press: New York, NY.
29
- 30 Keller EA. 1971. Areal sorting of bed-load material; the hypothesis of velocity reversal. *GSA*
31 *Bulletin* **82**: 753-756.
32
- 33 Keller EA, Florsheim JL. 1993. Velocity-Reversal Hypothesis: A Model Approach. *Earth*
34 *Surface Processes and Landforms* **18**: 733-740.
35
- 36 Knighton, A.D., 1998. *Fluvial Forms and Processes: A New Perspective*. Arnold, London.
37
- 38 Kondolf GM, Vick JC, Ramirez TM. 1996. Salmon spawning habitat rehabilitation on the
39 Merced River, California: An evaluation of project planning and performance. *Transactions*
40 *of the American Fisheries Society* **125**: 899-912.
41
- 42 Kondolf GM. 2000. Some suggested guidelines for geomorphic aspects of anadromous salmonid
43 habitat restoration proposals. *Restoration Ecology* **8**: 48-56.
44
- 45 Lacey RW, Millar RG. 2004. Reach scale hydraulic assessment of instream salmonid habitat
46 restoration. *Journal of the American Water Resources Association* **40**: 1631-1644.

- 1
2 Leclerc M, Boudreault A, Bechara JA, Corfa G. 1995. 2-Dimensional Hydrodynamic Modeling -
3 a Neglected Tool in the Instream Flow Incremental Methodology. *Transactions of the*
4 *American Fisheries Society* **124**: 645-662.
5
6 Leopold LB, Wolman MG, Miller JP. 1964. *Fluvial Processes in Geomorphology*. W.H.
7 Freeman and Company: San Francisco, CA.
8
9 Lisle T. 1979. A sorting mechanism for a riffle-pool sequence. *Geological Society of America*
10 *Bulletin* **11**: 1142-1157.
11
12 Lisle TE, Nelson JM, Pitlick J, Madej MA, Barkett BL. 2000. Variability of bed mobility in
13 natural, gravel-bed channels and adjustments to sediment load at local and reach scales.
14 *Water Resources Research* **36**: 3743-3755.
15
16 MacWilliams ML, Wheaton JM, Pasternack GB, Kitanidis PK, Street, RL. 2006. The Flow
17 Convergence-Routing Hypothesis for Pool-Riffle Maintenance in Alluvial Rivers. *Water*
18 *Resources Research* **42**, W10427, doi:10.1029/2005WR004391.
19
20 Maddock I, 1999. The importance of physical habitat assessment for evaluating river health.
21 *Freshwater Biology* **41**.
22
23 McBain, S, Trush, B. 2003. *Trinity River Coarse Sediment Management Plan*.
24
25 Merz JE, Setka JD. 2004. Evaluation of a Spawning Habitat Enhancement Site for Chinook
26 Salmon in a Regulated California River. *North American Journal of Fisheries Management*:
27 **24**, 397-407.
28
29 Merz JE, Pasternack GB, Wheaton JM. 2006. Sediment Budget for Salmonid Spawning Habitat
30 Rehabilitation in the Mokelumne River. *Geomorphology*: **76**, 207-228.
31
32 Merz JE, Pasternack GB, Wheaton JM. 2004. Predicting benefits of spawning habitat
33 rehabilitation to salmonid fry production in a regulated California river. *Canadian Journal of*
34 *Fisheries and Aquatic Science*, **61**:1433-1446.
35
36 Milhous RT, Updike M, Schneider DM. 1989. Physical Habitat Simulation System Reference
37 Manual - Version II. *U.S. Fish and Wildlife Service Instream Flow Information Paper No.*
38 *26*. Fort Collins, CO.
39
40 Milne, JE. 1982. Bed-material size and the riffle-pool sequence. *Sedimentology* **29**:
41 267-278.
42
43 Miller AJ, Cluer BL. 1998. Modeling considerations for simulation of flow in bedrock channels.
44 In *Rivers Over Rock: Fluvial Processes in Bedrock Channels*, Wohl EE, Tinkler KJ (Eds).
45 American Geophysical Union: Washington, DC; 61-104.
46

- 1 Moir, HJ, Pasternack, GB. 2008. Relationships between mesoscale morphological units, stream
2 hydraulics and Chinook salmon (*Oncorhynchus tshawytscha*) spawning habitat on the Lower
3 Yuba River, California. *Geomorphology*. doi:10.1016/j.geomorph.2008.02.001
4
- 5 Nation Research Council. 1992. *Restoration of Aquatic Ecosystems: Science, Technology, and*
6 *Public Policy*. NRC committee on Restoration of Aquatic Ecosystems, Washington, DC.
7
- 8 Nelson JM, Bennett JP, Wiele SM. 2003. Flow and sediment-transport modeling. In *Tools in*
9 *fluvial geomorphology*, Kondolf GM, Piégay, H (Eds). Wiley, Chichester, 539-576.
10
- 11 Newbury R, Gaboury M, Bates D. 1997. Restoring Habitats in Channelized or Uniform Streams
12 Using Riffle and Pool Sequences. Chapter 12. *Fish Habitat Rehabilitation Procedures.*
13 *Watershed Restoration Technical Circular No. 9*. British Columbia Watershed Restoration
14 Program: Vancouver, BC
15
- 16 Pasternack GB, Bounrisavong MK, Parikh KK. 2008. Backwater Control on Riffle-Pool
17 Hydraulics, Fish Habitat Quality, and Sediment Transport Regime in Gravel-Bed Rivers.
18 *Journal of Hydrology* **357**:1-2:125-139.
19
- 20 Pasternack GB, Wang CL, Merz J. 2004. Application of a 2D hydrodynamic model to reach-
21 scale spawning gravel replenishment on the lower Mokelumne River, California. *River*
22 *Research and Applications* **20**: 205-225.
23
- 24 Pasternack GB, Gilbert AT, Wheaton JM., Buckland EM. 2006. Error Propagation for Velocity
25 and Shear Stress Prediction Using 2D Models For Environmental Management. *Journal of*
26 *Hydrology* **328**:227-241.
27
- 28 Rathburn S. Wohl EE. 2003. Predicting Fine Sediment Dynamics along a Pool-Riffle
29 Mountain Channel. *Geomorphology* **55**: 111-124.
30
- 31 Richards KS. 1976. The Morphology of Riffle-Pool Sequences. *Earth Surface Processes and*
32 *Landforms* **1**: 71-88.
33
- 34 Rosgen D. 1996. *Applied River Morphology*. Wildland Hydrology: Pagosa Springs, CO.
35
- 36 Saldi-Caromile K, Bates K, Skidmore P, Barenti J, Pineo D. 2004. *Stream Habitat Restoration*
37 *Guidelines: Final Draft*. Washington Departments of Fish and Wildlife and Ecology and the
38 U.S. Fish and Wildlife Service: Olympia, Washington.
39
- 40 Sawyer AM, Pasternack GB, Merz JE, Escobar M., Senter AE. 2008. Construction constraints on
41 geomorphic-unit rehabilitation on regulated gravel-bed rivers. *River Research and*
42 *Applications*. doi: 10.1002/rra.1173.
43
- 44 Sear DA. 1996. Sediment transport processes in pool-riffle sequences. *Earth Surface Processes*
45 *and Landforms* **21**: 241-262.
46

- 1 Shields FD, Copeland RR, Klingeman PC, Doyle, MW, Simon A. 2003. Design for stream
2 restoration. *Journal of Hydraulic Engineering* **129**: 575-584.
3
- 4 Slaney PA, Zaldokas D. 1997. *Fish Habitat Rehabilitation Procedures*. Watershed Restoration
5 Technical Circular No. 9. Ministry of Environment Lands and Parks: Vancouver, BC
6
- 7 Smith AK. 1973. Development and application of spawning velocity and depth criteria for
8 Oregon salmonids. *Transactions of the American Fisheries Society* **102**: 312-316.
9
- 10 Thompson DM, Wohl EE, Jarrett RD. 1999. Velocity reversals and sediment sorting in pools and
11 riffles controlled by channel constrictions. *Geomorphology* **27**: 229-241.
12
- 13 Torgersen CE, Price DM, Li HW, McIntosh BA. 1999. Multiscale thermal refugia and stream
14 habitat associations of chinook salmon in northeastern Oregon. *Ecological Applications* **9**:
15 301-319.
16
- 17 Trush WJ, McBain SM, Leopold L B. 2000. Attributes of an alluvial river and their relation to
18 water policy and management. *Proceedings of the National Academy of Sciences of the*
19 *United States of America* **97**: 11858-11863.
20
- 21 United States Fish and Wildlife Service. 1997. *Microhabitat Suitability for Anadromous*
22 *Salmonids of the Trinity River*. U.S. Fish and Wildlife Service: Sacramento, CA.
23
- 24 United States Fish and Wildlife Service. 1999. *Trinity River Flow Evaluation*. Department of the
25 Interior: Washington, DC.
26
- 27 Walker DR, Millar RG, Newbury RW. 2004. *Hydraulic design of riffles in gravel-cobble bed*
28 *ivers*. *Journal of River Basin Management* **2**: 291-299.
29
- 30 Wheaton JM, Pasternack GB, Merz JE. 2004a. Spawning Habitat Rehabilitation - 1. Conceptual
31 Approach & Methods. *International Journal of River Basin Management* **2**: 3-20.
32
- 33 Wheaton JM, Pasternack GB, Merz JE. 2004b. Spawning Habitat Rehabilitation - 2. Using
34 hypothesis development and testing in design, Mokelumne River, California, U.S.A.
35 *International Journal of River Basin Management* **2**: 21-37.
36
- 37 Wheaton JM, Pasternack GB, Merz JE. 2004c. Use of habitat heterogeneity in salmonid
38 spawning habitat rehabilitation design. In *Fifth International Symposium on Ecohydraulics:*
39 *Aquatic Habitats: Analysis and Restoration*. IAHR-AIRH: Madrid, Spain.
40
- 41 Whiting, P. J., and W. E. Dietrich, 1991, [Convective accelerations and boundary shear stress](#)
42 [over a channel bar](#). *Water Resources Research* **27**: 783-796.
43
- 44 Wilcock PR, Kondolf GM, Matthews WV, Barta AF. 1996a. Specification of sediment
45 maintenance flows for a large gravel-bed river. *Water Resources Research* **32**:2911-2921.
46

- 1 Wilcock PR, Barta AF, Shea CC, Kondolf GM, Matthews WVG, Pitlick JC. 1996b.
2 Observations of flow and sediment entrainment on a large gravel-bed river. *Water Resources*
3 *Research* **32**: 2897–2909.
4
- 5 Wilkinson SN, Keller RJ, Rutherford ID. 2004. Phase-shifts in shear stress as an explanation for
6 the maintenance of pool-riffle sequences. *Earth Surface Processes and Landforms* **29**: 737-
7 753.
8
- 9 Wohl EE, Kirk VR, Merritts DJ. 1993. Pool and Riffle Characteristics in Relation to Channel
10 Gradient. *Geomorphology* **6**: 99-110.
11
- 12 Yalin MS. 1977. *Mechanics of Sediment Transport*. 2nd Ed. Peragamon.
13
- 14 Yen, B.C. 1991. *Channel Flow Resistance: Centennial of Manning's Formula*. Water
15 Resources Publications: Littleton, CO.
16
- 17 Yoshiyama RM, Fisher FW, Moyle PB. 1998. Historical abundance and decline of chinook
18 salmon in the Central Valley region of California. *North American Journal of Fisheries*
19 *Management* **18**: 487-521.
20
- 21 Zeh, M., and Donni, W. 1994. Restoration of Spawning Grounds for Trout and Grayling in the
22 River High-Rhine. *Aquatic Sciences* **56**: 59-69.
23
24
25

LIST OF FIGURES

Figure 1. Design topography for test design (a) one and (b) two. Note that in Design 1 the riffle crests are convergent being mostly orthogonal to the channel bank, while in Design 2 the topography of the riffle crest is transverse to the channel banks.

Figure 2. Definition sketch of geomorphic slope (Modified after Knighton, 1998). In this study the geomorphic slope is defined as the elevation difference between an upstream feature and the next downstream feature of the same type (i.e. bar or pool) divided by the distance between them.

Figure 3. Design 1 comparison plots of (a) velocity and (b) water surface elevation at $8.5 \text{ m}^3/\text{s}$ and (c) velocity and (d) water surface elevation at $170 \text{ m}^3/\text{s}$. The morphological unit labels refer to cross section locations where section-averaged values of depth, velocity, Shields stress, and GHSI were compared. Lightly shaded area is region within 2 standard deviations of the mean value predicted by the 2D model. Dark shaded area represents the region under the river bed.

Figure 4. Design 1, 2D model plots of (a) depth, (b) velocity, (c) Shields stress, and (d) GHSI for $8.5 \text{ m}^3/\text{s}$. The vertical lines refer to cross section locations where section-averaged values of depth, velocity, Shields stress, and GHSI were compared.

Figure 5. Design 1, 2D model plots of (a) depth, (b) velocity, (c) Shields stress for $170 \text{ m}^3/\text{s}$. The vertical lines refer to cross section locations where section-averaged values of depth, velocity, Shields stress, and GHSI were compared.

Figure 6. Design 2 comparison plots of (a) velocity and (b) water surface elevation at $8.5 \text{ m}^3/\text{s}$ and (c) velocity and (d) water surface elevation at $170 \text{ m}^3/\text{s}$. The morphological unit labels refer to cross section locations where section-averaged values of depth, velocity, Shields stress, and GHSI were compared. Lightly shaded area is region within 2 standard deviations of the mean value predicted by the 2D model. Dark shaded area represents the region under the river bed.

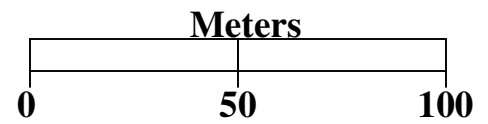
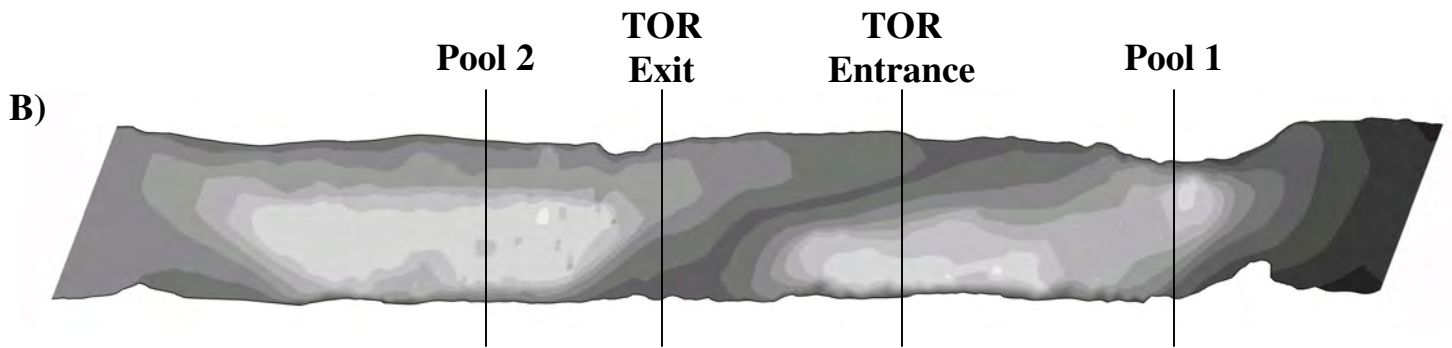
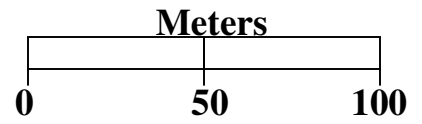
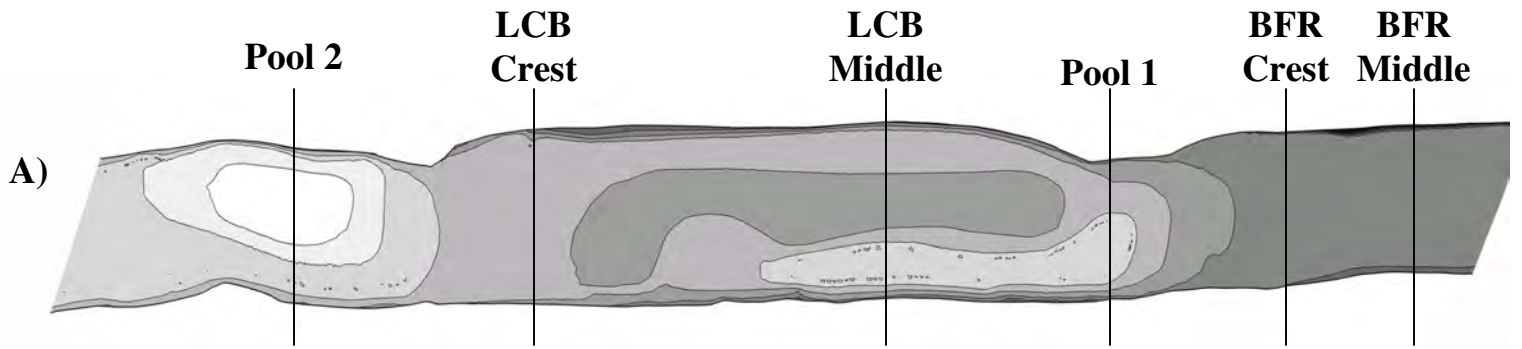
Figure 7. Design 2, 2D model plots of (a) depth, (b) velocity, (c) Shields stress, and (d) GHSI for $8.5 \text{ m}^3/\text{s}$. The vertical lines refer to cross section locations where section-averaged values of depth, velocity, Shields stress, and GHSI were compared.

Figure 8. Design 2, 2D model plots of (a) depth, (b) velocity, (c) Shields stress for $170 \text{ m}^3/\text{s}$. The vertical lines refer to cross section locations where section-averaged values of depth, velocity, Shields stress, and GHSI were compared.

Figure 9. Design 1 comparison of Shields stress at (a) $8.5 \text{ m}^3/\text{s}$ and (b) $170 \text{ m}^3/\text{s}$. The “*” results were calculated using a smoothed averaging method based on mean depth and velocity as opposed to values obtained from nodal averaging.

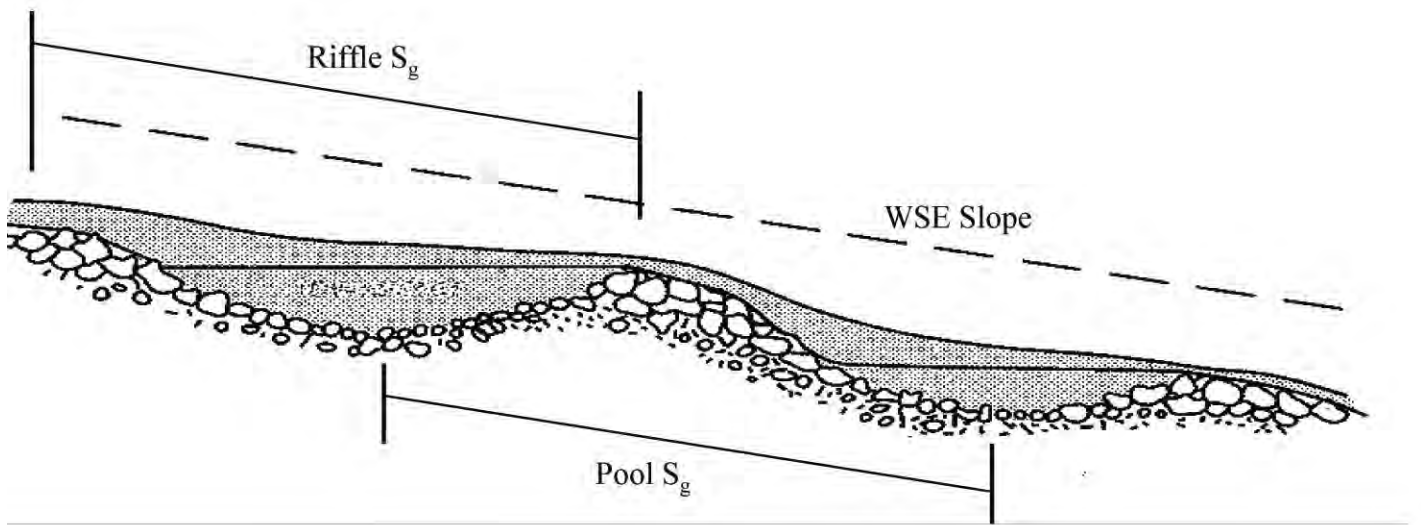
- 1 Figure 10. Design 2 comparison of Shields stress at (a) $8.5 \text{ m}^3/\text{s}$ and (b) $170 \text{ m}^3/\text{s}$. The “*”
2 results were calculated using a smoothed averaging method based on mean depth and
3 velocity as opposed to values obtained from nodal averaging.
4
- 5 Figure 11. GHSI comparison of (a) design 1 and (b) design 2 at $8.5 \text{ m}^3/\text{s}$. The “*” results were
6 calculated using a smoothed averaging method based on mean depth and velocity as opposed
7 to values obtained from nodal averaging.
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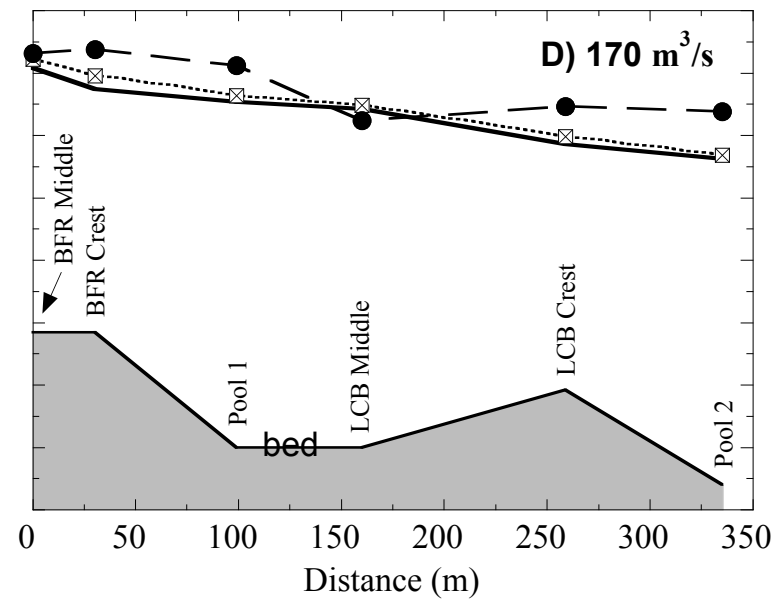
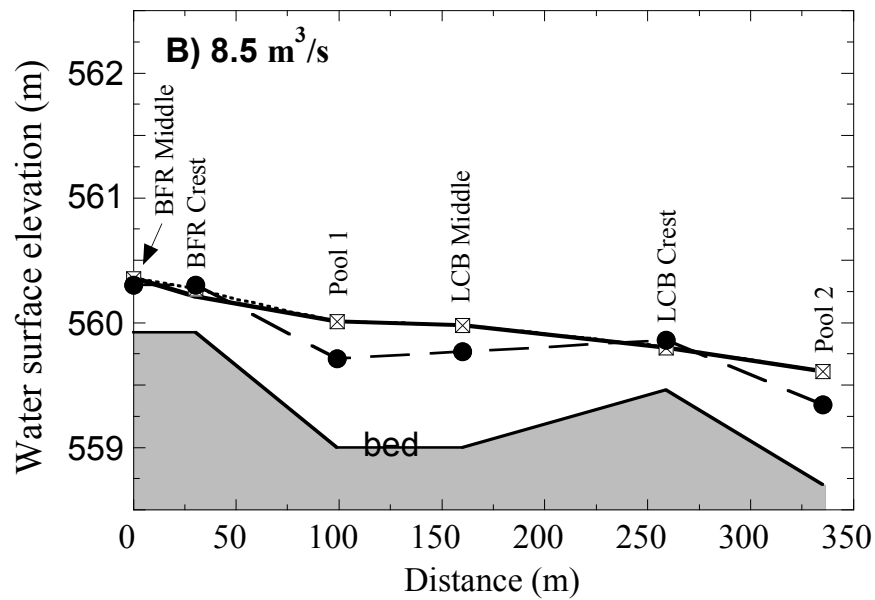
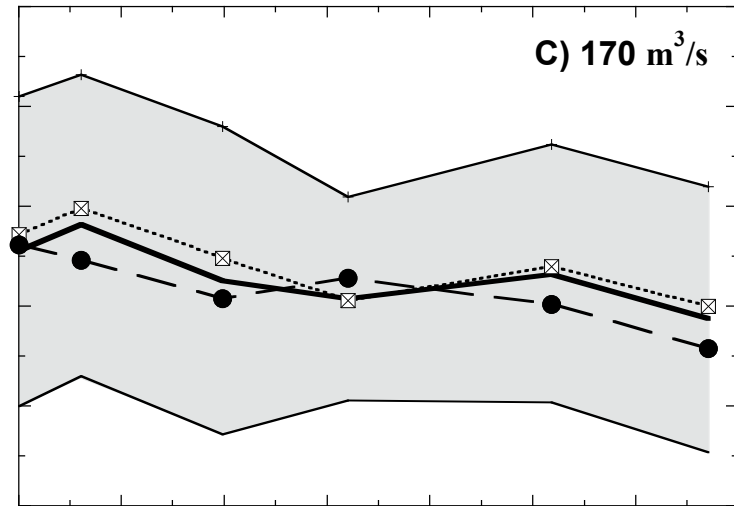
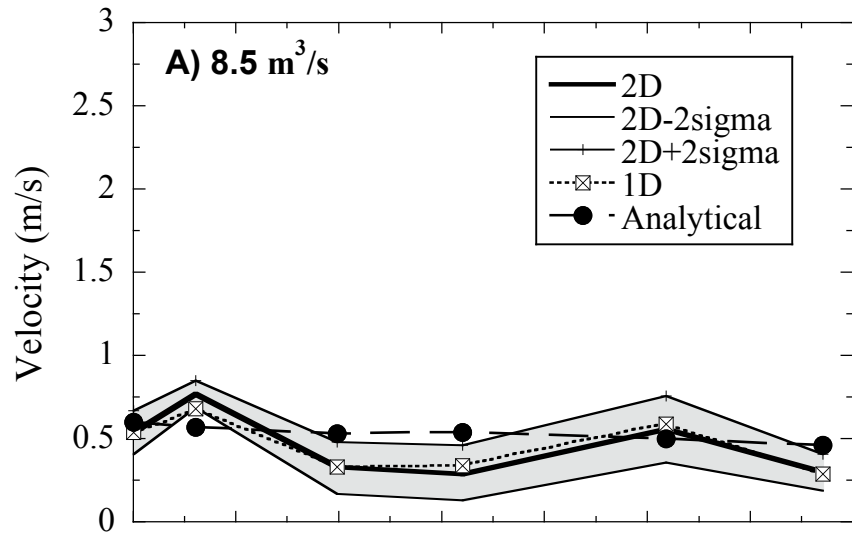
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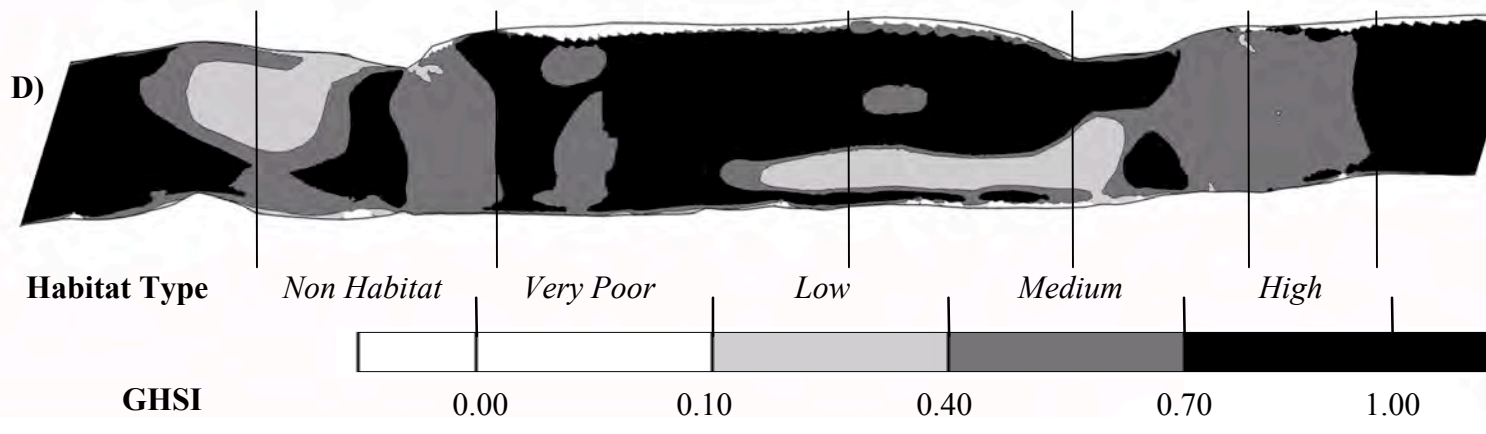
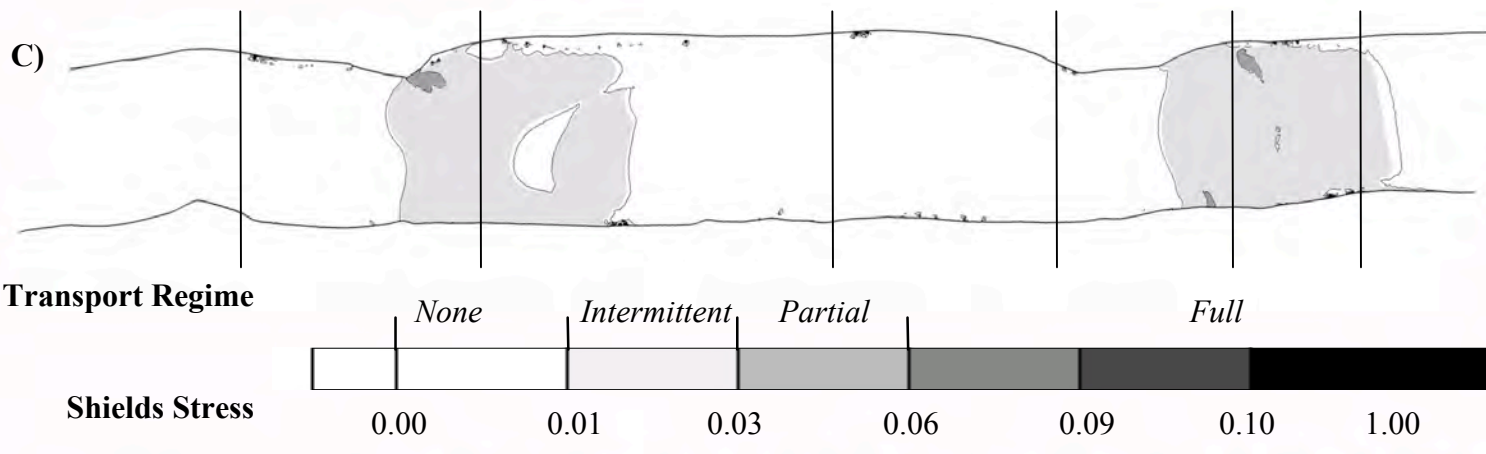
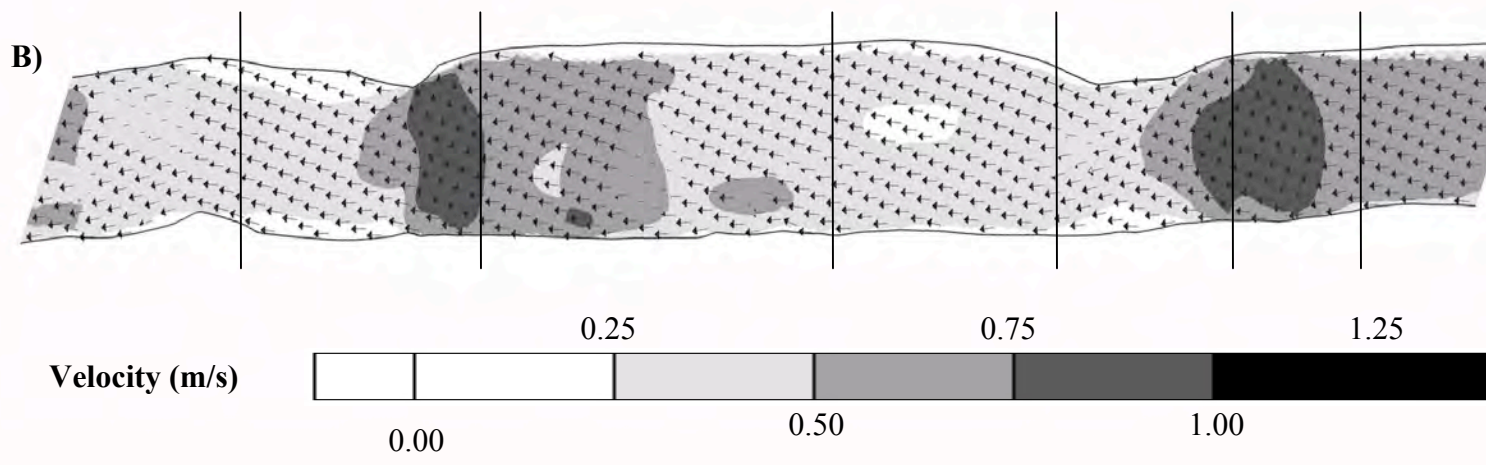
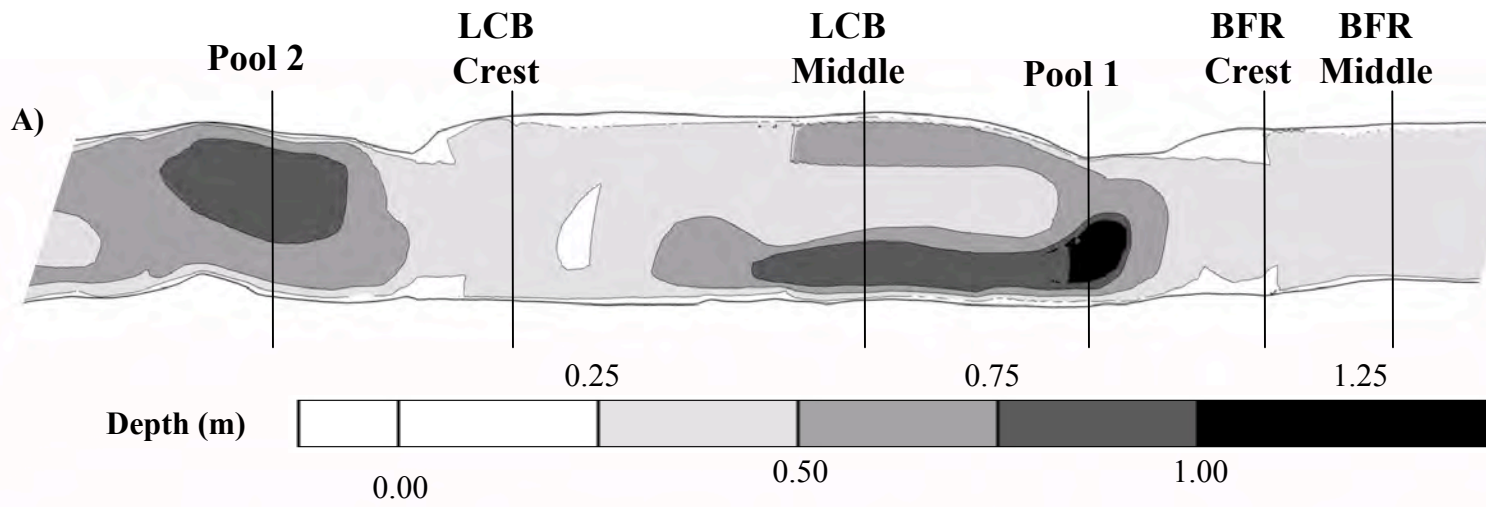


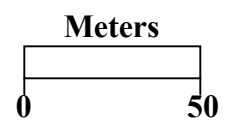
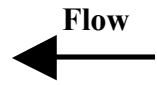
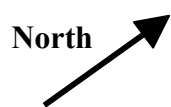
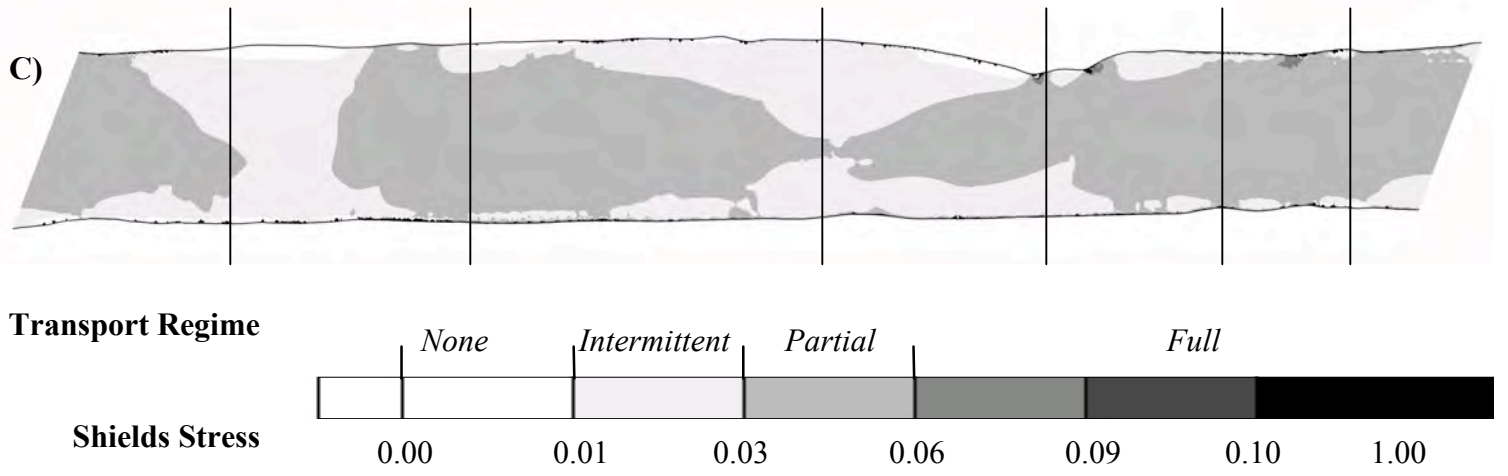
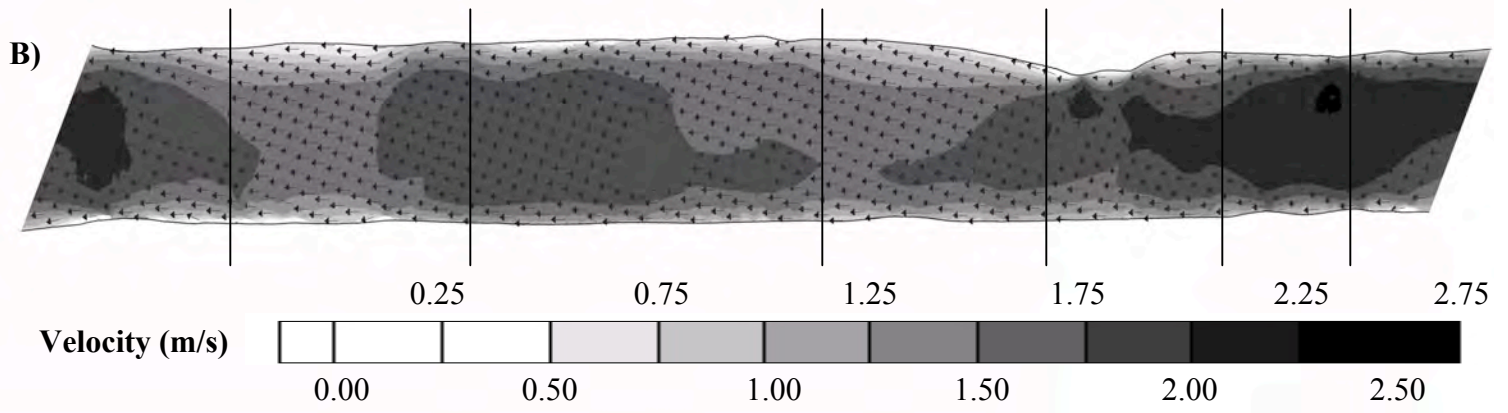
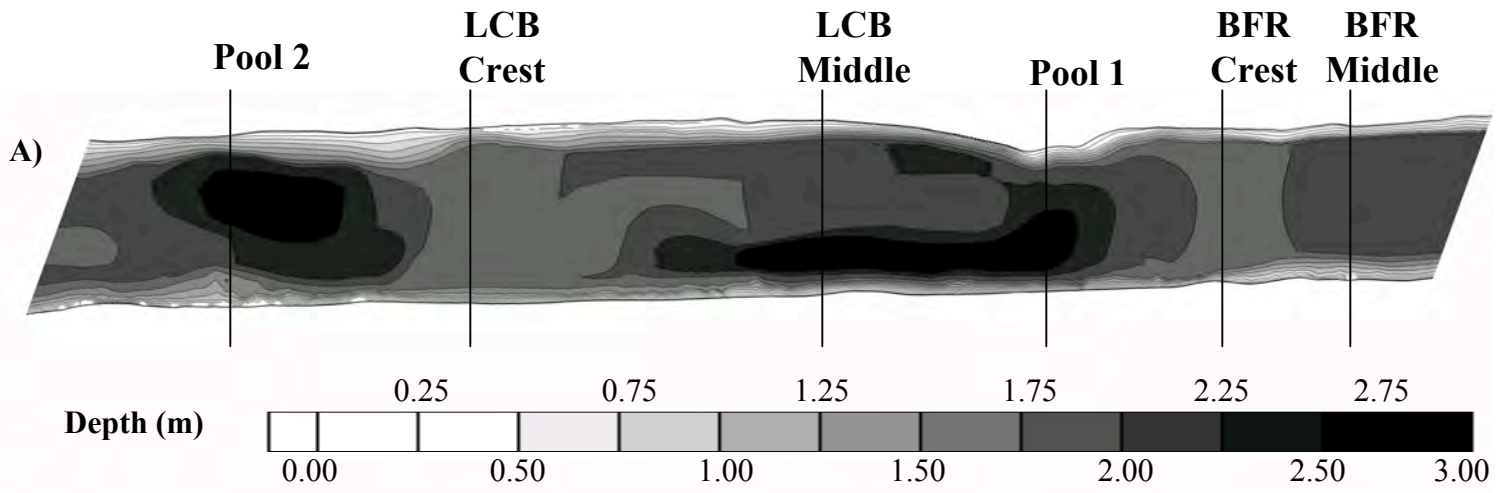
Elevation (m)

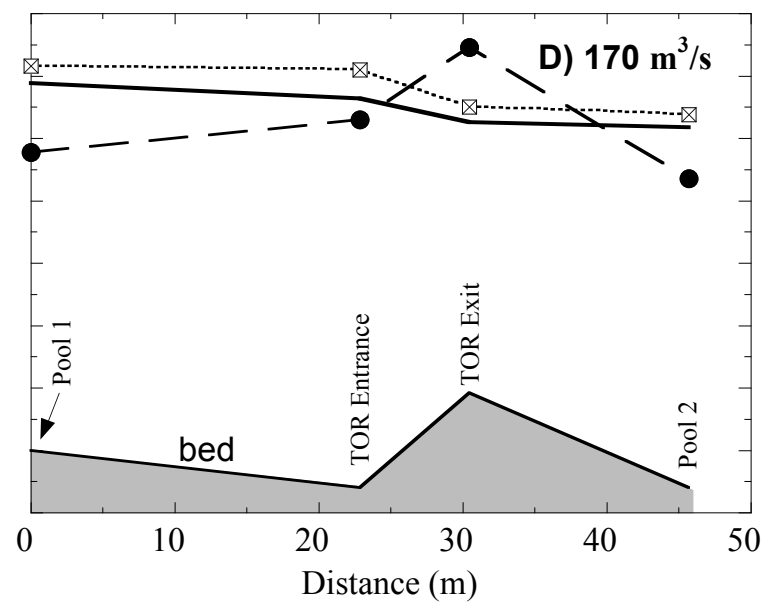
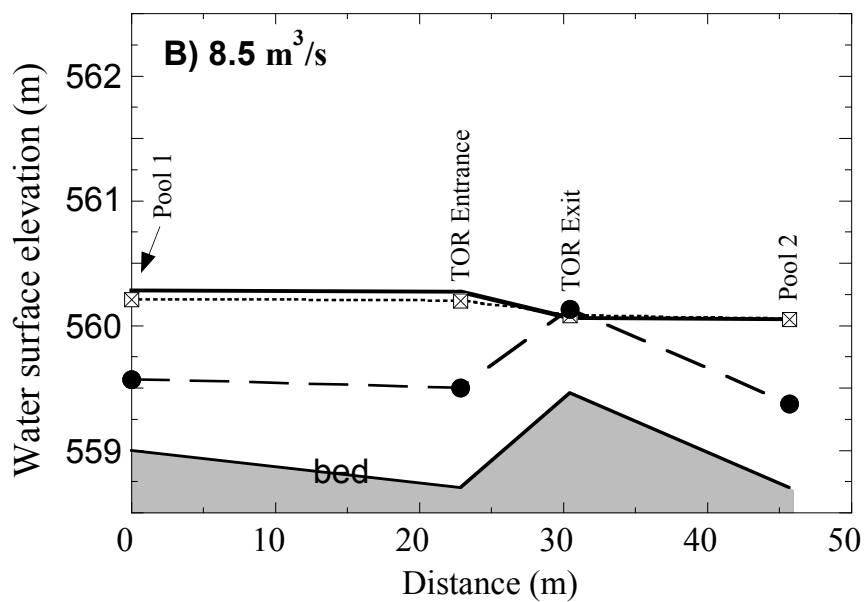
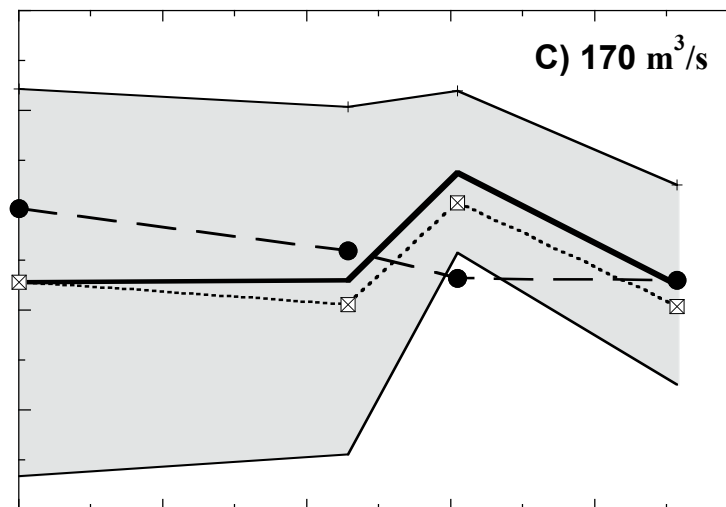
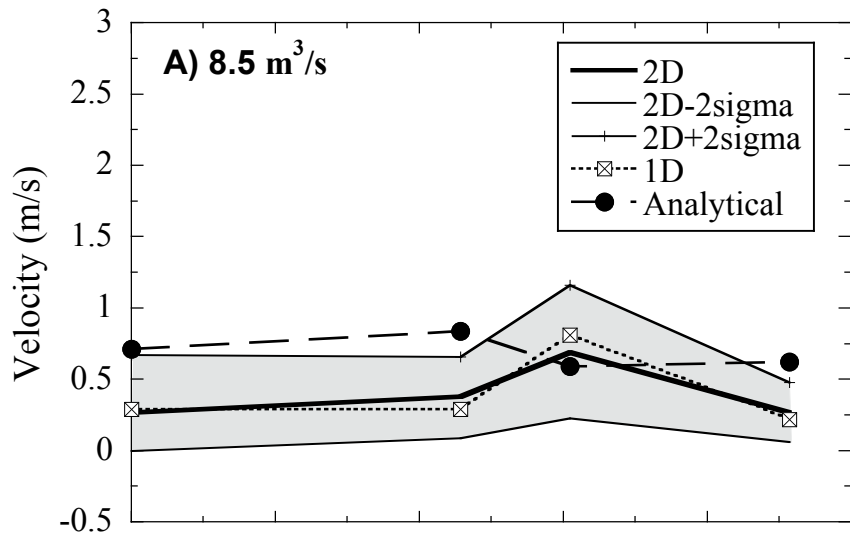




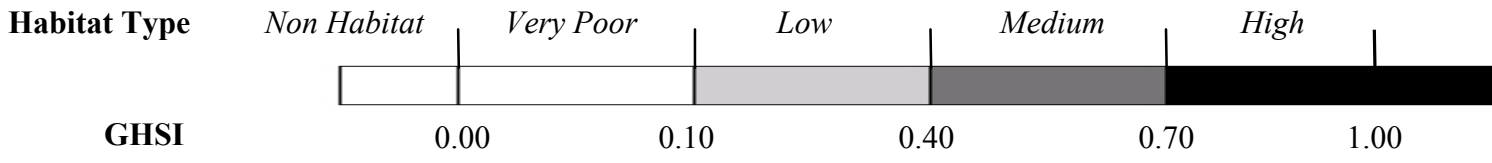
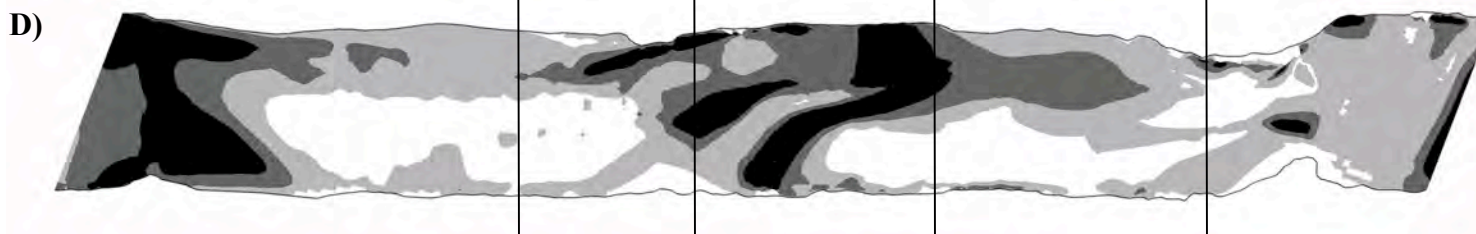
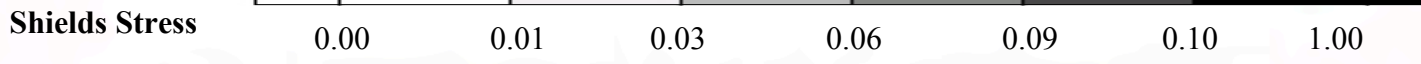
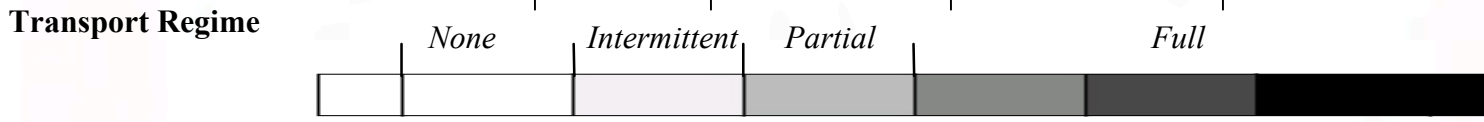
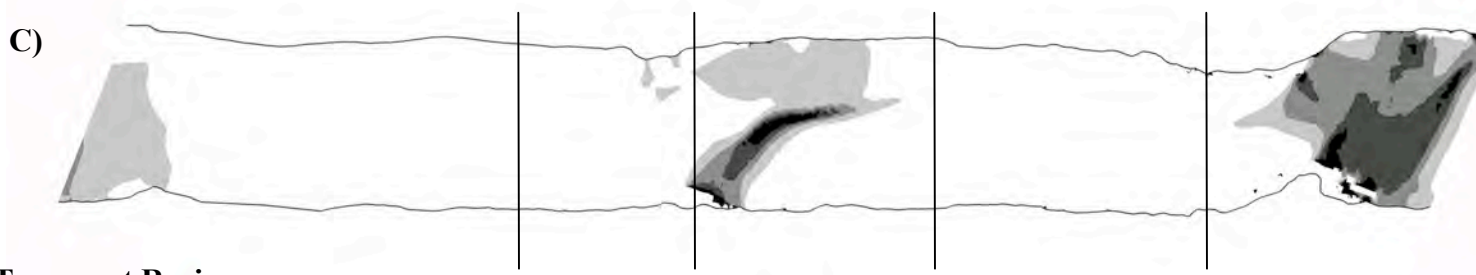
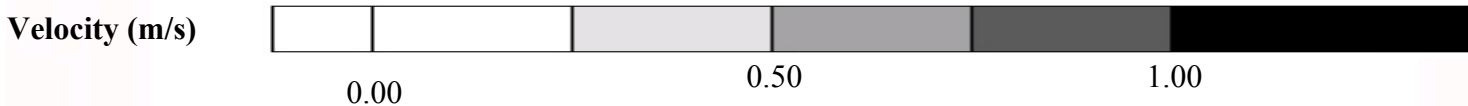
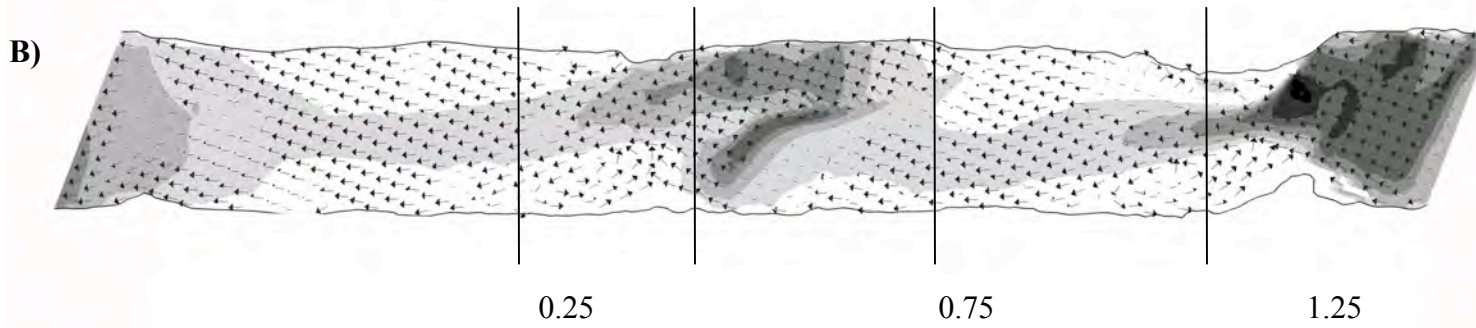
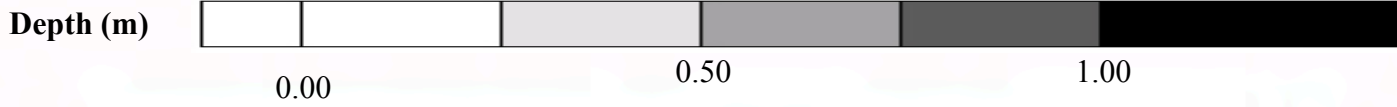
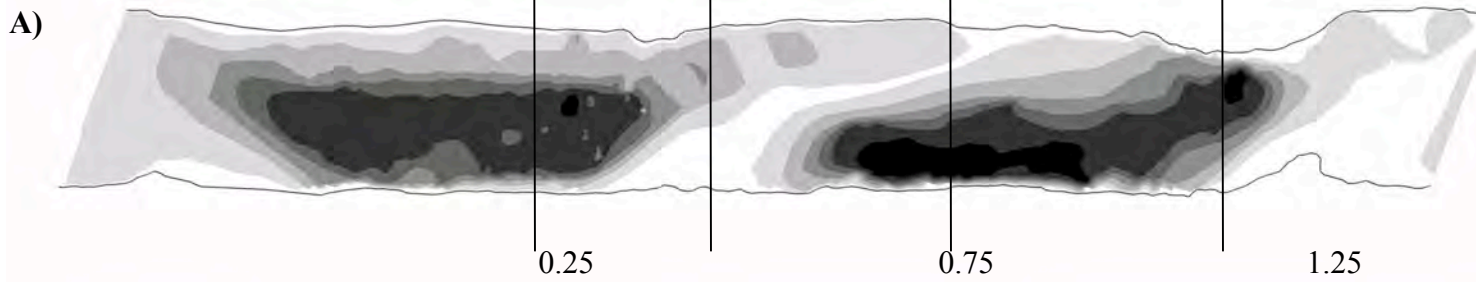








Pool 2 **TOR Exit** **TOR Entrance** **Pool 1**



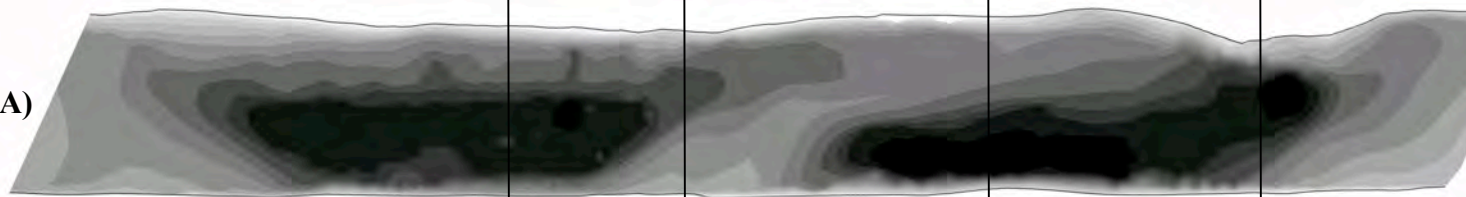
Pool 2

**TOR
Exit**

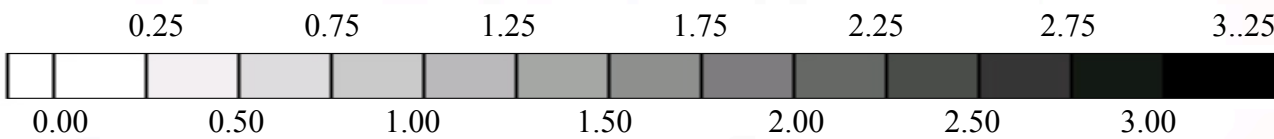
**TOR
Entrance**

Pool 1

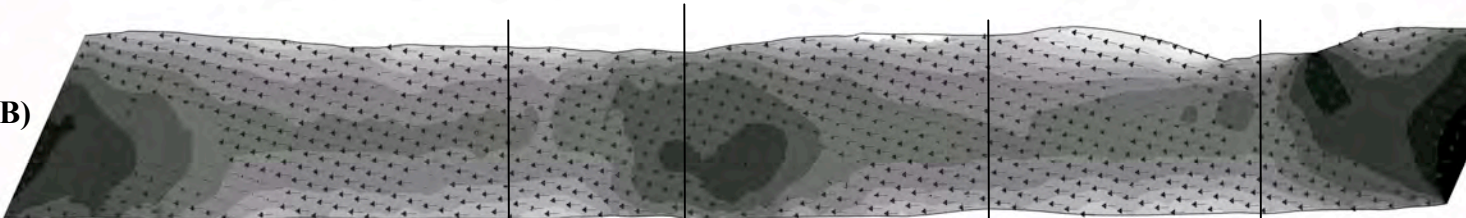
A)



Depth (m)



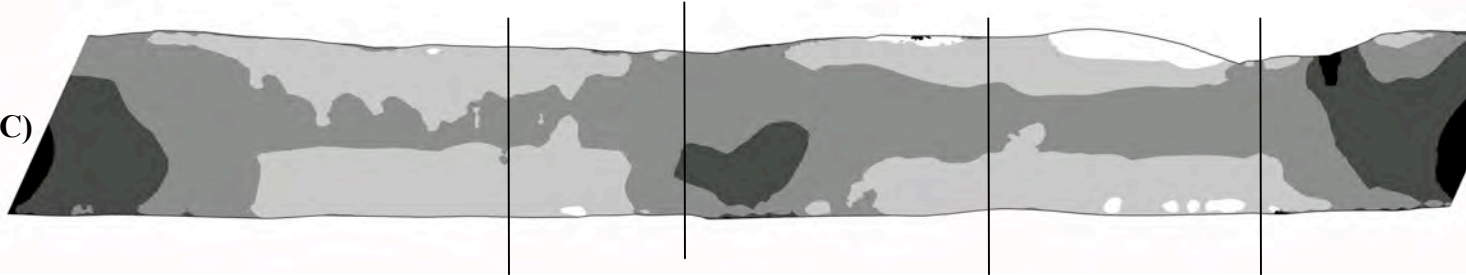
B)



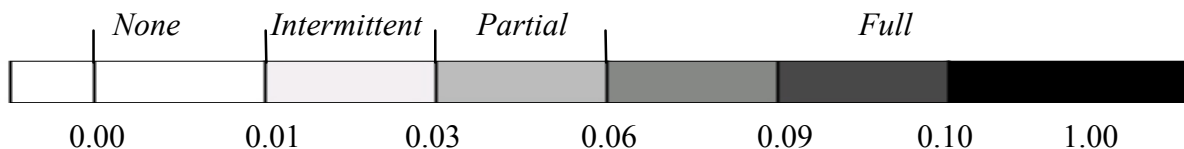
Velocity (m/s)



C)



Transport Regime



Shields Stress

