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Long-term integrated crop-livestock grazing stimulates soil ecosystem carbon flux, increasing subsoil carbon storage in California perennial agroecosystems

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2 **storage in California perennial agroecosystems**

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25 **ABSTRACT**

26 The strategic use of ruminant grazing in perennial cropland is steadily increasing throughout Mediterranean
27 perennial agroecosystems. Integrated sheep-vineyard (ISV) management, where small ruminant livestock graze on
28 understory vegetation, is viewed by some practitioners as a feasible transition opportunity to facilitate less
29 petrochemically intensive vineyard understory management. However, our knowledge of soil carbon dynamics
30 associated with grazing in perennial integrated crop-livestock (ICL) agroecosystems is notably limited, especially
31 within Mediterranean contexts. Here, we use a representative set of on-farm paired surveys to assess soil ecosystem
32 habitat and resource conditions related to SOC flux and storage in vineyards utilizing sheep-integration (ISV) and
33 conventional understory management techniques (CONV). Our results show that long-term grazing increased the
34 quantity of active, labile, and soluble carbon (C) within ISV soils, with much higher quantities of microbial biomass
35 carbon (MBC). Vineyard soils with sheep grazing also showed increases in phospholipid fatty acid (PLFA)
36 biomarkers, particularly amongst core functional groups related to decomposition. Soil microbial communities under
37 ISV had higher C mineralization rates as well as higher carbon use-efficiency, as indicated by less CO₂-C respired
38 relative to the size of the MBC pool. Whereas inorganic soil nitrogen (N) and phosphorous (P) were also higher
39 under ISV, microbial communities showed distinct metabolic investment strategies related to nutrient acquisition,
40 with lower P-cycling enzyme activity and higher N-cycling enzyme activity. Additionally, ISV resulted in an
41 increase in subsoil SOC storage, including higher quantities of physicochemical stabilization in the mineral-
42 associated organic carbon (MAOC) pool of the deepest measured subsoil layer (30–45 cm). We observed no
43 differences in soil structure indicators between treatments nor differences in the carbon fractions associated with four
44 distinct aggregate size categories. We propose a framework to explain observed shifts in SOC dynamics of perennial
45 ICL systems that include i) deposition of C and nutrient inputs with higher lability and solubility; ii) ruminant-
46 induced decoupling of C from N and P, resulting in increased nutrient bioavailability; and iii) altered soil microbial
47 metabolic strategies and more efficient biomass accumulation. These findings show strong potential of strategically
48 applied ICL grazing to enhance soil functioning and increase SOC storage in Mediterranean perennial
49 agroecosystems.

50

51 **Key words:** Soil organic carbon; Soil carbon storage; Grazer-plant-soil interactions; Microbial ecology; Integrated
52 crop-livestock; Perennial agriculture

53 1. INTRODUCTION

54 Agricultural resource conservation incentives provide new opportunities to explore underutilized farming methods
55 for their potential agronomic and environmental benefits. One such incentive is to utilize croplands for sequestering
56 additional soil carbon. Increasing the soil organic carbon (SOC) of global croplands has compounding GHG
57 mitigation benefits where it is facilitated by farming methods that rely less on mechanization and heavy use of
58 petrochemical inputs (Minasny et al., 2017). The production of pesticides, synthetic mineral fertilizers, and other
59 inputs produced using large quantities of petroleum and other fossil fuels are well understood to substantially
60 contribute to global GHG emissions (Walling and Vaneckhaute, 2020; Woods et al., 2010). One historically
61 foundational, yet scientifically understudied management strategy with proposed potential toward these coordinated
62 efforts is the integration of animals and crops within the same production system (Brewer and Gaudin, 2020; Garrett
63 et al., 2017; Russelle et al., 2007). Practitioners of integrated crop-livestock (ICL) management rely on extensively
64 developed local and indigenous knowledge systems, with regional specificity and complexity (Altieri, 1992; Altieri
65 et al., 2012; Garrett et al., 2020; Sekaran et al., 2021b). Whereas much of the world's crop and livestock components
66 are still managed in coordination, with animals providing diverse grazing-based services for crop production, the
67 intensification of U.S. agriculture has resulted in highly specialized and de-coupled farm-scale crop and livestock
68 components (Baur and Iles, 2022; Entz et al., 2005; Garrett et al., 2020; Sanderson et al., 2013).

69
70 While more rigorous inquiry is necessary, crop-livestock de-coupling is currently understood to contribute to poor
71 nutrient cycling within and between agricultural operations and an increased environmental footprint of both crop
72 and livestock production (Garrett et al., 2017; Lemaire et al., 2014). As such, the re-integration of crop and livestock
73 system components has been proposed as a strategy toward improving the environmental conservation and resource
74 efficiency outcomes of agricultural landscapes (Russelle et al., 2007). Perennial ICL management, where ruminant
75 livestock forage on understory plant communities during prolonged periods of vegetative growth, provide
76 opportunity as a feasible agroecological alternative to current petrochemically-intensive practices for management
77 of understory plant communities (such as mowing, herbicides, and tillage). Current and developing models of ICL

78 systems employ diverse, adaptable, and feasible management practices that can be strategically implemented across
79 various scales and crop production systems (Bell et al., 2014; Garrett et al., 2020, 2017; Lemaire et al., 2014).
80 However, the potential of this re-integration, with respect to the comparative benefits and/or trade-offs of this
81 agroecological approach relative to conventional practices, is highly underexplored (Garrett et al., 2017), especially
82 within highly industrialized agricultural contexts (i.e. modern, intensive methods of farming crops and animals for
83 mass production). California, with its extremely diverse agricultural landscapes, provides a unique opportunity for
84 scientific exploration into the potential of this re-integration.

85
86 A core proposition of agroecological farming models is to increase the utilization efficiency of externally-applied
87 resources, through retention and (re)cycling processes that are endogenous to the agroecosystem's design and,
88 therefore, *internally regulated* (Altieri et al., 2015; Garcia-Franco et al., 2018; Lipper et al., 2014; Lovell et al.,
89 2010; Snapp, 2017; Wagg et al., 2020). This internal regulation of agroecosystems – defined as its capacity to tightly
90 couple energy and nutrient (re)cycling (King and Hofmockel, 2017; Prommer et al., 2020; Tamburini et al., 2020) –
91 depends upon the dynamic functioning of many complex energy and nutrient transformation processes (Lal, 2016;
92 Power, 2010; Snapp, 2017; Tamburini et al., 2020; Xu et al., 2020). Soils play an essential role in agroecosystem
93 internal regulation – most notably related to the flux of soil carbon (energy) and nutrients, as well as the eventual
94 formation and stabilization of soil organic carbon (SOC).

95
96 While not yet explored under integrated crop-livestock (ICL) system contexts, the strategic integration of diverse
97 plant and animal communities across multiple spatial (e.g., field, farm, and landscape) and temporal (e.g., inter- and
98 intra-seasonal) scales have otherwise been shown to promote resource use-efficiency through internal regulation
99 pathways (Altieri et al., 2015; Griesser et al., 2022; Lange et al., 2015; Prommer et al., 2020; Snapp, 2017; Wagg et
100 al., 2020). This may occur through increasing synchronicity in the utilization of carbon (C) and applied nutrients
101 within soil ecosystems (Griesser et al., 2022; Lange et al., 2015; Prommer et al., 2020; Snapp, 2017; Wagg et al.,
102 2020). Some key underlying mechanisms relate to the facilitation of interconnected ecological interactions that
103 include, but are not limited to, niche partitioning and niche complementarity (Altieri et al., 2019; Garland et al.,
104 2020; Petersen-Rockney et al., 2021; Ponisio et al., 2015; Snapp, 2017).

105

106 The impacts of grazing will affect multiple ecosystem processes related to nutrient utilization and the pathways
107 regulating SOC flux and storage (Figure 1) (Brewer and Gaudin, 2020; de Faccio Carvalho et al., 2010; Jarvis, 2009;
108 Lemaire et al., 2014; Rumpel et al., 2015). Specifically, animal re-integration into cropland can alter carbon and
109 nutrient flows *directly* through (1) transformation of aboveground residues into soluble, nutrient-rich, and labile
110 dung and urine, where carbon and nutrients are more stoichiometrically decoupled (Jarvis, 2009; Jung and Allen,
111 1995; Rumpel et al., 2015); (2) biomass removal that triggers shifts in forage productivity and the reallocation of
112 resources above- and belowground (Dawson et al., 2009); and (3) residue deposition and incorporation due to the
113 trampling effect and hoof action of animal traffic (Acosta-Martínez et al., 2004; Greenwood and McKenzie, 2001;
114 Wei et al., 2021) (Figure 1). It may also alter carbon and nutrient flows *indirectly* via (4) inter- and intra-seasonal
115 shifts in plant community composition (Chen et al., 2018) and (5) changes in soil structure, which alter transport and
116 spatial distribution of soil carbon and nutrients as well as their physical protection from continual degradation
117 through occlusion within aggregates (Erktan et al., 2020; Lavalley et al., 2020; Six et al., 2000).

118

119 These modifications to agroecosystem carbon and nutrient inputs have substantial impacts on the size and
120 composition of resource pools and organisms present within cropland soil ecosystems. This may have consequences
121 for microbial ecological processes such as community assembly, substrate utilization and use-efficiency, and
122 microbial energy investment strategies related to resource acquisition, stress responses, and growth rate optimization
123 (Malik et al., 2020). Since the partitioning of SOC into different biochemical and physical pools is mediated by the
124 quantity, quality, and spatial distribution of substrates entering the soil ecosystem (Lavalley et al., 2020; Rasse et al.,
125 2005; Schmidt et al., 2011; Sokol et al., 2019; Sokol and Bradford, 2019), it is likely that the plant-grazer-soil
126 interactions associated with ICL adoption have significant implications for cropland carbon storage dynamics.
127 Ultimately, this partitioning will largely regulate the fate of carbon within soil ecosystems – particularly with respect
128 to how it will be utilized by soil microbes and if it will persist as long-term sequestered C (Cotrufo et al., 2013;
129 Liang et al., 2019; Schmidt et al., 2011; Zhu et al., 2020).

130

131 The strategic design of ICL systems for improving SOC storage and other conservation outcomes will require
132 improved understanding and consideration of carbon and nutrient flows through the agroecosystem (Brewer and
133 Gaudin, 2020) (Figure 1). Microbial communities largely drive these biogeochemical flows within soils, as they rely
134 on energy from soil carbon decay channels to power nutrient cycling processes (Janzen, 2006; Kopittke et al., 2022;
135 Zhu et al., 2020). Increasing the flux rate of carbon (energy) through soil ecosystems is therefore necessary to
136 increase rates of ecological functioning. Since microbial utilization of SOC for biomass growth (anabolism) is
137 generally associated with respiration and, therefore, CO₂ efflux (catabolism), it has been argued that storing soil
138 carbon is inherently in tension with increasing microbial functioning (Janzen, 2006). However, the accumulation of
139 stable and long-term stabilized mineral-associated organic carbon (MAOC) has increasingly been shown to
140 necessitate the formation of microbial necromass, and is therefore dependent upon the continual pulse and turnover
141 of labile and accessible carbon through the microbial food web (Dynarski et al., 2020; Lehmann and Kleber, 2015;
142 Six et al., 2004, 2002) – a concept known as the microbial carbon pump (MCP). In fact, recent studies have shown
143 that the MCP and its microbially-derived anabolic compounds are the predominate source of stabilized MAOC
144 (Basile-Doelsch et al., 2020; Kallenbach et al., 2016; Liang et al., 2019). Stabilized soil carbon may therefore be
145 viewed predominately as a reservoir of previously processed microbial products, with stored chemical energy which
146 may be accessed later by soil microbial communities when the habitat and resource conditions of the soil ecosystem
147 are altered (Erktan et al., 2020) – a process understood as soil carbon priming (Kuzyakov, 2010). As such, the
148 evaluation and interpretation of soil carbon storage as static and relatively inert *stocks* must be complimented by an
149 understanding of soil carbon *flows* – as flux processes and an energy source for biological functions. These stocks
150 and flows are, of course, related with the production of microbial biomass and accumulation of microbial necromass
151 being critical to building long-term SOC storage.

152
153 The goal of this study was to evaluate the soil carbon flux and storage dynamics of perennial integrated crop-
154 livestock systems within the context of working landscapes, with farmer implementation by early adopters who have
155 integrated grazing for multiple years. The use of precision grazing in perennial cropland is steadily increasing
156 throughout California and beyond, particularly within Mediterranean integrated sheep-vineyard (ISV) systems
157 (Ryschawy et al., 2021). However, mechanistic understandings of the legacy effects associated with grazing on

158 nutrient cycling and SOC flux and storage within perennial cropland are lacking. This is especially true in semi-arid
159 regions, which have low precipitation and high temperature features that regulate carbon flows and limit storage
160 pathways (Brewer and Gaudin, 2020; Garcia-Franco et al., 2018; Hoyle et al., 2016, 2013). Conducting on-farm
161 studies across a representative sample of early-adopter ISV practitioners facilitates the important endeavor of
162 evaluating the longer-term impacts of perennial cropland grazing practices within working landscapes (Garrett et al.,
163 2017; Ryschawy et al., 2021) and deepens our understanding of the SOC storage potential associated with perennial
164 crop-livestock re-integration.

165
166 We established an on-farm survey study across a spectrum of ISV early-adoption systems to evaluate the longer-
167 term impacts of perennial cropland grazing practices on SOC fluxes and measurable benefits and/or trade-offs
168 related to soil carbon storage. We explored the hypotheses that sheep grazing of winter soil covers will *1) increase*
169 *the quantity of carbon most readily available for processing by the soil food web*; and *2) increase the flux rate of*
170 *soil organic carbon turnover*; which should *3) increase soil organic carbon storage dynamics, especially the*
171 *fraction stored stably clay mineral surfaces as mineral-associated organic carbon (MAOC)*. The objectives of this
172 study were to better understand cropland grazing impacts on SOC and biogeochemical processes for the purpose of
173 developing best management practices (Garrett et al., 2017; Niles et al., 2018), informing adoption through
174 identification of potential benefits and tradeoffs, and improving our understanding of the climate change mitigation
175 and soil C sequestration potential of perennial cropland grazing (Brewer and Gaudin, 2020).

176

177 **2. MATERIALS AND METHODS**

178 *Ethics statement*

179 Permission for site access was previously granted by landowners. All sites were privately owned and no permits
180 were required.

181

182 *2.1 Study region and management characteristics*

183 A soil survey of paired vineyard sites was conducted in the Northern California coastal foothills at three locations in
184 2018 and one location in 2021, between the months of January and March (Figure 2). The paired sites sampled in

185 2021 were added to strengthen and validate the initial findings from the 2018 sampling. Paired vineyard sites at each
186 sampling location consisted of one ‘*non-integrated*’ vineyard (interrow vegetation managed through mowing;
187 CONV) and one adjacent ‘*integrated*’ vineyard (interrow vegetation managed through grazing for 10+ years; ISV) –
188 with one location in Sonoma County (home to Wappo and Patwin native peoples), two in Lake County (home to
189 Pomo, Lake Miwok, and Patwin native peoples), and one in southern Mendocino County (home to Pomo and Yuki
190 native peoples) (8 paired vineyards across 4 locations) (Figure 2). This Mediterranean climate is classified as a semi-
191 arid Köppen-type *Csc* (Beck et al., 2018). It is characterized by mild cool winters, warm and dry summers, and
192 seasonal mean annual precipitation that is lower than the regional evapotranspiration (ET) potential. The annual
193 regional precipitation for sites in 2017 and 2020 was 739 mm and 368 mm, respectively. The mean maximum and
194 minimum temperatures were 21.9°C and 5.7°C in 2017 and 25.8°C and 5.3°C in 2020, respectively. This
195 contributed to an annual potential evapotranspiration (ET_o) of 1145 mm (2017) and 1406 mm (2020) and, therefore,
196 a 1.5x (2017) and 3.8x (2020) higher potential water demand than the regional precipitation supply. As was the case
197 for all vineyard sites within this study, the vast majority of Northern California regional vineyard systems utilize
198 micro-irrigation, especially surface drip systems (Tindula et al., 2013), to match vine ET demand during the warm,
199 dry vine growing season (Prichard, 2000). This growing season generally begins in March (bud break) and goes
200 throughout August to October (harvest), depending on the winegrape varietal. Irrigation was not applied to the
201 interrow space, where ISV grazing predominately occurs (Niles et al., 2018; Ryschawy et al., 2021). The understory
202 vegetation growing season is instead limited to periods of sustained precipitation, which typically occurs between
203 November and April, during which 91% of all regional rainfall has occurred over the last 20 years
204 (<http://www.cimis.water.ca.gov/>). This is also the period in which the vast majority of sheep grazing in regional
205 vineyards occurs, including the ISV sites utilized within this study.

206

207 Most typically, sheep-vineyard grazing within this region is used as an understory plant growth termination
208 methodology, similarly to the application of mowing, and is most often implemented immediately before vine bud
209 break (Niles et al., 2018; Ryschawy et al., 2021). Though less common, sheep grazing sometimes occurs multiple
210 times across the understory growing (vine dormancy) and vine growing seasons, where it is strategically applied to
211 achieve additional management benefits such as vine leaf thinning and removal of suckering trunks (Niles et al.,

212 2018; Ryschawy et al., 2021). The grazing strategies on all four of the integrated vineyards in this study were
213 characterized as *high-density, short-duration rotational grazing* management (de Faccio Carvalho et al., 2010). This
214 rotational grazing strategy incorporate small paddocks that are grazed with high animal density and rotated
215 frequently amongst larger sections of the overall landscape. This strategy facilitates longer rest periods and
216 increased competition amongst grazing ruminants (Teague et al., 2008). This has been found to lower the duration of
217 grazing per unit of land area and reduce grazing selectivity and the spatial heterogeneity of grazing pressure (Teague
218 et al., 2008; Teague and Dowhower, 2003). Briefly described, temporary electrical fencing was erected to establish
219 1-acre sized grazing paddocks, where ~250 ewes were grazed for 1-2 days within each paddock before rotating to
220 the next temporary paddock. Grazing generally occurred once during vine dormancy. The timing of grazing events
221 varied with precipitation and understory plant growth rates, but generally occurred sometime between early March
222 through late April, before vine bud break and a coinciding decrease in regional precipitation rates. Each grazing
223 event aims to remove roughly 80% of understory biomass from the vineyard as a seasonal “termination” of the
224 vineyard understory plant community, which remained dormant throughout the warm, dry vineyard growing season.
225 Grazing sometimes occurred two or more times during the dormant season (November to April), when forage
226 productivity was substantially high. Before transitioning to sheep grazing of understory vegetation, sites were
227 managed by a mixture of mowing and herbicides, which generally occurred at the same time of year as grazing. The
228 undervine row of all four mowed (CONV) vineyards and one grazed vineyard (ISV; Site 3) were managed using
229 synthetic herbicide applications (Table 1). Both the ISV and CONV vineyards at Site 1 used conservation tillage,
230 with shallow (<10 cm depth) tillage of every other row in alternating years (Table 1).

231

232 ***2.2 Site selection and participatory engagement***

233 Study sites were first selected based on identifying early adopters of ISV management – vineyards that had a long-
234 term legacy of grazing sheep – and selection of participating vineyards was confined to the Northern California
235 coastal foothill region to reduce climate and soil variability (Table 1). While perceptions of ISV management are
236 increasingly favorable among adopters and non-adopters alike (Ryschawy et al., 2021), vineyard grazing is still
237 considered a niche production system compared to the dominant technological regimes (Garrett et al., 2020;
238 Ryschawy et al., 2021) and early adopters of ISV practices in California are rare. Producers utilizing ISV

239 management were identified using directories from the *LandSmart* collaborative (<http://landsmart.org>) and grower
240 networks from the Community Alliance with Family Farmers (CAFF). Four integrated vineyards (ISV) were
241 selected based on grower knowledge of long-term sheep grazing and co-management legacy. These ISV growers
242 expressed interest in participating and worked with the study's authors to identify the adjacent non-integrated
243 vineyards (CONV) for comparison. A participatory survey was conducted to collect management information for
244 each vineyard (vine/rootstock varietal, understory vegetation management, external amendments, irrigation, and
245 tillage) to assess and minimize variability between paired sites (Table 1). Soil type and topography data was
246 collected from the USDA-NRCS *SoilWeb* app (<https://casoilresource.lawr.ucdavis.edu>). Regional soils are
247 Inceptisols (USDA-NRCS), with textures ranging from clay loam to loam and an average clay content of $27\% \pm 2\%$
248 (Table 1). At the time of the study, two of the ISV vineyards (Site 1 & 4) were organic certified through the
249 California Certified Organic Farmers (CCOF) program.

250
251 Sampling plots (2 ha area) within each paired integrated (ISV) and non-integrated (CONV) vineyard site were
252 selected to maximize the edaphic and co-management similarity of paired sites and isolate the impact of sheep
253 grazing. Wine vineyards provide a unique opportunity for soil surveying, in that tightly controlling management
254 variability, especially related to water and soil fertility, is essential for improving wine grape quality. Vineyard
255 growers strategically limit irrigation and N uptake at certain stages of vine phenology to control vegetative growth
256 and mitigate various perceived tradeoffs between vine vigor and wine grape quality (Spayd et al., 1994; Wheeler and
257 Pickering, 2003; White et al., 2007). As such, excessive water and N availability is generally avoided by reducing
258 inputs (Gaiotti et al., 2017; Lazcano et al., 2020). When inputs are utilized they are most often applied in small
259 doses, timed only when vine demand is highest, and delivered directly under the vine (Peter Christensen et al., 1994;
260 Spayd et al., 1994). Consequently, while cover crops are increasingly utilized to prevent soil erosion and stabilize
261 soil quality (Novara et al., 2019; Rodrigo-Comino, 2018), wine vineyards are otherwise often low-input
262 agroecosystems, especially within the vineyard interrow, and thus have less co-management variability than other
263 perennial agroecosystems.

264

265 **2.3 Soil collection and processing**

266 Soil sampling occurred once per vineyard site and was timed before the seasonal understory forage termination
267 event(s) to maximize the soil acclimation period between the last graze or mow events. Soil samples were collected
268 from eight randomly selected points per 2 ha plot in a “W” pattern (Moebius-Clune et al., 2016). Sub-plots (1 m²)
269 were set-up at each sampling point, surface residues were removed, and three soil cores (5 cm diameter) were taken
270 at three depths (0–15 cm, 15–30 cm, and 30–45 cm) in the vineyard interrow. Samples were not taken from the
271 undervine row to minimize management variability due to seasonal irrigation, fertility, and herbicide applications.
272 Samples were weighed in the field, homogenized and composited for each sub-plot, and placed in a cooler for
273 transport. Bulk soils were processed promptly and stored at 4 °C until further analysis, except for ~75 g of soil that
274 was separated and stored at –80°C for PLFA and enzyme assays. Approximately 10 g of field moist soil was sieved
275 (2 mm) and oven dried (105°C) to a constant weight to determine soil gravimetric water content (GWC). Surface
276 soil (0–15 cm) bulk density (BD) was determined for each soil core using mass of oven-dried soil (105 °C, 24 h or
277 until consistent weight) and total volume of each soil core (Blake and Hartge, 1986). Another 250 g of soil was
278 subsampled for chemical analysis and ~100 g was used to determine texture and soil aggregate characteristics.

279

280 *2.4 Soil chemical properties*

281 A subsample of ~300 g was sent to a certified laboratory (Ward Laboratories – Kearney, NE) for analyses of soil
282 texture (sand:silt:clay) by hydrometer; pH (1:1 v/v method); soil salinity by electrical conductivity (EC; dS/m);
283 available P (mg kg⁻¹) via Olsen bicarbonate extraction; and cation exchange capacity (CEC) (Meq 100 g⁻¹) based on
284 ammonium acetate extraction and pH. All soil depth fractions were dried to constant mass, ball-milled, and weighed
285 for total elemental C and N using dry combustion (Costech ESC 4010 Elemental Analyzer – Valencia, CA, USA).
286 Soil NH₄⁺ and NO₃⁻ were extracted from 5 g of fresh soil with 20 ml 2 M KCl solution and measured using
287 colorimetric assays on a BioTek Synergy HTX (BioSPX B.V. – The Netherlands) microplate reader. Mineral
288 nitrogen is the sum of soil NH₄⁺ and NO₃⁻. Soil organic carbon (SOC) and nitrogen (SON) were measured by
289 subtracting HCl inorganic C extraction measurements and inorganic N pools (NO₃⁻ and NH₄⁺) from total elemental C
290 and N analyses described above.

291

292 *2.5 Soil aggregate size distribution and aggregate-associated carbon*

293 Aggregate size categorization was performed on air-dried soils by wet sieving to separate water-stable aggregates
294 into four size categories: large macro-aggregates (2000 μm), small macro-aggregates (250-2000 μm), micro-
295 aggregates (530-250 μm), and the silt and clay fraction (<53 μm) (Cambardella and Elliott, 1993; Kemper and
296 Rosenau, 1986). Each soil was submerged in deionized water for 10 minutes before wet-sieving, and a sub-sample
297 was taken to assess soil gravimetric water content (g g^{-1}) after saturation. A 40 g sub-sample of saturated soil was
298 then transferred to a vibratory sieving tower with rainfall simulator (Fritsch Analysette 3 Pro – Idar-Oberstein,
299 Germany), with vibration amplitude set at 0.1 μm and frequency at 50 Hz. Sieving lasted until the deionized water
300 used to wash soils on the sieve was flowing clear, which was generally around 60 seconds. The remaining fractions
301 on each sieve (2000, 250, and 53 μm), as well as the soil-water suspension passed through the 53 μm sieve (<53
302 μm), were dried at 60° C to dry until reaching constant weight. The mass recovery threshold was set between 0.98 –
303 1.02% and was calculated as follows (1):

304

305 (1) $Mass\ recovery(\%) = \frac{bulk\ soil(g) - \sum A_i}{bulk\ soil(g)} \cdot 100$

306

307 where bulk soil (g) is the mass of the soil used for wet-sieving of each sample; GWC (g g^{-1}) is the gravimetric water
308 content of the soil used for wet-sieving; and A_i is the oven-dry weight (g) of each aggregate fraction. When samples
309 did not meet the mass recovery threshold, the samples were repeated. Mean weight diameter (MWD), a weighted-
310 average index of aggregate stability (van Bavel, 1950), was calculated as follows (2):

311

312 (2) $MWD = \sum_{i=1}^4 X_i \cdot A_i$

313

314 where X_i is the average diameter (μm) for particles of each i -level aggregate fraction and A_i is the weight percentage
315 of the fraction in the bulk soil. SOC content was measured for each aggregate fraction using combustion analysis
316 (Costech ESC 4010 Elemental Analyzer – Valencia, CA, USA). The proportional concentration of SOC in each
317 aggregate fraction was calculated as follows (3):

318

319
$$(3) M_i = \frac{A_i \cdot SOC_i}{\sum_{i=1}^4 A_i \cdot SOC_i}$$

320

321

322 where M_i is the relative SOC concentration of each i -level aggregate fraction (%); A_i is the oven-dry weight (g) of

323 each aggregate fraction; and SOC_i is the relative SOC concentration of each i -level aggregate fraction. The SOC

324 stock for each aggregate fraction was calculated as follows (4):

325

326
$$(4) M_i = C_i \cdot SOC_i \cdot BD \cdot H \cdot 10^{-1}$$

327

328 where M_i is the SOC stock of each i -level aggregate fraction ($t \text{ hm}^{-2}$); C_i and SOC_i are the relative fraction and SOC

329 concentration each of i -level aggregate fraction, respectively; BD is the soil bulk density (g cm^{-3}) and H is the

330 thickness of soil layer, which was 15 cm for this measurement.

331

332 **2.6 Soil organic carbon size fractionation**

333 Soil organic carbon was separated into particulate organic carbon (POC), and mineral-associated organic carbon

334 (MAOC) using aggregate dispersion, wet sieving, and particle-size fractionation method (Six et al., 1998). In short,

335 20 g of air-dried soil was dispersed with 100 ml of 5% (w/v) sodium hexametaphosphate ($\text{Na}_6(\text{PO}_3)_6$) and an 18-hour

336 rotary shaking for sufficient dispersion. Dispersed soils were washed through a 53 μm sieve on a vibratory sieve

337 shaker (Fritsch Analysette 3 Pro – Idar-Oberstein, Germany) as described in Section 2.5. The fraction retained on

338 the sieve was considered as POC, while the finer fraction that passed through the sieve was considered as MAOC.

339 Both POC and MAOC fractions were dried at 60° C until reaching constant weight, then ground, and analyzed for

340 total C on an elemental analyzer (Costech ESC 4010 Elemental Analyzer – Valencia, CA, USA). The final content

341 of POC or MAOC in bulk soil was calculated based on the recovered mass. For example, POC (mg C g^{-1} bulk soil)

342 was calculated with two formulas as follows (5) and (6):

343

344
$$(5) \text{ Mass recovery } (\%) = [(POC(g) + MAOC(g)) / \text{bulk soil}(g)] \times 100$$

345

346
$$(6) POC(mg C g^{-1} \text{ bulk soil}) = \frac{POC(g) \cdot POC(mg C g^{-1})}{\text{bulk soil}(g) \cdot (\text{mass recovery} \cdot 0.01)}$$

347

348 where bulk soil (g) is the mass of the soil used for wet-sieving of each sample; POC and MAOC (g) are the masses
349 of the POC and MAOC fractions recovered after the wet-sieving, respectively; and POC (mg C g⁻¹) is the C
350 concentration measured in the POC fraction.

351

352 **2.7 Soil microbial biomass and dissolved organic carbon**

353 Soil microbial biomass carbon (MBC) was measured using the fumigation-extraction method (Horwath and Paul,
354 1994). Fresh soil was sieved to 4 mm and two replicates of 6 g were weighed into glass vials. One replicate was
355 fumigated for 24 h with chloroform (CHCl₃) and the other sample (unfumigated) was immediately extracted using
356 30 ml 0.5 M K₂SO₄. After chloroform fumigation, the fumigated sample was also extracted using 30 ml 0.5 M
357 K₂SO₄. The extracted solutions were filtered with Q5 filter paper and then analyzed for TOC and TN (Elementar
358 TOC/TNb – Ronkonkoma, NY, USA). Microbial biomass C was calculated as the difference in dissolved C
359 concentration between fumigated and unfumigated samples, using a K_e conversion factor of 0.35. The dissolved C
360 content in the unfumigated samples represent the dissolved organic C (DOC) fraction (Jones and Willett, 2006). The
361 microbial quotient (Q_{mic}) represents the ratio of MBC relative to SOC (ug MBC/ug SOC) (Sun et al., 2020).

362

363 **2.8 Carbon mineralization**

364 A 35-day (Haney et al., 2008) incubation was conducted to determine C mineralization over time, as well as
365 potentially mineralizable C (PMC; the flush of CO₂ during a 72-hour incubation (Wade et al., 2018)) using rewetted
366 air-dried soils. Three technical replicates were run per sample to account for potential methodological variability
367 (Wade et al., 2018). Briefly, 15g of air-dried soil (0-15 cm depth) was sieved (4 mm) and weighed into 50-ml glass
368 beakers. Each sample was rewetted from above to 50% water-filled pore space and placed into 1-quart mason jars.
369 Each sample was placed inside a 0.4 L mason jar, capped with a metal lid and a rubber septum, and incubated at 25

370 °C for 35 days. Respired CO₂ was determined by sampling the headspace gas using a continuous-flow CO₂/H₂O gas
371 analyzer (LI850 – LI-COR Biosciences, Lincoln, NE, USA) at 1, 3, 7, 14, 21, 28, and 35 days. Jars were opened
372 every 7 days to equilibrate with the atmosphere and allow replenishment of oxygen, but otherwise remained sealed
373 with no air flow. Soil water loss from evaporation was low. The soil evaporation was replaced on days 14 and 21 by
374 re-weighing soils and adding deionized water accordingly. Respiration for each sampling date was calculated as the
375 difference between a sample and a control, using the ideal gas law and adjusting for the total headspace. Net
376 respiration was calculated as the sum of the respiration measurements up to each sampling date. Soil C
377 mineralization was expressed both in terms of cumulative C mineralization (C_{min_{total}}; ug CO₂-C g soil⁻¹ day⁻¹)
378 (Grunwald et al., 2016) and relative to the SOC content of each sample (C_{min_{soc}}; mg CO₂-C g SOC⁻¹ day⁻¹) (Zhang
379 et al., 2007). The metabolic quotient (Q_{met}) represents the ratio of microbial respiration to microbial biomass and was
380 determined by dividing the 24-hour basal respiration (M_{resp}; ug CO₂-C g⁻¹ dry soil) by MBC (ug MBC g⁻¹ dry soil)
381 (Anderson and Domsch, 1993).

382

383 ***2.9 Microbial community structure and exocellular soil enzymes***

384 Soil microbial community structure was characterized using phospholipid fatty acid (PLFA) analysis (Ward
385 Laboratories – Kearney, NE) using a chloroform-methanol extraction and gas chromatograph with a 25 m Ultra 2
386 (5%-phenyl)-methylpolysiloxane column (Bossio and Scow, 1998). Community structure bioindicators for PLFA
387 were distinguished into bacterial groups including Gram-positive (Gram(+)) and Gram-negative (Gram(-)) bacteria,
388 actinomycetes, and fungal groups including saprophytic fungi and arbuscular mycorrhizal fungi (AMF). The
389 Gram(+)/Gram(-) and fungal/bacterial (F/B) ratios represent the relative distribution of Gram(+)-to-Gram(-)
390 bacterial biomass and fungal-to-bacterial biomass, respectively.

391

392 Exocellular enzyme potentials included BG: β-Glucosidase (glycoside hydrolysis), CB: β-D-cellubiosidase
393 (cellulose decomposition), LAP: L-aminopeptidase (peptide hydrolysis), and PHOS: Alkaline-Phosphatase
394 (phosphate hydrolysis) were measured using fluorescence microplate assays (Bell et al., 2013). Briefly, 2.75g of soil
395 was blended with 91 ml of 50 mM sodium acetate buffer and pH adjusted to the average pH of soil samples from a
396 given vineyard. The soil slurry was then mixed on a stir plate as 800 μl were transferred into a deep 96-well plates.

397 Substrate concentrations and incubation time were determined based on calibration tests to capture the maximum
398 potential enzyme activity. 600 μM of fluorescently labeled substrates were added for all enzymes assayed, except
399 LAP, where 400 μM were added. A 200 μL aliquot of substrate was pipetted into the sample and incubated for 3 h at
400 25°C. Standard curves were prepared for each sample using 4-methylumbelliferone or 7-amino-4-methylcoumarin
401 for LAP. After incubation, assays were centrifuged for 3 min at 1500 rpm and 250 μL of supernatant was pipetted
402 into black 96-well plates. Substrate fluorescence was measured on a BioTek Synergy HTX microplate reader
403 (BioSPX B.V. – The Netherlands) at wavelengths 365 nm (excitation) and 450 nm (emission). Urease (urea
404 hydrolysis) was measured using a standard colorimetric assay method (Kandeler and Gerber, 1988). The enzyme
405 activity was calculated based on the soil dry weight and incubation time (unit: $\text{nmol g}^{-1} \text{h}^{-1}$).

406

407 ***2.10 Statistical analysis and mixed model selection***

408 Statistical analyses were conducted using the statistical software R, version 3.6.3 (Team, 2021) Linear mixed-effect
409 regression models were used to measure univariate treatment effects across all vineyard pairs, as well as difference
410 between treatments within each vineyard pair. Models were fit using fixed effects for ‘treatment’ (ISV vs. CONV)
411 and ‘location’ as well as their interaction term (treatment(x)location) to estimate differing treatment effects on
412 response variables across locations with the *lmer* and *lmerTest* packages (Bates et al., 2015; Kuznetsova et al.,
413 2017). We accounted for our sampling design by using a nested ‘plot’ (vineyard sampling zone nested within
414 location) as a random effect, which yielded the lowest Akaike information criterion (AIC) score when
415 included. Depth was not included as a factor in the model. Instead, each depth was analyzed independently using the
416 model described above. Given the degree of multiple comparison testing associated with this univariate approach,
417 MANOVA was conducted for soil carbon response variables as a false discovery rate (FDR) controlling approach
418 using the *dplyr* package, in order to correct for random events that falsely appear significant as revealed by our
419 univariate assessments (Benjamini and Hochberg, 1995). Before conducting MANOVA, the *mvnrmtest* package
420 was used to conduct a Shapiro-Wilk test for multivariate normality and the absence of multicollinearity was checked
421 by conducting correlations among the response variables, which all measured $R^2 \leq 0.80$ and therefore presented no
422 concern. Our MANOVA revealed similar patterns in soil carbon response variables to our univariate approach.

423

424 We further tested factors associated with the climatic and edaphic differences (i.e., MAP, MAT, %clay) among
425 vineyards as covariates. These covariates were left out from final models as none of the environmental factors were
426 significant or strongly influenced our main model effects. Residuals were checked for normality and homogeneity of
427 variance. When variables were non-normally distributed or had unequal variance, data were log or square root
428 transformed prior to calculation of means and back-transformed for visualization. Fixed effects were investigated
429 with means comparisons and considered p -value < 0.001 (***) as highly significant, p -value < 0.01 (**) as
430 significant, and p -value < 0.05 (*) as marginally significant. Non-significant means comparisons with p -value < 0.10
431 were considered as a 'trend' (Hurlbert and Lombardi, 2009; Wasserstein et al., 2019). Tukey's pairwise
432 comparisons were used to assess differences between each treatment within each of the four study locations. Values
433 in tables and graphs are reported as comparisons within each site, whereas values reported in the results section are
434 averages across sites. Box plots were graphed using the *ggplot2* package in R. The horizontal line is the mean, and
435 upper and lower sectors are the first and third quartiles, respectively. Upper and lower 'whiskers' extend to the
436 highest or lowest value, respectively, within 1.5 times the inter-quartile range (the distance between the first and
437 third quartiles).

438

439 3. RESULTS

440 3.1 Soil physicochemical habitat

441 Management treatment (ISV vs. CONV) had a significant effect on several key soil physicochemical indicators
442 (Table 2). On average across all vineyards, the ISV treatment increased dissolved P (19.8 ± 1.7 vs. 11.5 ± 1.3 $\mu\text{g g}^{-1}$;
443 $p < 0.001$), Total N (TN; 2.1 ± 0.08 vs. 1.6 ± 0.06 g kg^{-1} ; $p = 0.013$), and salt content (EC; 0.21 ± 0.02 vs. 0.12 ± 0.006 dS
444 m^{-1} ; $p = 0.058$) in surface soils (0–15 cm) compared to CONV management. The increase in P was significant in all
445 four paired vineyards and EC values were significantly higher in three out of four of the paired vineyard surface
446 soils. While dissolved P, TN, and EC were not significantly different in the subsoil (15–30 and 30–45 cm), the
447 mineral N fraction ($\text{NH}_4^+ + \text{NO}_3^-$) was significantly higher under ISV management at the 30–45 cm depth (5.1 ± 0.4
448 vs. 3.3 ± 0.3 ; $p = 0.042$). While mineral N values were higher under ISV in the 0–15 cm (14.1 ± 1.2 vs. 8.6 ± 0.8 ;
449 $p = 0.183$) and 15–30 cm (6.3 ± 0.5 vs. 4.1 ± 0.3 ; $p = 0.211$) depths, these increases were non-significant. Grazing did not

450 significantly affect soil pH or CEC at any depth zone, although treatment effects for CEC varied between locations
451 (treatment(x)location $p=0.041$).

452

453 The physical characteristics of surface soils (0–15 cm) as indicated by compaction (BD; 1.32 ± 0.03 vs. 1.37 ± 0.03 g
454 cm^{-3} ; $p=0.634$; Table 2) and aggregate stability (MWD; 1.47 ± 0.13 vs. 1.44 ± 0.12 ; $p=0.953$; Supplementary Figure 1)
455 were not affected by management across any of the four locations. There was also no difference in the relative size
456 distribution of surface soil (0–15 cm) aggregates (Supplementary Figure 1) between ISV and CONV treatments for
457 *macroaggregates* (>2000 μm ; $21.2\pm 2.5\%$ vs. $19.9\pm 2.3\%$; $p=0.959$), *large microaggregates* (250–2000 μm ;
458 $32.6\pm 2.4\%$ vs. $36.4\pm 2.7\%$; $p=0.791$), *small microaggregates* (53–250 μm ; $22.5\pm 1.9\%$ vs. $20.2\pm 1.6\%$; $p=0.835$), or
459 the *silt and clay* fraction (<53 μm ; $23.7\pm 2.3\%$ vs. $23.5\pm 2.4\%$; $p=0.635$). Soil water content was also similar in both
460 treatments at all depth zones: 0–15 cm (0.21 ± 0.01 vs. 0.18 ± 0.1 g g^{-1} ; $p=0.798$), 15–30 cm (0.21 ± 0.01 vs. 0.19 ± 0.01
461 g g^{-1} ; $p=0.777$), and 30–45 cm (0.22 ± 0.01 vs. 0.23 ± 0.01 g g^{-1} ; $p=0.851$).

462

463 3.2 Soil microbial community structure and enzymatic activity

464 The surface soil (0–15 cm) abundance (ng g^{-1}) of total phospholipid fatty acid (PLFA) biomarkers, an indicator of
465 viable microbial biomass, significantly responded to long-term grazing (Table 3; $p<0.001$). Individual vineyard
466 pairs ranged from 24.6% to 64.9% higher total PLFA abundance under ISV (3543 ± 304) than CONV (2543 ± 234)
467 management, with significantly higher total PLFA abundance in three out of four paired vineyards. Structural
468 biomarkers showed higher abundance of both bacterial (1497 ± 149 vs. 956 ± 102 ; $p<0.001$) and fungal (385 ± 49 vs.
469 263 ± 33 ; $p=0.004$) groups in ISV vineyards across sites. Mole percent distribution (mol%) also indicated an increase
470 in bacteria abundance ($38.9\pm 1.8\%$ vs. $34.5\pm 1.6\%$; $p<0.001$), but no changes for fungi under long-term grazing
471 ($9.1\pm 1.0\%$ vs. $9.1\pm 1.2\%$; $p=0.991$). While total bacterial biomarkers increased under ISV in three out of four paired
472 vineyards, total fungal biomarkers treatment response did not differ across locations (Table 3), despite a significant
473 main treatment effect. One location (Site 4) showed very low total fungal biomarker abundance values in both the
474 ISV (7 ± 6 ng g^{-1}) and CONV (12 ± 9 ng g^{-1}) treatments. The relative abundance of bacterial and fungal groups was not
475 affected by grazing (F/B; 0.22 ± 0.02 vs. 0.22 ± 0.02 ; $p=0.887$).

476

477 The main treatment effect varied for specific fungal and bacterial functional groups, with higher abundances under
478 ISV for saprophytic fungi (261 ± 35 vs. 170 ± 23 ; $p=0.004$) and actinomycete (232 ± 21 vs. 141 ± 16 ; $p<0.001$), but no
479 significant difference between ISV and CONV for arbuscular mycorrhizal (AM) fungi (124 ± 15 vs. 91 ± 17 ; $p=0.544$).
480 There was also no significant main treatment effect between ISV and CONV for Gram(+)/Gram(-) ratio (1.49 ± 0.13
481 vs. 1.37 ± 0.10 ; $p=0.384$) and saturated/unsaturated fatty acid ratio (3.93 ± 0.79 vs. 3.85 ± 0.72 ; $p=0.909$), although the
482 saturated-to-unsaturated fatty acid ratio treatment(x)location interaction ($p=0.021$) indicated variability in the
483 treatment response between paired vineyards.

484
485 The surface soil (0-15 cm) exocellular enzymatic activity potentials (nmols g OD soil⁻¹ hour⁻¹; Figure 3) related to
486 nitrogen cycling were higher under ISV compared to CONV management for both L-aminopeptidase (peptide
487 hydrolysis) (30.9 ± 2.6 vs. 21.1 ± 2.1 ; $p<0.001$; Figure 3A) and Urease (urea hydrolysis) (26.6 ± 1.8 vs. 14.7 ± 1.2 ;
488 $p=0.001$; Figure 3B). Pairwise comparisons for Urease activity showed significant treatment effects at three out of
489 four paired vineyards, while L-aminopeptidase was higher in the ISV treatment at only one paired vineyard.
490 Phosphatase (phosphate hydrolysis) was lower under ISV management compared to CONV (79.2 ± 8.1 vs. 94.9 ± 8.8 ;
491 $p=0.017$; Figure 3C), with significant effects at only one site. There was no significant treatment effect for enzymes
492 related to carbon cycling – β -Glucosidase (glycoside hydrolysis) (54.1 ± 6.1 vs. 38.0 ± 4.8 ; $p=0.524$; Figure 3D) and β -
493 D-cellubiosidase (cellulose decomposition) (10.3 ± 1.2 vs. 9.8 ± 1.2 ; $p=0.639$; Figure 3E).

494

495 ***3.3 Soil carbon flux pools and metabolic activity indicators***

496 Soil carbon flux pools, indicative of labile and active soil carbon (Figure 1), were strongly impacted by animal
497 grazing. Although soil carbon flux pool values were generally highest in surface soils (0–15 cm) across both
498 treatments, ISV management was most significantly impactful in the subsoil (Figure 4). Microbial biomass C (MBC;
499 $\mu\text{g g}^{-1}$) was significantly higher under ISV management compared to CONV at all depth zones: 0–15 cm (454 ± 30 vs.
500 245 ± 22 ; $p=0.050$; Figure 4A), 15–30 cm (174 ± 15 vs. 94 ± 9 ; $p=0.008$; Figure 4B), and 30–45 cm (150 ± 12 vs. 70 ± 7 ;
501 $p<0.001$; Figure 4C). While there was no significant main treatment variation in dissolved organic C (DOC; $\mu\text{g g}^{-1}$)
502 or potentially mineralizable C (PMC; $\mu\text{g g}^{-1}$) at all depths, DOC contents trended higher in the 30–45 cm depth of

503 the ISV treatment (98 ± 5 vs. 76 ± 5 ; $p=0.118$; Figure 4F). PMC also trended higher at both the 15–30 cm (10.9 ± 0.8
504 vs. 6.7 ± 0.3 ; $p=0.064$) and 30–45 cm (5.5 ± 0.3 vs. 4.4 ± 0.3 ; $p=0.081$) depths (Figure 4H-I).

505

506 The soil microbial quotient (Q_{mic} ; MBC:SOC) was significantly higher in ISV surface soils (0–15 cm) compared to
507 CONV (0.018 ± 0.001 vs. 0.012 ± 0.001 ; $p=0.034$; Figure 5A). The Q_{mic} main treatment effect was non-significant in
508 both subsoil depths: 15–30 cm (0.016 ± 0.002 vs. 0.011 ± 0.001 ; $p=0.241$; Figure 5B) and 30–45 cm (0.024 ± 0.003 vs.
509 0.016 ± 0.002 ; $p=0.269$; Figure 5C). The soil metabolic quotient (Q_{met} ; M_{resp} :MBC) was also impacted by long-term
510 grazing similarly across sites. ISV management significantly lowered Q_{met} values compared to CONV in 0–15 cm
511 surface soils (0.023 ± 0.003 vs. 0.056 ± 0.019 ; $p=0.049$; Figure 5D). While Q_{met} values also trended lower in the 30–45
512 cm subsoil depth of ISV vineyards (0.021 ± 0.003 vs. 0.036 ± 0.005 ; $p=0.104$), Q_{met} was not significantly affected by
513 treatment in the 15–30 cm depth 0.038 ± 0.020 vs. 0.44 ± 0.006 ; $p=0.240$) (Figure 5E-F). Cumulative soil carbon
514 mineralization ($C_{min_{total}}$) rates, measured via respiration over a 35-day incubation period, showed similar trends
515 between ISV and CONV in surface soils (0–15 cm) with a cumulative rate of 1.28 ± 0.11 vs. 1.09 ± 0.12 $\mu\text{g CO}_2\text{-C g}$
516 $\text{soil}^{-1} \text{day}^{-1}$, respectively ($p=0.853$; Figure 6A). The $C_{min_{total}}$ rate was significantly higher under ISV in the 15–30 cm
517 subsoil depth (1.89 ± 0.29 vs. 1.46 ± 0.18 $\mu\text{g CO}_2\text{-C g soil}^{-1} \text{day}^{-1}$; $p=0.048$; Figure 6B), though not significantly
518 different at 30–45 cm (0.37 ± 0.04 vs. 0.33 ± 0.03 $\mu\text{g CO}_2\text{-C g soil}^{-1} \text{day}^{-1}$; $p=0.397$; Figure 6C).

519

520 **3.4 Soil carbon stabilization and storage pools**

521 Total SOC (g kg^{-1}) was highest in the surface soils (0–15 cm) and decreased in total quantity with increasing depth
522 in both ISV and CONV treatments, (Figure 7A-C). The main treatment effect on soil carbon storage pools was
523 generally most significantly impactful in the subsoil (15–30 and 30–45 cm) and we observed no significant
524 differences in total SOC content of surface soils (0–15 cm; 26.1 ± 1.2 vs. 21.4 ± 0.9 ; $p=0.197$; Figure 7A). Within
525 subsoils, SOC content trended higher in the 15–30 cm depth (12.4 ± 0.6 vs. 8.9 ± 0.5 ; $p=0.063$) and was significantly
526 higher at 30–45 cm (8.3 ± 0.6 vs. 6.2 ± 0.6 ; $p=0.003$) under ISV management (Figure 7B-C). Similarly, the main
527 treatment effect for MAOC was not significant between ISV and CONV treatments at both 0–15 cm (15.8 ± 0.6 vs.
528 13.3 ± 0.6 ; $p=0.168$; Figure 7D) and the 15–30 cm subsoil depth (11.0 ± 0.5 vs. 8.6 ± 0.5 ; $p=0.185$; Figure 7E), but did
529 MAOC significantly increased under ISV compared to CONV management in the 30–45 cm subsoil depth (8.1 ± 0.6

530 vs. 5.9 ± 0.5 ; $p < 0.001$; Figure 7F). The POC fraction showed opposite trends with a significant main treatment effect
531 at 0–15 cm and higher surface soil POC under CONV management (3.5 ± 0.3 vs. 6.2 ± 0.5 ; $p = 0.012$; Figure 7G), but
532 no significant effects in either the 15–30 (1.0 ± 0.1 vs. 1.0 ± 0.2 ; $p = 0.620$) or 30–45 cm (0.4 ± 0.04 vs. 0.5 ± 0.06 ;
533 $p = 0.472$) subsoil (Figure 7H-I).

534
535 There was no impact of long-term ISV grazing on absolute SOC values (g kg^{-1}) in surface soils (0–15 cm) for the
536 macroaggregate (> 2 mm; 6.7 ± 0.9 vs. 5.7 ± 0.8 ; $p = 0.964$; Figure 8A), large microaggregate (250–2000 μm ; 8.6 ± 0.7
537 vs. 8.6 ± 0.7 ; $p = 0.706$; Figure 8B), and silt and clay fractions (< 53 μm ; 5.4 ± 0.4 vs. 4.2 ± 0.3 ; $p = 0.219$; Figure D)
538 across sites. The ISV treatment did trend higher than CONV for the aggregate-associated C content within small
539 microaggregates (53–250 μm ; 5.1 ± 0.5 vs. 3.4 ± 0.3 ; $p = 0.107$; Figure 8C). Management treatment also did not shift
540 the relative distribution (% of total C; Supplemental Figure 2) of surface soil (0–15 cm) SOC across
541 *macroaggregates* (> 2000 μm ; $25.3 \pm 2.9\%$ vs. $21.8 \pm 2.2\%$; $p = 0.447$), *large microaggregates* (250–2000 μm ; $31 \pm 2.4\%$
542 vs. $39.7 \pm 2.1\%$; $p = 0.621$), *small microaggregates* (53–250 μm ; $20.9 \pm 2.1\%$ vs. $15.9 \pm 1.2\%$; $p = 0.149$), and the *silt and*
543 *clay* fraction (< 53 μm ; $22.6 \pm 2.3\%$ vs. $22.6 \pm 2.6\%$; $p = 0.871$).

544

545 4. DISCUSSION

546 Conducting on-farm studies with early-adopter ISV practitioners facilitates the important endeavor of evaluating the
547 impacts of perennial cropland grazing practices within the context of their on-the-ground application. In lieu of
548 developing extensive perennial ICL monitoring trials, survey studies provide our best opportunities to explore long-
549 term comparisons of these systems characteristics and potential. However, an important context consideration when
550 interpreting these results is the limitation in study sites and variation in characteristics between sites (Table 2). Soil
551 texture and clay content (%clay) – important parameters related to soil carbon cycling – were relatively similar
552 amongst sites. However, Sites 2 and 3 occurred on sloping hills (15-30% slopes) relative to the flat bottomlands at
553 Sites 1 and 4. The duration under current management practices was heterogeneous across sites, largely due to more
554 variation in the length of CONV managed vineyards. The ISV vineyards had relatively similar management
555 durations, providing comparable periods of acclimation to grazing across sites.

556

557 Climatic and management characteristics were qualitatively similar amongst sites. Most notably, the vineyard
558 interrow (where grazing and soil sampling both occurred) received no external water or nutrient applications at any
559 site. While similar co-management parameters between treatments (ISV vs. CONV) were tightly controlled for
560 within each site, some variation in co-management occurred between sites. Site 1 had the most variation from other
561 sites, with the lowest %clay and highest soil disturbance. Whereas most sites were managed without tillage, Site 1
562 received conservation tillage (i.e. shallow tillage of every other row in alternating years) and showed a much lower
563 distribution of >2000 um macroaggregates relative to other sites (Supplemental Figure 1). Further, while all sites
564 received planted cover crops, the seeding mixture composition slightly differed between years and sites. There is a
565 possibility that this co-management variation resulted in distinct and divergent outcomes between sites. However,
566 treatment(x)location interactions indicate notable similarities across sites, with soil exo-enzymes as the only
567 parameters showing significant treatment response variation between locations (Figure 3). These research
568 considerations are well understood and persistent challenges for developing ICL research platforms, especially given
569 the situational variability in on-the-ground cropland grazing applications (Tanaka et al., 2008).

570

571 Given this context, our study sought to evaluate the soil carbon flux dynamics and storage potential of long-term
572 perennial integrated crop-livestock management on working farms. We were particularly interested if, and to what
573 degree, ICL in perennial systems affects the partitioning of SOC into distinct biochemical and physical pools.
574 Further, we sought to understand whether long-term perennial ICL legacy effects impact microbial ecological
575 characteristics such as community structure, soil carbon utilization, and investment strategies related to biomass
576 accumulation, stress tolerance, and the production of soil exo-enzymes. We provide here supporting evidence that
577 small ruminant grazing can increase stable carbon storage within perennial cropland soils, especially when
578 accounting for the subsoil. Our findings also show that the continuous year-after-year use of perennial cropland
579 grazing altered soil carbon quality across four distinct vineyards, with a higher total quantity and greater relative
580 proportion of soil carbon allocated toward biologically active carbon flux pools – the carbon most readily available
581 and utilized by soil microbial communities. We argue that this stimulation in soil carbon flux is likely driven by
582 increased deposition of labile, soluble compounds (i.e. animal excreta and rhizodeposition) resulting from
583 successive years of high intensity rotational grazing events. This also presents a viable explanation as the dominant

584 driver of potential SOC accumulation in perennial ICL systems, due to enhanced production of microbial biomass,
585 accumulation of microbial necromass, and, therefore, increased stabilization as MAOC. These results indicate strong
586 potential of perennial ICL management to stimulate carbon (energy) flows and invigorate internal agroecosystem
587 processes, with significant relevance for climate change mitigation and adaptation goals in Mediterranean perennial
588 croplands.

589
590 Simultaneously,

591
592 Observing strong trends across sites and contexts is noteworthy, especially when considering the additional
593 variability tha

594

595 ***4.1 Perennial cropland grazing increased the active flux of soil carbon throughout the soil profile***

596 Our study shows that the introduction of sheep grazing increased the pool of actively fluxed soil carbon across four
597 distinct paired vineyards, with most significant impacts observed in subsoils. This was the case both in terms of the
598 total size of labile SOC pools and the relative proportion (% of total SOC) allocated toward labile carbon flux pools,
599 especially within the microbial biomass carbon (MBC) pool. This corroborates findings from other studies across
600 ICL systems, where positive impacts of cropland grazing on the size of soil microbial communities are commonly
601 reported (Acosta-Martínez et al., 2004; Bansal et al., 2022; da Silva et al., 2015; Franzluebbbers and Stuedemann,
602 2008; Sekaran et al., 2021a; Silva et al., 2022; Tracy and Zhang, 2008). The benefits of sheep grazing for soil
603 microbial growth were observed at all depths, with an average MBC increase of 82%, 65%, and 99% at the 0–15,
604 15–30, and 30–45 cm depths, respectively (Figure 4B-C). While the MBC pool comprised a notably higher relative
605 proportion of total SOC at the surface soil (0–15 cm) under ISV (+49% Q_{mic} ; Figure 5A), the MBC pool was
606 otherwise most significantly impacted by ISV management in the subsoil.

607

608 Increases in MBC may be attributed to shifts in the grazed plant community's composition, productivity, and the
609 allocation of energy and nutrients above- and belowground (Bardgett and Wardle, 2003; Cong et al., 2014; Dawson
610 et al., 2009; Rumpel et al., 2015; Tian et al., 2016). While cropland-specific impacts are less understood, research

611 throughout diverse grazed ecosystems show that feedbacks between grazing *intensity* (density and duration) and
612 *periodicity* (seasonality and frequency) exert unique selective pressure on plant communities and the rate and quality
613 of carbon influxes (Figure 1). Numerous studies have documented higher rates of rhizodeposition immediately
614 following high-intensity grazing events (Dawson et al., 2009; Gavrichkova et al., 2008; Hamilton et al., 2008;
615 Hamilton and Frank, 2001), which increases the availability of labile and soluble carbon substrates and facilitates
616 preferential and efficient utilization by soil microbial communities (Cheng and Kuzyakov, 2015; Gavrichkova et al.,
617 2008; Ota et al., 2013; Wilson et al., 2018). Observed increases in MBC under ISV management may also be
618 facilitated by the mineralization of aboveground plant residues within the ruminant of grazing animals – where
619 significant quantities of recalcitrant plant structural compounds such as cellulose and lignin are fragmented,
620 depolymerized, and returned to the soil as more labile, soluble, and nutrient-dense excreta (dung and urine) (Faissal
621 et al., 2017; Jarvis, 2009; Soussana and Lemaire, 2014). Ruminant mineralization and trampling of plant residues
622 also likely explain the lower quantities of POC found in the surface soils of the ISV treatment (Figure 7G).

623

624 The effects of ISV management on the availability and active microbial utilization of labile soil carbon were
625 generally most pronounced in the subsoil, which may result from increased leaching of soluble DOC compounds
626 deeper into the soil profile. This assumption is supported by our observations of increased trends in dissolved
627 organic carbon (DOC) content (Figure 4F) and potentially mineralizable carbon (PMC; Figure 4H-I) in subsurface
628 soil depths under ISV relative to CONV understory management, as well as other studies reporting higher DOC
629 concentrations under various ICL management systems (Sekaran et al., 2021a; Tian et al., 2010). Due to their
630 enhanced transport within soil solution, compounds in the DOC pool are generally more spatially accessible to
631 microbial processing (Erktan et al., 2020; Nakhavali et al., 2021; Neff and Asner, 2001; Ota et al., 2013; Rumpel
632 and Kögel-Knabner, 2011). The lower molecular weight and activation energy requirements of compounds in the
633 DOC pool also facilitate quick microbial assimilation and utilization than the complex and less nutrient-rich
634 structural compounds associated with the POC pool (Blagodatskaya et al., 2011; Kallenbach et al., 2016, 2015; Kok
635 et al., 2022; Lavalley et al., 2020; Shahbaz et al., 2017; Weiss et al., 1991). Soil carbon mineralization rates did not
636 vary between treatments at any depth relative to the SOC resource pool size of each vineyard ($C_{min,soc}$; Figure 6D-
637 F), indicating similarities in overall SOC quality regardless of differences in SOC pool size. Higher total soil carbon

638 flux rates in the 15–30 cm subsoil depth under ISV management (Figure 6B), measured via increased cumulative
639 rates of soil carbon mineralization ($C_{min_{total}}$), is likely related to overall larger labile soil carbon pool sizes under
640 grazed ISV compared to CONV vineyards. The larger MBC pool within the ISV treatment may also more
641 positioned to rapidly utilize labile and physically accessible soil carbon pools under ideal soil moisture and
642 temperature conditions relative to the CONV managed soils (Geyer et al., 2020). This may partially explain the
643 higher PMC and rates of $C_{min_{total}}$, despite a lack of observed differences in DOC content between treatments at both
644 0–15 and 15–30 cm soil depths.

645
646 Whereas nutrients from excreta may be more readily transformed and assimilated by soil microbes than ungrazed
647 plant residues (Kooch et al., 2020; Wang et al., 2018), ruminant mineralization of plant residues has also been
648 shown to increase the bioavailability of soil N and P for plant and microbial uptake as a result of high-intensity
649 grazing disturbance events (Costa et al., 2014; Tracy and Frank, 1998; Wu et al., 2011; Zhang et al., 2020). We
650 observed substantially higher quantities of extractable P at surface soil depths and soluble mineral N content (NH_4
651 and NO_3) in subsoils at all ISV sites (Table 2). The depth stratification of significant treatment differences between
652 these nutrient pools likely reflect differences in solubility of N and P and their physical transport pathways within
653 soil solution. Future research should explore the impacts of perennial ICL management on net primary productivity
654 (NPP) of the understory plant community. Positive feedback mechanisms between soil N and P bioavailability and
655 enhanced plant productivity are well established. However, it is unclear whether, and the degree to which, ISV
656 grazing ecophysiology and alterations in nutrient bioavailability (via animal-derived manure and urine deposition)
657 stimulate understory NPP and increase plant-derived carbon inputs relative to the ungrazed vineyards. In either case,
658 the increased bioavailability of soil N and P, coupled with intra-ruminal conversion of POC (*celluloses, hemi-*
659 *celluloses, lignin, etc.*) to DOC and altered C influx pulses from rhizodeposition, are all potential mechanisms
660 underlying observed increases in soil microbial biomass and their rates of carbon mineralization under ISV
661 management.

662
663 **4.2 Grazing shifted the resource investment strategy of surface soil microbial communities toward efficient**
664 ***biomass accumulation***

665 Our results indicate that perennial cropland grazing stimulated microbial growth efficiency and biomass
666 accumulation. A general framework to interpret this observation is with consideration of how soil habitat and
667 resource conditions orient microbial energy investment strategies (Ho et al., 2017). As soil ecosystem conditions
668 shift in response to agroecosystem management disturbances, this influences microbial metabolic processes related
669 to tolerance of stressors, the acquisition of resources, and capacity for growth. When less energy is required to
670 maintain functionality under limited resources and inhospitable conditions, soil microbial communities are more
671 likely to allocate energy toward biomass accumulation (Malik et al., 2020). Increased microbial biomass carbon
672 (MBC) in surface soils under ISV was paired with remarkably higher total quantities of PLFA biomarkers – a 47%
673 higher value relative to the CONV treatment (Table 3). We further observed 41% lower average metabolic quotient
674 (Q_{met}) values (Figure 5C) in surface soils of grazed vineyards, which indicates higher microbial carbon use-
675 efficiency (CUE) as less CO_2 -C is respired relative to the size of the MBC pool (Dilly and Munch, 1998; Sinsabaugh
676 et al., 2017). Increased availability of soil N and P relative to C have been shown to lower Q_{met} values across
677 climatic and soil management gradients (Xu et al., 2017). Increased microbial growth efficiency and investment in
678 biomass accumulation under ISV management is further supported by the substantially higher proportion of MBC
679 relative to total SOC (+49% Q_{mic}) within the 0–15 cm depth (Figure 5A). Higher Q_{mic} values are associated with a
680 greater potential for soil microbes to transform energy sources via increased availability of soil carbon and nutrients
681 (Sparling, 1992; Sun et al., 2020). Trends toward higher PMC values are also indicative of increased energy source
682 availability in the grazed vineyards, as this pool measures the reservoir of readily available soil carbon that drives
683 microbial functions (catabolism) and biomass accumulation (anabolism) (Levi-Minzi et al., 1990). These findings
684 corroborate another recent ICL study, which showed higher microbial biomass in grazed cropland that was similarly
685 attributed to increased availability of carbon and nutrient substrates ((Sekaran et al., 2021a)). Given the role of
686 microbial necromass in the formation of stable MAOC, these efficiency and growth indicators suggest a higher net
687 SOC storage potential under ISV management.

688
689 As another indicator of microbial investment, the production of metabolically-costly extra-cellular enzymes
690 represent shifts in resource acquisition strategies in response to growth factor limitations through altering energy and
691 nutrient availability within the near-cell soil environment (Nannipieri et al., 2002; Zheng et al., 2020). While the

692 availability of soil P was higher in surface soils of the ISV treatment (Table 2), we observed significantly lower
693 enzymatic activity related to P cycling (phosphatase; Figure 3C). The composition of sheep excreta has high
694 inorganic P content (Arnuti et al., 2020) and soil applications of inorganic P have been shown to reduce phosphatase
695 activity (Oshima et al., 1996), as this enzyme cleaves phosphate (PO_4) groups from proteins and becomes an
696 increasingly unnecessary investment under high P availability conditions. These observations indicate an increased
697 accessibility of soil P and reduced microbial investment in P acquisition (Nannipieri et al., 2011) in perennial
698 croplands with grazing. At the same time, we observed significantly higher N cycling enzymatic activity in the ISV
699 treatment as measured by aminopeptidase (peptide hydrolysis) and urease (urea hydrolysis) enzymes (Figure 3A-B).
700 The increase in urease activity corroborates previous findings across various grazed ecosystems (Acosta-Martínez et
701 al., 2010; McNaughton et al., 1997; Sekaran et al., 2021a) as a result of urea degradation, the dominant N
702 constituent found in urine (Bristow et al., 1992). The release of aminopeptidase enzymes could indicate potential
703 limitation in microbial N availability and an increased metabolic investment in N acquisition (Schimel and Bennett,
704 2004). However, both the mineral N (NO_3 and NH_4) pool and the higher total N content of ISV surface soils (0–15
705 cm) do not indicate reduced availability of soil N (Table 2). Alternatively, recent research has suggested that soil
706 microbial communities also use aminopeptidase enzymes as a means to access protein-derived carbon – a potent
707 energetic resource for cellular growth (Norman et al., 2020).

708
709 Shifts in microbial processes may reflect differences in the abundance of core functional groups such as those related
710 to decomposition (Bhatti et al., 2017; Setälä and McLean, 2004), which were higher in the ISV treatment across
711 both fungal (+53% saprophytic fungi) and bacterial (+64% actinomycete) groups (Table 3). While neither
712 fungal/bacterial ratios nor the mole percent distribution (mol%) of fungal PLFAs were significantly different
713 between treatments, we did observe a significant increase in mol% of bacterial PLFAs. Traditional soil food web
714 models assume distinct and preferential utilization of recalcitrant (slow energy channel) and labile (fast energy
715 channel) carbon substrates by fungal and bacterial groups, respectively (Hunt et al., 1987). Under this view, changes
716 in the quantity and quality of organic inputs should therefore induce shifts in soil fungal/bacterial ratios (Wardle et
717 al., 2004). Alternatively, emerging empirical evidence has shown that multi-channel omnivores are the dominant
718 constituency of both fungal and bacterial communities and the presence of these omnivores help to stabilize soil

719 food web communities (Kramer et al., 2016; Wolkovich, 2016). Conceptual models also indicate that fungal and
720 bacterial communities can coexist in a stable state with the presence of large labile carbon pulse inputs, such as the
721 input of dung and urine over the course of a grazing event (de Vries and Caruso, 2016), though bacteria still likely
722 hold a competitive advantage in utilizing these substrates (Ho et al., 2017; Xun et al., 2018). While we observed
723 benefit for both bacteria and fungi with vineyard grazing, relative increases in bacterial groups under ISV
724 nevertheless suggest a soil ecosystem (habitat and/or resource) shift that preferentially benefit consumers of labile
725 substrates and fast energy channels.

726

727 Within bacterial communities, the Gram(+)/Gram(-) ratio is thought a useful indicator of environmental disturbance
728 along the r-K-strategist spectrum. Gram(+) bacteria are generally more adapted to heavily disturbed soil
729 environments (habitat and/or resource limitation) and often less dependent than Gram(-) bacteria on the continuous
730 input of labile carbon compounds (De Vries and Shade, 2013; Fanin et al., 2019). We observed no significant
731 variation in the Gram(+)/Gram(-) ratio between treatments which, when analyzed in tandem with fungal/bacterial
732 ratios, suggests no variation in the stress response of soil microbial communities as a result of cropland grazing. This
733 is further supported by other measured soil physicochemical indicators such soil water content, pH, and compaction
734 (bulk density; BD) – which were similar amongst treatments and indicate physical habitat conditions that are
735 relatively alike (Table 2).

736

737 *.3 Integrated crop-livestock grazing increased perennial cropland soil carbon storage in subsoils*

738 Total SOC storage trended higher in the ISV treatment at both subsurface depth zones. The treatment effect was
739 more significant with increasing depth and resulted in a 39% and 34% increase in SOC under ISV compared to the
740 CONV treatment at 15–30 and 30–45 cm, respectively (Figure 7B-C). The introduction of sheep grazing into
741 perennial cropland increased the physicochemical stabilization of soil carbon within the mineral matrix (MAOC) of
742 the deepest measured subsoil layer (30–45 cm) by 37% compared to the CONV treatment (Figure 7F). The larger
743 soil carbon storage response in the subsoil from grazing reflect changes in soil carbon flux pools (Figure 9B) and are
744 likely related to the increased solubility of deposited animal excreta and deeper spatial distribution of DOC and
745 nutrient substrates (Gross and Harrison, 2019; Rumpel and Kögel-Knabner, 2011). The higher CUE of grazed soils,

746 as indicated by lower Q_{met} values under ISV in the 0–15 and 30–45 cm depths (Figure 5E-F), should theoretically
747 build the capacity for SOC storage (Figure 1B). Where both the rate and efficiency of microbial carbon utilization is
748 higher for labile urine and manure inputs than plant structural compounds (POC) (Cai et al., 2016; Hossain et al.,
749 2017), these properties have been shown to facilitate a more direct pathway toward long-term MAOC stabilization
750 and persistence (Cotrufo et al., 2015, 2013; Dynarski et al., 2020; Haddix et al., 2016; Lavalley et al., 2020;
751 Liebmann et al., 2020). The incorporation of litter-derived POC by animal trampling has also recently been shown
752 to increase its microbial utilization, expediting its decomposition rate and promoting increased physiochemical
753 stabilization of MAOC (Wei et al., 2021).

754
755 However, the introduction of ruminant grazers did not result in significant alterations to surface soil (0–15 cm)
756 aggregation (Figure 3), despite both the physical disturbance of animal trampling and grazing-induced reductions in
757 POC content (Figure 7G). This is notable, given that POC is essential as a nucleus in the formation and stability of
758 soil macroaggregates (Six et al., 2000). We also did not show significant differences in surface soil (0–15 cm)
759 aggregate-associated C between ISV and CONV treatments, as represented by both the total C content (g kg^{-1})
760 associated with soil aggregate size fractions (Figure 8A-D) and the relative distribution (% of total C) across those
761 size fractions (Supplementary Figure 2). As a measurement of the SOC pool associated with four distinct aggregate
762 size categories, aggregate-associated C is an indicator of SOC physical protection via occlusion within aggregates.
763 As such, this study does not indicate that aggregate occlusion of SOC is strongly impacted by perennial ICL grazing.

764
765 Our results corroborate some previous ICL research findings, where increases in SOC have been reported under
766 crop-livestock integration across a spectrum of crop production systems (Acosta-Martínez et al., 2010, 2004; Bansal
767 et al., 2022; Da Silva et al., 2014; de Faccio Carvalho et al., 2010; Franzluebbers et al., 2014; Fultz et al., 2013;
768 Maughan et al., 2009). Given the variability in ICL grazing intensity (density and duration) and its interactions with
769 climatic, edaphic, and co-management components across agricultural systems, other studies across diverse ICL
770 systems have also found negligible (Fernández et al., 2011; Liebig et al., 2020; Tian et al., 2010) and even negative
771 (Tobin et al., 2020) SOC storage benefits associated with cropland grazing. As empirical evidence increasingly
772 shows the positive relationship between soil microbial growth and SOC formation and stabilization (Bradford et al.,

2013; Kallenbach et al., 2016, 2015; Wang et al., 2021), agroecosystem design and management characteristics that stimulate microbial biomass formation and necromass preservation may be central toward increasing SOC storage (Lange et al., 2015; Liang et al., 2019; Prommer et al., 2020; Wilson et al., 2018). When metabolic investment trade-offs (i.e. less energy to invest elsewhere) are satisfied through ameliorating soil habitat and nutrient limitations, this may facilitate efficient microbial biomass accumulation strategies (Malik et al., 2020). The increased allocation of soil carbon toward active microbial pools, with higher use-efficiency, may indicate a determinant mechanism necessary for increasing SOC accumulation within grazed perennial cropland (Figure 9).

780

781 5. CONCLUSION

782 Minimal attention has been paid toward perennial ICL systems and even less so toward the differing dynamics of the
783 surface soils and subsoils within these systems. As such, this research provides some of the first insights into the
784 potential SOC storage benefits associated with perennial cropland grazing, particularly within subsoils. Our study
785 results provide early evidence that *high-density, short-duration rotational grazing* management in perennial
786 croplands holds significant potential to increase SOC storage. We propose increased rates and efficiency of
787 microbial carbon accrual as underlying mechanisms to explain observed gains in SOC. However, these outcomes
788 differ across soil depths and are strongly influenced by the intensity and periodicity (seasonality and frequency) of
789 grazing events as well as site-specific edaphic and climatic limitations. Soil biogeochemical outcomes may therefore
790 be highly variable under different agroecosystem and grazing management regimes and how they are synchronized
791 across time and space. Considerations of co-management strategies such as understory species composition and
792 tillage regime will partially determine soil habitat and resource conditions. The strategic application of grazing in
793 coordination with knowledgeable shepherding practitioners is likely necessary to optimize potential soil benefits.
794 Where perennial cropland agroecosystem design and management may be easily altered to facilitate both spatial and
795 temporal livestock re-integration, better understanding of the mechanistic pathways between grazing disturbances
796 and cropland SOC cycling will be useful toward strategically improving the internal regulation of soil functions and
797 increasing longer-term SOC storage. As such, we aim to highlight certain benefits of using updated soil carbon
798 conceptual frameworks for linking SOC flux and storage indicators within applied agricultural research contexts.
799 However, measurements of soil carbon flux are extremely dynamic across time and space. Future research should

800 focus on monitoring trials and capturing multiple sampling points over shorter durations, to better elicit specific
801 cropland soil responses to grazing disturbance events.

802

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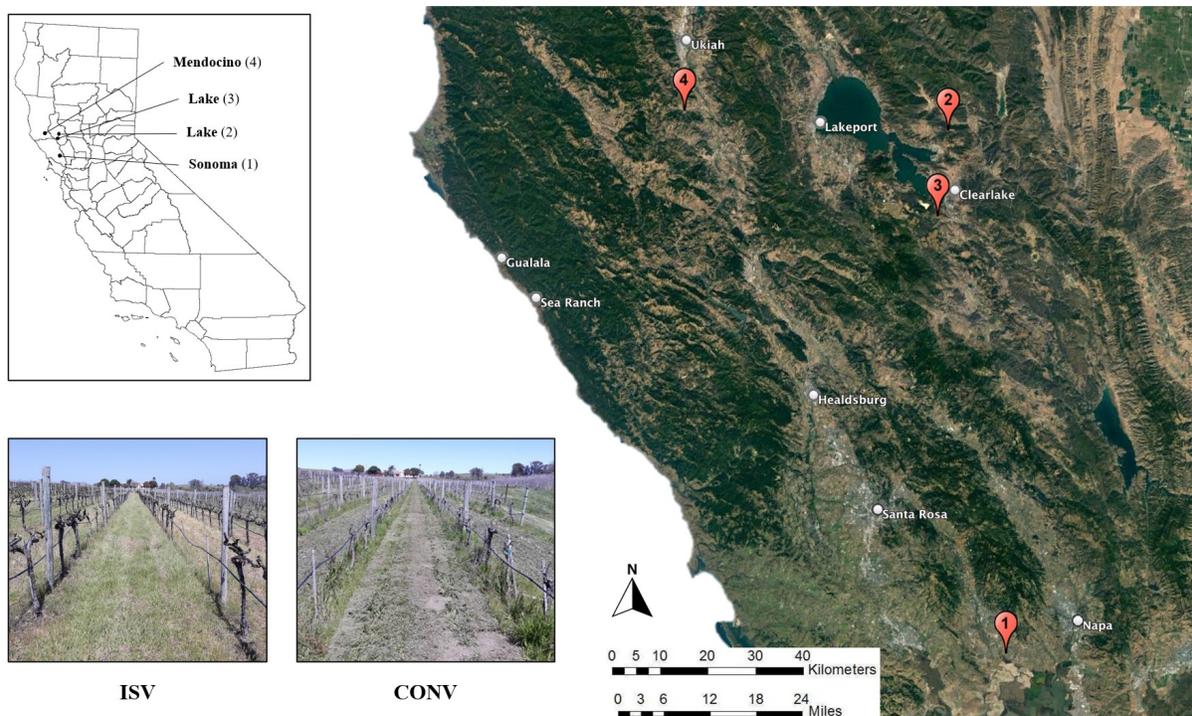
1347 **TABLES AND FIGURES**

1348 **Figure 1: Conceptual framework for soil carbon flux and storage pathways in perennial integrated crop-**
1349 **livestock systems**

1363 C content. Though lower in total C content compared to plant residues, the carbon quality of animal excreta is
 1364 characterized by a higher proportion of soluble, nutrient-rich, and labile dissolved organic carbon (DOC) substances.
 1365 These DOC substances are more reactive, easily diffusible throughout the soil profile, and readily available for
 1366 assimilation as microbial biomass carbon (MBC). The higher microbially carbon use-efficiency of DOC compared
 1367 to particulate organic carbon (POC) substances may potentially facilitate a more direct pathway toward stabilized
 1368 mineral-associated organic carbon (MAOC) storage, and therefore higher total soil organic carbon (SOC) storage.

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1370 **Figure 2: Map of study region**



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1372 Paired vineyard sites (4 locations) consisted of one ‘*non-integrated*’ vineyard (understory vegetation managed
 1373 through mowing and herbicides; CONV) and one adjacent ‘*integrated*’ vineyard (understory vegetation managed
 1374 through grazing; ISV), with one location in Sonoma County (1), two in Lake County (2 & 3), and one in southern
 1375 Mendocino County (4), California, USA. Photos show vineyard comparisons immediately after occurrence of
 1376 grazing (ISV) and mowing (CONV) events. Aerial imagery of the study area was derived from Google Earth Pro.

1377 **Table 1: ISV survey site characteristics of 8 paired vineyards in Sonoma, Lake, and Mendocino Counties,**
 1378 **California, USA**

Site	Understory management treatment	Length of current management (years)	Vine [varietal/rootstock] (vineyard plant date)	Soil texture ^a (% clay)	Soil type ^b (% slope)	Soil disturbance ^c	Synthetic herbicide application ^d	CCOF ^e organic certification status
1	ISV	17	Pinot Noir [UCD 12 / 1103P] (2001)	Loam (22%)	Haire (0-9%)	High	No	Yes
	CONV	21	Pinot Noir [UCD 13 / 1103P] (1997)	Loam (21%)	Haire (0-9%)	High	Yes	No
2	ISV	14	Cabernet Sauvignon [337 / 1103] (2001)	Clay Loam (28%)	Benridge-Sodabay (15-30%)	Moderate	No	No
	CONV	8	Cabernet Sauvignon [15 / 1103] (1998)	Clay Loam (28%)	Benridge-Sodabay (15-30%)	Low	Yes	No
3	ISV	14	Cabernet Sauvignon [08 / 110R] (1992)	Clay Loam (33%)	Sobrante-Guenoc-Hambright (15-30%)	Moderate	No	No
	CONV	21	Cabernet Sauvignon [CS7 / 110R] (1996)	Clay Loam (29%)	Sobrante-Guenoc-Hambright (15-30%)	Low	Yes	No
4	ISV	17	Chardonnay [76 / 5C] (2001)	Loam (25%)	Cole (2-5%)	Low	No	Yes
	CONV	10	Chardonnay [76 / 5C] (2003)	Loam (27%)	Cole (2-5%)	Low	Yes	No

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1380 ^a Based from hydrometer measured sand, silt, and clay particle content (0–15 cm)

1381 ^b Haire clay loam (0 to 9% slope): clayey, mixed, thermic Typic Haploxerults; Benridge-Sodabay loams (15 to 30%

1382 slope): (Benridge) fine, mixed, thermic Mollic Palexeralfs / (Sodabay) fine-loamy, mixed, thermic Mollic

1383 Haploxeralfs; Sobrante-Guenoc-Hambright complex (15 to 30% slope): (Sobrante) fine-loamy, mixed, thermic

1384 Mollic Haploxeralfs / (Gueonic) fine, kaolinitic, thermic Typic Rhodoxeralfs / (Hambright) loamy-skeletal, mixed,

1385 thermic Lithic Haploxerolls. Collected from the USDA-NRCS *SoilWeb* app

1386 (<https://casoilresource.lawr.ucdavis.edu/>)

1387 ^c Soil disturbance is categorized as *Low* (infrequent mow or graze), *Moderate* (combination of infrequent mow +
1388 graze), and *High* (combination of infrequent mow or graze + conservation tillage [i.e. shallow tillage of every other
1389 row in alternating years])

1390 ^d Applications occurring in vineyard undervine row only. Vineyard interrow was managed exclusively with grazing
1391 or mowing

1392 ^e California Certified Organic Farmers (CCOF) – United States Department of Agriculture certifying agency

1393 * *All vineyards contained planted cover crops in vineyard interrow, but not in undervine row*

1394 * *No vineyards contained either organic or synthetic fertilizer amendments within vineyard interrow*

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1414 **Table 2: Physicochemical properties from integrated sheep-vineyard (ISV) and conventional understory**
 1415 **(CONV) managed soils**

Depth (cm)	Location (Site #)								Treatment (p-value)	Treatment x Location (p-value)
	Sonoma (1)		Lake (2)		Lake (3)		Mendocino (4)			
	ISV	CONV	ISV	CONV	ISV	CONV	ISV	CONV		
Total N (g kg⁻¹) [Total C:N]										
0-15	2.0* [10.4]	1.3 [10.6]	1.8 [14.1]	1.6 [15.8]	2.3* [13.5]	1.6 [16.5]	2.2 [12.5]	1.8 [12.5]	0.013* [0.649]	0.620 [0.954]
15-30	1.2* [9.9]	0.8 [8.3]	0.7 [13.6]	0.8 [13.3]	0.9 [14.6*]	0.6 [12.1]	1.3 [11.7]	0.9 [12.4]	0.228 [0.593]	0.885 [0.750]
30-45	0.9 [8.8]	0.7 [7.1]	0.5 [12.6***]	0.5 [8.8]	0.5 [12.6***]	0.5 [9.4]	1.1 [12.3]	0.9 [12.2]	0.694 [0.084]	0.939 [0.430]
Mineral N (NH₄⁺ + NO₃⁻) (mg kg⁻¹)										
0-15	16.6	15.1	13.5***	5.9	21.2***	8.4	5.2	5.0	0.183	0.975
15-30	9.2*	5.2	4.2	4.0	7.7***	2.7	4.4	4.1	0.211	0.672
30-45	7.6**	4.7	2.4	1.4	5.2*	2.7	5.0	4.3	0.042*	0.750
Extractable P (μg g⁻¹)										
0-15	26.4* (3.5)	21.2 (3.4)	11.6* (2.6)	6.8 (0.8)	19.24* (1.2)	10.8 (1.6)	16.2** (1.3)	7.2 (0.9)	<0.001 ***	0.443
15-30	9.8 (3.5)	9.3 (2.7)	4.8 (0.3)	4.3 (1.0)	11.0*** (1.3)	3.8 (0.7)	8.0 (0.9)	8.0 (0.9)	0.309	0.985
30-45	7.1 (2.0)	6.1 (2.0)	4.8 (0.4)	4.6 (0.8)	9.1** (1.3)	3.5 (0.2)	6.5 (0.5)	10.5 (2.7)	0.371	0.789
Bulk density (g cm⁻³)										
0-15	1.54 (0.04)	1.64 (0.02)	1.28 (0.04)	1.25 (0.03)	1.20 (0.01)	1.30 (0.02)	1.28 (0.04)	1.27 (0.04)	0.634	0.966
Soil water content (g g⁻¹)										
0-15	0.25 (0.01)	0.19 (0.01)	0.21 (0.004)	0.16 (0.01)	0.30 (0.01)	0.27 (0.01)	0.08 (0.01)	0.08 (0.004)	0.798	0.840
15-30	0.20 (0.01)	0.18 (0.01)	0.24 (0.004)	0.20 (0.02)	0.30 (0.01)	0.28 (0.01)	0.11 (0.01)	0.09 (0.01)	0.777	0.844
30-45	0.22 (0.01)	0.22 (0.02)	0.27 (0.02)	0.27 (0.02)	0.27 (0.01)	0.31 (0.01)	0.12 (0.01)	0.11 (0.01)	0.851	0.931
pH										
0-15	7.28 (0.04)	7.49 (0.09)	6.79 (0.08)	6.88 (0.08)	6.43 (0.10)	6.51 (0.17)	6.69 (0.10)	6.60 (0.10)	0.738	0.655
15-30	7.00 (0.10)	6.98 (0.12)	6.96 (0.10)	7.01 (0.08)	6.61 (0.17)	6.69 (0.09)	6.48 (0.04)	6.81 (0.08)	0.678	0.254
30-45	7.01 (0.07)	6.93 (0.07)	7.00 (0.15)	7.14 (0.05)	6.80 (0.22)	6.83 (0.08)	6.60 (0.03)	6.51 (0.13)	0.416	0.325
EC (dS m⁻¹)										
0-15	0.25*** (0.03)	0.14 (0.01)	0.16* (0.005)	0.11 (0.01)	0.22*** (0.03)	0.11 (0.01)	0.12 (0.02)	0.10 (0.005)	0.058	0.153
15-30	0.21*** (0.02)	0.09 (0.01)	0.12* (0.02)	0.08 (0.01)	0.08 (0.01)	0.10 (0.02)	0.07 (0.01)	0.06 (0.004)	0.121	0.043*
30-45	0.27*** (0.03)	0.14 (0.03)	0.11 (0.01)	0.09 (0.01)	0.09 (0.01)	0.12 (0.02)	0.06 (0.004)	0.06 (0.003)	0.353	0.138
CEC (me 100g⁻¹)										
0-15	16.6*** (1.8)	11.2 (0.6)	14.8 (0.5)	14.3 (0.4)	15.8 (0.8)	18.1 (0.8)	18.2 (0.7)	17.8 (0.4)	0.270	0.041*
15-30	18.3*** (1.4)	11.9 (0.9)	13.7 (0.9)	13.2 (0.5)	13.8 (0.9)	15.1 (0.9)	15.3 (1.2)	19.6 (0.7)	0.333	0.203
30-45	22.0*** (1.7)	15.5 (2.6)	13.5 (0.8)	14.1 (0.9)	13.3 (0.8)	16.3 (0.3)	19.5 (1.7)	19.4 (0.07)	0.637	0.278

1416
 1417 Soils were sampled across 8 paired vineyards (4 locations). Soils cores were separated into three depths zones (0–15
 1418 cm, 15–30 cm, and 30–45 cm) and measured for total N (TN), mineral N (NH₄⁺-N plus NO₃⁻-N), extractable
 1419 phosphorous (P), bulk density, soil water content, pH, electrical conductivity (EC), and cation exchange capacity
 1420 (CEC). Means are followed by standard error in parentheses. Shown are the treatment and treatment(x)location
 1421 statistical significance across sites (n=64). For each location, a Tukey-Kramer means (n=16) comparison was used to
 1422 evaluate significant pairwise difference between each treatment. Asterisks (*) denote significant treatment
 1423 differences at each depth increment.

1424 * p <0.05; ** p <0.01; *** p < 0.001.

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1427 **Table 3: PLFA biomarkers and biological ratios from integrated sheep-vineyard (ISV) and conventional**
 1428 **understory (CONV) surface soils (0–15 cm depth)**

	Location (Site #)								Treatment (<i>p</i> -value)	Treatment x Location (<i>p</i> -value)
	Sonoma (1)		Lake (2)		Lake (3)		Mendocino (4)			
	ISV	CONV	ISV	CONV	ISV	CONV	ISV	CONV		
Total PLFAs										
ng g ⁻¹	3628* (632)	2560 (252)	5408* (227)	4341 (206)	3594* (470)	2338 (396)	1540 (535)	934 (263)	<0.001 ***	0.613
Fungi										
ng g ⁻¹	457 (119)	252 (45)	625 (23)	454 (55)	451 (88)	332 (54)	7 (6)	12 (9)	0.004**	0.062
mol%	11.8 (1.7)	9.3 (1.4)	11.7 (0.5)	10.4 (1.1)	12.5 (1.5)	15.8 (3.7)	0.6 (0.4)	0.8 (0.5)	0.991	0.749
AM fungi										
ng g ⁻¹	153 (34)	66 (12)	192 (9)	145 (17)	143 (25)	146 (67)	8 (6)	7 (6)	0.544	0.519
mol%	4.0 (0.6)	2.4 (0.4)	3.6 (0.2)	3.3 (0.3)	3.9 (0.3)	3.8 (0.7))	0.7 (0.5)	0.5 (0.3)	0.647	0.382
Saprophytic fungi										
ng g ⁻¹	304 (89)	187 (34)	433* (23)	309 (42)	308 (65)	186 (38)	0 (0)	0 (0)	0.004 **	0.197
mol%	7.9 (1.2)	6.9 (1.1)	8.1 (0.5)	7.1 (0.9)	8.6 (1.2)	7.9 (1.2)	0 (0)	0 (0)	0.276	0.576
Bacteria										
ng g ⁻¹	1734* (318)	1095 (117)	2380* (124)	1737 (102)	1528* (212)	788 (142)	346 (107)	205 (51)	<0.001***	0.161
mol%	47.3 (1.1)	42.6 (0.7)	43.9 (0.8)	40.0 (1.3)	42.0* (1.4)	33.6 (3.3)	22.2 (2.6)	21.7 (2.8)	0.002**	0.484
Actinomycetes										
ng g ⁻¹	271* (45)	151 (18)	350 (25)	281 (14)	217** (31)	83 (15)	91 (21)	50 (12)	<0.001***	0.270
mol%	7.5 (0.2)	6.0 (0.4)	6.5 (0.3)	6.5 (0.4)	6.0* (0.3)	3.6 (0.5)	6.4 (1.1)	5.4 (0.6)	0.223	0.781
Fungi:bacteria ratio										
	0.25 (0.03)	0.22 (0.03)	0.27 (0.01)	0.26 (0.03)	0.30 (0.03)	0.34 (0.03)	0.07 (0.05)	0.06 (0.03)	0.887	0.680
Gram(+):gram(-) bacteria ratio										
	1.22 (0.12)	1.51 (0.34)	1.05 (0.04)	1.07 (0.07)	1.22 (0.10)	1.06 (0.09)	2.48 (0.43)	1.82 (0.22)	0.384	0.021*
Saturated:unsaturated ratio										
	1.54 (0.17)	2.27*** (0.47)	1.47 (0.06)	1.69 (0.12)	1.74 (0.20)	1.62 (0.11)	10.96 (1.96)	9.84 (1.99)	0.909	0.319

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 1430 Phospholipid fatty acid (PLFA) profiles indicative of fungi, arbuscular mycorrhizal (AM) fungi, saprophytic fungi,
 1431 bacteria, and actinomycetes, as well as the relative ratios of fungi-to-bacteria and stress indicator ratios gram(+)-to-
 1432 gram(-) bacteria and saturated-to-unsaturated fatty acids. Ratios are unitless, while PLFAs are given in both ng g
 1433 soil⁻¹ and mole percent distribution (mol%). Means are followed by standard error in parentheses. Shown are the
 1434 treatment and treatment(x)location statistical significance across sites (n=64). For each location, a Tukey-Kramer
 1435 means (n=16) comparison was used to evaluate significant pairwise difference between each treatment. Asterisks (*)
 1436 denote significant treatment differences at each depth increment.

1437 * p <0.05; ** p <0.01; *** p < 0.001.

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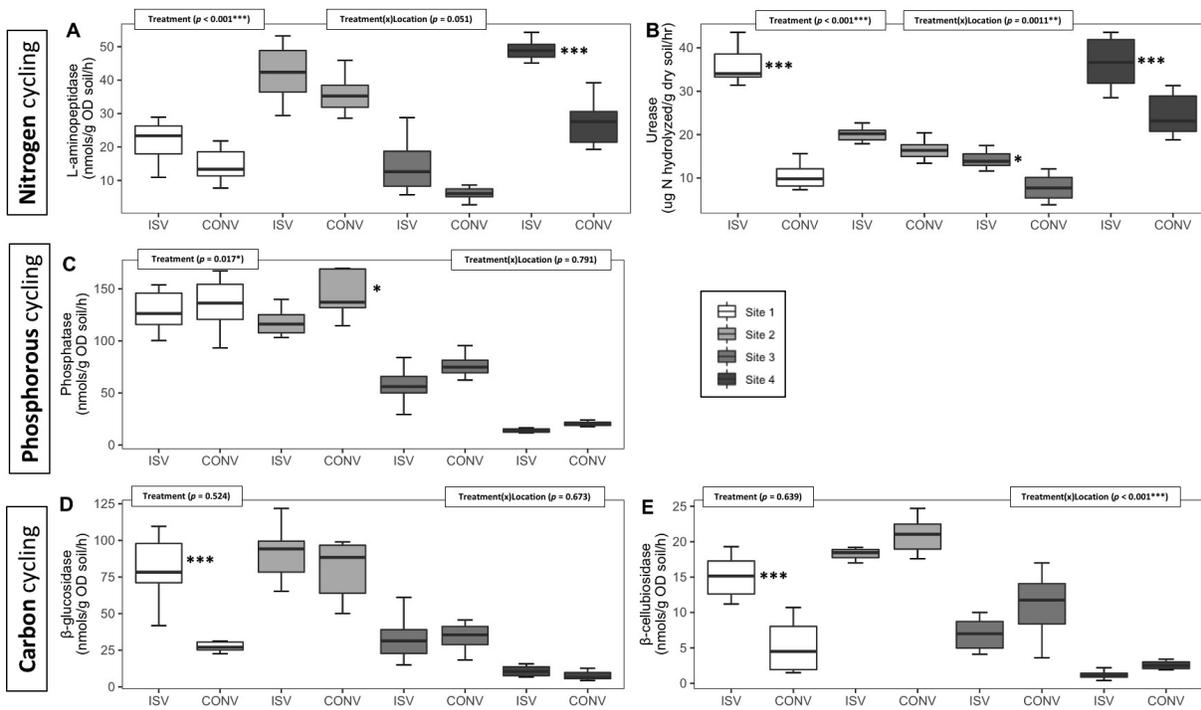
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1444 **Figure 3: Management impact on soil exo-enzyme activity potential from surface soil (0–15 cm depth)**



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1446 Impact of integrated sheep-vineyard (ISV) and conventional vineyard understory (CONV) management on surface
1447 soil (0 – 15 cm depth) exo-enzyme synthesis potential for (A) L-aminopeptidase, (B) urease, (C) phosphatase, (D) β -
1448 glucosidase, and (E) β -cellubiosidase. Shown are the treatment and treatment(x)location statistical significance
1449 across sites (n=64). For each location, a Tukey-Kramer means (n=16) comparison was used to evaluate significant
1450 pairwise difference between each treatment. Error bars represent standard error. Asterisks (*) denote significant
1451 treatment differences at each depth increment.

1452 * $p < 0.05$; ** $p < 0.01$; *** $p < 0.001$.

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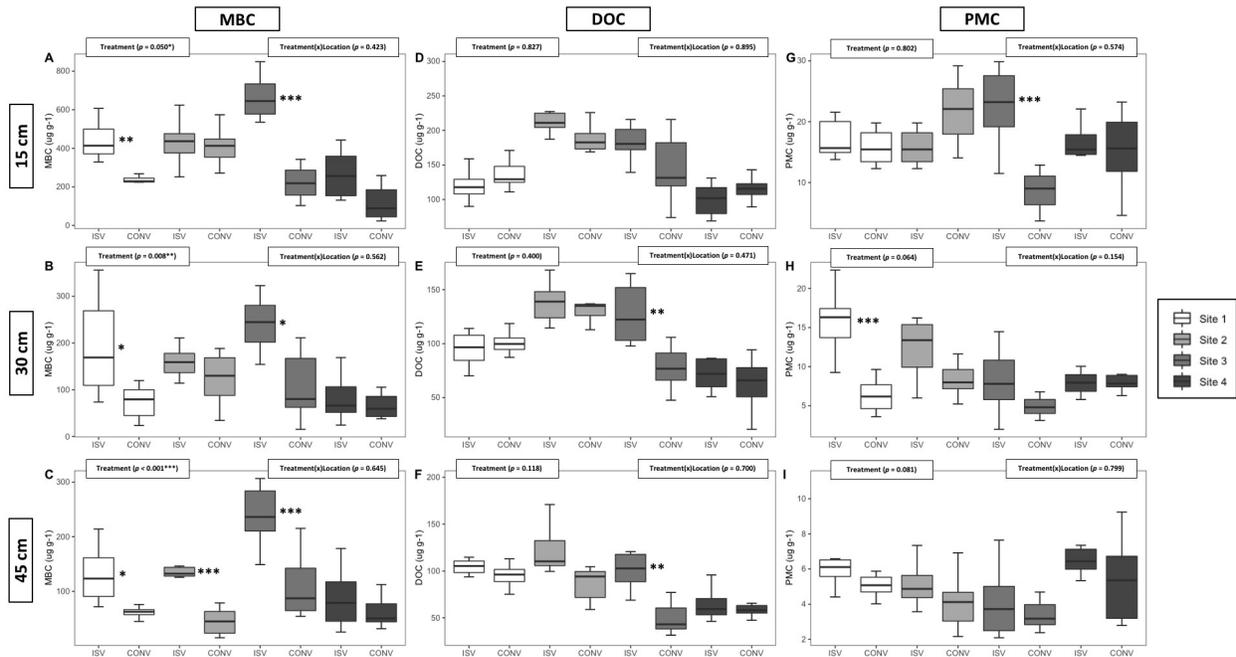
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1461 **Figure 4: Soil carbon flux pools in integrated sheep-vineyard (ISV) and conventional understory (CONV) managed soils**

1462 **managed soils**



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1464 (A-C) Microbial biomass C (MBC), (D-F) dissolved organic C (DOC), and (G-I) 3-day potentially mineralizable C

1465 (PMC) were measured at three depths zones (0–15cm, 15–30cm, and 30–45cm) from integrated sheep-vineyard

1466 (ISV) and conventional vineyard understory (CONV) managed soils. Shown are the treatment and

1467 treatment(x)location statistical significance across sites (n=64). For each location, a Tukey-Kramer means (n=16)

1468 comparison was used to evaluate significant pairwise difference between each treatment. Error bars represent

1469 standard error. Asterisks (*) denote significant treatment differences at each depth increment.

1470 * p < 0.05; ** p < 0.01; *** p < 0.001.

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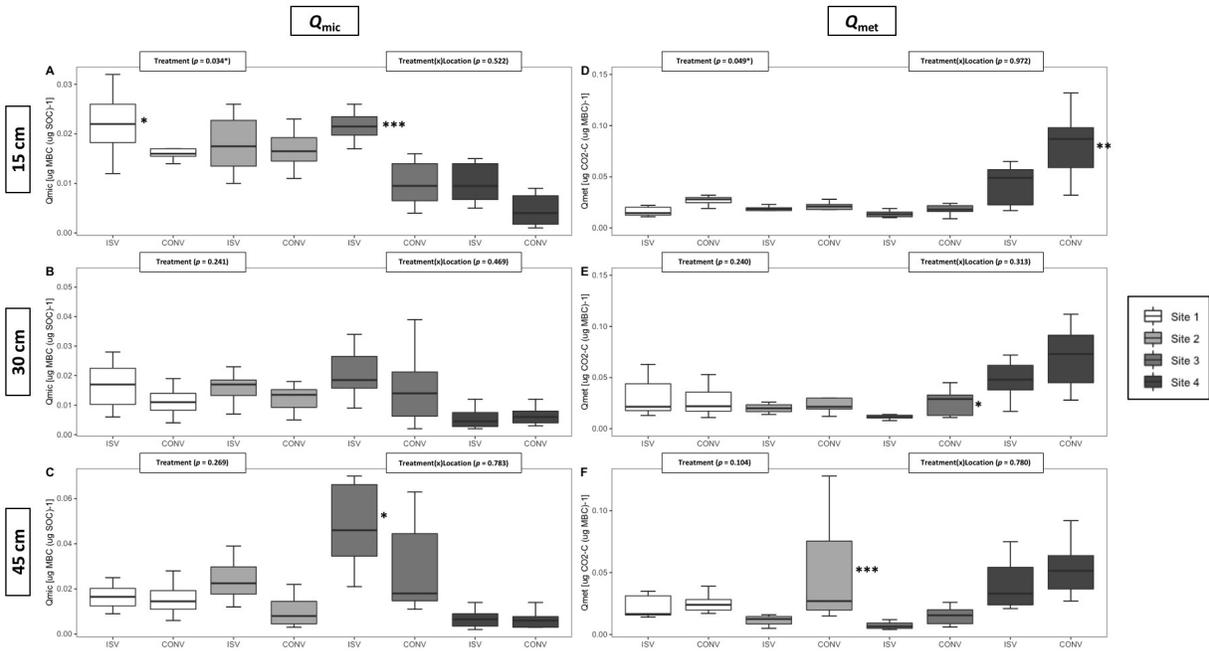
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Figure 5: Microbial quotient (Q_{mic}) and metabolic quotient (Q_{met}) of integrated sheep-vineyard (ISV) and conventional understory (CONV) managed soils



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(A-C) Microbial quotient (Q_{mic}) and (D-F) metabolic quotient (Q_{met}) were measured at three depths zones (0–15cm, 15–30cm, and 30–45cm) from integrated sheep-vineyard (ISV) and conventional vineyard understory (CONV) managed soils. Shown are the treatment and treatment(x)location statistical significance across sites (n=64). For each location, a Tukey-Kramer means (n=16) comparison was used to evaluate significant pairwise difference between each treatment. Error bars represent standard error. Asterisks (*) denote significant treatment differences at each depth increment.

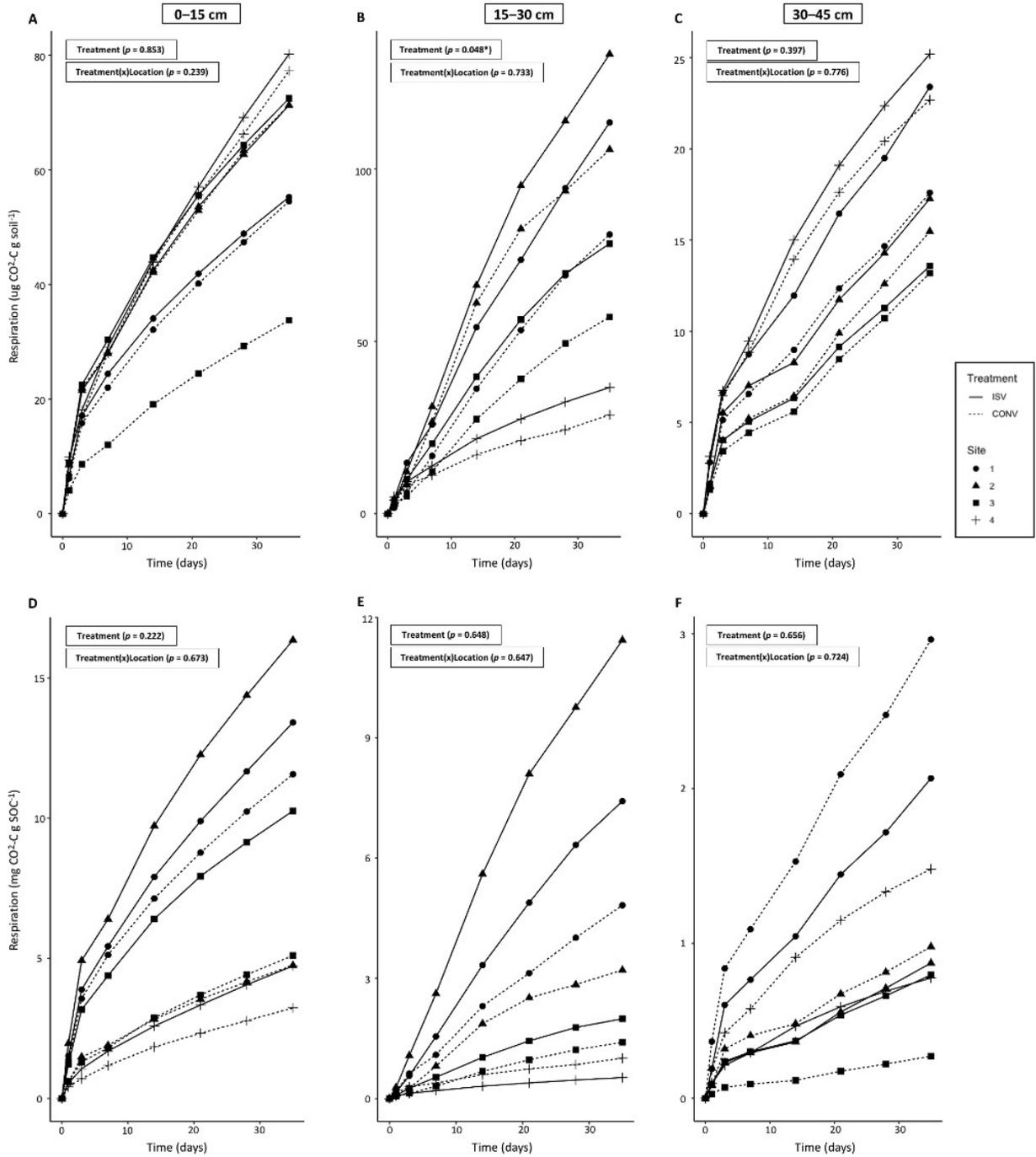
* $p < 0.05$; ** $p < 0.01$; *** $p < 0.001$.

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1495 **Figure 6: Soil carbon mineralization rates over a 35-day incubation from integrated sheep-vineyard (ISV)**

1496 **and conventional understory (CONV) managed soils**



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1498 Soils were incubated and measured for cumulative C mineralization ($C_{min_{total}}$; $\mu\text{g CO}_2\text{-C g soil}^{-1} \text{ day}^{-1}$) via microbial
1499 respiration rates at seven time points (1, 3, 7, 14, 21, 28, and 35 days) in soils from three depth zones (A) 0–15cm,
1500 (B) 15–30cm, and (C) 30–45cm. Soil C mineralization was also expressed relative to total SOC content ($C_{min_{soc}}$;
1501 $\text{mg CO}_2\text{-C g SOC}^{-1} \text{ day}^{-1}$) for depth zones (D) 0–15cm, (E) 15–30cm, and (F) 30–45cm. Treatment and
1502 treatment(x)location significance was calculated across sites (n=64).

1503 * p <0.05; ** p <0.01; *** p < 0.001.

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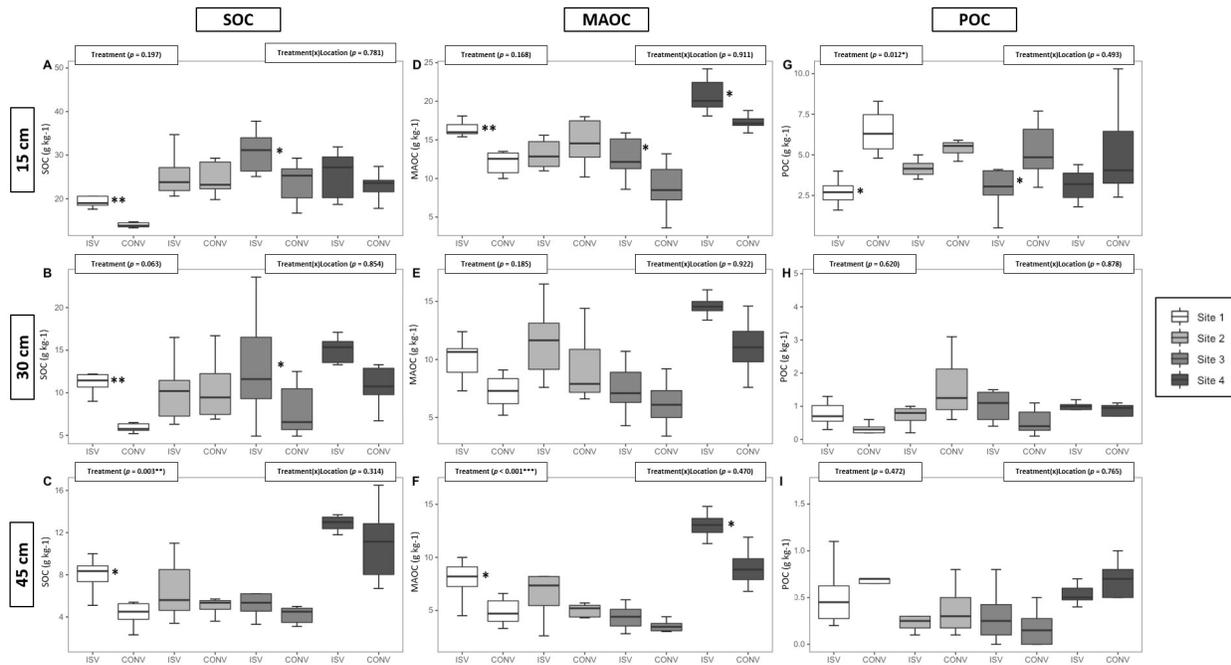
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1527 **Figure 7: Soil carbon stabilization and storage pools in integrated sheep-vineyard (ISV) and conventional**

1528 **understory (CONV) managed soils**



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1530 (A-C) Total soil organic C (SOC), (D-F) mineral-associated organic C (MAOC), and (G-I) particulate organic C

1531 (POC) were measured at three depths zones (0–15 cm, 15–30 cm, and 30–45 cm) from integrated sheep-vineyard

1532 (ISV) and conventional vineyard understory (CONV) managed soils. Shown are the treatment and

1533 treatment(x)location statistical significance across sites (n=64). For each location, a Tukey-Kramer means (n=16)

1534 comparison was used to evaluate significant pairwise difference between each treatment. Error bars represent

1535 standard error. Asterisks (*) denote significant treatment differences at each depth increment.

1536 * p < 0.05; ** p < 0.01; *** p < 0.001.

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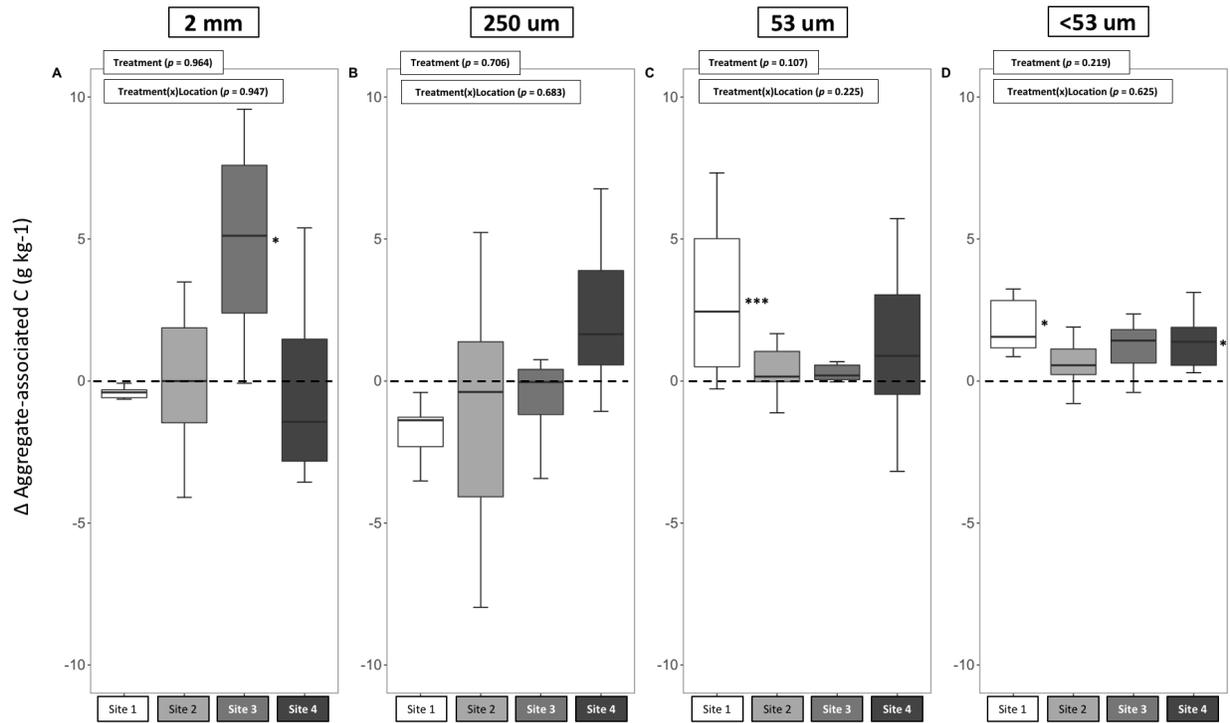
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1544 **Figure 8: Aggregate-associated C pools in integrated sheep-vineyard (ISV) and conventional understory**

1545 **(CONV) managed surface soils (0–15 cm depth)**



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1547 Surface soils (0–15 cm depth) were measured for the total C content associated with four soil aggregate physical
1548 size fractions (>2000 μ m, 250-2000 μ m, 53-250 μ m, and <53 μ m; A-D). For A-D, the dashed bar represents the CONV

1549 treatment mean for each paired site. Boxes above or below the line represent the relative increase or decrease in ISV
1550 values (Δ) for each site. Shown are the treatment and treatment(x)location statistical significance across sites (n=64).

1551 For each location, a Tukey-Kramer means (n=16) comparison was used to evaluate significant pairwise difference
1552 between each treatment. Error bars represent standard error. Asterisks (*) denote significant treatment differences
1553 within site at each size fraction.

1554 * p < 0.05; ** p < 0.01; *** p < 0.001.

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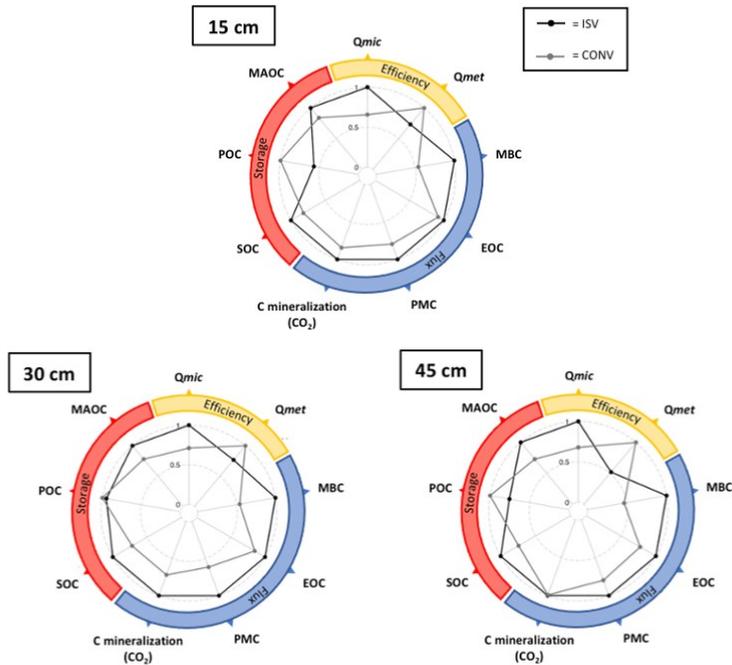
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1560 **Figure 9: Proposed linkages between soil carbon flux and storage as impacted by perennial crop-livestock**

1561 **integration**



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1563 Central radar plots with mean normalized values of soil carbon storage pools, flux pools, and microbial use-

1564 efficiency indicators from integrated sheep-vineyard (ISV) and conventional vineyard understory (CONV) managed

1565 soils at three depth zones (0–15, 15–30, and 30–45 cm).

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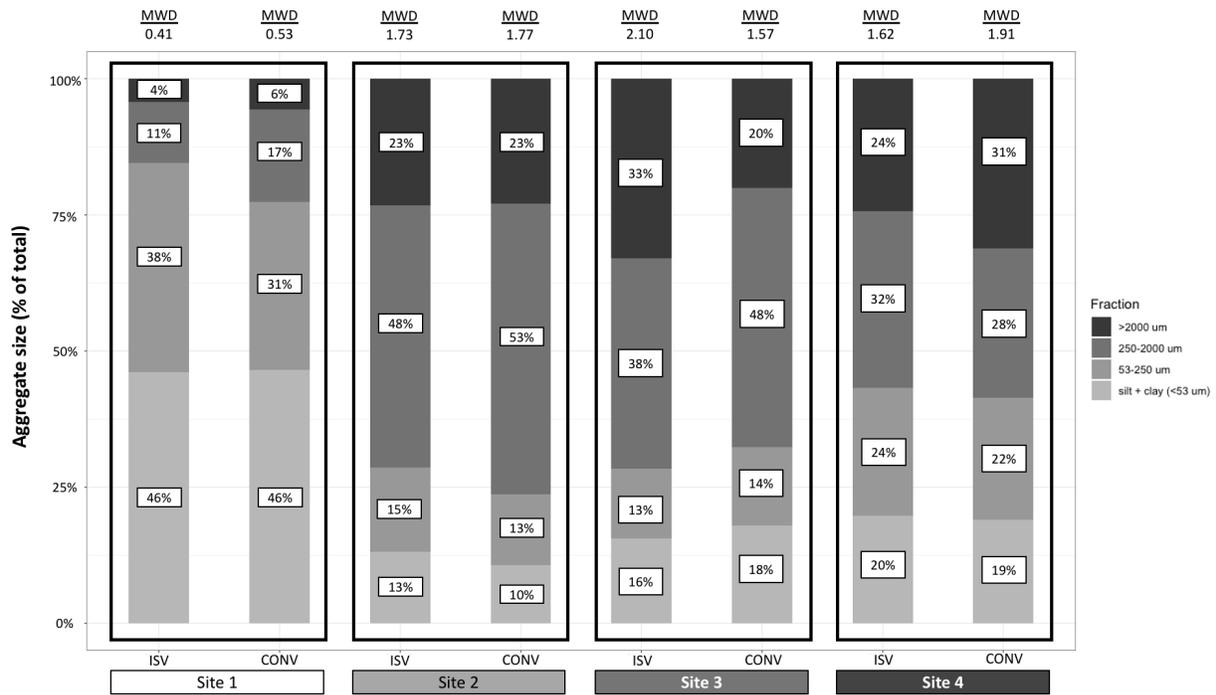
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1577 SUPPLEMENTAL TABLES AND FIGURES

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1579 Supplemental Figure 1: Aggregate size distribution from integrated sheep-vineyard (ISV) and conventional

1580 understory (CONV) managed surface soils (0–15 cm depth)



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1582 Relative distribution of four soil aggregate physical size fractions (>2000um, 250-2000um, 53-250um, and <53um).

1583 Shown are mean values for each size category and the mean weight diameter (MWD). For each site, a Tukey-

1584 Kramer means (n=16) comparison was used to evaluate significant pairwise difference between each treatment.

1585 Treatment and treatment(x)location significance was also calculated across sites (n=64). For MWD, both the

1586 Treatment (p=0.953) and treatment(x)location interaction (p=0.979) effects were non-significant. Treatment and

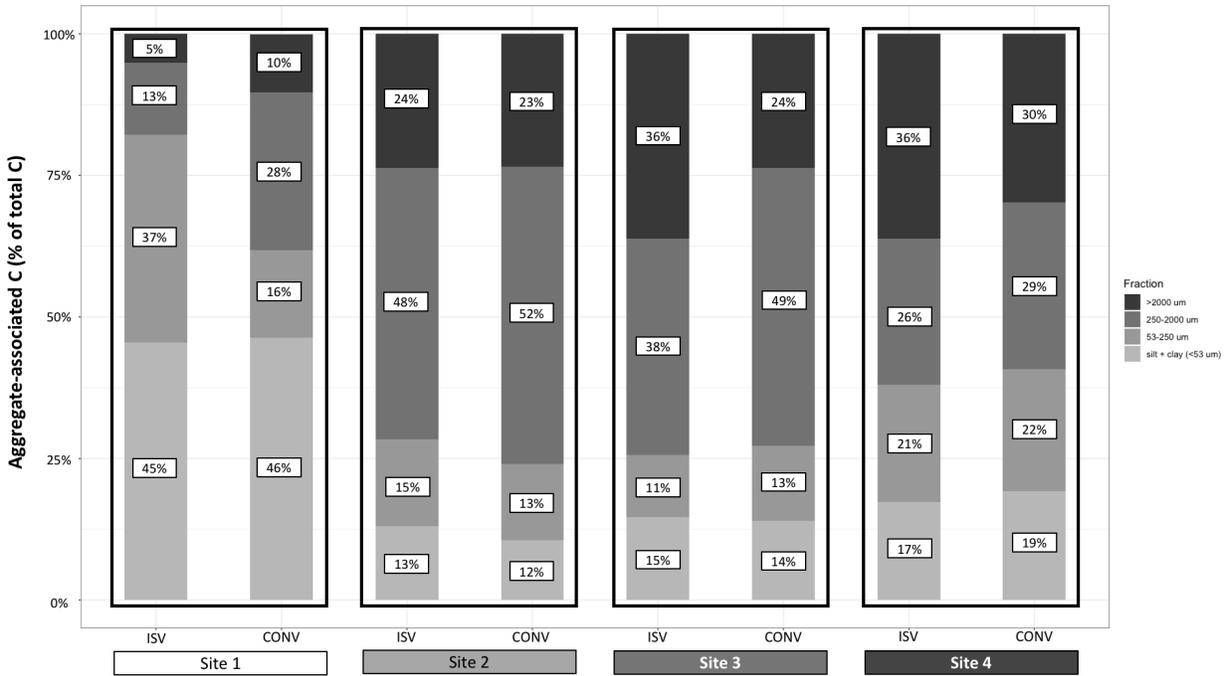
1587 treatment(x)location effects were also non-significant for all size fractions.

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Supplemental Figure 2: Relative distribution of aggregate-associated C pools in integrated sheep-vineyard (ISV) and conventional understory (CONV) managed surface soils (0–15 cm depth)



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The relative distribution (% of total C) across four surface soil (0–15 cm depth) aggregate physical size fractions (>2000um, 250-2000um, 53-250um, and <53um). For each site, a Tukey-Kramer means (n=16) comparison was used to evaluate significant difference between each treatment. Treatment and treatment(x)location significance was also calculated across sites (n=64). Treatment and treatment(x)location effects were non-significant for all size fractions.