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Estimation of stress and stress-induced permeability change in a geological nuclear waste repository in a thermo-hydrologically coupled simulation

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1 **Estimation of stress and stress-induced permeability change in a**
2 **geological nuclear waste repository in a thermo-hydrologically**
3 **coupled simulation**

4
5 **Abstract**

6 Geologic disposal is a promising solution for a safe permanent isolation of accumulated high-level nuclear
7 waste from nuclear powerplants. The behavior of host rock is highly coupled thermally, hydromechanically,
8 and chemically. Numerical simulations of such coupled phenomena for the extremely long term (> 100,000
9 years) and large length scale (> 1 km) of geologic disposal remain to be computationally challenging. In
10 this study, a methodology has been developed to approximate stress and stress-induced permeability change
11 in host rock using only thermo-hydrological (TH) variables. A coupled thermo-hydromechanical (THM)
12 simulation is carried out using TOUGH-FLAC simulator to model THM behaviors of a generic nuclear
13 waste repository, in order to evaluate the performance of the developed methodology, which is implemented
14 in a coupled TH simulation. Results show that stress and permeability change estimated by the developed
15 methodology in the TH-coupled simulation match those calculated in the THM-coupled simulation over
16 the simulated timeframe of over 10,000 years. Details about errors in stress and permeability estimates
17 accrued by the developed methodology are discussed in this paper. The developed methodology will help
18 incorporate stress-induced permeability change into existing TH simulators for the long-term radionuclide
19 transport in geologic disposal.

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1. Introduction

The United States has accumulated, as of end of 2018, over 84,000 metric tons of high-level nuclear waste [1], which must be stored remotely and carefully such that the waste will be permanently isolated from human beings and the biosphere [2]. Geological disposal is considered for safe permanent isolation of high-level nuclear waste, and hence such possibility is being extensively investigated in many countries considering various host rocks [3], including the past studies at the Yucca Mountain, Nevada, USA [4]. For each deep disposal site, the safety is to be assessed on the basis of a safety case developed from the scientific and technical basis for geological disposal [5]. The core of the safety case consists of safety and performance assessments. In a safety assessment, the estimated consequences of any releases from the repository are compared with the appropriate safety criteria, whereas in a performance assessment (PA), the evolution and performance of the isolation barriers for hundreds of thousands of years is estimated [6].

One type of host rock formations considered for geologic nuclear waste disposal is clay/shale formations, because of their favorable properties (e.g., low permeability, slow diffusion, high retention of radionuclides, self-healing of fractures) to curtail migration of radionuclides if released from a waste package. As a result, many studies have been conducted to assess the feasibility of nuclear waste storage in clay/shale formations in Europe [7]–[10] as well as in the US [11], [12]. Another feature of clay/shale formation is their relatively low thermal conductivity, which maybe a disadvantage because it can lead to high temperatures that need to be managed by, for example, storing the waste for a long time on the surface before underground disposal.

In the US, the properties of clay/shale formations are currently being investigated as potential feasible host rock formations for nuclear waste disposal [13] and generic (non-site specific) PA model simulations are developed and tested [14]. PFLOTRAN [15] is a state-of-the-art massively parallel subsurface flow and reactive transport code that is being applied for such PA modeling exercises. However, PA calculations

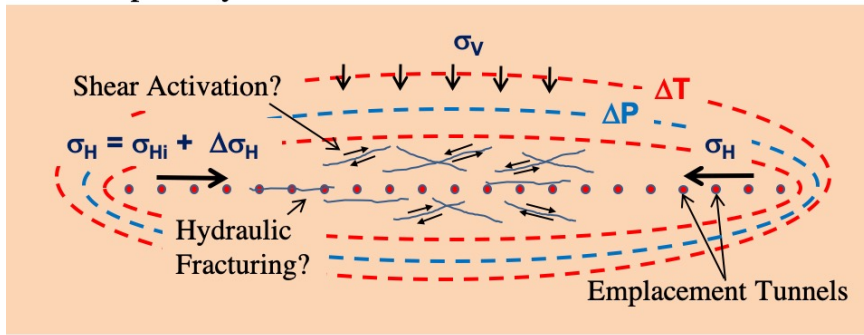
47 with uncertainty quantification will require a large number of simulations, while at the same time need to
48 consider the evolution of the system over hundreds of thousands of years. This includes potential changes
49 in the transport properties, such as the permeability of host rock, induced by coupled thermo-hydro-
50 mechanical-chemical (THMC) processes.

51
52 Field experiments at underground rock laboratories as well as numerical simulation have shown that
53 strongly coupled THM processes is likely to occur in a clay/shale repository [16]–[19] that could also be
54 impacted by chemical processes [20], [21]. For example, so-called thermal pressurization and thermal stress
55 could potentially lead to formation fracturing or shear activation, as well as high stress concentration and
56 damage around emplacement tunnels (Figure 1) [22]. Moreover, increasing horizontal stress will act on the
57 repository tunnels and that, through stress concentration around the tunnel openings, could cause
58 compressive spalling failure or tensile failure in different parts of the tunnel walls (Figure 1).

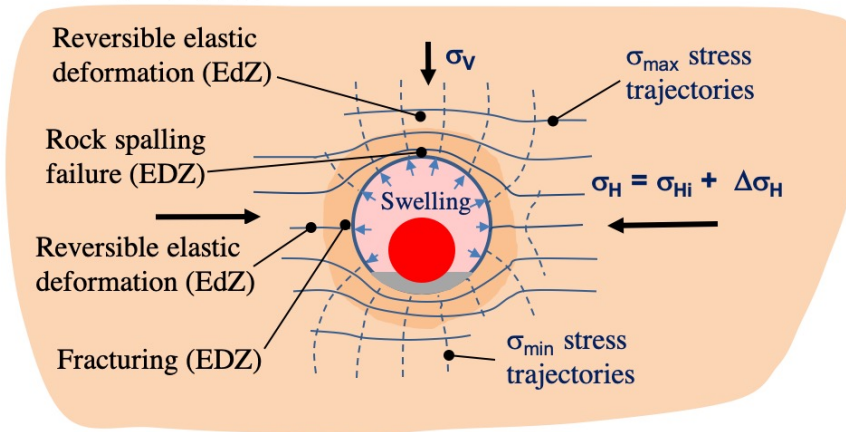
59
60 Stresses and potential failure around tunnel openings will impact the formation and evolution of the
61 disturbed zone around the emplacement tunnels. Generally, the disturbance may be divided into excavation
62 disturbed (EdZ) and damaged zones (EDZ). The EdZ is a zone with hydro-mechanical and geochemical
63 modifications, without major changes in flow and transport properties, whereas the EDZ is a zone with
64 hydro-mechanical and geochemical modifications, inducing significant changes in flow and transport
65 properties [23]. The THM evolution within the bentonite buffer will also impact the EDZ evolution as the
66 re-saturation and swelling of the bentonite buffer provides mechanical support to the EDZ [19]. In turn, the
67 swelling pressure may be affected by chemical changes within the bentonite buffer, thus leading to the
68 necessity for such coupled THMC evolution to be assessed over the regulatory repository period [20].

69

At Repository Scale



At Tunnel Scale



70

71 **Figure 1 Schematic of repository scale coupled thermo-hydro-mechanical responses and their**
72 **impact on emplacement tunnels (Modified from Rutqvist, 2020).**

73

74 Coupled THMC numerical analyses are computationally challenging considering the extremely long
75 timeframe that is required for geological disposal, which is beyond hundreds of thousands of years, in
76 addition to the spatial and temporal resolutions required to sufficiently resolve the geomechanics. Among
77 the THMC processes, the mechanical part is often the most time-consuming as complex constitutive
78 equations must be solved to simulate elasto-plastic and failure response of the rock as well as of the buffer
79 material filled between the rock and nuclear waste canister. Hence, there is a need for a method which can
80 approximate the mechanical behavior of the repository and its effect on flow properties of the rock (i.e.,
81 permeability) from thermal, hydrological, and/or chemical behaviors of the repository. Such a method is
82 beneficial for existing radionuclide transport simulators, such as PFLOTRAN, which are usually only

83 thermo-hydrologically (TH) coupled with radionuclide transport, as mechanically-induced permeability
84 change can be incorporated so as to improve the accuracy of the radionuclide flow simulation.

85

86 In this study, a coupled THM simulation is conducted to model the behavior of a generic clay/shale
87 geological repository for nuclear waste. A method to approximate the mechanical response of the repository
88 is then developed which only requires TH variables to estimate stress and stress-induced permeability
89 change. The performance of the approximation method is evaluated in comparison between the coupled
90 THM and TH simulations in calculating stress and stress-induced permeability change in the rock as well
91 as in the buffer during heating from the nuclear waste canister. The objectives of this study are as follows:

92

- 93 • Estimate stress and stress induced-permeability change in a generic clay/shale geological
94 repository for nuclear waste in the long term
- 95 • Develop a methodology to approximate the mechanical response of the rock and buffer around
96 the nuclear waste canister from TH variables

97

98 TOUGH-FLAC simulator (Rutqvist, 2011; Rinaldi et al., 2018; Luu, 2020) is employed to model THM
99 behaviors of the generic nuclear waste repository. Recent TOUGH-FLAC versions consist of the newly
100 updated TOUGH3 multiphase fluid and thermal transport code [25] and the commercial FLAC3D code
101 [26], in which TOUGH3 is the master code (Blanco-Martín et al., 2016; Kim et al., 2011; Rinaldi et al.,
102 2018). TOUGH3 solves the thermo-hydrological equations and passes TH variables (i.e., pore pressure, gas
103 saturation, capillary pressure, and temperature in the case of equation of state #4 (EOS4)) to FLAC3D at
104 each converged TOUGH3 time step. FLAC3D is then called to solve the mechanical equations under
105 drained conditions to calculate stresses, which are then used to update TH variables and parameters in
106 TOUGH3 (i.e., permeability, porosity, and capillary pressure). A TOUGH-FLAC simulation runs in this

107 sequential manner to model various THM coupled phenomena in geological formation [19], [29], [30]. The
108 Barcelona Basic Model (BBM) [31] is used to simulate the mechanical response of the buffer, which is
109 assumed to be composed of swelling bentonite. In the following section, further details about the
110 methodology of this study are provided.

111

112 **2. Methods**

113 **2.1 Model geometry**

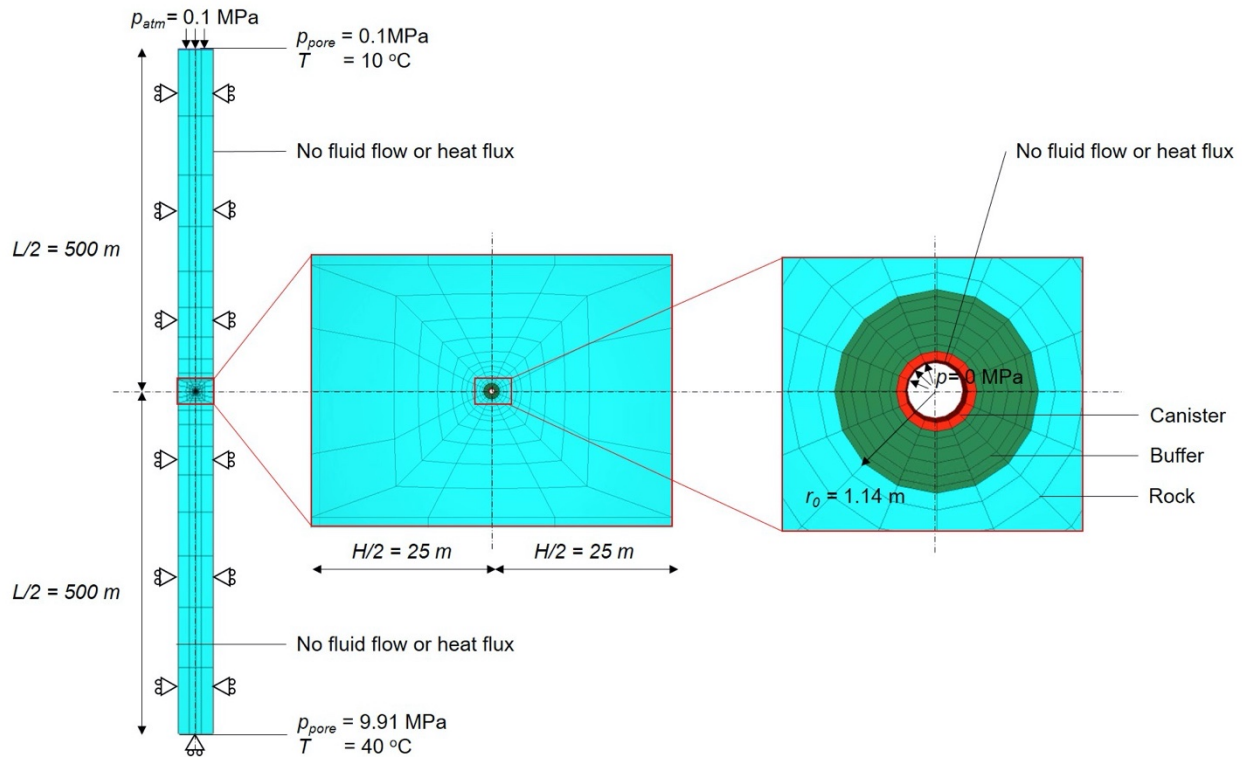
114 Figure 2 shows the geometry and mesh of the generic nuclear waste repository. The nuclear waste disposal
115 tunnel is located at the center of the model, which is at the depth of 500 m from the ground surface and 25
116 m away horizontally from the lateral boundaries (i.e., mid-pillar). The inner and outer radii of the canister
117 are 0.35 m and 0.45 m, respectively, whereas the radius of the tunnel is 1.14 m. This model is three
118 dimensional but the plane-strain condition is applied in the out-of-plane direction where the model thickness
119 is 1 m. The entire model is discretized into 360 elements (i.e., 248 for the rock, 96 for the buffer, and 16
120 for the canister) and the size of the mesh is set finer around the center of the tunnel than the rest of the
121 model domain.

122

123 Boundary conditions are also shown in Figure 2. Thermo-hydrological boundary conditions are set as
124 follows: constant pore pressure and temperature on the top and bottom boundaries, and no fluid flow or
125 heat flux along the lateral and canister interior boundaries. A time-varying heat source (Figure 3) is
126 prescribed in each of the sixteen canister elements to model the heat flux from the nuclear waste package.
127 Mechanical boundary conditions are set as follows: zero displacement perpendicular to the bottom and
128 lateral boundaries (displacement parallel to the boundary surfaces is allowed) and in the out-of-plane
129 direction, and constant pressure on the top and canister interior boundaries. The model geometry and
130 boundary conditions are to represent conditions and THM evolution for an emplacement tunnel at the inner

131 part of a repository, where the highest temperature changes and thermal impact could be expected due to
 132 symmetry assumptions at the tunnel scale (Figure 1).

133



134

135 **Figure 2 The geometry and mesh of the generic nuclear waste repository model.**

136

137 **2.2 Material behaviors**

138 **2.2.1 Thermo-hydrological behaviors**

139 The thermo-hydrological (TH) behaviors of the model are computed by the TOUGH3 simulator. The
 140 equation-of-state module #4 (EOS4) is used. In EOS4, the model components are water and air, which can
 141 be either in the gas or liquid phase, and the primary variables are gas pressure, gas saturation, and air partial

142 pressure for two-phase conditions, whereas they are gas pressure, temperature, and air partial pressure for
 143 single-phase conditions.

144

145 Darcy's law is used to simulate the multiphase flow of the fluid components, which is expressed as follows:

146

$$q_i = -\frac{k_{ri}k}{\mu_i} \nabla P_i \quad (1)$$

147

148 where q_i is the flow rate, k_{ri} is the relative permeability, k is the absolute permeability, μ_i is the dynamic
 149 viscosity, and ∇P_i is the pressure gradient. The subscript i indicates that the parameters are for phase i . The
 150 following relative permeability function (i.e., van Genuchten-Mualem model) is used for the rock:

151

$$k_{rl} = \begin{cases} \sqrt{S^*} \left(1 - \left(1 - (S^*)^{\frac{1}{\lambda_k}} \right)^{\lambda_k} \right)^2 & \text{if } S_l < S_{ls} \\ 1 & \text{if } S_l \geq S_{ls} \end{cases} \quad (2)$$

$$k_{rg} = \begin{cases} 1 - k_{rl} & \text{if } S_{gr} = 0 \\ (1 - \hat{S})^2 (1 - \hat{S}^2) & \text{if } S_{gr} > 0 \end{cases}$$

152

153 where

154

$$S^* = \frac{S_l - S_{lr}}{S_{ls} - S_{lr}} \quad (3)$$

$$\hat{S} = \frac{S_l - S_{lr}}{1 - S_{lr} - S_{gr}} \quad (4)$$

155

156 where k_{rl} is relative permeability of the liquid phase, k_{rg} is relative permeability of the gas phase, λ_k is a
 157 van Genuchten parameter (in the original notation [32] λ_k is expressed as m), S_l is liquid phase saturation,
 158 S_{lr} is residual (i.e., minimum) liquid phase saturation, S_{ls} is saturated (i.e., maximum) liquid phase
 159 saturation, and S_{gr} is residual gas phase saturation. It is noted that $0 \leq k_{rl} \leq 1$ and $0 \leq k_{rg} \leq 1$ are
 160 imposed.

161

162 The following relative permeability function is used for the bentonite buffer:

163

$$\begin{aligned} k_{rl} &= \hat{S}^3 \\ k_{rg} &= 1 \end{aligned} \quad (5)$$

164

165 As to the calculation of capillary pressure, the following equation (i.e., van Genuchten model) is used for
 166 both rock and bentonite buffer:

167

$$P_{cap} = -P_0 \left((S^*)^{\frac{-1}{\lambda_k}} - 1 \right)^{1-\lambda_k} \quad (6)$$

168

169 where P_0 is a parameter analogous to the air entry pressure. The maximum capillary pressure value, P_{max} ,
 170 is specified such that $-P_{max} \leq P_{cap} \leq 0$.

171

172 Fourier's law and Fick's law are used to simulate thermal conduction and multi-phase diffusion,
173 respectively.

174

175 Table 1 shows the TH parameter values (i.e., TOUGH3 input parameter values) employed in this study for
176 the rock, buffer, and canister. It is noted that the residual liquid saturation value for the capillary pressure
177 function is set smaller than that for the relative permeability function. This is done so as to avoid incurring
178 infinite capillary pressure values when the liquid phase becomes immobile. In addition, all phases are
179 assumed perfectly mobile ($k_{rl} = k_{rg} = 1$) in the canister and identical capillary pressure function and
180 parameter values that are employed for the buffer are used for the canister. This is done to alleviate
181 convergence issues for the flow calculation at the buffer-canister interface. The hydrological behavior of
182 the buffer would not be erroneously affected by that of the canister as the porosity of the canister is set to a
183 value significantly smaller than that of the rest of the rock and bentonite buffer. The inside part of the
184 canister, where nuclear waste is supposed to be stored, is ignored in the model as such part would not have
185 mechanical resistance or affect fluid flow or heat flux calculations as long as an equivalent heat source is
186 assigned to the canister elements. The canister elements are treated as a porous medium due to the
187 requirements of TOUGH3. The porosity of the canister elements is set to a value that is orders of magnitude
188 smaller than that of the bentonite and rock, in order to mitigate the inconsistency.

189

190 **Table 1 Thermo-hydrological parameter values of the model.**

	Rock	Buffer	Canister
Grain density (kg/m ³)	2700	2700	7800
Porosity (-)	0.15	0.41	0.001
Absolute permeability (m ²)	$5 \cdot 10^{-20}$	$2 \cdot 10^{-21}$	$2 \cdot 10^{-21}$
Thermal conductivity (W/(m·°C))	2.2	1.26	12.0

Grain specific heat (J/(kg·°C))	900	800	500
Pore compressibility (1/Pa)	$1 \cdot 10^{-9}$	$5 \cdot 10^{-8}$	0
Pore expansivity (1/°C)	0	$1 \cdot 10^{-4}$	0
Diffusion coefficient for water vapor (m ² /s)	$1.73 \cdot 10^{-5}$	$1.73 \cdot 10^{-5}$	$1.73 \cdot 10^{-5}$
van Genuchten parameter, λ_k (-)	0.41	0.32	0.32
Residual liquid saturation for relative permeability, S_{lr} (-)	0.2	-	-
Residual liquid saturation for capillary pressure, S_{lr} (-)	0.1	0.1	0.1
Saturated liquid saturation, S_{ls} (-)	1.0	1.0	1
Residual gas saturation, S_{gr} (-)	0.01	0	0
Capillary pressure parameter, P_0 (Pa)	$4.76 \cdot 10^7$	$3 \cdot 10^7$	$3 \cdot 10^7$
Maximum capillary pressure, P_{max} (Pa)	$1 \cdot 10^8$	$5 \cdot 10^9$	$5 \cdot 10^9$

191

192 **2.2.2 Mechanical behaviors**

193 The mechanical behavior is modelled by the FLAC3D simulator in the THM coupled simulation case,
 194 whereas they are modelled analytically in the TH coupled simulation case, which will be explained in detail
 195 in later sections. In either case, the stress-strain relation (i.e., constitutive model) is described in terms of
 196 the effective stress, which is defined as follows:

197

$$\sigma'_{ij} = \sigma_{ij} - \alpha_{BW} \cdot p_{pore} \delta_{ij} \tag{7}$$

$$p_{pore} = \max(p_g, p_l)$$

198

199 where σ'_{ij} is the effective stress (tensor), σ_{ij} is the total stress (tensor), α_{BW} is the Biot-Willis coefficient,
 200 δ_{ij} is the Kronecker delta, p_{pore} is pore pressure, p_g is gas phase pressure, and p_l is the liquid phase
 201 pressure. If the pore space is fully saturated with the liquid phase (i.e., water), pore pressure is the liquid
 202 phase pressure, whereas it is the gas phase pressure if the pore space is partially saturated.

203
 204 The linear isotropic elastic constitutive relation is used to model the mechanical behavior of the rock and
 205 canister, where only two elastic parameters (e.g., Young's modulus and Poisson's ratio) are required.

206
 207 The Barcelona Basic Model (BBM) [31] is used to model the mechanical behavior of the bentonite buffer.
 208 This model is effective in simulating the stress and strain change of swelling clays during monotonic
 209 saturation/desaturation, which is expected to occur in the bentonite buffer. The elastic mechanical behavior
 210 of the BBM is nonlinear and plastic yielding and plastic hardening/softening are included. More advanced
 211 models [33], [34] are available in case the stress and strain change during cyclic saturation/desaturation
 212 needs to be simulated. The mechanical parameter values used for the model are listed in Table 2.

213
 214 **Table 2 Mechanical parameter values of the model.**

	Rock	Buffer	Canister
Young's modulus (Pa)	$5.0 \cdot 10^9$	-	$200 \cdot 10^9$
Poisson's ratio (-)	0.3	0.4	0.3
Biot-Willis coefficient, α_{BW} (-)	1.0	1.0	1.0
Linear thermal expansion coefficient ($1/^\circ\text{C}$)	$1.0 \cdot 10^{-5}$	$1.5 \cdot 10^{-4}$	$1.0 \cdot 10^{-5}$
Gradient of swelling line for stress, κ_{PS0} (-)	-	0.05	-
Gradient of swelling line for suction, κ_{SP0} (-)	-	0.25	-

Swelling gradient adjusting parameter, α_{PS} (1/Pa)	-	$-3.0 \cdot 10^{-9}$	-
Swelling gradient adjusting parameter, α_{SS} (1/Pa)	-	0	-
Swelling gradient adjusting parameter, α_{SP} (-)	-	-0.161	-
Reference pressure, p_{ref} (Pa)	-	$0.01 \cdot 10^6$	-
Gradient of compression line for stress, λ_{PS0} (-)	-	0.15	-
Gradient of compression line for suction, λ_{SP0} (-)	-	0.5	-
Compression gradient adjusting parameter, r_λ (-)	-	0.925	-
Compression gradient adjusting parameter, β_λ (1/Pa)	-	$0.1 \cdot 10^{-6}$	-
Tensile strength at zero suction, p_{s0} (Pa)	-	0	-
Tensile strength gradient, k_s (-)	-	0.1	-
Tensile strength gradient adjusting parameter, ρ_s (1/°C)	-	0	-
Reference mean net stress, p^c (Pa)	-	$0.5 \cdot 10^6$	-
Critical state frictional constant, M (-)	-	1.0	-
Pre-consolidation mean net stress, p_0^* (Pa)	-	$12.0 \cdot 10^6$	-
Porosity (-)	-	0.41	-
Non-associated plastic flow parameter, α_a (-)	-	0.53	-

215

216 **2.3 Simulation stages**

217 **2.3.1 In situ conditions**

218 The in situ conditions of the model prior to the excavation of the nuclear waste disposal tunnel are specified
219 as follows: the temperature is linearly increased with depth from 10°C on the top boundary to 40°C on the
220 bottom boundary, and pore pressure and stresses are also linearly increased with gradients corresponding
221 to the weight of pore water and rock, respectively. The in situ rock stresses are assumed either isotropic or

222 anisotropic in this study, which are explained in detail in later sections. At the in situ state, the volume
223 inside the tunnel is occupied by rock elements while the buffer and canister elements are deactivated.

224

225 **2.3.2 Tunnel excavation stage**

226 The TH part of the tunnel excavation stage is simulated by TOUGH3 as follows: constant pore pressure of
227 0.1 MPa and constant temperature of 25°C are specified in the buffer and canister elements and the
228 TOUGH3 simulation is run for 1.5 years. The pore pressure and temperature profiles at the end of the
229 simulated timeframe are used as the initial pore pressure (p_{pore0}) and temperature (T_0) in the following
230 canister heating stage.

231

232 The mechanical part of the tunnel excavation stage is simulated by FLAC3D as follows: first the stiffness
233 of the buffer and canister elements is set to small values (smaller by orders of magnitudes than that of the
234 rock) so as not to restrict the displacement of the tunnel boundary, and then pore pressure and temperature
235 values at each converged TOUGH3 time step are imported and used as boundary conditions in FLAC3D to
236 compute the corresponding displacements, strains, and stresses. The stress distributions at the end of the
237 1.5-year simulation timeframe are used as the initial stress values for the THM simulation case of the
238 canister heating stage. For the TH simulation case, on the other hand, the initial stresses after tunnel
239 excavation are calculated analytically as described in a later section.

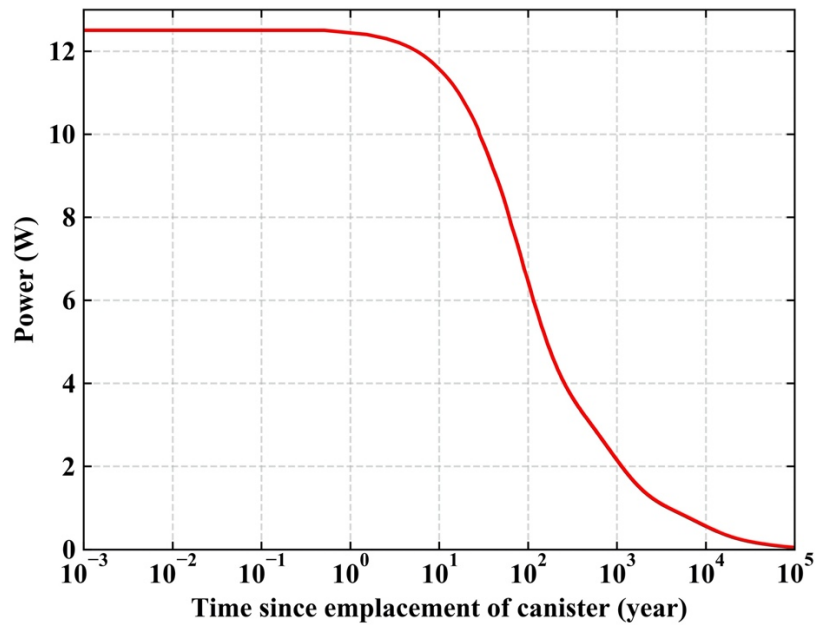
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241 **2.3.3 Canister heating stage**

242 The canister heating stage is simulated by introducing a time-varying heat source in the canister elements
243 in TOUGH3. Figure 3 shows the heat rate assigned to the canister elements. The power curve shown in
244 Figure 3 corresponds to a 4-PWR element waste package which is assumed to be emplaced after 60 years
245 of interim storage with the center-to-center spacing of 9 m between individual waste canisters along the

246 tunnel axis. The initial power is specified as 200 (W/m) where the unit W/m means power per unit length
247 along the axis of the tunnel. The power is divided by the number of the canister elements (i.e., sixteen),
248 which renders the initial power of 12.5 (W/m) for each canister element as shown in Figure 3. Also, the
249 mechanical parameter values of the buffer and canister are changed from the fictitious values used in the
250 excavation stage to the ones listed in Table 2. The initial stresses in the buffer and canister are set as
251 follows: $\sigma_{xx} = \sigma_{yy} = \sigma_{zz} = 0.15$ MPa (compression positive) and $\tau_{xx} = \tau_{yy} = \tau_{zz} = 0$ MPa. The initial
252 pore pressure (gas pressure) and temperature in the buffer and canister are set to 0.1 MPa and 25°C.
253 Finally, the initial liquid phase saturation in the buffer and canister is set to 0.65. The canister heating
254 stage is simulated up to 100,000 years.

255



256

257

Figure 3 Heat rate assigned to the canister elements.

258

259

260 **2.4 Stress calculation methodologies**

261 **2.4.1 Stress approximation using TH variables**

262 In this section, a methodology is developed to estimate stresses from TH variables, such as pore pressure
263 (i.e., gas pressure), temperature, and suction (i.e., the absolute value of capillary pressure). The stress
264 estimates can then be used to calculate stress-induced permeability changes. It is noted that compressive
265 stress is expressed in positive values in the following derivation of the stresses.

266

267 **2.4.1.1 Buffer stresses**

268 The buffer stress in the direction normal to the tunnel surface, p_B , is calculated by the Barcelona Basic
269 Model (BBM) under the assumption that the volume of the buffer (i.e., the space in between the canister
270 and the waste disposal tunnel) remains constant ($d\epsilon_{total} = d\epsilon_{mechanical} + d\epsilon_{suction} + d\epsilon_{thermal} = 0$)
271 and that the strain increments are isotropic (no shear components). The stress-strain relation of BBM is
272 nonlinear, hence the buffer stress has to be calculated iteratively as follows:

273

$$\begin{aligned} p'_{B,i+1} &= p'_{B,i} + \Delta p'_{B,i} \\ \Delta p'_{B,i} &= K'_i (\Delta \epsilon_{s,i} + \Delta \epsilon_{th,i}) \end{aligned} \tag{8}$$

274

275 where p'_B is the mean effective stress in the buffer, $\Delta p'_B$ is the mean effective stress increment in the buffer,
276 K' is the bulk modulus of the buffer, $\Delta \epsilon_s$ is the suction-induced volumetric strain increment, and $\Delta \epsilon_{th}$ is
277 the thermally-induced volumetric strain increment. The subscript i indicates the step number. The bulk
278 modulus and strain increments are calculated as follows:

279

$$K'_i = \frac{v p'_{B,i}}{\kappa_{PS,i}} \quad (9)$$

$$\Delta \epsilon_{s,i} = \frac{-\kappa_{SP,i}}{v(s_i + p_{atm})} (s_{B,i+1} - s_{B,i}) \quad (10)$$

$$\Delta \epsilon_{th,i} = 3\alpha_B (T_{B,i+1} - T_{B,i}) \quad (11)$$

280

281 where v is specific volume of the buffer, κ_{PS} is a stiffness coefficient of the buffer against mean effective
 282 stress change, κ_{SP} is a stiffness coefficient of the buffer against suction change, p_{atm} is the atmospheric
 283 pressure, s_B is suction in the buffer (i.e., the absolute value of capillary pressure), α_B is the linear thermal
 284 expansion coefficient of the buffer, and T_B is temperature in the buffer. The stiffness coefficients are
 285 suction- and/or mean effective stress-dependent as shown below:

286

$$\kappa_{PS,i} = \kappa_{PS0} (1 + \alpha_{PS} s_{B,i}) \quad (12)$$

$$\kappa_{SP,i} = \kappa_{SP0} \left(1 + \alpha_{SP} \ln \frac{p'_{B,i}}{p_{ref}} \right) \exp(\alpha_{SS} s_{B,i}) \quad (13)$$

287

288 where κ_{PS0} is the stiffness coefficient against mean effective stress change at zero suction, κ_{SP0} is the
 289 stiffness coefficient against suction change at zero mean effective stress and zero suction, p_{ref} is a
 290 reference pressure value, and α_{PS} , α_{SP} , α_{SS} are stiffness-adjusting parameters.

291

292 Finally, the mean total stress in the buffer (i.e., buffer stress in the direction normal to the tunnel surface)
 293 is obtained as follows:

294

$$p_{B,i} = p'_{B,i} + p_{pore,i} \quad (14)$$

295

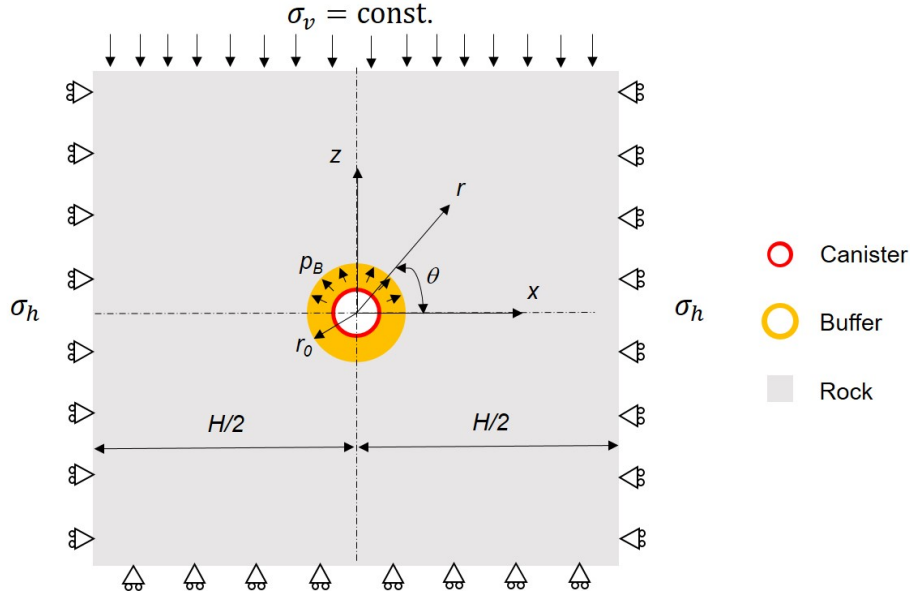
296 where p_{pore} is pore pressure in the buffer.

297

298 **2.4.1.2 Rock stresses**

299 In order to calculate stresses in the formation around the nuclear waste storage tunnel, Kirsch equations are
300 utilized under the analytical model domain and boundary conditions assumed as shown in **Error!**
301 **Reference source not found.** The horizontal displacement on the lateral boundaries are restricted while
302 constant pressure loads, whose magnitude corresponds to the vertical total stress level at the depth of the
303 tunnel, are specified on the top and bottom boundaries. The gravity is ignored in mechanical calculations
304 (but not in thermo-hydraulic calculations) in the model domain. The vertical length of the model is arbitrary
305 but it should be sufficiently larger than the diameter of the tunnel and on the same order as the horizontal
306 length of the model.

307



308

309

Figure 4 Analytical model of the proximity of the nuclear waste disposal tunnel.

310

311 The total stress on the lateral boundary (i.e., mid-pillar) after the emplacement of the canister is calculated

312 from Hooke's law for linear isotropic elastic material under the plane-strain ($d\epsilon_{H,total} = d\epsilon_{H,mech} +$

313 $d\epsilon_{H,thermal} = 0$) and zero-lateral strain ($d\epsilon_{h,total} = d\epsilon_{h,mech} + d\epsilon_{h,thermal} = 0$) conditions as follows:

314

$$\sigma_h = \frac{2\mu}{\lambda + 2\mu} \left(p_{pore} \Big|_{r=\frac{H}{2}, \theta=0} - p_{pore0} \Big|_{r=\frac{H}{2}, \theta=0} \right) + \frac{3\lambda\mu + 2\mu^2}{\lambda + 2\mu} 2\alpha \left(T \Big|_{r=\frac{H}{2}, \theta=0} - T_0 \Big|_{r=\frac{H}{2}, \theta=0} \right) + \sigma_{h0} \quad (15)$$

315

316 where σ_h is the lateral boundary stress, p_{pore} is pore pressure, T is temperature, σ_{h0} , p_{pore0} , and T_0 are the

317 initial lateral boundary stress, pore pressure and temperature, respectively, at the start of the canister heating

318 stage, α is the linear thermal expansion coefficient of the rock. The subscript $\Big|_{r=\frac{H}{2}, \theta=0}$ indicates the

319 specific location of the model from which the pore pressure and temperature values are extracted to be

320 substituted in the equation. The definitions of r, θ, H are provided in Figure 4. The parameter λ and μ are
 321 the Lamé constants, which are defined as follows:

$$\lambda = \frac{E\nu}{(1+\nu)(1-2\nu)}$$

$$\mu = \frac{E}{2(1+\nu)}$$
(16)

323
 324 where E is Young's modulus and ν is Poisson's ratio. The radial and circumferential stresses in the rock
 325 are calculated from the Kirsch equations as follows:

$$\sigma_{rr} = \frac{\sigma_v + \sigma_h}{2} \left(1 - \left(\frac{r_0}{r}\right)^2\right) + \left(\frac{r_0}{r}\right)^2 p_B - \frac{\sigma_v - \sigma_h}{2} \left(1 - 4\left(\frac{r_0}{r}\right)^2 + 3\left(\frac{r_0}{r}\right)^4\right) \cos 2\theta$$
(17)

$$\sigma_{\theta\theta} = \frac{\sigma_v + \sigma_h}{2} \left(1 + \left(\frac{r_0}{r}\right)^2\right) - \left(\frac{r_0}{r}\right)^2 p_B + \frac{\sigma_v - \sigma_h}{2} \left(1 + 3\left(\frac{r_0}{r}\right)^4\right) \cos 2\theta$$

$$+ \frac{E\alpha \left(T - T\big|_{r=\frac{H}{2}, \theta=0}\right)}{(1-\nu)} + \frac{E \left(p_{pore} - p_{pore}\big|_{r=\frac{H}{2}, \theta=0}\right)}{3K(1-\nu)}$$
(18)

$$\tau_{r\theta} = \frac{\sigma_v - \sigma_h}{2} \left(1 + 2\left(\frac{r_0}{r}\right)^2 - 3\left(\frac{r_0}{r}\right)^4\right) \sin 2\theta$$
(19)

329
 330 where r_0 is the radius of the tunnel, and $\sigma_v = \rho_{rock}gL/2$ is the constant vertical stress (ρ_{rock} is the rock
 331 bulk density, and g is the gravitational acceleration). It is noted that the out-of-plane shear stresses

332 $(\tau_{ry}, \tau_{\theta y})$ are assumed to be negligible. The out-of-plane normal stress is calculated from Hooke's law for
 333 linear isotropic elastic material under the plane-strain condition ($d\epsilon_{yy,total} = d\epsilon_{yy,mech} + d\epsilon_{yy,thermal} =$
 334 0) as follows:

$$\begin{aligned} \sigma_{yy} = & \frac{\lambda}{2(\lambda + \mu)} \left((\sigma_{rr} - \sigma_{rr0}) + (\sigma_{\theta\theta} - \sigma_{\theta\theta0}) - 2(p_{pore} - p_{pore0}) - 2\lambda\alpha(T - T_0) \right) \\ & + (\lambda + 2\mu)\alpha(T - T_0) + (\sigma_{yy0} - p_{pore0}) + p_{pore} \end{aligned} \quad (20)$$

336
 337 where p_{pore} is pore pressure, and $T_0, p_{pore0}, \sigma_{rr0}, \sigma_{\theta\theta0}, \sigma_{yy0}$ are the initial temperature, pore pressure,
 338 radial stress, circumferential stress, and out-of-plane stress, respectively, at the start of the canister heating
 339 stage.

340
 341 The initial stresses ($\sigma_{h0}, \sigma_{rr0}, \sigma_{\theta\theta0}, \tau_{r\theta0}, \sigma_{yy0}$) are those after the excavation of the nuclear waste disposal
 342 tunnel, which are calculated as follows:

$$\begin{aligned} \sigma_{h0} = & \frac{2\mu}{\lambda + 2\mu} \left(p_{pore0} \Big|_{r=\frac{H}{2}, \theta=0} - p_{pore, in situ} \Big|_{r=\frac{H}{2}, \theta=0} \right) \\ & + \frac{3\lambda\mu + 2\mu^2}{\lambda + 2\mu} 2\alpha \left(T_0 \Big|_{r=\frac{H}{2}, \theta=0} - T_{in situ} \Big|_{r=\frac{H}{2}, \theta=0} \right) + \sigma_{h, in situ} \end{aligned} \quad (21)$$

344

$$\begin{aligned} \sigma_{rr0} = & \frac{\sigma_v + \sigma_{h0}}{2} \left(1 - \left(\frac{r_0}{r} \right)^2 \right) + \left(\frac{r_0}{r} \right)^2 p_{tunnel} \\ & - \frac{\sigma_v - \sigma_{h0}}{2} \left(1 - 4 \left(\frac{r_0}{r} \right)^2 + 3 \left(\frac{r_0}{r} \right)^4 \right) \cos 2\theta \end{aligned} \quad (22)$$

345

$$\begin{aligned} \sigma_{\theta\theta} = & \frac{\sigma_v + \sigma_{h0}}{2} \left(1 + \left(\frac{r_0}{r} \right)^2 \right) - \left(\frac{r_0}{r} \right)^2 p_{tunnel} + \frac{\sigma_v - \sigma_{h0}}{2} \left(1 + 3 \left(\frac{r_0}{r} \right)^4 \right) \cos 2\theta \\ & + \frac{E\alpha \left(T_0 - T_0 \Big|_{r=\frac{H}{2}, \theta=0} \right)}{(1-\nu)} + \frac{E \left(p_{pore0} - p_{pore0} \Big|_{r=\frac{H}{2}, \theta=0} \right)}{3K(1-\nu)} \end{aligned} \quad (23)$$

346

$$\tau_{r\theta} = \frac{\sigma_v - \sigma_{h0}}{2} \left(1 + 2 \left(\frac{r_0}{r} \right)^2 - 3 \left(\frac{r_0}{r} \right)^4 \right) \sin 2\theta \quad (24)$$

347

$$\begin{aligned} \sigma_{yy0} = & \frac{\lambda}{2(\lambda + \mu)} \left((\sigma_{rr0} - \sigma_{rr, in situ}) + (\sigma_{\theta\theta0} - \sigma_{\theta\theta, in situ}) - 2(p_{pore0} - p_{pore, in situ}) \right. \\ & \left. - 2\lambda\alpha(T_0 - T_{in situ}) \right) + (\lambda + 2\mu)\alpha(T_0 - T_{in situ}) \\ & + (\sigma_{yy, in situ} - p_{pore, in situ}) + p_{pore0} \end{aligned} \quad (25)$$

348

349 where p_{tunnel} is the (atmospheric) pressure inside the tunnel, and the subscript $in situ$ indicates the in situ

350 (i.e., hydrostatic, geostatic) state prior to the excavation of the disposal tunnel. The in situ stresses are

351 calculated as follows:

352

$$\begin{pmatrix} \sigma_{rr,in\ situ} \\ \sigma_{\theta\theta,in\ situ} \\ \tau_{r\theta,in\ situ} \\ \sigma_{yy,in\ situ} \end{pmatrix} = \begin{pmatrix} \sigma_{xx} \cos^2 \theta + \sigma_{zz} \sin^2 \theta \\ \sigma_{xx} \sin^2 \theta + \sigma_{zz} \cos^2 \theta \\ (-\sigma_{xx} + \sigma_{zz}) \sin \theta \cos \theta + \tau_{xz} \\ \sigma_{yy} \end{pmatrix} \quad (26)$$

353

354 where

355

$$\begin{pmatrix} \sigma_{xx} \\ \sigma_{zz} \\ \tau_{xz} \\ \sigma_{yy} \end{pmatrix} = \begin{pmatrix} K_0 (\rho_{rock} \cdot g \cdot (L/2 - z) - p_{pore,in\ situ}) + p_{pore,in\ situ} \\ \rho_{rock} \cdot g \cdot (L/2 - z) \\ 0 \\ K_0 (\rho_{rock} \cdot g \cdot (L/2 - z) - p_{pore,in\ situ}) + p_{pore,in\ situ} \end{pmatrix} \quad (27)$$

356

357 where ρ_{rock} is the bulk density of rock, g is the gravitational acceleration, z is the vertical distance from
 358 the center of the tunnel (upward is positive), $L/2$ is the depth from the ground surface to the center of the
 359 tunnel, and K_0 is the lateral earth pressure coefficient. The out-of-plane shear stresses (τ_{xy}, τ_{yz}) are
 360 assumed to be nonexistent. It is noted that $\sigma_{h,in\ situ}$ is assumed equal to σ_{xx} at the depth of the center of the
 361 tunnel as shown below:

362

$$\sigma_{h,in\ situ} = K_0 (\rho_{rock} \cdot g \cdot (L/2) - p_{pore,in\ situ}) + p_{pore,in\ situ} \quad (28)$$

363

364 The in situ (i.e., hydrostatic) pore pressure and (i.e., thermostatic) temperature prior to the tunnel excavation
 365 stage are calculated as follows:

366

$$p_{pore,in\ situ} = \rho_{water} \cdot g \cdot (L/2 - z) \quad (29)$$

$$T_{in\ situ} = \frac{T_{bottom} - T_{top}}{L} (L/2 - z) + T_{top} \quad (30)$$

367

368 where ρ_{water} is the density of pore water, and T_{top} and T_{bottom} are the static temperature at $z = \frac{L}{2}, -\frac{L}{2}$,
 369 respectively.

370

371 **2.4.2 Stress calculation using THM coupling**

372 The THM coupled simulation is implemented by the TOUGH-FLAC simulator, in order to evaluate the
 373 accuracy of stresses estimated by the TH stress approximation method introduced earlier. In this THM
 374 coupled stress calculation method, the stresses are calculated numerically by FLAC3D.

375

376 **2.5 Simulation cases**

377 **2.5.1 Isotropic rock stress case**

378 In this simulation case, in situ rock stresses prior to the tunnel excavation stage are assumed isotropic as
 379 shown in the equation below:

380

$$\sigma'_{xx} = \sigma'_{yy} = \sigma'_{zz} = (\rho_{rock} - \rho_{water})g(L/2 - z) \quad (31)$$

381

382 where $\sigma'_{xx}, \sigma'_{yy}, \sigma'_{zz}$ are the effective stresses in the horizontal, out-of-plane, and vertical directions of the
 383 model, respectively, ρ_{rock} is the rock bulk density, ρ_{water} is the water bulk density, g is the gravitational
 384 acceleration, and z is the vertical distance from the center of the tunnel. The shear stresses are assumed

385 nonexistent. The rock stresses are assumed isotropic because the modelled nuclear waste repository is
 386 generic (no specific site is intended). The objective of this study is to demonstrate the effectiveness of the
 387 stress approximation method using TH variables, and hence the in situ rock stresses are assumed the
 388 simplest in this baseline simulation case. More realistic rock stresses (i.e., anisotropic rock stresses) are
 389 considered in the following simulation case described below.

390

391 **2.5.2 Anisotropic rock stress case**

392 The anisotropic stress state of the clay/shale formation is referred from Mont Terri Rock Laboratory in
 393 Switzerland [35]. The rock at Mont Terri Rock Laboratory consists of the Opalinus Clay, which is one of
 394 the most ideal materials for nuclear waste repositories due to its extremely low permeability. A gallery
 395 tunnel at Mont Terri was excavated in 1998, which was located at the depth of roughly 250 m from the
 396 ground surface, in order to investigate the effect of excavation on the fracture development around the
 397 tunnel in the Opalinus Clay. The in situ stress state around the tunnel was measured by triaxial strain cells
 398 and the measurement indicated the following stress state: $\sigma'_{xx} = -1.4 \sim 0.2$ (MPa), $\sigma'_{yy} = 2.0$ (MPa), $\sigma'_{zz} =$
 399 4.5 (MPa) [36], [37]. The horizontal stress estimate (σ'_{xx}) is uncertain because the tunnel was located in the
 400 vicinity of a valley. In this study, $\sigma'_{xx} = 0.2$ (MPa) is assumed. As a result, the lateral earth pressure
 401 coefficients are obtained as follows: $K_{0,yy} = \sigma'_{yy}/\sigma'_{zz} = 0.44$, $K_{0,xx} = \sigma'_{xx}/\sigma'_{zz} = 0.044$. The
 402 extrapolation of this stress state at Mont Terri to the generic repository model employed in this study is
 403 carried out as shown below:

404

$$\begin{aligned}
 \sigma'_{zz} &= (\rho_{rock} - \rho_{water})g(L/2 - z) \\
 \sigma'_{xx} &= K_{0,xx}\sigma'_{zz} \\
 \sigma'_{yy} &= K_{0,yy}\sigma'_{zz}
 \end{aligned}
 \tag{32}$$

405

406 The shear stresses are assumed nonexistent. The above stresses are used as the in situ stress state for the
407 anisotropic rock stress case.

408

409 **2.6 Permeability profiles in rock near the tunnel**

410 In this study, rock permeability near the disposal tunnel is calculated by the following equation:

411

$$k = \begin{cases} (k_r + \Delta k_{max} \exp(\beta_1 \sigma'_m)) \cdot \exp(\gamma \langle \sigma_d - \sigma_{d,crit} \rangle) & \text{if } r \leq 2r_0 \\ k_r & \text{if } r > 2r_0 \end{cases} \quad (33)$$

412

413 where k is the absolute permeability, Δk_{max} is the increment in absolute permeability due to mean effective
414 stress change, $\beta_1 (< 0)$ is a parameter correlating the absolute permeability and mean effective stress
415 change, σ'_m is the mean effective stress (= mean total stress – pore pressure), $\gamma (> 0)$ is a parameter
416 correlating the absolute permeability and deviatoric stress change, σ_d is the deviatoric stress (= von Mises
417 equivalent stress), and $\sigma_{d,crit}$ is the critical deviatoric stress value. The $\langle \ \rangle$ operator is the MacCaulay
418 brackets, in which $\langle x \rangle = x$ if $x \geq 0$ and $\langle x \rangle = 0$ if $x < 0$. The mean effective stress and deviatoric stress
419 are defined as follows:

420

$$\sigma'_m = \frac{\sigma'_{xx} + \sigma'_{yy} + \sigma'_{zz}}{3} = \frac{\sigma'_{rr} + \sigma'_{\theta\theta} + \sigma'_{\theta\theta}}{3} \quad (34)$$

$$\sigma_d = \sqrt{\frac{1}{2} \left((\sigma_{xx} - \sigma_{yy})^2 + (\sigma_{yy} - \sigma_{zz})^2 + (\sigma_{zz} - \sigma_{xx})^2 + 6(\tau_{xy}^2 + \tau_{yz}^2 + \tau_{xz}^2) \right)}$$

$$= \sqrt{\frac{1}{2} \left((\sigma_{rr} - \sigma_{yy})^2 + (\sigma_{yy} - \sigma_{\theta\theta})^2 + (\sigma_{\theta\theta} - \sigma_{rr})^2 + 6(\tau_{ry}^2 + \tau_{y\theta}^2 + \tau_{r\theta}^2) \right)}$$
(35)

421

422 The excavation damaged zone (EDZ), where the permeability increases according to Equation 33, is
 423 assumed to be located within the area enclosed by twice the radius of the tunnel ($2r_0$). This size of EDZ is
 424 referred from experimental observations at Mont Terri Rock Laboratory [35]. Figure 5a and b show the
 425 absolute permeability profiles in the rock around the tunnel after the tunnel excavation stage for the
 426 isotropic and anisotropic rock stress cases, respectively. The permeability values are calculated with the
 427 parameter values listed in Table 3. The permeability increases near the tunnel due mainly to deviatoric
 428 stress increase for the isotropic rock stress case and to the combination of mean stress decrease and
 429 deviatoric stress increase for the anisotropic rock stress case. The size of EDZ is assumed constant in the
 430 canister heating stage.

431

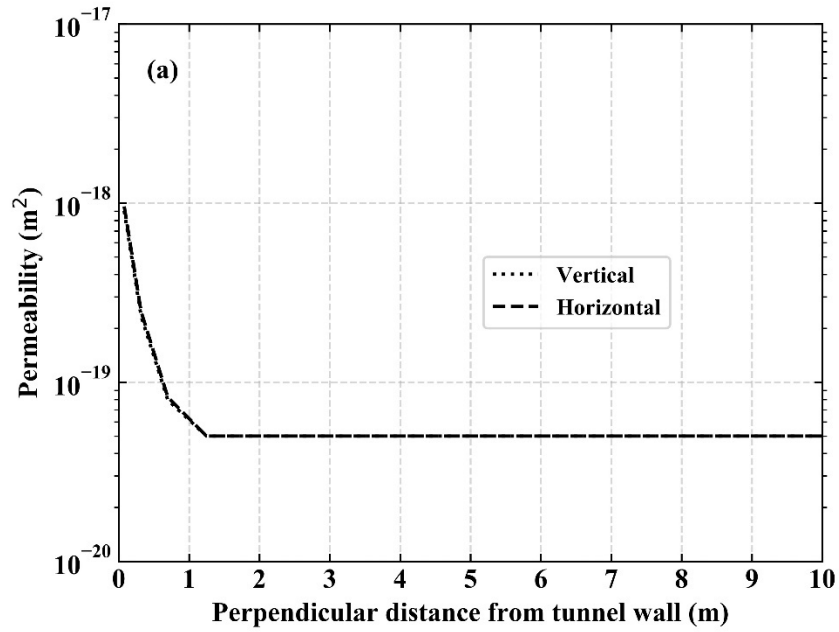
432

Table 3 Parameter values for the rock permeability function.

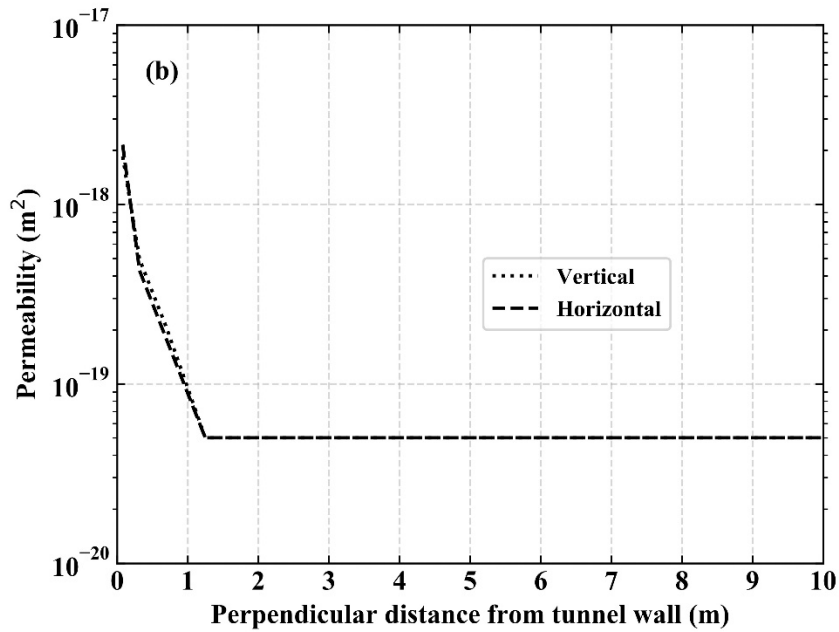
Residual permeability, k_r (m^2)	$5 \cdot 10^{-20}$
Permeability increment, Δk_{max} (m^2)	$1 \cdot 10^{-17}$
Coefficient for mean effective stress, β_l (1/Pa)	$-1 \cdot 10^{-6}$
Coefficient for deviatoric stress, γ (1/Pa)	$3 \cdot 10^{-7}$
Critical deviatoric stress, $\sigma_{d, crit}$ (Pa)	$5 \cdot 10^6$

433

434



435



436

437 **Figure 5 Absolute permeability profiles around the tunnel after excavation: (a) isotropic rock stress**

438

case; (b) anisotropic rock stress case.

439

440

3. Results

441

3.1 TH variables in the buffer

442

Figure 6 shows the evolution of TH variables (i.e., temperature, pore pressure, saturation, and capillary

443

pressure) in the buffer during the canister heating stage. These time-varying profiles are taken from the

444

buffer element located at the mid-thickness along the vertical axis ($r = 0.8 \text{ m}, \theta = 90^\circ$).

445

446

Results are found identical between the TH and THM simulation cases (i.e., the solid (THM) and dashed

447

(TH) lines are on top of each other), which indicates that the variance in stress-induced permeability

448

change in the rock between the two cases has negligible effects on the fluid flow and heat transfer in the

449

buffer. No difference is found between the isotropic and anisotropic rock stress cases either. Hence, only

450

the result of the isotropic case is shown in Figure 6.

451

452

The maximum temperature in the buffer is approximately 85°C which is reached between 10 and 100

453

years since the canister emplacement. Around the same time, the buffer becomes fully saturated with

454

water and pore pressure starts to increase significantly. Thermal expansion of the pore water (i.e., thermal

455

pressurization) in the host rock causes the pore pressure buildup. Pore pressure still keeps increasing

456

while temperature starts to decrease from its peak value at roughly 100 years. During this time, the pore

457

pressure buildup is driven by the thermal pressurization of the surrounding rock, where the temperature

458

peak occurs later than in the buffer due to slow heat transfer. Pore pressure peaks at slightly over 11 MPa,

459

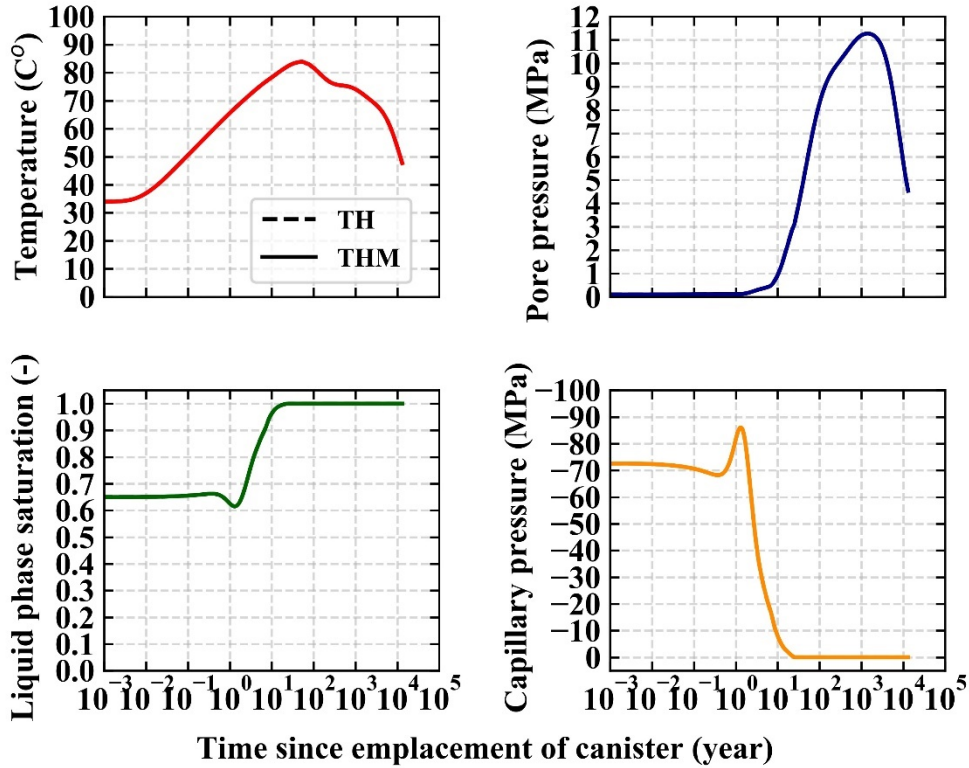
which is significantly greater than the hydrostatic pore pressure level of 4.5 MPa at this depth, and it starts

460

to decrease when thermal depressurization starts to occur in the surrounding rock at approximately 1000

461

years.



463

464 **Figure 6** The evolution of TH variables (i.e., temperature, pore pressure, saturation, and capillary
465 pressure) in the buffer during the canister heating stage.

466

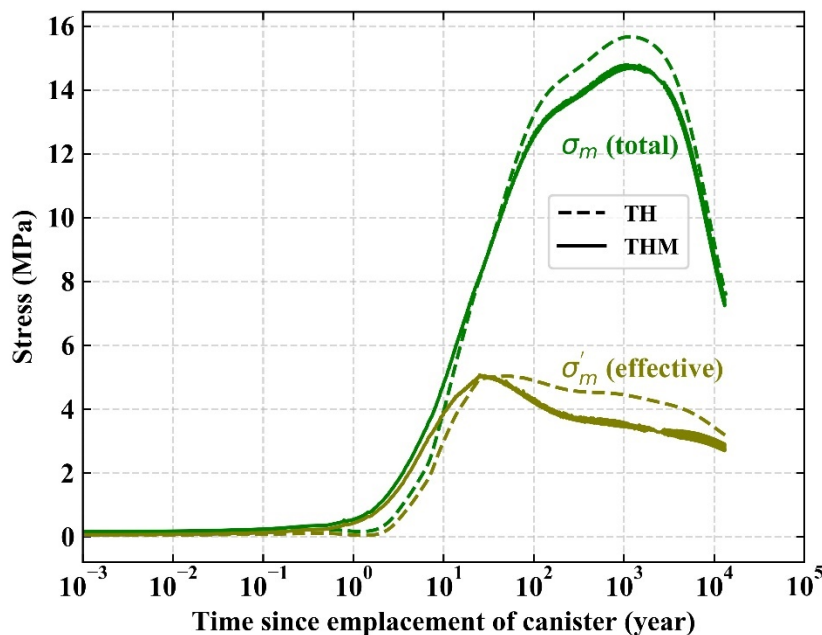
467 **3.2 Buffer stresses**

468 Figure 7 shows the evolution of the mean total and effective stresses in the buffer during the canister heating
469 stage (positive stress values indicate compression). The match between the TH and THM simulation cases
470 is excellent, validating the stress approximation methodology developed for the TH simulation case. The
471 mean effective stress change is driven by the suction-induced swelling, and hence it corresponds with the
472 change in saturation (capillary pressure). Thermal expansion and contraction also contribute to the mean
473 effective stress change but the impact of the thermal effect is secondary to that of the swelling effect. The
474 peak total stress level of 15 MPa is reached at roughly 1000 years, which is slightly above the in situ total

475 stress level at the tunnel depth before tunnel excavation. It is noted that no difference in the mean total and
476 effective stress levels is calculated between the isotropic and anisotropic rock stress cases, hence only the
477 result of the isotropic case is presented in Figure 7.

478
479 The deviation in buffer stress values between the TH and THM cases is caused primarily by the selection
480 of buffer elements whose TH variables are used to calculate buffer stresses in the TH case. In this study,
481 one buffer element at the mid-length in the thickness direction is chosen for simplicity reasons; averaging
482 over two elements (one located by the tunnel wall and the other by the canister wall) did not improve the
483 accuracy of buffer stress estimation. This is an acceptable simplification since radionuclide transport
484 models (e.g., PFLOTRAN models) cannot usually afford fine meshing for the buffer (e.g., one or two
485 elements in the thickness direction). Thus, selecting one buffer element for its TH variables to calculate
486 buffer stresses is suitable.

487



488

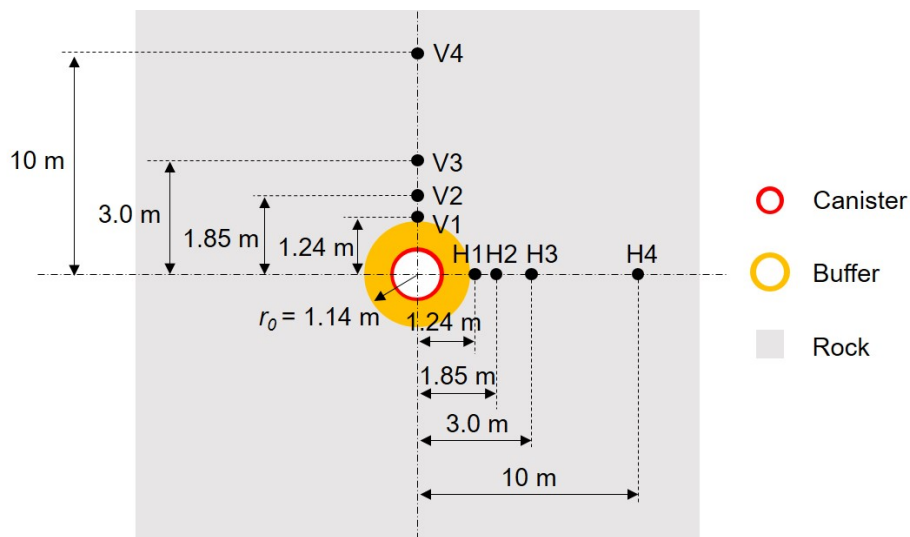
489 **Figure 7 The mean total and effective stress change in the buffer during the canister heating stage.**

490

491 3.3 Rock stresses

492 Rock stresses (i.e., radial effective stress, σ'_{rr} , circumferential effective stress, $\sigma'_{\theta\theta}$, and out-of-plane
493 effective stress, σ'_{yy}) are compared at eight different locations near the tunnel: four locations on the vertical
494 axis (V1, V2, V3, V4) and the other four locations on the horizontal axis (H1, H2, H3, H4). These locations
495 are shown in Figure 8. It is noted that positive stress values indicate compression.

496



497

498 **Figure 8 Locations in the rock where stress data are extracted for comparison.**

499

500 Figure 9a and b show the evolution of the effective stresses in the rock during the canister heating stage
501 for the isotropic and anisotropic rock stress cases, respectively. The in-plane shear stress change is not
502 provided because its magnitude at these locations are found to be a couple of orders of magnitude smaller
503 than that of the other stresses.

504

505 Excellent matches are achieved between the coupled THM simulation case and the developed stress
506 approximation method in the TH simulation case, in both the isotropic and anisotropic rock stress cases.
507 This shows the effectiveness of the developed stress approximation method to estimate stress changes in
508 the rock in a geologic nuclear waste repository using only TH variables.

509

510 Largest errors are calculated at the location V1 and H1. For example, in the isotropic stress case, the
511 largest error occurs in approximately two years since canister emplacement where -56%, 12% and 1%
512 errors are generated in the radial, circumferential, and out-of-plane effective stress estimates, respectively,
513 at both V1 and H1 locations. These errors correspond to errors in buffer stress estimates where mean total
514 and mean net stresses are miscalculated by -73% and -87%, respectively. However, such errors diminish
515 in 10 years, after which is more critical to PA and radionuclide transport than the earlier timeframe.
516 Therefore, these errors are not significant.

517

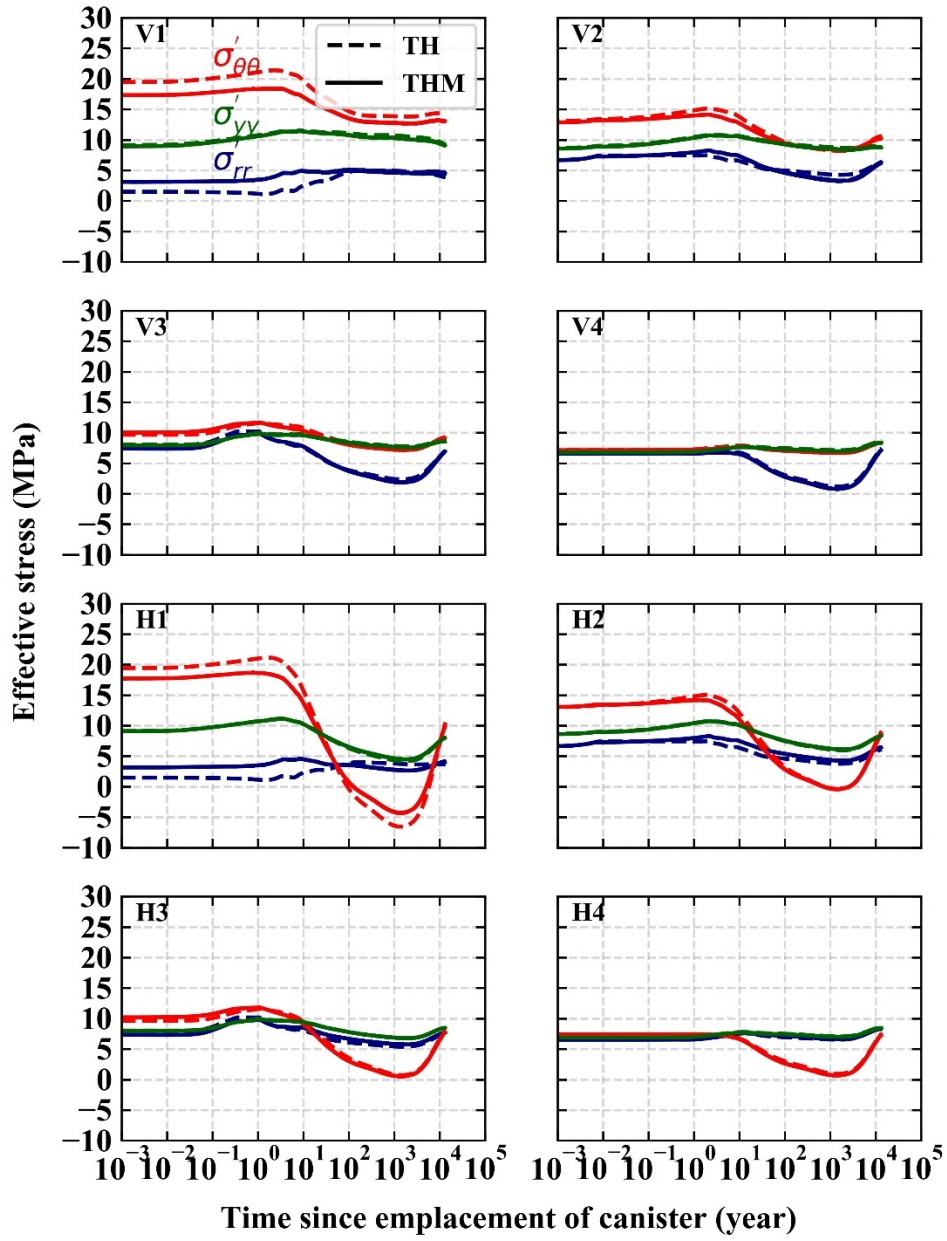
518 The stresses are initially anisotropic near the tunnel for the isotropic rock stress case ($\sigma'_{rr} = 3$ MPa,
519 $\sigma'_{yy} = 10$ MPa, $\sigma'_{\theta\theta} = 17$ MPa at location V1), whereas they are close to an isotropic stress state for the
520 anisotropic rock stress case ($\sigma'_{rr} = 0.5$ MPa, $\sigma'_{yy} = 2$ MPa, $\sigma'_{\theta\theta} = 3$ MPa at location V1). This is induced
521 by the cylindrical cavity contraction during the tunnel excavation stage; the tunnel wall relaxes toward the
522 center of the tunnel during excavation, which reduces the radial effective stress, while at the same time
523 increases the circumferential effective stress.

524

525 During the canister heating stage, on the other hand, the cavity expansion occurs as the bentonite buffer
526 swells in the confined space in the tunnel. The cavity expansion, combined with thermal pressurization,
527 causes a significant reduction in the circumferential effective stress level at the H1 location. The radial
528 and out-of-plane effective stress levels are not affected significantly as cavity expansion increases the

529 radial and out-of-plane total stress levels which compensates for the pore pressure buildup from thermal
530 pressurization.

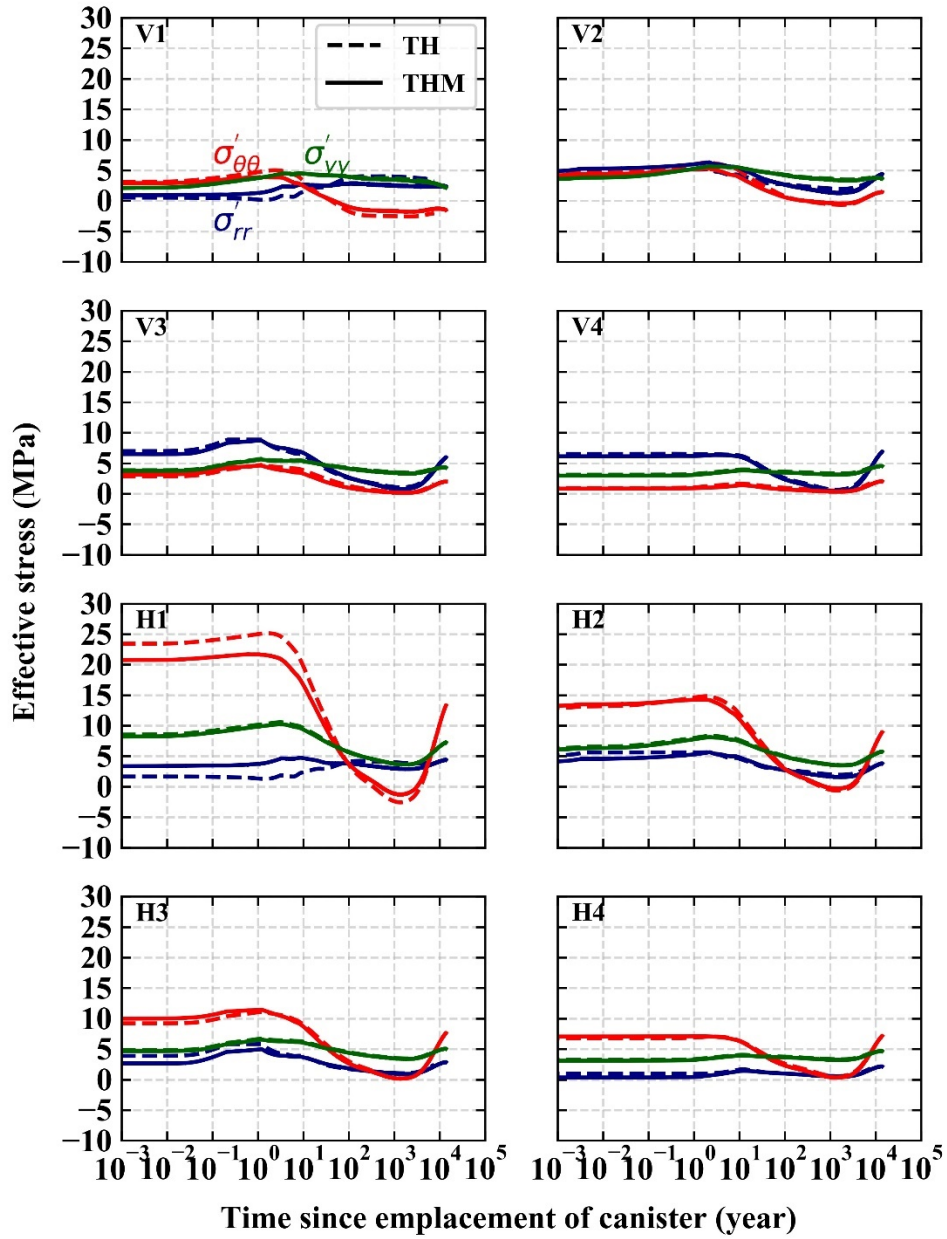
531



532

533

(a)



534

535

(b)

536

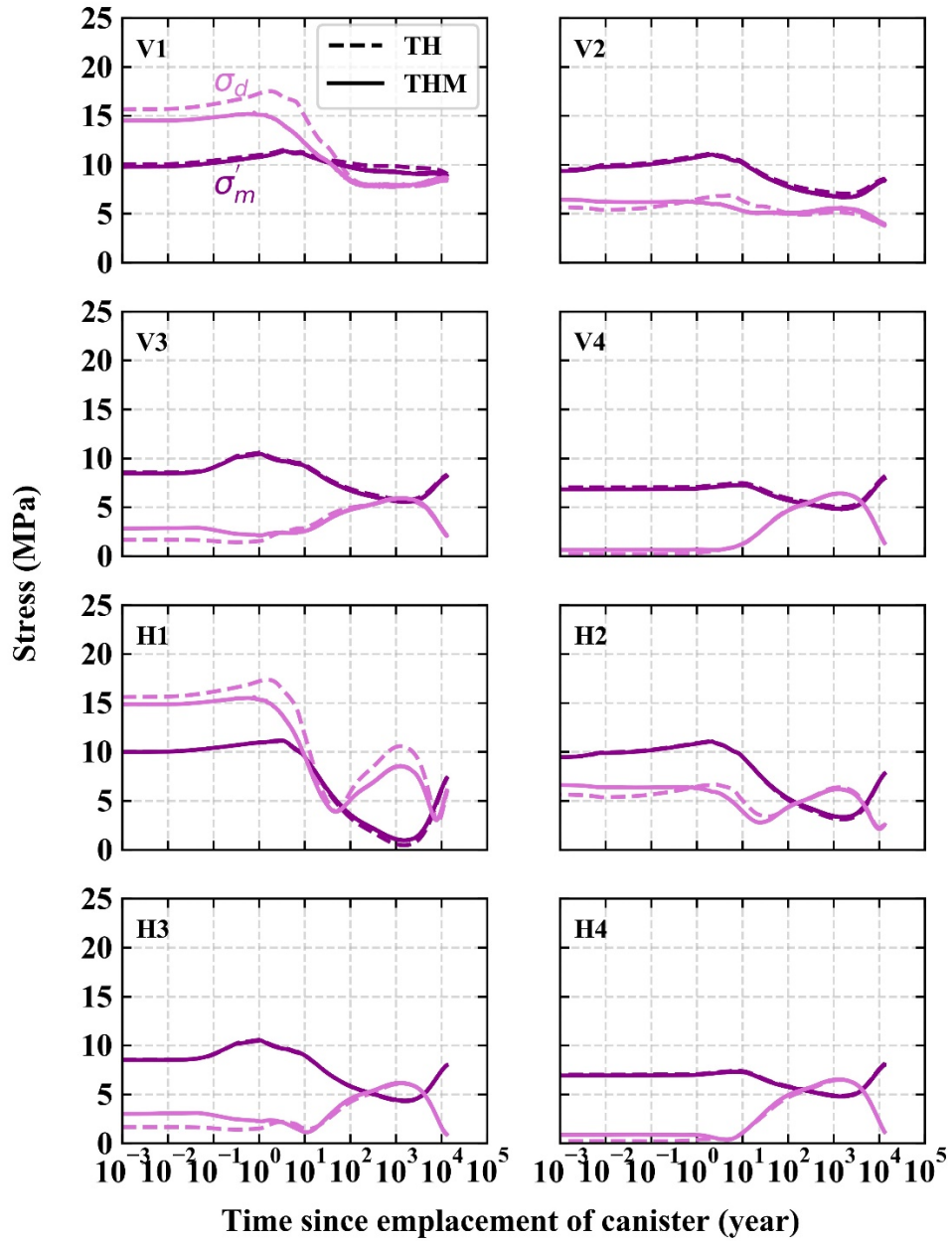
Figure 9 Effective stress changes in the rock near the tunnel during the canister heating stage: (a)

537

isotropic rock stress case; (b) anisotropic rock stress case.

538

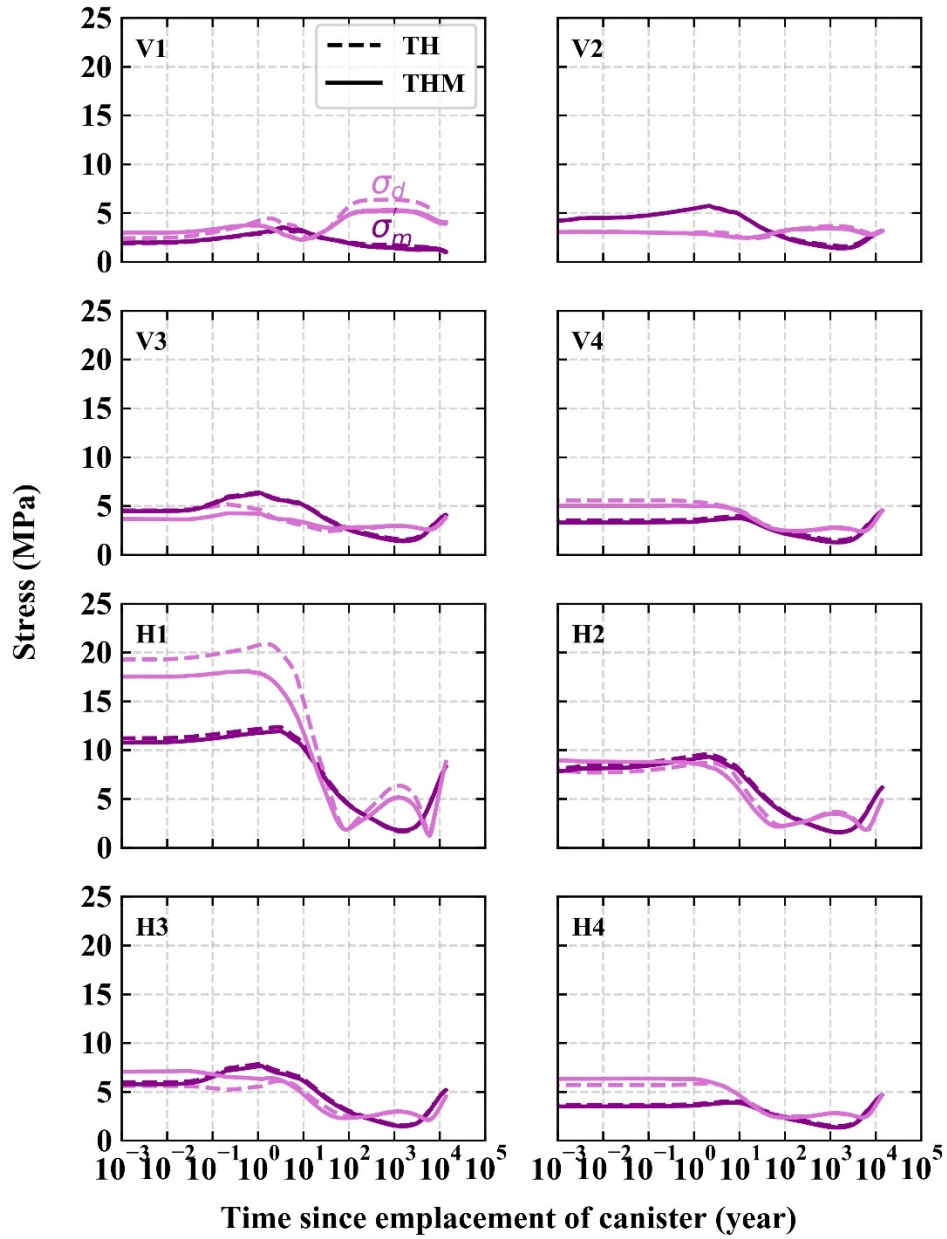
539 Figure 10a and b show the change in mean effective and deviatoric stresses during the canister heating
540 stage for the isotropic and anisotropic rock stress cases, respectively. The match between the coupled
541 THM simulation case and developed stress approximation method in the TH simulation case is excellent
542 in both cases. The mean effective stress change in the TH case, for example, seems to match perfectly
543 with that in the THM case at all examined locations in the rock, whereas greater errors are generated in
544 the deviatoric stress change near the tunnel wall (location V1, V2, H1, H2). This is because errors in the
545 radial, circumferential, and out-of-plane effective stresses are averaged out in the calculation of mean
546 effective stress, whereas they are magnified in the calculation of deviatoric stress. Nevertheless, errors in
547 the deviatoric stress change diminish in 10 years, and hence such errors at the early timeframe would not
548 be critical for the long-term radionuclide transport.
549



550

551

(a)



552

553

(b)

554

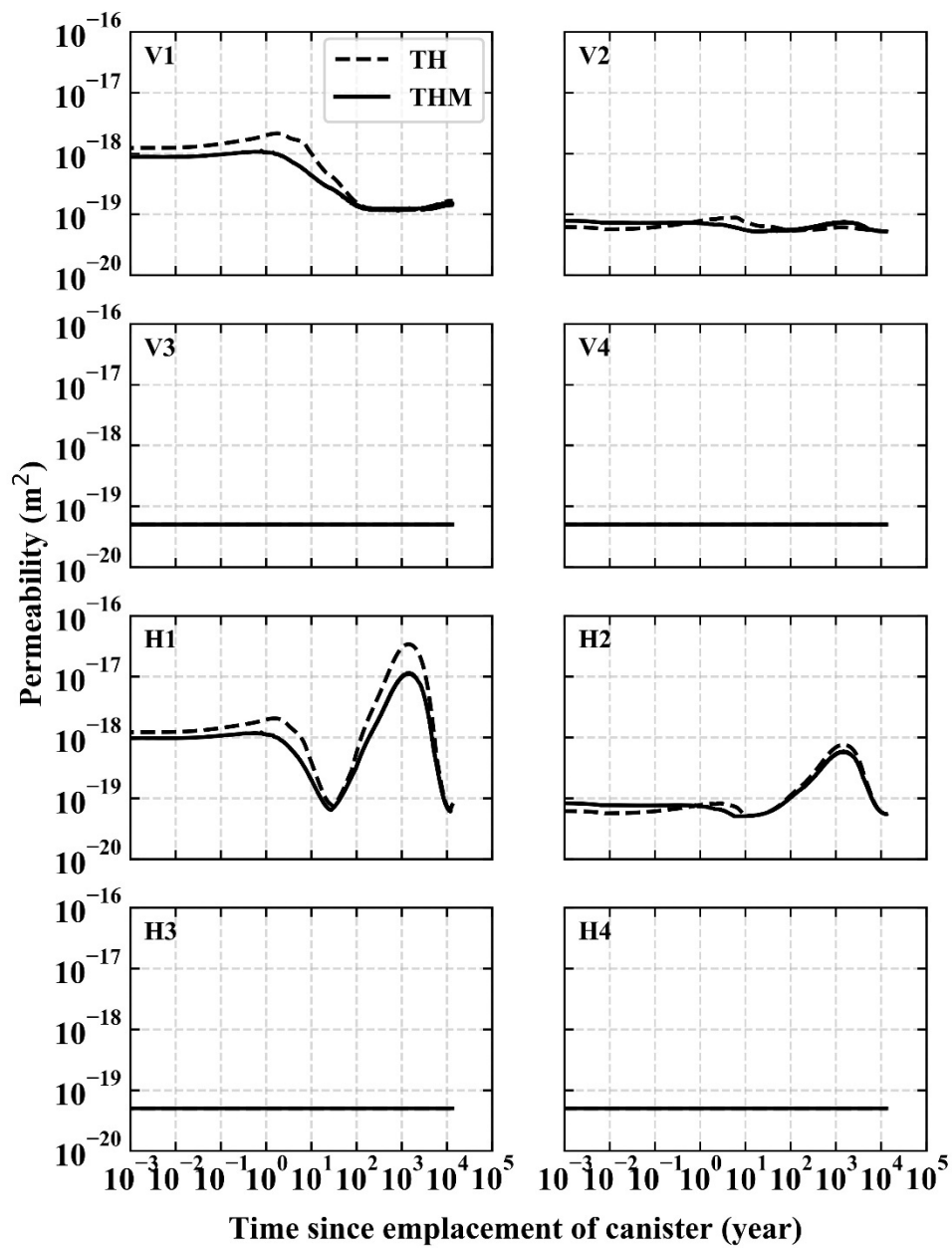
Figure 10 Mean effective and deviatoric stress changes in the rock near the tunnel during the

555

canister heating stage: (a) isotropic rock stress case; (b) anisotropic rock stress case.

556

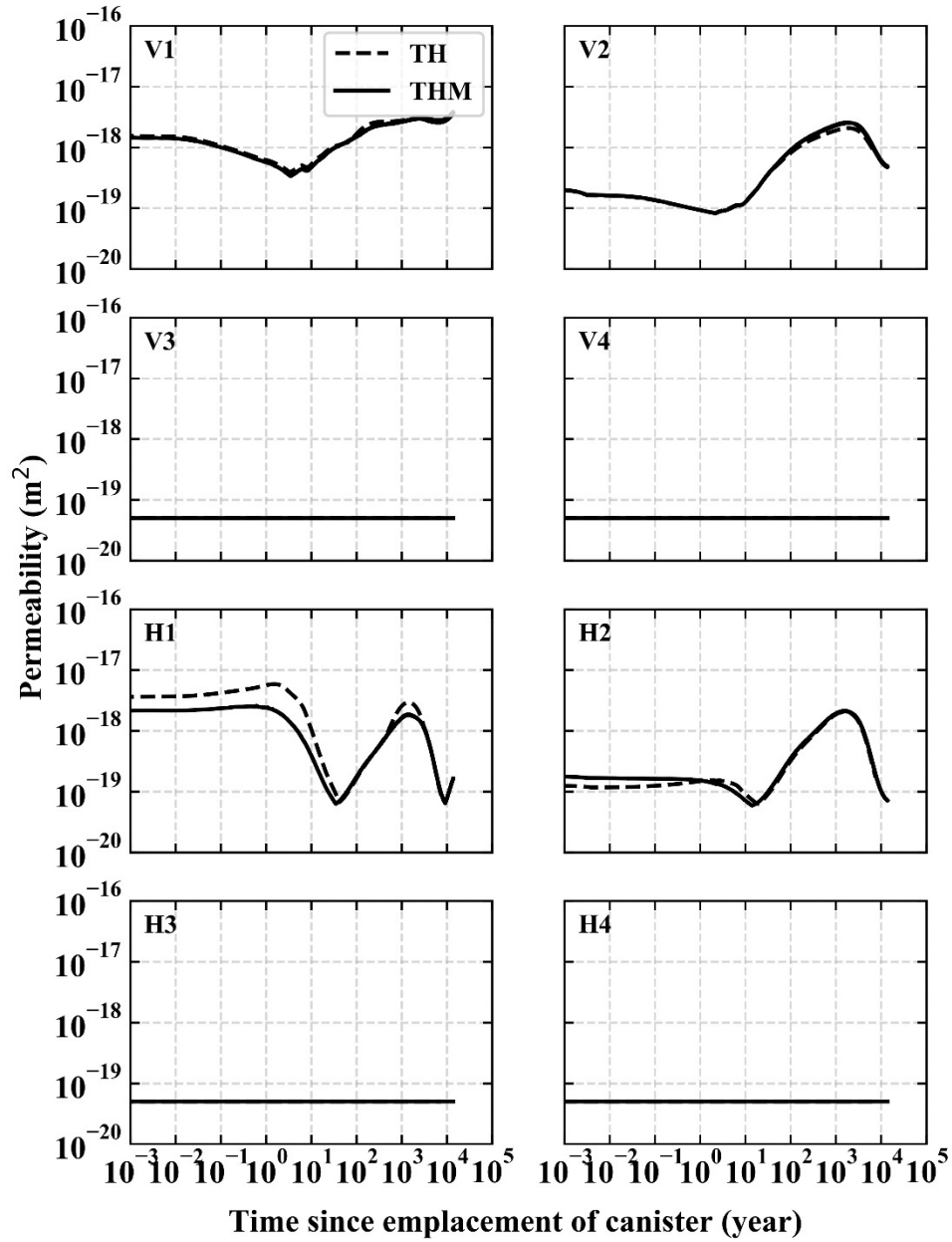
3.4 Rock permeability



558

559

(a)



560

561

(b)

562 **Figure 11 Absolute permeability changes in the rock near the during the canister heating stage: (a)**

563

isotropic rock stress case; (b) anisotropic rock stress case.

564

565 Figure 11a and b show stress-induced absolute permeability changes in the rock near the tunnel for the

566

isotropic and anisotropic rock stress cases, respectively. Stress-induced permeability changes occur only at

567 locations within the excavation damaged zone (V1, V2, H1, H2), which are defined as the area inside twice
568 the radius of the tunnel in this study. At these locations, permeability values are calculated higher than the
569 in situ rock permeability ($5 \cdot 10^{-20} \text{ m}^2$) due to stress-induced fracture opening.

570

571 The permeability value at the H1 location decreases initially by a couple of magnitudes (1-20 years) but it
572 then increases by two orders of magnitude (20-1000 years) in the isotropic rock stress case. The initial
573 permeability decrease is caused by the buffer swelling-induced pressure increase normal to the tunnel wall,
574 which decreases the deviatoric stress level via cavity expansion, whereas the following permeability
575 increase is due to the combination of thermal pressurization-induced mean effective stress decrease and
576 further cavity expansion which increases the deviatoric stress level. Similar permeability changes occur in
577 the anisotropic rock stress case according to the same mechanism at the H1 location.

578

579 At the V1 location, the permeability change is governed by the deviatoric stress change in the isotropic rock
580 stress case, while it is controlled by the mean effective stress change in the anisotropic rock stress case.
581 This is because the deviatoric stress level increases (from the initial zero value) and exceeds the critical
582 level of 5 MPa after tunnel excavation in the isotropic stress case, whereas it decreases (from the initial
583 nonzero value) and drops below the critical level in the anisotropic stress case. As a result, thermal
584 pressurization-induced mean effective stress is more critical to the permeability change in the anisotropic
585 rock stress case than in the isotropic case at the V1 location.

586

587 The match in rock permeability between the THM simulation and developed TH stress approximation
588 method is found adequate, as relatively large errors of over 100% are generated near the tunnel wall in the
589 horizontal direction (i.e., location H1) at approximately 2 years since canister emplacement. The error is
590 caused by the overestimation of the deviatoric stress in the developed TH method. The permeability

591 function employed in this study (Equation 33) is an exponential function of deviatoric stress above the
592 critical value (5 MPa in this study), hence errors in deviatoric stress estimate above 5 MPa are exponentially
593 magnified in the permeability calculation. These errors, however, are not critical as they diminish in roughly
594 10 years, after which is the target timeframe for long-term PA and radionuclide transport simulations. Also,
595 these errors are still small compared to the change in the permeability level during the canister heating
596 stage, which is a couple of orders of magnitude.

597

598

599 **4. Discussion**

600 Errors in the permeability estimate have been shown insignificant in the previous section, but such errors
601 will be affected by specific parameter values (Table 3) of the function employed in this study (Equation 33)
602 and by the permeability function itself. Therefore, further studies are necessary to evaluate the effect of
603 different permeability functions (such as Li & Liu, 2013; Rutqvist & Tsang, 2003) on the accuracy of the
604 permeability change estimated by the developed TH stress approximation method.

605

606 The excavation damaged zone (EDZ) is assumed to remain constant in this study. This might not be
607 appropriate as the calculated deviatoric stress levels outside the EDZ (locations V3, V4, H3, H4) tend to
608 increase significantly in a later timeframe (~1,000 years) during the canister heating stage (Figure 10).

609 This deviatoric stress increase may cause shear fracture development outside the EDZ and hence the
610 permeability profile around the tunnel may be altered. Tensile fracture, on the other hand, would be
611 unlikely to develop outside the EDZ as the calculated effective stresses tend to remain in compression
612 (Figure 9).

613

614 Also, the host rock formation is assumed a linear elastic material in this study. The actual formation,
615 especially in the EDZ, may exhibit plastic and visco-elastic (creep) behaviors, which could affect the
616 stress profile around the tunnel significantly. The modelling of such plastic and visco-elastic rock
617 behaviors is outside the scope of this study, and hence it may be addressed in future studies. This may
618 also include longer-term healing and sealing created fractures in the EDZ, whose occurrence will depend
619 on the site-specific host rock properties. The approach developed and tested in this study, which should
620 involve calibration and validation of the stress-permeability function against site specific field
621 experiments, is a way to implicitly consider these effects in the simplified elastic model simulations.

622

623

624 **5. Conclusions**

625 In this study, a coupled THM simulation of the long-term behavior of a generic nuclear waste repository
626 in clay/shale formation has been conducted, in order to estimate stress and stress-induced permeability
627 change in the formation around the disposal tunnel. A methodology has been developed for
628 approximating the stress and stress-induced permeability change by using only TH variables, in order to
629 facilitate the incorporation of stress-induced permeability change into existing TH simulators for
630 radionuclide transport simulations in the PA of a repository site. The TOUGH-FLAC simulator was used
631 to model the coupled THM response of the entire repository during canister heating while Barcelona
632 Basic Model (BBM) is employed to model the mechanical behavior of swelling bentonite buffer. Results
633 have provided the following findings:

634

- 635 • Absolute permeability in clay/shale formation in the proximity of the disposal tunnel could
636 change by two orders of magnitude due to stress-induced fracture opening and closing during the
637 first 10,000 years of geologic disposal.

- 638 • Stress and stress-induced permeability change estimated by the developed stress approximation
639 methodology in the TH simulation match those calculated by the coupled THM simulation.
- 640 • The maximum error in the permeability estimate is found to be a fraction of the range of
641 calculated permeability change, which is a couple of orders of magnitude.
- 642 • Such errors are found to diminish in 10 years since the canister emplacement and the match in the
643 long term is found to be excellent.

644

645 Therefore, the accuracy of the developed stress approximation methodology in estimating stress and
646 stress-induced permeability change has been shown adequate. Future studies mentioned in the preceding
647 section will improve the accuracy and reliability of the developed methodology so as to increase the
648 confidence in applying it for the PA and radionuclide transport simulations.

649

650

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652

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662 **References**

- 663 [1] D. Vinson and C. J.T., “Spent Nuclear Fuel and High-Level Radioactive Waste Inventory Report.
664 U.S. Department of Energy, Spent Fuel and Waste Disposition. FCRD-NFST-2013-000263, Rev.
665 6,” 2019.
- 666 [2] M. E. Kraft, “Nuclear power and the challenge of high-level waste disposal in the United States,”
667 *Polity*, vol. 45, no. 2, pp. 265–280, 2013, doi: 10.1057/pol.2013.4.
- 668 [3] B. Faybishenko, J. Birkholzer, D. Sassani, and S. P., “International Approaches for Nuclear Waste
669 Disposal in Geological Formations: Geological Challenges in Radioactive Waste Isolation—Fifth
670 Worldwide Review. Prepared for the U.S. Department of Energy Spent Fuel and Waste Science
671 and Technology R&D Campaign,” 2016.
- 672 [4] G. S. Bodvarsson, W. Boyle, R. Patterson, and D. Williams, “Overview of scientific investigations
673 at Yucca Mountain—the potential repository for high-level nuclear waste,” *J. Contam. Hydrol.*,
674 vol. 38, no. 1–3, pp. 3–24, 1999, doi: 10.1016/S0169-7722(99)00009-1.
- 675 [5] IAEA, “Scientific and Technical Basis for the Geological Disposal of Radioactive Wastes,
676 Technical Report Series No. 413,” Vienna, 2003.
- 677 [6] IAEA, “Spent Fuel and High Level Waste: Chemical Durability and Performance under Simulated
678 Repository Conditions,” Vienna, 2007.
- 679 [7] O. Fouché, H. Wright, J.-M. Le Cléac’h, and P. Pellenard, “Fabric control on strain and rupture of
680 heterogeneous shale samples by using a non-conventional mechanical test,” *Appl. Clay Sci.*, vol.
681 26, no. 1–4, pp. 367–387, 2004, doi: 10.1016/j.clay.2003.12.014.
- 682 [8] D. Patriarche, E. Ledoux, R. Simon-Coinçon, J.-L. Michelot, and J. Cabrera, “Characterization and
683 modeling of diffusion process for mass transport through the Tournemire argillites (Aveyron,
684 France),” *Appl. Clay Sci.*, vol. 26, no. 1–4, pp. 109–122, 2004, doi: 10.1016/j.clay.2003.10.005.
- 685 [9] P. Meier, T. Trick, P. Blumling, and G. Volckaert, “Self-healing of fractures within the EDZ at the
686 Mont Terri Rock Laboratory: results after one year of experimental work,” in *Proceedings of the*

- 687 *international workshop on geomechanics, hydromechanical and thermomechanical behavior of*
688 *deep argillaceous rocks: theory and experiments, 2000.*
- 689 [10] J. D. Barnichon and G. Volckaert, “Observations and predictions of hydromechanical coupling
690 effects in the Boom clay, Mol Underground Research Laboratory, Belgium,” *Hydrogeol. J.*, vol.
691 11, no. 1, pp. 193–202, 2003, doi: 10.1007/s10040-002-0240-6.
- 692 [11] S. Gonzales and K. S. Johnson, “Shales and other argillaceous strata in the United States,” Athens,
693 GA (USA), 1985.
- 694 [12] F. D. Hansen *et al.*, “Shale Disposal of U . S . High-Level Radioactive Waste,” Albuquerque, New
695 Mexico, 2010.
- 696 [13] P. Dobson and J. Houseworth, “Inventory of Shale Formations in the US, Including Geologic,
697 Geochemical, Hydrological, Mechanical, and Thermal Characteristics. Prepared for U.S.
698 Department of Energy, Used Fuel Disposition Campaign, FCRD-UFD-2014-000512,” 2013.
- 699 [14] P. . Mariner *et al.*, “Progress in Deep Geologic Disposal Safety Assessment in the U.S. since 2010.
700 Prepared for U.S. Department of Energy, Spent Fuel and Waste Disposition. M2SF-
701 19SN010304041, SAND2019-12001 R,” 2019.
- 702 [15] G. E. Hammond, P. C. Lichtner, and R. T. Mills, “Evaluating the performance of parallel
703 subsurface simulators: An illustrative example with PFLOTRAN,” *Water Resour. Res.*, vol. 50,
704 no. 1, pp. 208–228, 2014, doi: 10.1002/2012WR013483.
- 705 [16] G. Chen, X. Sillen, J. Verstricht, and X. Li, “ATLAS III in situ heating test in boom clay: Field
706 data, observation and interpretation,” *Comput. Geotech.*, vol. 38, no. 5, pp. 683–696, 2011, doi:
707 10.1016/j.compgeo.2011.04.001.
- 708 [17] B. François, L. Laloui, and C. Laurent, “Thermo-hydro-mechanical simulation of ATLAS in situ
709 large scale test in Boom Clay,” *Comput. Geotech.*, vol. 36, no. 4, pp. 626–640, 2009, doi:
710 10.1016/j.compgeo.2008.09.004.
- 711 [18] G. Armand, F. Bumbieler, N. Conil, de la V. R., J.-M. Bosgiraud, and M.-N. Vu, “Main outcomes
712 from in situ thermo-hydro-mechanical experiments programme to demonstrate feasibility of

- 713 radioactive high-level waste disposal in the Callovo-Oxfordian claystone,” *J. Rock Mech. Geotech.*
714 *Eng.*, vol. 9, no. 3, pp. 415–427, 2017, doi: 10.1016/j.jrmge.2017.03.004.
- 715 [19] J. Rutqvist, L. Zheng, F. Chen, H. H. Liu, and J. Birkholzer, “Modeling of coupled thermo-hydro-
716 mechanical processes with links to geochemistry associated with bentonite-backfilled repository
717 tunnels in clay formations,” *Rock Mech. Rock Eng.*, vol. 47, no. 1, pp. 167–186, 2014, doi:
718 10.1007/s00603-013-0375-x.
- 719 [20] L. Zheng, J. Rutqvist, J. T. Birkholzer, and H. H. Liu, “Coupled THMC models for bentonite in an
720 argillite repository for nuclear waste: illitization and its effect on swelling stress under high
721 temperature,” *Eng. Geol.*, vol. 230, pp. 118–129, 2017, doi: 10.1016/j.enggeo.2017.10.002.
- 722 [21] L. Zheng, J. Rutqvist, J. T. Birkholzer, and H. H. Liu, “On the impact of temperatures up to 200°C
723 in clay repositories with bentonite engineer barrier systems: a study with coupled thermal,
724 hydrological, chemical, and mechanical modeling,” *Eng. Geol.*, vol. 197, pp. 278–295, 2015, doi:
725 10.1016/j.enggeo.2015.08.026.
- 726 [22] J. Rutqvist, “Thermal Management Associated with Geologic Disposal of Large Spent Nuclear
727 Fuel Canisters in Tunnels with Thermally Engineered Backfill,” *Tunn. Undergr. Sp. Technol.*, vol.
728 102, pp. 1–13, 2020, doi: 10.1016/j.tust.2020.103454.
- 729 [23] C. F. Tsang, F. Bernier, and C. Davies, “Geohydromechanical processes in the Excavation
730 Damaged Zone in crystalline rock, rock salt, and indurated and plastic clays - In the context of
731 radioactive waste disposal,” *Int. J. Rock Mech. Min. Sci.*, vol. 42, no. 1, pp. 109–125, 2005, doi:
732 10.1016/j.ijrmms.2004.08.003.
- 733 [24] K. Luu, “TOUGH3-FLAC for Dummies,” Berkeley, CA, 2020.
- 734 [25] Y. Jung, G. Shu Heng Pau, S. Finsterle, and C. Doughty, “TOUGH3 User’s Guide,” Berkeley,
735 CA, 2018.
- 736 [26] Itasca Consulting Group, “FLAC3D — Fast Lagrangian Analysis of Continua in Three-
737 Dimensions, Ver. 7.0,” Minneapolis, 2020.
- 738 [27] J. Kim, H. A. Tchelepi, and R. Juanes, “Stability and convergence of sequential methods for

- 739 coupled flow and geomechanics: fixed-stress and fixed-strain splits,” *Comput. Methods Appl.*
740 *Mech. Eng.*, vol. 200, no. 13–16, pp. 1591–1606, 2011, doi: 10.1016/j.cma.2010.12.022.
- 741 [28] L. Blanco-Martín, R. Wolters, J. Rutqvist, K. H. Lux, and J. T. Birkholzer, “Thermal-hydraulic-
742 mechanical modeling of a large-scale heater test to investigate rock salt and crushed salt behavior
743 under repository conditions for heat-generating nuclear waste,” *Comput. Geotech.*, vol. 77, pp.
744 120–133, 2016, doi: 10.1016/j.compgeo.2016.04.008.
- 745 [29] J. Rutqvist, Y.-S. Wu, C.-F. Tsang, and G. Bodvarsson, “A modeling approach for analysis of
746 coupled multiphase fluid flow, heat transfer, and deformation in fractured porous rock,” *Int. J. Rock*
747 *Mech. Min. Sci.*, vol. 39, no. 4, pp. 429–442, 2002, doi: 10.1016/S1365-1609(02)00022-9.
- 748 [30] J. Rutqvist, “Status of the TOUGH-FLAC simulator and recent applications related to coupled
749 fluid flow and crustal deformations,” *Comput. Geosci.*, vol. 37, no. 6, pp. 739–750, 2011, doi:
750 10.1016/j.cageo.2010.08.006.
- 751 [31] E. E. Alonso, A. Gens, and A. Josa, “A constitutive model for partially saturated soils,”
752 *Géotechnique*, vol. 40, no. 3, pp. 405–430, 1990, doi: 10.1680/geot.1991.41.2.273.
- 753 [32] M. T. van Genuchten, “A Closed-form Equation for Predicting the Hydraulic Conductivity of
754 Unsaturated Soils1,” *Soil Science Society of America Journal*, vol. 44, no. 5, p. 892, 1980, doi:
755 10.2136/sssaj1980.03615995004400050002x.
- 756 [33] E. E. Alonso, J. Vaunat, and A. Gens, “Modelling the mechanical behaviour of expansive clays,”
757 *Eng. Geol.*, vol. 54, no. 1–2, pp. 173–183, 1999, doi: 10.1016/S0013-7952(99)00079-4.
- 758 [34] M. Sánchez, A. Gens, L. do Nascimento Guimarães, and S. Olivella, “A double structure
759 generalized plasticity model for expansive materials,” *Int. J. Numer. Anal. Methods Geomech.*,
760 vol. 29, no. 8, pp. 751–787, 2005, doi: 10.1002/nag.434.
- 761 [35] P. Bossart, P. M. Meier, A. Moeri, T. Trick, and J. C. Mayor, “Geological and hydraulic
762 characterisation of the excavation disturbed zone in the Opalinus Clay of the Mont Terri Rock
763 Laboratory,” *Eng. Geol.*, vol. 66, no. 1–2, pp. 19–38, 2002, doi: 10.1016/S0013-7952(01)00140-5.
- 764 [36] A. G. Corkum and C. D. Martin, “The mechanical behaviour of weak mudstone (Opalinus Clay) at

- 765 low stresses,” *Int. J. Rock Mech. Min. Sci.*, vol. 44, no. 2, pp. 196–209, 2007, doi:
766 10.1016/j.ijrmms.2006.06.004.
- 767 [37] C. D. Martin and G. W. Lanyon, “Measurement of in-situ stress in weak rocks at Mont Terri Rock
768 Laboratory, Switzerland,” *Int. J. Rock Mech. Min. Sci.*, vol. 40, no. 7–8, pp. 1077–1088, 2003, doi:
769 10.1016/S1365-1609(03)00113-8.
- 770 [38] L. C. Li and H. H. Liu, “A numerical study of the mechanical response to excavation and
771 ventilation around tunnels in clay rocks,” *Int. J. Rock Mech. Min. Sci.*, vol. 59, pp. 22–32, 2013,
772 doi: 10.1016/j.ijrmms.2012.11.005.
- 773 [39] J. Rutqvist and C. F. Tsang, “Analysis of thermal-hydrologic-mechanical behavior near an
774 emplacement drift at Yucca Mountain,” *J. Contam. Hydrol.*, vol. 62–63, pp. 637–652, 2003, doi:
775 10.1016/S0169-7722(02)00184-5.
- 776