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Abstract: The aluminosilicate layers of the swelling clay mineral montmorillonite, and the saturated pores they delineate, control the mechanical properties and the transport of solutes in many natural and engineered environments. However, the structural basis of montmorillonite porosity remains poorly characterized due to the difficulty in visualizing hydrated samples in their native state. Here, we used cryogenic transmission electron microscopy (cryo-TEM) and cryo electron tomography (cryo-ET) to show that stacking defects within minimally altered, fully hydrated montmorillonite particles define multiscale porosity networks. Variations in layer lateral dimensions over tens to thousands of nanometers cause a range of topological and dynamic defects that generate pervasive curvature and introduce previously uncharacterized solute transport pathways. Observations of long-range rotational order between neighboring layers indicate that the layer-layer interactions that govern clay swelling involve three dimensional orienting forces that operate across nanoscale pores. These direct observations of the hierarchical structure of hydrated montmorillonite pore networks with nanoscale resolution reveal potentially general aspects of colloidal interactions in fluid-saturated clay minerals.

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1 Layer size polydispersity in hydrated montmorillonite creates multiscale porosity networks 2 Michael L. Whittaker,^{1,2†} Luis R. Comolli,^{1†} Benjamin Gilbert,¹ Jillian F. Banfield^{1,2*} 3 4 ¹Lawrence Berkeley National Laboratory, Life Sciences Division, Berkeley, California, USA. 5 6 ²Department of Earth and Planetary Science, University of California, Berkeley, California, 7 USA. * jbanfield@berkeley.edu 8 9 [†] These authors contributed equally to this work 10 11 12 Keywords: clay, montmorillonite, porosity, cryo-TEM, cryo electron tomography 13 14 Abstract 15 16 17 The aluminosilicate layers of the swelling clay mineral montmorillonite, and the saturated pores

18 they delineate, control the mechanical properties and the transport of solutes in many natural and 19 engineered environments. However, the structural basis of montmorillonite porosity remains 20 poorly characterized due to the difficulty in visualizing hydrated samples in their native state. 21 Here, we used cryogenic transmission electron microscopy (cryo-TEM) and cryo electron 22 tomography (cryo-ET) to show that stacking defects within minimally altered, fully hydrated 23 montmorillonite particles define multiscale porosity networks. Variations in layer lateral 24 dimensions over tens to thousands of nanometers cause a range of topological and dynamic 25 defects that generate pervasive curvature and introduce previously uncharacterized solute 26 transport pathways. Observations of long-range rotational order between neighboring layers 27 indicate that the layer-layer interactions that govern clay swelling involve three dimensional orienting forces that operate across nanoscale pores. These direct observations of the hierarchical 28 29 structure of hydrated montmorillonite pore networks with nanoscale resolution reveal potentially 30 general aspects of colloidal interactions in fluid-saturated clay minerals.

32 1. Introduction

33 Swelling clays such as smectites are among the most abundant inorganic nanomaterials in the lithosphere(Hochella et al. 2019), and play an outsize role in controlling the transport and 34 35 retention of water, CO_2 , nutrients and pollutants in both natural and engineered settings due to 36 their exceptionally high surface areas(Bourg and Ajo-Franklin 2017, Charlet et al. 2017). SWy is 37 an archetypical smectite whose clay fraction is almost exclusively montmorillonite (Mt) that 38 formed via devitrification of volcanic ash and tuffs through submarine weathering, with no 39 significant post-formational recrystallization(Cadrin et al. 1995). Ion binding 40 selectivities(Whittaker et al. 2019), permeabilities(Tournassat et al. 2016b), shear strength(Ikari 41 et al. 2015), and myriad other properties of Mt are highly dependent on the shape, size, and the 42 specifics of how smectite particles are arranged.

43 Smectite 2:1 layers are less than one nanometer thick and up to several microns in lateral 44 dimensions, and stack to form particles in various ways depending on the relative concentrations of water, clay, and electrolyte(Tournassat and Steefel 2015). While the average separation 45 46 between layers generally decreases with increasing concentration of clay or salt(Norrish 1954) 47 there is no consensus about which of the many potential microscopic clay colloid arrangements 48 are expected to form under specific conditions(Bergaya and Lagaly 2013). Fluid and solute 49 transport rates through natural SWy depend on the pore structures defined by layer stacking 50 motifs(Wenk et al. 2008), and are generally classified into two broad categories based on the 51 separation distance between adjacent Mt layers: nanopores and macropores. Nanopores include 52 clay interlayers that, because of the effective negative charge on smectite layers, generally 53 exclude anions and therefore facilitate ion-selective transport(Tournassat et al. 2016a). 54 Macropores are the larger and less well-defined spaces between particles through which anions,

cations, and even larger solutes like macromolecules and nanoparticles can diffuse(Tournassat et
al. 2016b). In both cases, experimental descriptions of pore geometries and connectivity that are
required for accurate transport models are lacking(Churakov and Gimmi 2011, Tournassat et al.
2016b).

59 High-resolution transmission electron microscopy (HR-TEM) has been utilized for 60 decades to reveal atomic- and nanoscale structures in non-hydrous clay particles(Veblen 1985, 61 Vali and Köster 1986), including stacking order between layers in smectite and 62 illite/smectite(Veblen 1990, Guthrie and Veblen 1989). However, there is evidence that the 63 native structure of smectites is disrupted(Dudek et al. 2002) during conventional sample 64 preparation(Gu et al. 2001). Low-dose transmission electron microscopy of cryogenically frozen 65 samples (cryo-TEM) is uniquely capable of characterizing hydrated clay structures over spatial 66 scales ranging from near-atomic resolution(Whittaker et al. 2019) to whole particle 67 aggregates(Whittaker et al. 2019, Gilbert et al. 2015, Segad et al. 2012). Water vitrifies without 68 crystallization upon rapid freezing for sample thicknesses(Deirieh et al. 2018) that are electron-69 transparent at the accelerating voltages commonly employed for cryo-TEM (200-300 kV), 70 preserving structures with minimal perturbation from their native-state(Cheng 2018). Increased 71 electron-dose robustness of cryo-frozen samples(Henderson and Glaeser 1985) allows for 3D 72 images to be reconstructed using cryo electron tomography (cryo-ET) from a series of images 73 taken at different tilt angles with minimal beam-induced damage. 74 Here, we use cryo-TEM to show that native pore structures in minimally altered, 75 hydrated SWy arise from disparately-sized layers that stack defectively and introduce pervasive 76 layer curvature. We employ cryo-ET to characterize the 3D structure of SWy pore networks that

cannot be resolved from 2D images alone and are not accounted for in commonly used models

78 based on x-ray diffraction and simulations that assume perfectly planar layers. We observe 79 rotational crystallographic ordering between adjacent layers separated by over 1 nm of interlayer 80 water, in disagreement with the common assertion that hydrated smectites are fully turbostratic. 81 82 2. Methods 83 84 Wyoming bentonite (SWy-1/SWy-2) from the Clay Minerals Society was suspended in deionized water or an aqueous solution of NaCl (200 mM) or MgCl₂ (100 mM) by manual 85 86 shaking for 2 minutes with no filtration, washing or prior separation of non-clay minerals. Suspensions were incubated overnight before cryo-TEM analysis. No significant differences in 87 88 clay structure were observed for samples suspended in NaCl versus MgCl₂ and images from both 89 electrolyte solutions are presented. 90 Imaging was performed on a JEOL-3100-FFC electron microscope (JEOL Ltd., 91 Akishima, Tokyo, Japan) equipped with a FEG electron source operating at 300 kV. An Omega 92 energy filter (JEOL) attenuated electrons with energy losses that exceeded 30 eV of the zero-loss 93 peak. A Gatan 795 4Kx4K CCD camera (Gatan Inc., Pleasanton, CA, USA) was mounted at the 94 exit of an electron decelerator maintained at a voltage of 200–250 kV. Aliquots of 5 µl were 95 taken directly from the suspensions, placed onto 200 mesh lacey carbon Formvar or 300 mesh 96 lacey carbon Cu-grids (Ted Pella, INC, #01881 and #01895 respectively) and manually flash-97 frozen in liquid ethane. Grids were rapidly transferred to liquid nitrogen, and either stored or 98 immediately transferred to the microscope submerged in liquid nitrogen in a cryo-transfer stage 99 maintained at 80 K during acquisition of all data sets.

100	Images and tomographic tilt series were acquired under with doses of ~125-250 e ⁻ /Å ² ,
101	$320 \text{ e}^{-}/\text{Å}^{2}$, $720 \text{ e}^{-}/\text{Å}^{2}$, or 1,000 e $^{-}/\text{Å}^{2}$ per image for images with pixel sizes of 0.112 nm, 0.07 nm,
102	0.047 nm, and 0.028 nm at the specimen, respectively. Tilt series images were taken at
103	increments of 2° between $+65^{\circ}$ and $-65^{\circ} \pm 5^{\circ}$ for a total of 61-66 images with the program Serial-
104	EM (http://bio3d.colorado.edu/). A dose of 33 e ⁻ /pixel was accumulated in each image, with
105	pixel size of 0.224 nm at the specimen, corresponding to a dose of 6.7 $e^{-}/Å^{2}/image$ and a total
106	dose of ~400 e ⁻ /Å ² . Underfocus values ranged between 750 \pm 100 nm to 1,200 \pm 250 mm. The
107	electron dose limit for the onset of quantifiably high-resolution damage at 4.6 at Å is
108	approximately 4,000 $e^{-}/Å^2$.
109	All tomographic reconstructions were performed with the program etomo
110	(http://bio3d.colorado.edu/). The package of imaging tools Priism (http://msg.ucsf.edu/IVE/) was
111	used for the noise reduction and contrast enhancement. The software ImageJ 1.38x (NIH,
112	http://rsb.info.nih.gov/ij/) was used for analysis and measurements of the 2D image projections.
113	The programs Imod (http://bio3d.colorado.edu/) and UCSF Chimera
114	(https://www.cgl.ucsf.edu/chimera/) were used for 3D renderings. All movies were created with
115	the open source package ffmpeg (http://www.ffmpeg.org/). Adobe Photoshop CS5.1 was used to
116	adjust contrast in the images and to insert calibrated scale bars into images.
117	
118	3. Results
119	3.1 Rotational Ordering in Hydrated Mt

Mt particles adopted a range of orientations that reflected their structure in suspension
immediately prior to plunge-freezing. Cryo-TEM images of these suspensions exhibited two
dominant modes of contrast that varied in relative intensity depending on the orientation of the

123	layers with respect to the electron-beam axis. In a <i>face-on</i> orientation, parallel to the TEM grid
124	supports, the particles were much thinner than they were wide, and contrast was generated
125	primarily by phase interference (Figure 1). Layers that were oriented <i>edge-on</i> (within $\sim 5^{\circ}$ of the
126	beam axis(Gilbert et al. 2015)) gave strong amplitude contrast that resulted from greater
127	thickness and increased scattering along the beam direction (Figures 2, 3). Layers at oblique
128	angles to the beam, or that curved away from the beam axis (Figure 4), generated intermediate
129	contrast in 2D projection images that required 3D tomographic reconstruction to be fully
130	resolved (Figure 5). In both amplitude and phase contrast images, orientational order was
131	observed between neighboring layers.
132	Multiple stacked layers contributed to phase image contrast, creating a Moiré interference
133	pattern that contained information about the number and relative orientation of the layers
134	(Figure 1). Contrast in phase images varied with microscope settings including defocus and
135	astigmatism, as well as sample-dependent geometrical factors such as tilt angle and curvature,
136	and therefore images could not be directly compared with atomic structural models. However,
137	spots at 2.2 and 3.8 nm ⁻¹ in Fourier transformation (FT) images (Figure 1, insets) confirmed that
138	fringes originated from Mt layers. These spots corresponded to lattice planes with spacings of a
139	= 0.46 nm (020), 0.45 nm (110)/($\overline{1}10$), and 0.26 nm (130)/(200) indicating 2.6 Å to be the
140	highest resolution achievable in this study. The presence of multiple sets of spots at each spacing
141	in FT images was an indication that multiple individual layers contributed to the image contrast.
142	In some particles, five or more layers were oriented within a fairly narrow (~10°) range of angles
143	(Figure 1A, C). This small angular rotation about the stacking axis with respect to neighboring
144	particles led to a Moiré interference pattern with defined periodicity along the directions in
145	which contrast is strongest (see arrows in Figure 1A).

146	Parts of Figure 1A exhibit Moiré fringes with a spacing of approximately $d = 10$ nm,
147	suggesting an interlayer angle of $d/a \approx 2.6^{\circ}$. This is consistent with the ~1-3° separation between
148	individual spots within a cluster in the FT image. However, regions containing more layers
149	and/or those that are arranged with a much wider spread in the relative interlayer orientations
150	exhibited more complex Moiré effects that were not readily interpretable (Figure 1B). Thus,
151	phase contrast images alone carried only limited information about the number of layers per
152	particle with the same orientation (from FT spots) and the relative orientation of those layers
153	with respect to one another (from FT spots and, in certain cases, Moiré patterns).
154	Amplitude contrast images allowed both (darker) layer stacking motifs and the (lighter)
155	pores between them to be visualized directly. Regular stacking of layers with a 1.9 nm
156	periodicity, consistent with approximately 0.9 nm thick layers bounding a 1.0-1.2 nm thick
157	hydrated interlayer region(Norrish 1954), was consistently observed (Figure 2). Fringes within
158	each layer, which occasionally spanned neighboring layers, corresponded to the same 0.46 nm
159	(020) or 0.45 nm (110)/($\overline{1}10$) lattice spacings that were observed in phase images, but in this case
160	viewed along an orthogonal axis. These fringes made angles of approximately 73°, 90°, or 103°
161	with respect to the horizontal axis of the layer, depending on its relative orientation, which
162	corresponded to the $\langle 110 \rangle$, $\langle 100 \rangle$ or $\langle \overline{1}10 \rangle$ zone axes. Immediately adjacent layers often
163	exhibited similar orientations, analogous to mica-like stacking. While this order rarely persisted
164	over more than two or three layers, the stacking sequence was clearly not completely random
165	(i.e., not turbostratic).

3.2 Hydrated Pore Structures

Lower-magnification images revealed interactions between SWy particles and aggregates of many particles (**Figure 3**). Particles consisted of stacks of between two to a few hundred layers that were connected through a continuous network of face-face contacts at the modal interlayer spacing of approximately 1.9 nm. Thus, we propose that a network of equivalent basal spacings that percolates between the bounding layers constitutes a consistent definition of an individual suspended Mt crystalline hydrate particle *in aqua*.

Most Mt particles imaged in this study contained considerably more layers than has been observed in studies of extensively treated clay samples. Following exfoliation, washing, sizeseparation and drying, re-suspended smectite particles typically show a log-normal layer number distribution centered at ~7-8(Dudek et al. 2002, Whittaker et al. 2019). That the simply resuspended particles observed in this study display such pronounced differences in shape and size is evidence that they are minimally altered prior to imaging.

180 In all particles examined, layer-size polydispersity led to a panoply of distinctive defects in layer stacking structure. The most pronounced was the variation in the number of layers per 181 182 particle, which ranged from two to hundreds, shown in order of increasing number of layers in 183 Figure 3. Large particles contained a number of defects arising from the geometrical 184 consequences of the distribution of layer sizes (Figure 3E). Two layers that terminated in close 185 proximity within the same plane created a gap defect between layers with an interlayer distance 186 of 2.8 nm, more than twice the equilibrium pore spacing (Figure 3F). A layer that terminated 187 within a particle formed a dislocation (Figure 3G), giving rise to a pore with 2.8 nm spacing in 188 the immediate vicinity of the terminated layer that decreased back to 1.9 nm as adjacent layers 189 locally deformed around the defect. A sufficiently small layer created a dislocation loop in which 190 the layer was entirely contained within the particle (Figure 3K). A layer that was far longer than

average was the source of a wedge disclination (Figure 3H), on which multiple dislocations
terminated. An exceptionally long layer also spanned two stacks of layers that otherwise would
not be part of the same particle (Figure 3I), participating in both a wedge disclination and a large
gap (Figure 3 J, K). At least one layer, and possibly more, even bridged larger stacks of tens to
hundreds of layers each (Figure 3L), which bent to accommodate the different orientations in
different regions of the particle.

197 A further stacking defect, observed in the smallest particles, is the presence of face-edge 198 and edge-edge contacts rather than face-face contacts (Figure 3A-D). Between layers arranged 199 this way the pore size increased continuously until either of the bounding layers terminated. 200 These pores may be unstable structures that may arise from layers pulling away from, or 201 attaching to, each other. Although cryo-imaging provides only static snapshots of particle 202 structures, recent observations suggest that dynamic delamination and restacking of layers occurs 203 in bulk solution(Whittaker et al. 2019). Thus, this type of porosity is likely transient, and is not 204 prevalent in particles consisting of more layers.

205

206 *3.3 Layer stacking defects induce curvature*

A survey of observed defect types is provided in **Figure 4**. Many of the defects identified in **Figure 3** were associated with some degree of curvature in the layers, necessary to accommodate layer size disparity and variable particle orientations. The direction of curvature seldom changed over the length of a layer, allowing the persistence length, $\xi = -\frac{L}{\ln(\cos\theta)}$ to be calculated from 2D images by determining θ , the angle between the layer normals at each end of a layer with length *L*(Boal 2012). For layers with the highest curvature (**Figure 3L**) $\xi = 1.6 \mu m$, but in most cases was in the range of millimeters. This means that Mt layers associated with particles were

214 only gently or locally bent, since they were considerably smaller than the minimum persistence 215 length. Nonetheless, defects did distort the layers from their otherwise planar geometry. For 216 example, the termination of a layer within a particle caused curvature in neighboring layers. Layers curved gently, with large radii of curvature characterized by $r_c = \left|\frac{\delta x}{\delta \theta}\right|$, where x is the 217 218 distance along the 2D cross-section of a layer. Curvature was common in 2D images, especially 219 near dislocations and disclinations, but also near loci of delamination and restacking. In the 220 vicinity of a dislocation the local radius of curvature approached 75 nm (Figure 3L), the smallest 221 radius of curvature identified in this study.

Quantification of persistence lengths in 3D was limited by the relatively small volumesthat could be reconstructed and rendered with the available computational resources.

224 Nonetheless, important information was obtained from 3D images that was not discernable from 225 2D images. For example, layers frequently curved along two or more different radii. Figure 5A 226 depicts a 2D cross section of a 3D reconstruction shown in Figure 5B. Size polydispersity in 227 layer dimensions led to the formation of a particle in which larger exterior layers bound smaller 228 interior layers, inducing curvature of the larger layers and leading to the formation of a large pore. Gaussian curvature is defined as $K = (r_{c,1}r_{c,2})^{-1}$ where $r_{c,1}$ and $r_{c,2}$ are the principal 229 230 curvatures in orthogonal directions. The upper layer in **Figure 5B** exhibits positive Gaussian 231 curvature as it gently curves in the plane of the image and around the pore space running 232 orthogonal the plane of the image. The pore is continuous throughout the reconstructed region 233 (~300 nm \times 500 nm), but changes size due both to curvature of the upper and lower layers and 234 the intrusion of other neighboring layers. While the pore is essentially contained within a 235 particle, it has a maximum diameter of approximately 40 nm, far larger than the average 236 interlayer separation. Although particles with Gaussian curvature can sometimes be inferred

from 2D images based on the presence or absence of intermediate contrast adjacent to an edgeon layer, 3D tomographic reconstructions make the geometry of both curvature and defects
explicit for all layers, and show that they play an important role in delimiting pore structures. Tilt
series movies and tomographic reconstruction animations from which the images in Figure 5

241 were taken are available as **Supplementary Movies 1-5**.

242 The termination of individual layers within a stack also increases pore connectivity. 243 Tomograms revealed that the internal porosity within a particle was much more connected than 244 would be expected from a 2D projection image alone. The edge dislocation highlighted in 245 Figure 5C is shown in 3D in Figure 5D, with a molecular model for comparison in Figure 5E. 246 The presence of dislocations within a particle provides a diffusional pathway orthogonal to the 247 stacking direction. Dislocations also played an important role in defining the shape of a particle. 248 Lenticular cross-sections (Figure 5C) were the result of smaller layers that terminated within 249 particles, while larger layers tended to be found near the exterior. The total length of edges that 250 were exposed to bulk solution was minimized and resulted in the curving of layers to 251 accommodate the internal defects. This contrasted with particles structures of extensively treated 252 Mt, in which the opposite is true and smaller layers were generally found on the exterior of 253 particles(Tester et al. 2016, Whittaker et al. 2019).

254

255 **4. Discussion**

Using cryo-preparation methods to minimize sample preparation artifacts(Deirieh et al. 2018) and beam-induced damage during imaging(Henderson and Glaeser 1985), in analogy to cryoEM of biological samples(Cheng 2018), provides confidence that structures observed by cryo-TEM faithfully represent the native state of hydrated SWy. The observed structures are

260 therefore likely the result of authigenic processes, because SWy does not appear to have 261 recrystallized after formation(Cadrin et al. 1995). Consequently, the morphology of single 262 particles and of their aggregates characterized here provides a basis for understanding the fabrics 263 of smectite-rich soils and rocks. For example, the high density of dislocations within particles 264 that give rise to lenticular cross-sections may be the result of terminated growth as the layers 265 above and below retard transport to layers inside. Uniaxial compaction during sedimentation 266 would not be expected to alter the distribution of topological defects like dislocations 267 appreciably, but diagenetic alteration at elevated temperatures or dissolution and reprecipitation 268 in response to chemical gradients would be expected to decrease the density of such high-energy 269 defects. Identifying defect distributions in smectite particles and the factors that alter these 270 distributions raises exciting new possibilities for characterizing the evolution of smectite 271 microfabrics.

272 This work challenges assumptions made about clay particle structures based on 273 macroscopic measurements. For example, this Mt sample has an tendency for near-274 crystallographic registry at approximately 0°, 60° or 120° relative rotation angles(Plançon et al. 275 1979, Viani, Gualtieri and Artioli 2002). These findings favor the semi-ordered rotational 276 stacking initially proposed for mica(Méring 1975) and argue strongly against the long-standing 277 assumption that smectites possess complete turbostratic disorder (Moore and Reynolds 1997). 278 However, ordering through a geometrical 'lock-and-key' mechanism involving the interlayer 279 cation observed in micas is not required to orient neighboring smectite layers. Rather, long-range 280 forces between layers, which could include electrostatic interactions involving counterions, van 281 der Waals torque(Zhang et al. 2017a), or the interlayer hydrogen bond network(Meleshyn and 282 Bunnenberg 2005, Zhang et al. 2017b), appear to generate a rotational energy landscape with

modulo 60° energy minima. No rotational ordering is observed between aggregated Mt particles,
in agreement with observation that hydrated colloidal particles(Michot et al. 2013) and the
particles in smectite-rich rocks(Lutterotti et al. 2009) show preferred intraparticle orientation but
are themselves randomly oriented.

287 Nanopore spaces inside particles that were much larger than the equilibrium interlayer 288 separation were frequently observed in cryo-TEM and cryo-ET images. Such defects may 289 explain features that are common in bulk x-ray scattering and diffraction data but for which the 290 structural origin has not been treated explicitly. Mt typically exhibits (00*l*), and particularly 291 (001), peaks that are asymmetric towards lower scattering vectors, (Ferrage 2005, Segad et al. 292 2012) consistent with our observation here of a distribution layer-layer spacings that are larger 293 than the mean value due to the presence of defects. Knowledge of the defect types and 294 prevalence in Mt nanostructures may therefore be used to improve bulk characterization 295 approaches.

296 The smallest radii of curvature reported here (~75 nm) are far larger than previous reports 297 of highly curved and bent particles and layers when prepared with epoxy resin, which are similar 298 to those found in Mt-polymer composites (Dudek et al. 2002, Fu et al. 2011). Observations of 299 curvature at large angles in those materials(Zbik et al. 2008, Zbik and Frost 2010, Mouzon, 300 Bhuiyan and Hedlund 2016, Du et al. 2019) are likely due to externally applied stress that induce 301 spontaneous curvature. Molecular simulations predict that the energy penalty for bending single 302 Mt layers between carbon support layers does not exceed the thermal energy until it reaches $r_c =$ 303 15 nm(Fu et al. 2011), suggesting that the geometries of the minimally treated Mt particles 304 observed here are dominated by colloidal interactions in the absence of external stress.

305 Currently, very few of the Mt structures observed here can be predicted from first 306 principles theories of interlayer interactions. In contrast to rigid and monodisperse colloidal 307 polyhedra, for which crystal structures can be predicted based on the shape of the 308 particle(Damasceno, Engel and Glotzer 2012), simulations with system sizes smaller than the 309 average Mt layer dimension, or that do not account for layer curvature(Underwood and Bourg 310 2020, Ebrahimi, Whittle and Pellenq 2016) identify a range of microstructures that differ from 311 those observed here. In particular, the size polydispersity of flexible Mt layers causes curvature, 312 changing the energy landscape of clay particles(Suter et al. 2007, Kunz et al. 2009, Castellanos-313 Gomez et al. 2012, Honorio et al. 2018) and the pore size distribution in ways that will likely 314 require coupled chemical-mechanical models to describe.

315

316 5. Conclusions

Microstructures of minimally altered, hydrated Wyoming smectite were imaged in two 317 318 and three dimensions, revealing a panoply of defects that govern clay layer arrangements. The 319 dominant feature of Mt particles was the polydispersity in layer dimensions, which gives rise to 320 defects via incommensurate stacking. Layers curve to accommodate stacking defects, creating 321 hierarchical pore networks that can vary greatly in size distribution and can be highly 322 interconnected. Some defects appear to be unstable and may reflect dynamic processes such as 323 the delamination and restacking of individual layers. Many layers also exhibit orientational order 324 at approximately 60° intervals, which suggests the presence of long-range torsional interactions 325 between neighboring layers within particles.

We anticipate that the nanoscale observations of defect microstructures in Mt presentedherein will help link microscopic structures to the macroscopic properties of clay-rich systems.

- 328 Pore spaces elucidated here may play a significant and underappreciated role in controlling the
- 329 transport of ionic and molecular constituents through clay-rich media, and their identification
- 330 with nanoscale resolution will help inform the prediction of properties such as solute
- 331 permeability and strength at larger scales.

332 6. References

- Bergaya, F. & G. Lagaly. 2013. *Handbook of Clay Science*. Amsterdam: Elsevier.
- Boal, D. 2012. *Mechanics of the Cell*. Cambridge University Press.
- Bourg, I. C. & J. B. Ajo-Franklin (2017) Clay, Water, and Salt: Controls on the Permeability of
 Fine-Grained Sedimentary Rocks. *Acc Chem Res*, 50, 2067-2074.
- Cadrin, A. A. J., T. K. Kyser, W. G. E. Caldwell & F. J. Longstaffe (1995) Isotopic and chemical
 compositions of bentonites as paleoenvironmental indicators of the Cretaceous
 Western Interior Seaway. *Palaeogeography, Palaeoclimatology, Palaeoecology,* 119,
 301-320.
- Castellanos-Gomez, A., M. Poot, A. Amor-Amorós, G. A. Steele, H. S. J. van der Zant, N. Agraït &
 G. Rubio-Bollinger (2012) Mechanical properties of freely suspended atomically thin
 dielectric layers of mica. *Nano Research*, 5, 550-557.
- Charlet, L., P. Alt-Epping, P. Wersin & B. Gilbert (2017) Diffusive transport and reaction in clay
 rocks: A storage (nuclear waste, CO 2, H 2), energy (shale gas) and water quality issue.
 Advances in Water Resources, 106, 39-59.
- Cheng, Y. (2018) Single-particle cryo-EM—How did it get here and where will it go. *Science*, 361,
 867-880.
- Churakov, S. V. & T. Gimmi (2011) Up-Scaling of Molecular Diffusion Coefficients in Clays: A
 Two-Step Approach. *The Journal of Physical Chemistry C*, 115, 6703-6714.
- Damasceno, P. F., M. Engel & S. C. Glotzer (2012) Predictive self-assembly of polyhedra into
 complex structures. *Science*, 337, 453-7.
- Deirieh, A., I. Y. Chang, M. L. Whittaker, S. Weigand, D. Keane, J. Rix, J. T. Germaine, D. Joester &
 P. B. Flemings (2018) Particle arrangements in clay slurries: The case against the
 honeycomb structure. *Applied Clay Science*, 152, 166-172.
- Du, M., J. Liu, P. Clode & Y.-K. Leong (2019) Microstructure and rheology of bentonite slurries
 containing multiple-charge phosphate-based additives. *Applied Clay Science*, 169, 120 128.
- Dudek, T., J. Srodon, D. Eberl, F. Elass & P. Uhlik (2002) Thickness distribution of illite crystals in
 shales. I: X-ray diffraction vs. high resolution transmission electron microscopy. *Clays and Clay Minerals*, 50, 562-577.
- Ebrahimi, D., A. Whittle & R. M. Pellenq (2016) Effect of Polydispersity of Clay Platelets on the
 Aggregation And Mechanical Properties of Clay at the Mesoscale. *Clays and Clay Minerals,* 64, 425-437.

Ferrage, E. (2005) Investigation of smectite hydration properties by modeling experimental X ray diffraction patterns: Part I. Montmorillonite hydration properties. *American Mineralogist*, 90, 1358-1374.

368 Fu, Y.-T., G. D. Zartman, M. Yoonessi, L. F. Drummy & H. Heinz (2011) Bending of Layered 369 Silicates on the Nanometer Scale: Mechanism, Stored Energy, and Curvature Limits. The 370 Journal of Physical Chemistry C, 115, 22292-22300. 371 Gilbert, B., L. R. Comolli, R. M. Tinnacher, M. Kunz & J. F. Banfield (2015) Formation and 372 Restacking of Disordered Smectite Osmotic Hydrates. Clays and Clay Minerals, 63, 432-373 442. 374 Gu, B. X., L. M. Wang, L. D. Minc & R. C. Ewing (2001) Temperature e. Journal of Nuclear 375 Materials, 297, 345-354. 376 Guthrie, G. D. & D. R. Veblen (1989) High-resolution transmission electron microscopy of mixed-377 layer illite/smectite: computer simulations. Clays and Clay Minerals, 37, 1-11. 378 Henderson, R. & R. M. Glaeser (1985) Quantitative analysis of image contrast in electron 379 micrographs of beam-sensitive crystals. *Ultramicroscopy*, 16, 139-150. 380 Hochella, M. F., Jr., D. W. Mogk, J. Ranville, I. C. Allen, G. W. Luther, L. C. Marr, B. P. McGrail, M. 381 Murayama, N. P. Qafoku, K. M. Rosso, N. Sahai, P. A. Schroeder, P. Vikesland, P. 382 Westerhoff & Y. Yang (2019) Natural, incidental, and engineered nanomaterials and 383 their impacts on the Earth system. Science, 363. 384 Honorio, T., L. Brochard, M. Vandamme & A. Lebee (2018) Flexibility of nanolayers and stacks: implications in the nanostructuration of clays. Soft Matter, 14, 7354-7367. 385 386 Ikari, M. J., Y. Ito, K. Ujiie & A. J. Kopf (2015) Spectrum of slip behaviour in Tohoku fault zone 387 samples at plate tectonic slip rates. Nature Geoscience, 8, 870-874. 388 Kunz, D. A., E. Max, R. Weinkamer, T. Lunkenbein, J. Breu & A. Fery (2009) Deformation 389 measurements on thin clay tactoids. Small, 5, 1816-20. 390 Lutterotti, L., M. Voltolini, H. R. Wenk, K. Bandyopadhyay & T. Vanorio (2009) Texture analysis 391 of a turbostratically disordered Ca-montmorillonite. American Mineralogist, 95, 98-103. 392 Meleshyn, A. & C. Bunnenberg (2005) The gap between crystalline and osmotic swelling of Na-393 montmorillonite: a Monte Carlo study. J Chem Phys, 122, 34705. 394 Méring, J. 1975. Smectites. In Soil Components, ed. J. Gieseking. New York: Springer-Verlag. 395 Michot, L. J., I. Bihannic, F. Thomas, B. S. Lartiges, Y. Waldvogel, C. Caillet, J. Thieme, S. S. Funari 396 & P. Levitz (2013) Coagulation of Na-montmorillonite by inorganic cations at neutral pH. 397 A combined transmission X-ray microscopy, small angle and wide angle X-ray scattering 398 study. Langmuir, 29, 3500-10. 399 Moore, D. M. & J. R. C. Reynolds. 1997. Xray diffraction and the identification and analysis of 400 clay minerals. New York, New York: Oxford University Press. 401 Mouzon, J., I. U. Bhuiyan & J. Hedlund (2016) The structure of montmorillonite gels revealed by 402 sequential cryo-XHR-SEM imaging. J Colloid Interface Sci, 465, 58-66. 403 Norrish, K. (1954) The Swelling of Montmorillonite. 120-134. 404 Plançon, A., G. Besson, C. Tchoubar, J. P. Gaultier & J. Mamy (1979) Qualitative and 405 Quantitative Study of a Structural Reorganization in Montmorillonite After Potassium Fixation. Developments in Sedimentology, 27, 45-54. 406 407 Segad, M., S. Hanski, U. Olsson, J. Ruokolainen, T. Åkesson & B. Jönsson (2012) Microstructural 408 and Swelling Properties of Ca and Na Montmorillonite: (In Situ) Observations with Cryo-409 TEM and SAXS. The Journal of Physical Chemistry C, 116, 7596-7601.

410	Suter, J. L., P. V. Coveney, H. C. Greenwell & MA. Thyveetil (2007) Large-Scale Molecular
411	Dynamics Study of Montmorillonite Clay: Emergence of Undulatory Fluctuations and
412	Determination of Material Properties. Journal of Physical Chemistry C, 111, 8248-8295.
413	Tester, C. C., S. Aloni, B. Gilbert & J. F. Banfield (2016) Short- and Long-Range Attractive Forces
414	That Influence the Structure of Montmorillonite Osmotic Hydrates. Langmuir, 32,
415	12039-12046.
416	Tournassat, C., I. Bourg, M. Holmboe, G. Sposito & C. Steefel (2016a) Molecular Dynamics
417	Simulations of Anion Exclusion in Clay Interlayer Nanopores. Clays and Clay Minerals, 64,
418	374-388.
419	Tournassat, C., S. Gaboreau, JC. Robinet, I. Bourg & C. I. Steefel. 2016b. Impact of
420	microstructure on anion exclusion in compacted clay media. 137-149.
421	Tournassat, C. & C. I. Steefel (2015) Ionic Transport in Nano-Porous Clays with Consideration of
422	Electrostatic Effects. Reviews in Mineralogy and Geochemistry, 80, 287-329.
423	Underwood, T. R. & I. C. Bourg (2020) Large-Scale Molecular Dynamics Simulation of the
424	Dehydration of a Suspension of Smectite Clay Nanoparticles. The Journal of Physical
425	Chemistry C.
426	Vali, H. & H. M. Köster (1986) Expanding behavior, structural disorder, regular and random
427	irregular interstratification of 2:1 layer silicates studied by high-resolution images of
428	transmission electron microscopy. Clay Minerals, 21, 827-859.
429	Veblen, D. R. (1985) Direct TEM Imaging of Complex Structures and Defects in Silicates. Annual
430	Review of Earth and Planetary Science, 13, 119-146.
431	(1990) High-Resolution Transmission Electron Microscopy and Electron Diffraction of Mixed-
432	Layer Illite/Smectite: Experimental Results. Clays and Clay Minerals, 38, 1-13.
433	Viani, A., A. F. Gualtieri & G. Artioli (2002) The nature of disorder in montmorillonite by
434	simulation of X-ray powder patterns. American Mineralogist, 87, 966-975.
435	Wenk, HR., M. Voltolini, M. Mazurek, L. R. v. Loon & A. Vinsot (2008) Preferred orientations
436	and anisotropy in shales: callovo-oxfordian shale (france) and opalinus clay
437	(switzerland). Clays and Clay Minerals, 56, 285-306.
438	Whittaker, M. L., L. N. Lammers, S. Carrero, B. Gilbert & J. F. Banfield (2019) Ion exchange
439	selectivity in clay is controlled by nanoscale chemical-mechanical coupling. Proceedings
440	of the National Academy of Sciences.
441	Zbik, M. S. & R. L. Frost (2010) Influence of smectite suspension structure on sheet orientation
442	in dry sediments: XRD and AFM applications. Journal of Colloid and Interface Science,
443	346, 311-316.
444	Zbik, M. S., W. N. Martens, R. L. Frost, YF. Song, YM. Chen & JH. Chen (2008) Transmission
445	X-ray Microscopy (TXM) Reveals the Nanostructure of a Smectite Gel. Langmuir, 24,
446	8954-8958.
447	Zhang, X., Y. He, M. Sushko, J. Liu, L. Luo, J. J. D. Yoreo, S. X. Mao, C. Wang & K. Rosso (2017a)
448	Direction-specific van der Waals attraction between rutile TiO2 nanocrystals. Science,
449	356, 434-437.
450	Zhang, X., Z. Shen, J. Liu, S. N. Kerisit, M. E. Bowden, M. L. Sushko, J. J. De Yoreo & K. M. Rosso
451	(2017b) Direction-specific interaction forces underlying zinc oxide crystal growth by
452	oriented attachment. <i>Nat Commun,</i> 8, 835.
453	

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459 Author Contributions

- 460 M. L. W. analyzed and interpreted data, and wrote the manuscript. L. C. collected and analyzed data. B.
- 461 G. and J. F. B. conceived the idea and wrote the manuscript.
- 462
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- 466

467 Supplementary Information

- 468 Supplementary Movie 1
- 469 Supplementary Movie 2
- 470 Supplementary Movie 3
- 471 Supplementary Movie 4
- 472 Supplementary Movie 5
- 473

Figure 1. High-resolution phase-contrast cryo-TEM images of SWy particles suspended in water and oriented face-on to the
electron beam. The associated Fourier transforms are given in the insets. (A) A region of a particle composed of multiple layers
that are rotationally oriented at similar angles, leading to a periodic Moiré interference pattern indicated by white arrows. (B)
Region with additional layers oriented over a wider range of angles relative to neighboring layers, displaying more complex
Moiré pattern. Scale bars in images represent 5 nm.

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Figure 2. Long-range ordering of Mt particles in NaCl. (A) A particle consisting of approximately ~15 layers. The cross fringes in the top nine layers exhibit two predominant orientations, either 73° or 103° relative to the plane of the adjacent layer, consistent with <110> or $\langle \bar{1}10 \rangle$ zone axes. The cross fringes in the bottom six layers exhibit orientations aligned along the <100> direction. Scale bar represents 2.5 nm. (B) FT of (A), showing the orientation of cross fringes from top nine (white dashed lines) or bottom six (orange dashed lines) layers. The periodicity along the stacking direction (arrows) is one quarter of the basal spacing (1.9 nm, bottom arrow), indicating the presence of symmetric interlayer contrast variation with a spacing of 0.48 nm. Scale bar 1 nm⁻¹.

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489 Figure 3. Cryo-TEM images of SWy-2 particle aggregates with increasing levels of hierarchical organization. (A) Aggregate of 490 ~24 layers in face-face, face-edge, and edge-edge contact (B)-(D). (B) Face-face delamination within a particle. (C) Edge-edge 491 contact joining two adjacent particles, the bottom of which contains only two layers. (D) Face-edge contact due to the termination 492 of a layer (dislocation) at an oblique angle. (E) Particle with ~20 edge-on layers and numerous defects. (F) A 'gap' defect. (G) A 493 dislocation. (H) A wedge disclination, with layers terminating on layer indicated with arrow. (I) Particle aggregate with tens to 494 hundreds of layers. (J) Large 'gap' defect spanning multiple layers. (K) Loop dislocation in which a layer is completely enclosed 495 within particle by neighboring layers. (L) Many particles, each containing tens to hundreds of layers, joined by a spanning layer, 496 or layers, that must curve to accommodate disparate particle orientations. (A-D) Dispersed in 100 mM MgCl₂; (E-L) 200 mM 497 NaCl. Scale bars are 50 nm.

498 Figure 4. Defects observed in hydrated Mt particles. Geometric defects lead to gaps and polydisperse particles. Topological defects result from layer size polydispersity, and lead to layer curvature. Dynamic defects are likely the result of unstable configurations adopted during delamination or restacking. Curvature manifests in a variety of ways, but is typically smooth because layers have persistence lengths that are far longer than the particle dimensions.

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Figure 5 Cryo-electron tomography of SWy particle aggregates and pore networks. A) Slice through the 3D reconstruction of a single smectite particle in in MgCl₂ solution. Scale bar represents 25 nm. (B) Isosurface rendering of the Mt layers contained in the region indicated in (A); see Supplement for animation. (C) Slice through tomographic reconstruction of particle with

- 507 508 lenticular cross section in NaCl solution. (D) Isosurface rendering of interlayer space, showing interlayer connectivity at dislocation. (E) Atomistic model of pore introduced by dislocation. Tilt series movies and tomographic reconstruction animations are available in the Supplementary Information. Scale bar represents 50 nm.

1 Layer size polydispersity in hydrated montmorillonite creates multiscale porosity networks 2 Michael L. Whittaker,^{1,2†} Luis R. Comolli,^{1†} Benjamin Gilbert,¹ Jillian F. Banfield^{1,2*} 3 4 ¹Lawrence Berkeley National Laboratory, Life Sciences Division, Berkeley, California, USA. 5 6 ²Department of Earth and Planetary Science, University of California, Berkeley, California, 7 USA. * jbanfield@berkeley.edu 8 9 [†] These authors contributed equally to this work 10 11 12 Keywords: clay, montmorillonite, porosity, cryo-TEM, cryo electron tomography 13 14 Abstract 15 16 17 The aluminosilicate layers of the swelling clay mineral montmorillonite, and the saturated pores 18 they delineate, control the mechanical properties and the transport of solutes in many natural and 19 engineered environments. However, the structural basis of montmorillonite porosity remains 20 poorly characterized due to the difficulty in visualizing hydrated samples in their native state. 21 Here, we used cryogenic transmission electron microscopy (cryo-TEM) and cryo electron 22 tomography (cryo-ET) to show that stacking defects within minimally altered, fully hydrated 23 montmorillonite particles define multiscale porosity networks. Variations in layer lateral dimensions over tens to thousands of nanometers cause a range of topological and dynamic 24 25 defects that generate pervasive curvature and introduce previously uncharacterized solute 26 transport pathways. Observations of long-range rotational order between neighboring layers 27 indicate that the layer-layer interactions that govern clay swelling involve three dimensional orienting forces that operate across nanoscale pores. These direct observations of the hierarchical 28 29 structure of hydrated montmorillonite pore networks with nanoscale resolution reveal potentially

30 general aspects of colloidal interactions in fluid-saturated clay minerals.

32 1. Introduction

33 Swelling clays such as smectites are among the most abundant inorganic nanomaterials in the lithosphere(Hochella et al. 2019), and play an outsize role in controlling the transport and 34 35 retention of water, CO_2 , nutrients and pollutants in both natural and engineered settings due to 36 their exceptionally high surface areas(Bourg and Ajo-Franklin 2017, Charlet et al. 2017). SWy is 37 an archetypical smectite whose clay fraction is almost exclusively montmorillonite (Mt) that 38 formed via devitrification of volcanic ash and tuffs through submarine weathering, with no 39 significant post-formational recrystallization(Cadrin et al. 1995). Ion binding 40 selectivities(Whittaker et al. 2019), permeabilities(Tournassat et al. 2016b), shear strength(Ikari 41 et al. 2015), and myriad other properties of Mt are highly dependent on the shape, size, and the 42 specifics of how smectite particles are arranged.

43 Smectite 2:1 layers are less than one nanometer thick and up to several microns in lateral 44 dimensions, and stack to form particles in various ways depending on the relative concentrations of water, clay, and electrolyte(Tournassat and Steefel 2015). While the average separation 45 46 between layers generally decreases with increasing concentration of clay or salt(Norrish 1954) 47 there is no consensus about which of the many potential microscopic clay colloid arrangements 48 are expected to form under specific conditions(Bergaya and Lagaly 2013). Fluid and solute 49 transport rates through natural SWy depend on the pore structures defined by layer stacking 50 motifs(Wenk et al. 2008), and are generally classified into two broad categories based on the 51 separation distance between adjacent Mt layers: nanopores and macropores. Nanopores include 52 clay interlayers that, because of the effective negative charge on smectite layers, generally 53 exclude anions and therefore facilitate ion-selective transport(Tournassat et al. 2016a). 54 Macropores are the larger and less well-defined spaces between particles through which anions,

cations, and even larger solutes like macromolecules and nanoparticles can diffuse(Tournassat et
al. 2016b). In both cases, experimental descriptions of pore geometries and connectivity that are
required for accurate transport models are lacking(Churakov and Gimmi 2011, Tournassat et al.
2016b).

59 High-resolution transmission electron microscopy (HR-TEM) has been utilized for 60 decades to reveal atomic- and nanoscale structures in non-hydrous clay particles(Veblen 1985, 61 Vali and Köster 1986), including stacking order between layers in smectite and 62 illite/smectite(Veblen 1990, Guthrie and Veblen 1989). However, there is evidence that the 63 native structure of smectites is disrupted(Dudek et al. 2002) during conventional sample 64 preparation(Gu et al. 2001). Low-dose transmission electron microscopy of cryogenically frozen 65 samples (cryo-TEM) is uniquely capable of characterizing hydrated clay structures over spatial 66 scales ranging from near-atomic resolution(Whittaker et al. 2019) to whole particle 67 aggregates(Whittaker et al. 2019, Gilbert et al. 2015, Segad et al. 2012). Water vitrifies without 68 crystallization upon rapid freezing for sample thicknesses(Deirieh et al. 2018) that are electron-69 transparent at the accelerating voltages commonly employed for cryo-TEM (200-300 kV), 70 preserving structures with minimal perturbation from their native-state(Cheng 2018). Increased 71 electron-dose robustness of cryo-frozen samples(Henderson and Glaeser 1985) allows for 3D 72 images to be reconstructed using cryo electron tomography (cryo-ET) from a series of images 73 taken at different tilt angles with minimal beam-induced damage. 74 Here, we use cryo-TEM to show that native pore structures in minimally altered, 75 hydrated SWy arise from disparately-sized layers that stack defectively and introduce pervasive 76 layer curvature. We employ cryo-ET to characterize the 3D structure of SWy pore networks that

cannot be resolved from 2D images alone and are not accounted for in commonly used models

78 based on x-ray diffraction and simulations that assume perfectly planar layers. We observe 79 rotational crystallographic ordering between adjacent layers separated by over 1 nm of interlayer 80 water, in disagreement with the common assertion that hydrated smectites are fully turbostratic. 81 82 2. Methods 83 84 Wyoming bentonite (SWy-1/SWy-2) from the Clay Minerals Society was suspended in deionized water or an aqueous solution of NaCl (200 mM) or MgCl₂ (100 mM) by manual 85 86 shaking for 2 minutes with no filtration, washing or prior separation of non-clay minerals. Suspensions were incubated overnight before cryo-TEM analysis. No significant differences in 87 88 clay structure were observed for samples suspended in NaCl versus MgCl₂ and images from both 89 electrolyte solutions are presented. 90 Imaging was performed on a JEOL-3100-FFC electron microscope (JEOL Ltd., 91 Akishima, Tokyo, Japan) equipped with a FEG electron source operating at 300 kV. An Omega 92 energy filter (JEOL) attenuated electrons with energy losses that exceeded 30 eV of the zero-loss 93 peak. A Gatan 795 4Kx4K CCD camera (Gatan Inc., Pleasanton, CA, USA) was mounted at the 94 exit of an electron decelerator maintained at a voltage of 200–250 kV. Aliquots of 5 µl were 95 taken directly from the suspensions, placed onto 200 mesh lacey carbon Formvar or 300 mesh 96 lacey carbon Cu-grids (Ted Pella, INC, #01881 and #01895 respectively) and manually flash-97 frozen in liquid ethane. Grids were rapidly transferred to liquid nitrogen, and either stored or 98 immediately transferred to the microscope submerged in liquid nitrogen in a cryo-transfer stage 99 maintained at 80 K during acquisition of all data sets.

100	Images and tomographic tilt series were acquired under with doses of ~125-250 e ⁻ /Å ² ,
101	$320 \text{ e}^{-}/\text{Å}^2$, $720 \text{ e}^{-}/\text{Å}^2$, or 1,000 e $^{-}/\text{Å}^2$ per image for images with pixel sizes of 0.112 nm, 0.07 nm,
102	0.047 nm, and 0.028 nm at the specimen, respectively. Tilt series images were taken at
103	increments of 2° between $+65^{\circ}$ and $-65^{\circ} \pm 5^{\circ}$ for a total of 61-66 images with the program Serial-
104	EM (http://bio3d.colorado.edu/). A dose of 33 e ⁻ /pixel was accumulated in each image, with
105	pixel size of 0.224 nm at the specimen, corresponding to a dose of 6.7 $e^{-}/Å^{2}/image$ and a total
106	dose of ~400 e ⁻ /Å ² . Underfocus values ranged between 750 \pm 100 nm to 1,200 \pm 250 mm. The
107	electron dose limit for the onset of quantifiably high-resolution damage at 4.6 at Å is
108	approximately 4,000 $e^{-}/Å^2$.
109	All tomographic reconstructions were performed with the program etomo
110	(http://bio3d.colorado.edu/). The package of imaging tools Priism (http://msg.ucsf.edu/IVE/) was
111	used for the noise reduction and contrast enhancement. The software ImageJ 1.38x (NIH,
112	http://rsb.info.nih.gov/ij/) was used for analysis and measurements of the 2D image projections.
113	The programs Imod (http://bio3d.colorado.edu/) and UCSF Chimera
114	(https://www.cgl.ucsf.edu/chimera/) were used for 3D renderings. All movies were created with
115	the open source package ffmpeg (http://www.ffmpeg.org/). Adobe Photoshop CS5.1 was used to
116	adjust contrast in the images and to insert calibrated scale bars into images.
117	
118	3. Results
119	3.1 Rotational Ordering in Hydrated Mt

Mt particles adopted a range of orientations that reflected their structure in suspension
immediately prior to plunge-freezing. Cryo-TEM images of these suspensions exhibited two
dominant modes of contrast that varied in relative intensity depending on the orientation of the

123	layers with respect to the electron-beam axis. In a <i>face-on</i> orientation, parallel to the TEM grid
124	supports, the particles were much thinner than they were wide, and contrast was generated
125	primarily by phase interference (Figure 1). Layers that were oriented <i>edge-on</i> (within $\sim 5^{\circ}$ of the
126	beam axis(Gilbert et al. 2015)) gave strong amplitude contrast that resulted from greater
127	thickness and increased scattering along the beam direction (Figures 2, 3). Layers at oblique
128	angles to the beam, or that curved away from the beam axis (Figure 4), generated intermediate
129	contrast in 2D projection images that required 3D tomographic reconstruction to be fully
130	resolved (Figure 5). In both amplitude and phase contrast images, orientational order was
131	observed between neighboring layers.
132	Multiple stacked layers contributed to phase image contrast, creating a Moiré interference
133	pattern that contained information about the number and relative orientation of the layers
134	(Figure 1). Contrast in phase images varied with microscope settings including defocus and
135	astigmatism, as well as sample-dependent geometrical factors such as tilt angle and curvature,
136	and therefore images could not be directly compared with atomic structural models. However,
137	spots at 2.2 and 3.8 nm ⁻¹ in Fourier transformation (FT) images (Figure 1, insets) confirmed that
138	fringes originated from Mt layers. These spots corresponded to lattice planes with spacings of a
139	= 0.46 nm (020), 0.45 nm (110)/($\overline{1}10$), and 0.26 nm (130)/(200) indicating 2.6 Å to be the
140	highest resolution achievable in this study. The presence of multiple sets of spots at each spacing
141	in FT images was an indication that multiple individual layers contributed to the image contrast.
142	In some particles, five or more layers were oriented within a fairly narrow (~10°) range of angles
143	(Figure 1A, C). This small angular rotation about the stacking axis with respect to neighboring
144	particles led to a Moiré interference pattern with defined periodicity along the directions in
145	which contrast is strongest (see arrows in Figure 1A).

146	Parts of Figure 1A exhibit Moiré fringes with a spacing of approximately $d = 10$ nm,
147	suggesting an interlayer angle of $d/a \approx 2.6^{\circ}$. This is consistent with the ~1-3° separation between
148	individual spots within a cluster in the FT image. However, regions containing more layers
149	and/or those that are arranged with a much wider spread in the relative interlayer orientations
150	exhibited more complex Moiré effects that were not readily interpretable (Figure 1B). Thus,
151	phase contrast images alone carried only limited information about the number of layers per
152	particle with the same orientation (from FT spots) and the relative orientation of those layers
153	with respect to one another (from FT spots and, in certain cases, Moiré patterns).
154	Amplitude contrast images allowed both (darker) layer stacking motifs and the (lighter)
155	pores between them to be visualized directly. Regular stacking of layers with a 1.9 nm
156	periodicity, consistent with approximately 0.9 nm thick layers bounding a 1.0-1.2 nm thick
157	hydrated interlayer region(Norrish 1954), was consistently observed (Figure 2). Fringes within
158	each layer, which occasionally spanned neighboring layers, corresponded to the same 0.46 nm
159	(020) or 0.45 nm (110)/($\overline{1}10$) lattice spacings that were observed in phase images, but in this case
160	viewed along an orthogonal axis. These fringes made angles of approximately 73°, 90°, or 103°
161	with respect to the horizontal axis of the layer, depending on its relative orientation, which
162	corresponded to the $\langle 110 \rangle$, $\langle 100 \rangle$ or $\langle \overline{1}10 \rangle$ zone axes. Immediately adjacent layers often
163	exhibited similar orientations, analogous to mica-like stacking. While this order rarely persisted
164	over more than two or three layers, the stacking sequence was clearly not completely random
165	(i.e., not turbostratic).

3.2 Hydrated Pore Structures

Lower-magnification images revealed interactions between SWy particles and aggregates of many particles (**Figure 3**). Particles consisted of stacks of between two to a few hundred layers that were connected through a continuous network of face-face contacts at the modal interlayer spacing of approximately 1.9 nm. Thus, we propose that a network of equivalent basal spacings that percolates between the bounding layers constitutes a consistent definition of an individual suspended Mt crystalline hydrate particle *in aqua*.

Most Mt particles imaged in this study contained considerably more layers than has been observed in studies of extensively treated clay samples. Following exfoliation, washing, sizeseparation and drying, re-suspended smectite particles typically show a log-normal layer number distribution centered at ~7-8(Dudek et al. 2002, Whittaker et al. 2019). That the simply resuspended particles observed in this study display such pronounced differences in shape and size is evidence that they are minimally altered prior to imaging.

180 In all particles examined, layer-size polydispersity led to a panoply of distinctive defects in layer stacking structure. The most pronounced was the variation in the number of layers per 181 182 particle, which ranged from two to hundreds, shown in order of increasing number of layers in 183 Figure 3. Large particles contained a number of defects arising from the geometrical 184 consequences of the distribution of layer sizes (Figure 3E). Two layers that terminated in close 185 proximity within the same plane created a gap defect between layers with an interlayer distance 186 of 2.8 nm, more than twice the equilibrium pore spacing (Figure 3F). A layer that terminated 187 within a particle formed a dislocation (Figure 3G), giving rise to a pore with 2.8 nm spacing in 188 the immediate vicinity of the terminated layer that decreased back to 1.9 nm as adjacent layers 189 locally deformed around the defect. A sufficiently small layer created a dislocation loop in which 190 the layer was entirely contained within the particle (Figure 3K). A layer that was far longer than

average was the source of a wedge disclination (Figure 3H), on which multiple dislocations
terminated. An exceptionally long layer also spanned two stacks of layers that otherwise would
not be part of the same particle (Figure 3I), participating in both a wedge disclination and a large
gap (Figure 3 J, K). At least one layer, and possibly more, even bridged larger stacks of tens to
hundreds of layers each (Figure 3L), which bent to accommodate the different orientations in
different regions of the particle.

197 A further stacking defect, observed in the smallest particles, is the presence of face-edge 198 and edge-edge contacts rather than face-face contacts (Figure 3A-D). Between layers arranged 199 this way the pore size increased continuously until either of the bounding layers terminated. 200 These pores may be unstable structures that may arise from layers pulling away from, or 201 attaching to, each other. Although cryo-imaging provides only static snapshots of particle 202 structures, recent observations suggest that dynamic delamination and restacking of layers occurs 203 in bulk solution(Whittaker et al. 2019). Thus, this type of porosity is likely transient, and is not 204 prevalent in particles consisting of more layers.

205

206 *3.3 Layer stacking defects induce curvature*

A survey of observed defect types is provided in **Figure 4**. Many of the defects identified in **Figure 3** were associated with some degree of curvature in the layers, necessary to accommodate layer size disparity and variable particle orientations. The direction of curvature seldom changed over the length of a layer, allowing the persistence length, $\xi = -\frac{L}{\ln(\cos\theta)}$ to be calculated from 2D images by determining θ , the angle between the layer normals at each end of a layer with length *L*(Boal 2012). For layers with the highest curvature (**Figure 3L**) $\xi = 1.6 \mu m$, but in most cases was in the range of millimeters. This means that Mt layers associated with particles were

214 only gently or locally bent, since they were considerably smaller than the minimum persistence 215 length. Nonetheless, defects did distort the layers from their otherwise planar geometry. For 216 example, the termination of a layer within a particle caused curvature in neighboring layers. Layers curved gently, with large radii of curvature characterized by $r_c = \left|\frac{\delta x}{\delta \theta}\right|$, where x is the 217 218 distance along the 2D cross-section of a layer. Curvature was common in 2D images, especially 219 near dislocations and disclinations, but also near loci of delamination and restacking. In the 220 vicinity of a dislocation the local radius of curvature approached 75 nm (Figure 3L), the smallest 221 radius of curvature identified in this study.

Quantification of persistence lengths in 3D was limited by the relatively small volumesthat could be reconstructed and rendered with the available computational resources.

224 Nonetheless, important information was obtained from 3D images that was not discernable from 225 2D images. For example, layers frequently curved along two or more different radii. Figure 5A 226 depicts a 2D cross section of a 3D reconstruction shown in Figure 5B. Size polydispersity in 227 layer dimensions led to the formation of a particle in which larger exterior layers bound smaller 228 interior layers, inducing curvature of the larger layers and leading to the formation of a large pore. Gaussian curvature is defined as $K = (r_{c,1}r_{c,2})^{-1}$ where $r_{c,1}$ and $r_{c,2}$ are the principal 229 230 curvatures in orthogonal directions. The upper layer in **Figure 5B** exhibits positive Gaussian 231 curvature as it gently curves in the plane of the image and around the pore space running 232 orthogonal the plane of the image. The pore is continuous throughout the reconstructed region 233 (~300 nm \times 500 nm), but changes size due both to curvature of the upper and lower layers and 234 the intrusion of other neighboring layers. While the pore is essentially contained within a 235 particle, it has a maximum diameter of approximately 40 nm, far larger than the average 236 interlayer separation. Although particles with Gaussian curvature can sometimes be inferred

from 2D images based on the presence or absence of intermediate contrast adjacent to an edgeon layer, 3D tomographic reconstructions make the geometry of both curvature and defects
explicit for all layers, and show that they play an important role in delimiting pore structures. Tilt
series movies and tomographic reconstruction animations from which the images in Figure 5

241 were taken are available as **Supplementary Movies 1-5**.

242 The termination of individual layers within a stack also increases pore connectivity. 243 Tomograms revealed that the internal porosity within a particle was much more connected than 244 would be expected from a 2D projection image alone. The edge dislocation highlighted in 245 Figure 5C is shown in 3D in Figure 5D, with a molecular model for comparison in Figure 5E. 246 The presence of dislocations within a particle provides a diffusional pathway orthogonal to the 247 stacking direction. Dislocations also played an important role in defining the shape of a particle. 248 Lenticular cross-sections (Figure 5C) were the result of smaller layers that terminated within 249 particles, while larger layers tended to be found near the exterior. The total length of edges that 250 were exposed to bulk solution was minimized and resulted in the curving of layers to 251 accommodate the internal defects. This contrasted with particles structures of extensively treated 252 Mt, in which the opposite is true and smaller layers were generally found on the exterior of 253 particles(Tester et al. 2016, Whittaker et al. 2019).

254

255 **4. Discussion**

Using cryo-preparation methods to minimize sample preparation artifacts(Deirieh et al. 2018) and beam-induced damage during imaging(Henderson and Glaeser 1985), in analogy to cryoEM of biological samples(Cheng 2018), provides confidence that structures observed by cryo-TEM faithfully represent the native state of hydrated SWy. The observed structures are

260 therefore likely the result of authigenic processes, because SWy does not appear to have 261 recrystallized after formation(Cadrin et al. 1995). Consequently, the morphology of single 262 particles and of their aggregates characterized here provides a basis for understanding the fabrics 263 of smectite-rich soils and rocks. For example, the high density of dislocations within particles 264 that give rise to lenticular cross-sections may be the result of terminated growth as the layers 265 above and below retard transport to layers inside. Uniaxial compaction during sedimentation 266 would not be expected to alter the distribution of topological defects like dislocations 267 appreciably, but diagenetic alteration at elevated temperatures or dissolution and reprecipitation 268 in response to chemical gradients would be expected to decrease the density of such high-energy 269 defects. Identifying defect distributions in smectite particles and the factors that alter these 270 distributions raises exciting new possibilities for characterizing the evolution of smectite 271 microfabrics.

272 This work challenges assumptions made about clay particle structures based on 273 macroscopic measurements. For example, this Mt sample has an tendency for near-274 crystallographic registry at approximately 0°, 60° or 120° relative rotation angles(Plançon et al. 275 1979, Viani, Gualtieri and Artioli 2002). These findings favor the semi-ordered rotational 276 stacking initially proposed for mica(Méring 1975) and argue strongly against the long-standing 277 assumption that smectites possess complete turbostratic disorder (Moore and Reynolds 1997). 278 However, ordering through a geometrical 'lock-and-key' mechanism involving the interlayer 279 cation observed in micas is not required to orient neighboring smectite layers. Rather, long-range 280 forces between layers, which could include electrostatic interactions involving counterions, van 281 der Waals torque(Zhang et al. 2017a), or the interlayer hydrogen bond network(Meleshyn and 282 Bunnenberg 2005, Zhang et al. 2017b), appear to generate a rotational energy landscape with

modulo 60° energy minima. No rotational ordering is observed between aggregated Mt particles,
in agreement with observation that hydrated colloidal particles(Michot et al. 2013) and the
particles in smectite-rich rocks(Lutterotti et al. 2009) show preferred intraparticle orientation but
are themselves randomly oriented.

287 Nanopore spaces inside particles that were much larger than the equilibrium interlayer 288 separation were frequently observed in cryo-TEM and cryo-ET images. Such defects may 289 explain features that are common in bulk x-ray scattering and diffraction data but for which the 290 structural origin has not been treated explicitly. Mt typically exhibits (00*l*), and particularly 291 (001), peaks that are asymmetric towards lower scattering vectors, (Ferrage 2005, Segad et al. 292 2012) consistent with our observation here of a distribution layer-layer spacings that are larger 293 than the mean value due to the presence of defects. Knowledge of the defect types and 294 prevalence in Mt nanostructures may therefore be used to improve bulk characterization 295 approaches.

296 The smallest radii of curvature reported here (~75 nm) are far larger than previous reports 297 of highly curved and bent particles and layers when prepared with epoxy resin, which are similar 298 to those found in Mt-polymer composites (Dudek et al. 2002, Fu et al. 2011). Observations of 299 curvature at large angles in those materials(Zbik et al. 2008, Zbik and Frost 2010, Mouzon, 300 Bhuiyan and Hedlund 2016, Du et al. 2019) are likely due to externally applied stress that induce 301 spontaneous curvature. Molecular simulations predict that the energy penalty for bending single 302 Mt layers between carbon support layers does not exceed the thermal energy until it reaches $r_c =$ 303 15 nm(Fu et al. 2011), suggesting that the geometries of the minimally treated Mt particles 304 observed here are dominated by colloidal interactions in the absence of external stress.

305 Currently, very few of the Mt structures observed here can be predicted from first 306 principles theories of interlayer interactions. In contrast to rigid and monodisperse colloidal 307 polyhedra, for which crystal structures can be predicted based on the shape of the 308 particle(Damasceno, Engel and Glotzer 2012), simulations with system sizes smaller than the 309 average Mt layer dimension, or that do not account for layer curvature(Underwood and Bourg 310 2020, Ebrahimi, Whittle and Pellenq 2016) identify a range of microstructures that differ from 311 those observed here. In particular, the size polydispersity of flexible Mt layers causes curvature, 312 changing the energy landscape of clay particles(Suter et al. 2007, Kunz et al. 2009, Castellanos-313 Gomez et al. 2012, Honorio et al. 2018) and the pore size distribution in ways that will likely 314 require coupled chemical-mechanical models to describe.

315

316 5. Conclusions

Microstructures of minimally altered, hydrated Wyoming smectite were imaged in two 317 318 and three dimensions, revealing a panoply of defects that govern clay layer arrangements. The 319 dominant feature of Mt particles was the polydispersity in layer dimensions, which gives rise to 320 defects via incommensurate stacking. Layers curve to accommodate stacking defects, creating 321 hierarchical pore networks that can vary greatly in size distribution and can be highly 322 interconnected. Some defects appear to be unstable and may reflect dynamic processes such as 323 the delamination and restacking of individual layers. Many layers also exhibit orientational order 324 at approximately 60° intervals, which suggests the presence of long-range torsional interactions 325 between neighboring layers within particles.

We anticipate that the nanoscale observations of defect microstructures in Mt presentedherein will help link microscopic structures to the macroscopic properties of clay-rich systems.

- 328 Pore spaces elucidated here may play a significant and underappreciated role in controlling the
- 329 transport of ionic and molecular constituents through clay-rich media, and their identification
- 330 with nanoscale resolution will help inform the prediction of properties such as solute
- 331 permeability and strength at larger scales.

332 6. References

- Bergaya, F. & G. Lagaly. 2013. *Handbook of Clay Science*. Amsterdam: Elsevier.
- Boal, D. 2012. *Mechanics of the Cell*. Cambridge University Press.
- Bourg, I. C. & J. B. Ajo-Franklin (2017) Clay, Water, and Salt: Controls on the Permeability of
 Fine-Grained Sedimentary Rocks. *Acc Chem Res*, 50, 2067-2074.
- Cadrin, A. A. J., T. K. Kyser, W. G. E. Caldwell & F. J. Longstaffe (1995) Isotopic and chemical
 compositions of bentonites as paleoenvironmental indicators of the Cretaceous
 Western Interior Seaway. *Palaeogeography, Palaeoclimatology, Palaeoecology,* 119,
 301-320.
- Castellanos-Gomez, A., M. Poot, A. Amor-Amorós, G. A. Steele, H. S. J. van der Zant, N. Agraït &
 G. Rubio-Bollinger (2012) Mechanical properties of freely suspended atomically thin
 dielectric layers of mica. *Nano Research*, 5, 550-557.
- Charlet, L., P. Alt-Epping, P. Wersin & B. Gilbert (2017) Diffusive transport and reaction in clay
 rocks: A storage (nuclear waste, CO 2, H 2), energy (shale gas) and water quality issue.
 Advances in Water Resources, 106, 39-59.
- Cheng, Y. (2018) Single-particle cryo-EM—How did it get here and where will it go. *Science*, 361,
 867-880.
- Churakov, S. V. & T. Gimmi (2011) Up-Scaling of Molecular Diffusion Coefficients in Clays: A
 Two-Step Approach. *The Journal of Physical Chemistry C*, 115, 6703-6714.
- Damasceno, P. F., M. Engel & S. C. Glotzer (2012) Predictive self-assembly of polyhedra into
 complex structures. *Science*, 337, 453-7.
- Deirieh, A., I. Y. Chang, M. L. Whittaker, S. Weigand, D. Keane, J. Rix, J. T. Germaine, D. Joester &
 P. B. Flemings (2018) Particle arrangements in clay slurries: The case against the
 honeycomb structure. *Applied Clay Science*, 152, 166-172.
- Du, M., J. Liu, P. Clode & Y.-K. Leong (2019) Microstructure and rheology of bentonite slurries
 containing multiple-charge phosphate-based additives. *Applied Clay Science*, 169, 120 128.
- Dudek, T., J. Srodon, D. Eberl, F. Elass & P. Uhlik (2002) Thickness distribution of illite crystals in
 shales. I: X-ray diffraction vs. high resolution transmission electron microscopy. *Clays and Clay Minerals*, 50, 562-577.
- Ebrahimi, D., A. Whittle & R. M. Pellenq (2016) Effect of Polydispersity of Clay Platelets on the
 Aggregation And Mechanical Properties of Clay at the Mesoscale. *Clays and Clay Minerals,* 64, 425-437.

Ferrage, E. (2005) Investigation of smectite hydration properties by modeling experimental X ray diffraction patterns: Part I. Montmorillonite hydration properties. *American Mineralogist*, 90, 1358-1374.

368 Fu, Y.-T., G. D. Zartman, M. Yoonessi, L. F. Drummy & H. Heinz (2011) Bending of Layered 369 Silicates on the Nanometer Scale: Mechanism, Stored Energy, and Curvature Limits. The 370 Journal of Physical Chemistry C, 115, 22292-22300. 371 Gilbert, B., L. R. Comolli, R. M. Tinnacher, M. Kunz & J. F. Banfield (2015) Formation and 372 Restacking of Disordered Smectite Osmotic Hydrates. Clays and Clay Minerals, 63, 432-373 442. 374 Gu, B. X., L. M. Wang, L. D. Minc & R. C. Ewing (2001) Temperature e. Journal of Nuclear 375 Materials, 297, 345-354. 376 Guthrie, G. D. & D. R. Veblen (1989) High-resolution transmission electron microscopy of mixed-377 layer illite/smectite: computer simulations. Clays and Clay Minerals, 37, 1-11. 378 Henderson, R. & R. M. Glaeser (1985) Quantitative analysis of image contrast in electron 379 micrographs of beam-sensitive crystals. *Ultramicroscopy*, 16, 139-150. 380 Hochella, M. F., Jr., D. W. Mogk, J. Ranville, I. C. Allen, G. W. Luther, L. C. Marr, B. P. McGrail, M. 381 Murayama, N. P. Qafoku, K. M. Rosso, N. Sahai, P. A. Schroeder, P. Vikesland, P. 382 Westerhoff & Y. Yang (2019) Natural, incidental, and engineered nanomaterials and 383 their impacts on the Earth system. Science, 363. 384 Honorio, T., L. Brochard, M. Vandamme & A. Lebee (2018) Flexibility of nanolayers and stacks: implications in the nanostructuration of clays. Soft Matter, 14, 7354-7367. 385 386 Ikari, M. J., Y. Ito, K. Ujiie & A. J. Kopf (2015) Spectrum of slip behaviour in Tohoku fault zone 387 samples at plate tectonic slip rates. Nature Geoscience, 8, 870-874. 388 Kunz, D. A., E. Max, R. Weinkamer, T. Lunkenbein, J. Breu & A. Fery (2009) Deformation 389 measurements on thin clay tactoids. Small, 5, 1816-20. 390 Lutterotti, L., M. Voltolini, H. R. Wenk, K. Bandyopadhyay & T. Vanorio (2009) Texture analysis 391 of a turbostratically disordered Ca-montmorillonite. American Mineralogist, 95, 98-103. 392 Meleshyn, A. & C. Bunnenberg (2005) The gap between crystalline and osmotic swelling of Na-393 montmorillonite: a Monte Carlo study. J Chem Phys, 122, 34705. 394 Méring, J. 1975. Smectites. In Soil Components, ed. J. Gieseking. New York: Springer-Verlag. 395 Michot, L. J., I. Bihannic, F. Thomas, B. S. Lartiges, Y. Waldvogel, C. Caillet, J. Thieme, S. S. Funari 396 & P. Levitz (2013) Coagulation of Na-montmorillonite by inorganic cations at neutral pH. 397 A combined transmission X-ray microscopy, small angle and wide angle X-ray scattering 398 study. Langmuir, 29, 3500-10. 399 Moore, D. M. & J. R. C. Reynolds. 1997. Xray diffraction and the identification and analysis of 400 clay minerals. New York, New York: Oxford University Press. 401 Mouzon, J., I. U. Bhuiyan & J. Hedlund (2016) The structure of montmorillonite gels revealed by 402 sequential cryo-XHR-SEM imaging. J Colloid Interface Sci, 465, 58-66. 403 Norrish, K. (1954) The Swelling of Montmorillonite. 120-134. 404 Plançon, A., G. Besson, C. Tchoubar, J. P. Gaultier & J. Mamy (1979) Qualitative and 405 Quantitative Study of a Structural Reorganization in Montmorillonite After Potassium Fixation. Developments in Sedimentology, 27, 45-54. 406 407 Segad, M., S. Hanski, U. Olsson, J. Ruokolainen, T. Åkesson & B. Jönsson (2012) Microstructural 408 and Swelling Properties of Ca and Na Montmorillonite: (In Situ) Observations with Cryo-409 TEM and SAXS. The Journal of Physical Chemistry C, 116, 7596-7601.

410	Suter, J. L., P. V. Coveney, H. C. Greenwell & MA. Thyveetil (2007) Large-Scale Molecular
411	Dynamics Study of Montmorillonite Clay: Emergence of Undulatory Fluctuations and
412	Determination of Material Properties. Journal of Physical Chemistry C, 111, 8248-8295.
413	Tester, C. C., S. Aloni, B. Gilbert & J. F. Banfield (2016) Short- and Long-Range Attractive Forces
414	That Influence the Structure of Montmorillonite Osmotic Hydrates. Langmuir, 32,
415	12039-12046.
416	Tournassat, C., I. Bourg, M. Holmboe, G. Sposito & C. Steefel (2016a) Molecular Dynamics
417	Simulations of Anion Exclusion in Clay Interlayer Nanopores. Clays and Clay Minerals, 64,
418	374-388.
419	Tournassat, C., S. Gaboreau, JC. Robinet, I. Bourg & C. I. Steefel. 2016b. Impact of
420	microstructure on anion exclusion in compacted clay media. 137-149.
421	Tournassat, C. & C. I. Steefel (2015) Ionic Transport in Nano-Porous Clays with Consideration of
422	Electrostatic Effects. Reviews in Mineralogy and Geochemistry, 80, 287-329.
423	Underwood, T. R. & I. C. Bourg (2020) Large-Scale Molecular Dynamics Simulation of the
424	Dehydration of a Suspension of Smectite Clay Nanoparticles. The Journal of Physical
425	Chemistry C.
426	Vali, H. & H. M. Köster (1986) Expanding behavior, structural disorder, regular and random
427	irregular interstratification of 2:1 layer silicates studied by high-resolution images of
428	transmission electron microscopy. Clay Minerals, 21, 827-859.
429	Veblen, D. R. (1985) Direct TEM Imaging of Complex Structures and Defects in Silicates. Annual
430	Review of Earth and Planetary Science, 13, 119-146.
431	(1990) High-Resolution Transmission Electron Microscopy and Electron Diffraction of Mixed-
432	Layer Illite/Smectite: Experimental Results. Clays and Clay Minerals, 38, 1-13.
433	Viani, A., A. F. Gualtieri & G. Artioli (2002) The nature of disorder in montmorillonite by
434	simulation of X-ray powder patterns. American Mineralogist, 87, 966-975.
435	Wenk, HR., M. Voltolini, M. Mazurek, L. R. v. Loon & A. Vinsot (2008) Preferred orientations
436	and anisotropy in shales: callovo-oxfordian shale (france) and opalinus clay
437	(switzerland). Clays and Clay Minerals, 56, 285-306.
438	Whittaker, M. L., L. N. Lammers, S. Carrero, B. Gilbert & J. F. Banfield (2019) Ion exchange
439	selectivity in clay is controlled by nanoscale chemical-mechanical coupling. Proceedings
440	of the National Academy of Sciences.
441	Zbik, M. S. & R. L. Frost (2010) Influence of smectite suspension structure on sheet orientation
442	in dry sediments: XRD and AFM applications. Journal of Colloid and Interface Science,
443	346, 311-316.
444	Zbik, M. S., W. N. Martens, R. L. Frost, YF. Song, YM. Chen & JH. Chen (2008) Transmission
445	X-ray Microscopy (TXM) Reveals the Nanostructure of a Smectite Gel. Langmuir, 24,
446	8954-8958.
447	Zhang, X., Y. He, M. Sushko, J. Liu, L. Luo, J. J. D. Yoreo, S. X. Mao, C. Wang & K. Rosso (2017a)
448	Direction-specific van der Waals attraction between rutile TiO2 nanocrystals. Science,
449	356, 434-437.
450	Zhang, X., Z. Shen, J. Liu, S. N. Kerisit, M. E. Bowden, M. L. Sushko, J. J. De Yoreo & K. M. Rosso
451	(2017b) Direction-specific interaction forces underlying zinc oxide crystal growth by
452	oriented attachment. <i>Nat Commun,</i> 8, 835.
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459 Author Contributions

- 460 M. L. W. analyzed and interpreted data, and wrote the manuscript. L. C. collected and analyzed data. B.
- 461 G. and J. F. B. conceived the idea and wrote the manuscript.
- 462
- 463 Supplementary Information accompanies this paper.464
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- 466

467 Supplementary Information

- 468 Supplementary Movie 1
- 469 Supplementary Movie 2
- 470 Supplementary Movie 3
- 471 Supplementary Movie 4
- 472 Supplementary Movie 5
- 473

Figure 1. High-resolution phase-contrast cryo-TEM images of SWy particles suspended in water and oriented face-on to the
electron beam. The associated Fourier transforms are given in the insets. (A) A region of a particle composed of multiple layers
that are rotationally oriented at similar angles, leading to a periodic Moiré interference pattern indicated by white arrows. (B)
Region with additional layers oriented over a wider range of angles relative to neighboring layers, displaying more complex
Moiré pattern. Scale bars in images represent 5 nm.

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Figure 2. Long-range ordering of Mt particles in NaCl. (A) A particle consisting of approximately ~15 layers. The cross fringes in the top nine layers exhibit two predominant orientations, either 73° or 103° relative to the plane of the adjacent layer, consistent with <110> or $\langle \bar{1}10 \rangle$ zone axes. The cross fringes in the bottom six layers exhibit orientations aligned along the <100> direction. Scale bar represents 2.5 nm. (B) FT of (A), showing the orientation of cross fringes from top nine (white dashed lines) or bottom six (orange dashed lines) layers. The periodicity along the stacking direction (arrows) is one quarter of the basal spacing (1.9 nm, bottom arrow), indicating the presence of symmetric interlayer contrast variation with a spacing of 0.48 nm. Scale bar 1 nm⁻¹.

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489 Figure 3. Cryo-TEM images of SWy-2 particle aggregates with increasing levels of hierarchical organization. (A) Aggregate of 490 ~24 layers in face-face, face-edge, and edge-edge contact (B)-(D). (B) Face-face delamination within a particle. (C) Edge-edge 491 contact joining two adjacent particles, the bottom of which contains only two layers. (D) Face-edge contact due to the termination 492 of a layer (dislocation) at an oblique angle. (E) Particle with ~20 edge-on layers and numerous defects. (F) A 'gap' defect. (G) A 493 dislocation. (H) A wedge disclination, with layers terminating on layer indicated with arrow. (I) Particle aggregate with tens to 494 hundreds of layers. (J) Large 'gap' defect spanning multiple layers. (K) Loop dislocation in which a layer is completely enclosed 495 within particle by neighboring layers. (L) Many particles, each containing tens to hundreds of layers, joined by a spanning layer, 496 or layers, that must curve to accommodate disparate particle orientations. (A-D) Dispersed in 100 mM MgCl₂; (E-L) 200 mM 497 NaCl. Scale bars are 50 nm.

498 Figure 4. Defects observed in hydrated Mt particles. Geometric defects lead to gaps and polydisperse particles. Topological defects result from layer size polydispersity, and lead to layer curvature. Dynamic defects are likely the result of unstable configurations adopted during delamination or restacking. Curvature manifests in a variety of ways, but is typically smooth because layers have persistence lengths that are far longer than the particle dimensions.

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Figure 5 Cryo-electron tomography of SWy particle aggregates and pore networks. A) Slice through the 3D reconstruction of a single smectite particle in in MgCl₂ solution. Scale bar represents 25 nm. (B) Isosurface rendering of the Mt layers contained in the region indicated in (A); see Supplement for animation. (C) Slice through tomographic reconstruction of particle with

- 507 508 lenticular cross section in NaCl solution. (D) Isosurface rendering of interlayer space, showing interlayer connectivity at dislocation. (E) Atomistic model of pore introduced by dislocation. Tilt series movies and tomographic reconstruction animations are available in the Supplementary Information. Scale bar represents 50 nm.

Highlights (for review)

Optional





Figure Click here to download high resolution image







wedge disclination



Bending detects principal curvature



Gaussian curvature



persistence





Layer size polydispersity in hydrated montmorillonite creates multiscale porosity networks

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Abstract

The aluminosilicate layers of the swelling clay mineral montmorillonite, and the saturated pores they delineate, control the mechanical properties and the transport of solutes in many natural and engineered environments. However, the structural basis of montmorillonite porosity remains poorly characterized due to the difficulty in visualizing hydrated samples in their native state. Here, we used cryogenic transmission electron microscopy (cryo-TEM) and cryo electron tomography (cryo-ET) to show that stacking defects within minimally altered, fully hydrated montmorillonite particles define multiscale porosity networks. Variations in layer lateral dimensions over tens to thousands of nanometers cause a range of topological and dynamic defects that generate pervasive curvature and introduce previously uncharacterized solute transport pathways. Observations of long-range rotational order between neighboring layers indicate that the layer-layer interactions that govern clay swelling involve three dimensional orienting forces that operate across nanoscale pores. These direct observations of the hierarchical structure of hydrated montmorillonite pore networks with nanoscale resolution reveal potentially general aspects of colloidal interactions in fluid-saturated clay minerals.

Declaration of interests

 \boxtimes The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

□The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

Author Contributions

M. L. W. analyzed and interpreted data, and wrote the manuscript. L. C. collected and analyzed data. B. G. and J. F. B. conceived the idea and wrote the manuscript.