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Title

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1 Prokaryotic diversity, distribution, and insights into their role in biogeochemical cycling in
2 marine basalts

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36 **Abstract**

37 We used molecular techniques to analyze basalts of varying ages that were collected from
38 the East Pacific Rise, 9 °N, from the rift axis of the Juan de Fuca Ridge, and from neighboring
39 seamounts. Cluster analysis of 16S rDNA Terminal Restriction Fragment Polymorphism data
40 revealed that basalt endoliths are distinct from seawater and that communities clustered, to some
41 degree, based on the age of the host rock. This age-based clustering suggests that alteration
42 processes may affect community structure. Cloning and sequencing of bacterial and archaeal
43 16S rRNA genes revealed twelve different phyla and sub-phyla associated with basalts. These
44 include the Gemmatimonadetes, Nitrospirae, the candidate phylum SBR1093 in the Bacteria, and
45 in the Archaea Marine Benthic Group B, none of which have been previously reported in basalts.
46 We delineated novel ocean crust clades in the gamma-Proteobacteria, Planctomycetes, and
47 Actinobacteria that are composed entirely of basalt associated microflora, and may represent
48 basalt ecotypes. Finally, microarray analysis of functional genes in basalt revealed that genes
49 coding for previously unreported processes such as carbon fixation, methane-oxidation,
50 methanogenesis, and nitrogen fixation are present, suggesting that basalts harbor previously
51 unrecognized metabolic diversity. These novel processes could exert a profound influence on
52 ocean chemistry.

53

54 **Keywords:** Archaea/Bacteria/biogeochemical cycling/functional genes/microbial

55 ecology/prokaryotic basalt alteration.

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59 **Introduction**

60 Oceanic basalts are one of the most abundant rock types on Earth, covering upwards of
61 60% of the Earth's surface. These rocks typically have high permeabilities, which enables
62 infiltration and circulation of large quantities of seawater (Fisher, 1998; Fisher and Becker,
63 2000). The rock-seawater interaction results in a significant flux of energy and solutes between
64 basalt crust and the overlying seawater (Fisher, 1998). Recent quantitative analyses revealed that
65 basalts harbor 6×10^5 to 4×10^6 and 3×10^6 to 1×10^9 cells per g rock (Einen et al., 2008;
66 Santelli et al., 2008). In fact, Einen et al. (2008) suggested that the total number of
67 microorganisms present in ocean crust exceeds the number present in seawater. These
68 observations raise intriguing questions about the role that microorganisms play in
69 biogeochemical cycling in basalts. Biological alteration of basalt by microorganisms has been
70 the focus of numerous studies, with compelling evidence suggesting that they do play a part in
71 this process (Thorseth et al., 1995; Giovannoni et al., 1996; Fisk et al., 1998; Torsvik et al.,
72 1998; Furnes and Staudigel, 1999; Furnes et al., 2001; Banerjee and Muehlenbachs, 2003; Fisk et
73 al., 2003; Furnes et al., 2004).

74 Alteration, whether abiotic or biotic, intrinsically changes the chemistry and mineralogy
75 of the rock. For example, alteration of reactive primary minerals to secondary minerals,
76 changing rock permeabilities, and changes in the oxidation state of the rocks alters the chemical
77 milieu in which endolithic microorganisms reside. These changes may result in shifts in the
78 microbial community. Analysis of prokaryotic communities associated with marine basalts
79 revealed that several clades appear to be cosmopolitan in their distribution, as they are associated
80 with globally distributed basalts, (Mason et al., 2007; Santelli et al., 2008) regardless of rock age
81 and degree of alteration. The ubiquity of certain clades, such as the alpha-Marine Group I ocean

82 crust clade IX delineated by Mason et al. (2007), regardless of the age of the host rock, suggests
83 that overall basalt microflora do not change on a temporal scale. However, Lysnes et al., (2004)
84 reported that basalts of varying ages support different microbial phyla and sub-phyla. For
85 example, the Actinobacteria were associated with older basalts, but were absent in recently
86 erupted material. Therefore, certain clades may, in fact, respond to alteration processes, which
87 could, for example, affect the available electron donors and acceptors.

88 Certain microbial taxa may be associated with rocks of varying ages, as suggested by
89 Lysnes et al., (2004); however, it is unclear what factors contribute to this habitat specificity.
90 Fresh basalts are ~ 8 % wt FeO and 2 % wt Fe₂O₃. The increasing oxidation of reduced iron
91 with time could lead to a shift in the microbial community from oxidizers to reducers. In fact,
92 Edwards et al. (2003b) demonstrated that chemolithoautotrophic, iron-oxidizing alpha- and
93 gamma-proteobacteria isolated from sulfides and metalliferous sediments are able to grow on
94 basalt glass. These isolates are capable of using oxygen and nitrate as electron acceptors. The
95 ability to use multiple electron acceptors would be requisite as basalt alteration progresses and
96 the *in situ* redox conditions change.

97 Alternately, the reduced iron available to iron-oxidizing prokaryotes, may become
98 hydrated during fluid-rock interactions. This reaction can evolve hydrogen (Janecky and
99 Seyfried Jr, 1986; Berndt et al., 1996), which can serve as an electron donor for numerous
100 microorganisms including methanogens and sulfate reducers. In fact, Bach and Edwards (2003)
101 estimated that autotrophic sulfate reduction and methanogenesis in marine basalts could result in
102 substantial prokaryotic biomass ($9 \pm 7 \times 10^{10}$ g C/yr and $3 \pm 2 \times 10^{10}$ g C/yr, respectively).

103 While the geological characteristics of basalts, such as the availability of FeO for
104 microbial iron-oxidation, discussed above, do provide some insight into potential metabolic

105 function in this environment, examination of the *in situ* metabolic diversity of prokaryotes by
106 cultivation efforts is limited to one study. Templeton et al. (2005) isolated Mn-oxidizing,
107 heterotrophic Bacteria from Loihi Seamount. Thus, there is a need to circumvent the lack of
108 cultured microorganisms using a molecular approach to determine metabolic diversity in basalts.
109 GeoChip is a molecular tool that does not rely on cultivation based methods to assay for
110 functional diversity. Specifically, it is a functional gene microarray that has 24 243
111 oligonucleotide probes covering >10 000 genes in >150 functional groups involved in nitrogen,
112 carbon, sulfur, and phosphorus cycling (He et al., 2007). GeoChip can provide significant
113 insight into metabolic potential in a given environment, such as in marine basalts.

114 In this study we used terminal restriction fragment polymorphism (T-RFLP), cloning and
115 sequencing, and microarray analysis of functional genes to 1) assess successional changes in the
116 microbial communities associated with basalts of varying ages and from different geographical
117 locations, 2) examine species composition and distribution, and 3) determine potential metabolic
118 function in basalts by examining functional genes.

119 Our analyses revealed that rock age, or degree of alteration, may, to some degree, play a
120 role in community succession. Additionally, we report previously unrecognized phyla in basalts
121 and several novel ocean crust clades of microorganisms that may represent basalt specialists.
122 Finally, examination of functional genes in basalt revealed the genetic potential for several novel
123 metabolic processes. This analysis provides insight into biogeochemical cycling in this ocean
124 crust environment.

125

126 **Material and methods**

127 **Sample collection**

128 Glassy pillow basalts were collected from areas of low (or no) sediment accumulation
129 using the *DSV Alvin* on two separate cruises to East Pacific Rise (9 °N) and to the CoAxial
130 segment of the Juan de Fuca Ridge (JdF) and neighboring seamounts (Table 1). Basalt samples
131 were collected and placed inside a collection box, or “biobox” which was designed to prevent
132 sample exposure to ambient seawater during the ascent to the surface. Prior to the dive, the box
133 was filled with either 0.2 µm filtered seawater or sterile Millipore water. During Alvin's descent,
134 residual airspace was replaced with seawater that passed through 0.2 µm filters embedded in the
135 lid. The biobox volume (16 liters) allowed for several liters of ambient deep seawater to be
136 collected with the basalts. Once on deck, the samples were removed from the biobox using
137 sterile (flamed) tongs and placed into separate freezer bags. Samples were immediately frozen at
138 -80 °C and remained frozen until shore-based analyses. To control for deep-sea planktonic
139 organisms that may have found their way into fractures and pores in the basalt samples, the
140 biobox water was filtered and the filters were frozen and analyzed along with the basalts (see
141 below).

142

143 **Nucleic acid extraction from the basalt samples**

144 For molecular analyses all rock sample handling and all extraction steps were performed
145 in a sterile laminar flow hood. Ceramic tumbling vessels, chisels, mortar and pestles were baked
146 at 220 °C for at least 24 hours. The outer rock surface was removed by tumbling the rock several
147 times for 20 minutes, replacing with sterile grit each time. The glassy rind was pared away with
148 a chisel and/or sterile rock splitter. Approximately 1 cm³ was powdered with a tungsten mortar
149 and pestle, and 2 ml of powder was used in each extraction. Two control DNA extractions, to
150 which either 2 ml of the grit from the last tumbling step or no rock or grit material were added,

151 were used to assess contamination introduced by the tumbling steps or from the DNA extraction
152 reagents, respectively.

153 The DNA extraction protocol was optimized for extracting DNA from basalts.
154 Specifically, each extraction tube contained 2 ml of rock powder and final concentrations of the
155 following: 4.5 mM Tris·HCl, pH 7.4; 185 mM EDTA, pH 8.0; 4.5% Chelex-100 (w/v); 0.7%
156 sodium dodecyl sulfate (w/v); 2 mg proteinase K (Qiagen Inc., Valencia, CA). Tubes were
157 placed in a 37 °C rotator with gentle agitation (180 rpm) overnight. The rock powder was
158 separated from the supernatant through low speed centrifugation and mixed with final
159 concentrations of the following: 923 mM NaCl and 1.3% CTAB (cetyltrimethylammonium
160 bromide; w/v). Samples were incubated at 65°C for 30 min, extracted once with an equal
161 volume of phenol/chloroform/isoamyl alcohol (25:24:1), pH 7.4, and then extracted twice with
162 an equal volume of chloroform/isoamyl alcohol (24:1). Nucleic acids were precipitated by
163 adding one volume of isopropanol and incubating the samples overnight at 4°C. Samples were
164 spun at 24,000 g for 1 hr at 4°C in a TL-100 ultracentrifuge (Beckman Instruments, Inc., Palo
165 Alto, CA) to pellet the precipitated nucleic acids. Pellets were washed with 70% ethanol (v/v),
166 dried in a laminar flow hood at room temperature, and suspended in sterile water. Replicate
167 extracts were combined (40 µl final vol.) and stored at -80°C.

168

169 **DNA extraction from the filtered biobox water samples**

170 At least 12 L of biobox water from each dive was filtered through a 142 mm 0.2 µm
171 Supor filter (Pall Gelman Laboratory, Ann Arbor, MI) in a polycarbonate filter holder (Geotech
172 Environmental Equipment, Inc. Denver, CO) that was connected to a peristaltic pump. Filters
173 were immediately preserved in 5 ml of sucrose lysis buffer (20 mM EDTA, 400 mM NaCl, 0.75

174 M sucrose, 50 mM Tris·HCl, pH 9.0) and stored at -80 °C. Total community nucleic acids were
175 extracted from the filters according to Giovannoni et al. (1990).

176

177 **Terminal-restriction fragment length polymorphism (T-RFLP) analysis**

178 T-RFLP analysis was used to compare the archaeal and bacterial communities from
179 several rock and corresponding biobox seawater samples according to Moeseneder et al. (1999)
180 with few modifications. The archaeal 16S rRNA genes were amplified using the primers
181 Arch20F (DeLong et al., 1999) and Arch915R (Stahl and Amann, 1991), with the forward primer
182 5' end-labeled with phosphoramidite fluorochrome 5-carboxy-fluorescein (6-FAM) and the
183 reverse primer labeled with 5-hexachlorofluorescein (5-HEX). Fifty PCR cycles were necessary
184 to amplify archaeal DNA, while a semi-nested approach was required to amplify bacterial 16S
185 rRNA genes from nearly all of the basalts, with primers 27F-B (5'-
186 AGRGTTYGATYMTGGCTCAG) and 1492RY (5'-GGYTACCTTGTTACGACTT) modified
187 from (Lane, 1991) used in the initial PCR reaction (30 cycles), and primers 27F-B-[FAM] and
188 1391R (Lane, 1991) used in the second reaction (20 cycles). Only the forward strand of this
189 PCR product was labeled for the T-RFLP analysis. For both archaeal and bacterial
190 amplifications, three replicate PCR reactions (50 µl) for each DNA sample contained final
191 concentrations of the following: 1µl of DNA extract; 1% (v/v) PCR buffer (+ NH₄SO₄; MBI
192 Fermentas, Hanover, MD); 0.2 mM each deoxynucleotide triphosphate; 0.2 µM each primer; 2
193 mM MgCl₂ (MBI Fermentas); 1.2 mg ml⁻¹ bovine serum albumin (non-acetylated, SIGMA); 1%
194 (wt/v) PVP (polyvinylpyrrolidone); 2.5 U Taq polymerase (MBI Fermentas). PCR cycling
195 consisted of denaturation at 94 °C for 1.5 min, annealing at 55 °C for 1.5 min, and extension at

196 72 °C for 1.5 min. The filtered biobox seawater samples from each dive were also analyzed
197 using the same cycling conditions as the basalts, except the number of cycles was reduced to 30.

198 PCR products (50 ng) were digested with 10 units of enzyme for 6 hrs at 37 °C with each
199 of three separate restriction enzymes: AluI, BsuRI (HaeIII), and Hin6I (HhaI) (MBI Fermentas).
200 Samples were run on an ABI 3100 (Applied Biosystems, Inc. (ABI), Foster City, CA). The
201 fingerprint patterns for the rock and seawater communities were compared according to
202 Moeseneder et al. (1999); however, only peaks longer than 70 bp in length were included in the
203 analysis. Data were standardized by inclusion of peaks that represented >1% of the total peak
204 height for each fingerprint and were then converted to binary matrices. Binary data were
205 analyzed by the unweighted pair group with mathematical averages (UPGMA) method in
206 PAUP* (Phylogenetic Analysis Using Parsimony *(and Other Methods)) version 4.0 b10
207 (Swofford, 1998) using the site distance matrix method of Nei and Li (1979) according to
208 Moeseneder et al. (1999).

209

210 **PCR amplification and cloning of prokaryotic 16S rRNA genes**

211 Data from the UPGMA analysis was used to select three basalt samples that differed in
212 age and community structure (D3718B, 9 °N EPR, D3815F and D3823M Juan de Fuca) for
213 cloning and sequencing of archaeal and bacterial 16S rRNA genes. The archaeal communities
214 were amplified according to the PCR conditions described above for T-RFLP analysis, except
215 the primers were not fluorescently labeled. To amplify archaeal 16S rDNA from D3815F a
216 semi-nested approach was employed using the primers Arch20F and 1492RY in the initial
217 reaction and Arch20F and Arch915R in the semi-nested reaction. The bacterial community from
218 D3718B was amplified using the semi-nested approach described above. Amplification of

219 bacterial 16S rDNA from D3823M did not require a semi-nested approach. As with the T-RFLP
220 analysis, the corresponding seawater samples from each dive were also cloned for comparison.
221 PCR reactions (50 µl vol) were cloned into the pGEM®-T Easy vector (Promega Corp.,
222 Madison, WI). Clone libraries were constructed and screened according to the methods of
223 Vergin et al. (2001). Briefly, clones were assigned to clone families based upon shared patterns
224 for two separate restriction digests. Digested PCR products were resolved on a 3% agarose gel.
225 One clone from each unique RFLP pattern was sequenced using an ABI 3730 capillary
226 sequencer. Full-length sequences were obtained for clones representing each phylotype. Clones
227 with restriction patterns that only appeared once in the library were designated “unique.”
228 Percent coverage was calculated based upon the number of unique clones versus total clones
229 according to the method of Good (1953). Chimeric sequences were identified with the
230 CHECK_CHIMERA program (Maidak et al., 1997; Maidak et al., 1999) and Mallard (Ashelford
231 et al., 2006).

232

233 **Phylogenetic analysis**

234 The phylogenies of microorganisms from D3718B and D3823M were extensively
235 reviewed in Mason et al. (2007). Clones from these libraries are presented here in phylogenetic
236 dendrograms only if they are part of novel ocean crust clades delineated here, or if they are
237 highly similar to clones from D3815F. However, clones from all libraries were analyzed during
238 phylogenetic reconstruction. Phylogenetic analyses and clade delineations were carried out
239 according to Mason et al. (2007), using the Greengenes database (DeSantis et al., 2006). Briefly,
240 neighbor-joining, maximum parsimony, and maximum-likelihood trees of near full-length
241 sequences were generated in ARB (Ludwig et al., 2004). Maximum-likelihood trees were

242 generated using Tree-Puzzle, (Schmidt et al., 2002) with the Hasegawa-Kishino-Yano model
243 (Hasegawa et al., 1985). Quartet-puzzling (QP) reliability values are not shown at bifurcations if
244 they are below 50%. In determining clades QP values from 90 to 100% are strongly supported;
245 however, QP values less than 70% can also be trusted (Schmidt *et al.*, 2002). Clades delineated
246 here with QP values lower than 70% were analyzed relative to QP support values of the other
247 branches in the tree (Schmidt *et al.*, 2002).

248

249 **Nucleotide sequence accession numbers**

250 The 16S rRNA gene sequences for the archaeal and bacterial clones were submitted to
251 the GenBank database and have been assigned the following accession numbers: DQ070750 to
252 DQ070835 (D3718B and D3823M) and FJ024305-FJ024341 (D3815F).

253

254 **Functional genes**

255 Basalt sample D3815F was selected for functional gene analysis because it had several
256 clades that have not been previously reported from this environment, particularly the Marine
257 Benthic Group B. We hypothesized that this diversity of species would be mirrored in the
258 diversity of functional genes. Second, thin sections of this sample showed textures that suggest
259 bioalteration; therefore, analysis of functional genes in this sample would provide insight into the
260 biological processes that may result in these textural features. Functional genes were assayed for
261 using the GeoChip 2.0 (He et al., 2007) microarray following previously described methods (Wu
262 et al., 2006; He et al., 2007). Briefly, DNA from D3815F was amplified in triplicate using a
263 Templiphi 500 amplification kit (Amersham Biosciences, Piscataway, NJ) following the
264 manufacturer's protocol. To facilitate amplification spermidine ($0.1 \mu\text{g } \mu\text{l}^{-1}$) and single-strand

265 binding protein (0.04 mM) were added to the reaction mixture. Amplified DNA was
266 fluorescently labeled with Cy5. Hybridizations were performed using a HS4800Pro
267 Hybridization Station (TECAN, US, Durham, NC) overnight at 42 °C. Microarrays were
268 scanned using a ProScanArray (PerkinElmer, Waltham, MA). Images were then analyzed using
269 ImaGene 6.0 (BioDiscovery, El Segundo, CA) to designate the identity of each spot and to
270 determine spot quality. Data was processed as described by Wu et al (2006). Briefly, raw data
271 from Imagen was analyzed using a GeoChip data analysis pipeline. A signal to noise ratio of \geq
272 3 was considered a positive signal. A positive signal in at least 1/3 of the probes for a particular
273 gene (minimum of 2 probes) was required for a gene to be considered positive. Each gene had 1,
274 2, or 3 probes per array based on the number of probes available meeting the criteria described
275 by He et al. (2007)

276

277 **Results and Discussion**

278 **T-RFLP**

279 UPGMA cluster analysis of T-RFLP data revealed that the archaeal and bacterial
280 communities were distinct from deep seawater communities (Figure 1). Further, there was
281 striking congruency in the UPGMA clustering patterns for the four oldest JdF samples. These
282 old samples, ranging from a few thousand to about three million years in age, clustered together,
283 while the younger basalts from 9 °N (from an eruption in 1991) clustered with one JdF sample of
284 a similar age (D3826U, from the 1993 lava flow). This clustering is evidence that there are
285 differences in microbial communities present in recently erupted basalts compared to older, more
286 weathered rocks. The observed clustering is supported, to some degree, by phylogeny. For
287 example, the Planctomycetes ocean crust clade XIV members (see below) are from recently

288 erupted to medium-aged basalts. Overall, however, there is distinct overlap in the microbial
289 communities regardless of rock age. For example, the basalt specific ocean crust clade presented
290 here, such as the gamma-Proteobacteria ocean crust clade XII, is composed of microorganisms
291 from young, fresh basalts to 3.3 Ma year old basalts. This pattern suggests that basalt microflora
292 are largely associated with rocks of varying ages, but that a minority may reside in, for example,
293 younger, less altered rocks to the exclusion of older, more weathered rocks. This finding is
294 consistent with that of Lysnes et al. (2004), who reported that specific bacterial species are found
295 only in rocks of a certain age.

296

297 **Phylogenetic analysis**

298 A total of 547 bacterial and archaeal 16S rDNA clones were analyzed and 173 unique
299 clones were sequenced (Table 2). This analysis revealed that Gemmatimonadetes, Nitrospirae,
300 SBR1093, and in the Archaea the Marine Benthic Group B (Figures 2, 3, and 4) were present in
301 basalt samples. None of these clades have been previously reported in marine basalts.
302 Additionally, microorganisms in the alpha-, delta-, and gamma-Proteobacteria, Acidobacteria,
303 Actinobacteria, Bacteroidetes, and Planctomycetes in the bacterial domain are reported (Figures
304 2 and 3). The most prevalent microorganisms were Proteobacteria (56%), the majority of which
305 were gamma- (25%), alpha- (15%), and delta- (13%), followed by the Bacteroidetes (10%),
306 Actinobacteria (9%), Planctomycetes (7%), Acidobacteria (6%), and Gemmatimonadetes (3%).
307 The remaining clades were observed in a single rock sample.

308 Our observations are consistent with those reported by Santelli et al. (2008) who analyzed
309 basalts from the East Pacific Rise and from Hawaii and found 68%/66% (EPR%/Hawaii%)
310 Proteobacteria, 8%/5% Planctomycetes, 7%/8% Actinobacteria, 4%/1% Bacteroidetes, and

311 3%/4% Acidobacteria. The similarity in bacterial communities associated with basalts from a
312 broad geographic distribution suggests cosmopolitan distributions of these clades, which is in
313 agreement with findings presented by Mason et al. (2007) and Santelli et al. (2008).

314 Phylogenetic reconstruction revealed three novel ocean crust clades composed entirely of
315 microorganisms associated with basalt. These new clades are the gamma-Proteobacteria ocean
316 crust clade XII (Figure 2), Actinobacteria ocean crust clade XIII (Figure 3), and Planctomycetes
317 ocean crust clade XIV (Figure 3). These clades are comprised of Bacteria sampled from Juan de
318 Fuca (this study), East Pacific Rise, 9 °N (this study and Santelli et al., 2008) and Hawaiian
319 (Santelli et al., 2008) basalts. These cosmopolitan basalt clades may represent ecotypes of
320 Bacteria that are specifically adapted to this environment.

321 Cloning and sequencing of Archaeal 16S rDNA revealed that Marine Benthic Group B
322 (MBGB) were present in basalts (Figure 4). This is the first report of this clade in this
323 environment, as previous studies that examined the archaeal communities in basalts revealed
324 only Marine Group I Crenarchaeota (MGI) (Thorseth et al., 2001; Fisk et al., 2003; Lysnes et al.,
325 2004; Mason et al., 2007). Recently, quantitative analyses of the microbial communities in
326 basalts revealed that Archaea comprise 4-12% and 0.02% or less of the prokaryotic communities
327 (Einen et al., 2008; Santelli et al., 2008), respectively. While these estimates are disparate they
328 do reveal that Archaea are a minor component in the overall microbial communities that reside in
329 basalt. Although Archaea are less prevalent they are ubiquitous in basalts and have been
330 reported in all studies that assayed for their presence (Thorseth et al., 2001; Fisk et al., 2003;
331 Lysnes et al., 2004; Mason et al., 2007; Einen et al., 2008; Santelli et al., 2008). Further, as
332 discussed previously a clade of Marine Group I Archaea appear to be endemic to basalt (Mason
333 et al., 2007). This habitat specificity and global distribution indicates that some Archaea, while

334 less abundant than Bacteria, are particularly adapted to life in basalt and likely play a role in
335 biogeochemical cycling.

336

337 **Functional genes**

338 GeoChip (He et al., 2007) microarray analysis of functional genes in basalt sample
339 D3815F revealed the presence of genes coding for metabolic processes previously unrecognized
340 in this environment. In this analysis a total of 604 probes of the 24 243 total probes present on
341 GeoChip were positive. Specifically, we found genes coding for carbon fixation, methane
342 production and oxidation, nitrogen fixation, ammonium-oxidation, nitrate and nitrite reduction,
343 dissimilatory sulfate reduction, and iron reduction (see Supplementary Table 1 for a complete
344 list).

345 Here we report genes coding for carbon fixation. Basalts lacking a sediment layer are
346 considered to be an oligotrophic, low carbon environment (Edwards et al., 2003a), thus carbon
347 cycling in this habitat is particularly significant. The oligotrophic nature of this environment
348 suggests that carbon fixation would be paramount in this habitat. In fact, chemolithoautotrophic
349 processes in marine subsurface ridge flank hydrothermal environments have been theoretically
350 shown to provide energy that could result in significant microbial biomass ($\sim 1 \times 10^{12}$ g C yr⁻¹)
351 (Bach and Edwards, 2003). Therefore, chemolithoautotrophic processes occurring *in situ* could
352 serve to underpin a basalt hosted biosphere. One such process is methanogenesis, where
353 hydrogen can serve as the electron donor to reduce carbon dioxide, evolving methane. During
354 fluid-rock interactions when the basalt minerals olivine and pyroxene react with water, hydrogen
355 may be evolved (Janecky and Seyfried Jr, 1986; Berndt et al., 1996). Thus the requisite electron
356 donor may be present as a result of this abiotic reaction.

357 Here we report that genes coding for methanogenesis are present in basalt. Methanogens
358 have not been reported in molecular analyses of Archaea in basalts conducted to date (Thorseth
359 et al., 2001; Fisk et al., 2003; Lysnes et al., 2004; Mason et al., 2007). However, Lysnes et al.,
360 (2004) reported that methane was evolved in enrichment cultures inoculated with marine basalts.
361 Although the Marine Benthic Group B clade currently lacks a cultured representative (Knittel et
362 al., 2005) they are frequently associated with environments dominated by methane,
363 methanogens, and methanotrophs (Knittel et al., 2005; Kendall and Boone, 2006; Kendall et al.,
364 2007). The role of this clade in the environment is unknown, but it is plausible that they are
365 involved in methane biogeochemical cycling. Although no known methanogens were observed
366 in our study the diversity of *mcr* genes (Supplementary Table 1) in conjunction with a clade
367 typically observed in methane rich environments suggests that this metabolic process may be
368 occurring in basalts. Methane resulting from biological processes could serve as a carbon and
369 energy source for heterotrophic processes. In fact, we found genes coding for methane-
370 oxidation. Methane cycling in marine basalts would have a direct impact on the overlying
371 hydrosphere.

372 As discussed above, basalts are not carbon replete. Similarly they are composed of only
373 a small amount of nitrogen, averaging approximately 2 ppm (Marty et al., 1995). Therefore, the
374 detection of genes coding for nitrogen fixation is intriguing. Cowen et al. (2003), and more
375 recently Huber et al. (2006), investigated ocean crust fluids and reported elevated levels of
376 ammonium compared to seawater. Cowen et al. (2003) suggested that nitrogen fixation may
377 serve as the source of this excess ammonium. Mehta et al. (2005) attributed nitrogen fixation in
378 crustal fluids and in deep seawater to non-methanogenic Archaea, which are the only known
379 archaeal nitrogen fixers. In that study, *nifH* genes were detected in crustal fluids. Nitrogen

380 fixation may also be taking place in the host rocks themselves given the presence of *nifH* genes
381 in our basalt sample.

382 Nitrogen fixation could augment the low nitrogen concentrations in basalts and may
383 ultimately support ammonium oxidizing microorganisms. This hypothesis is supported by the
384 presence of genes that code for ammonium-oxidation in our basalt sample. As reported by
385 Mason et al. (2007) (see Figures 3 and 4), basalt sequences similar to *Nitrosococcus oceani* (89%
386 similar) and *Nitrospira multiformis* (96% similar), both of which are known ammonium-
387 oxidizing microorganisms (Watson, 1965; Watson et al., 1971), were derived from basalts from
388 Juan de Fuca and 9 °N (this study), and Mohns Ridge (Einen et al., 2006). Thus phylogenetic
389 and functional gene analyses both suggest that ammonium-oxidation may be occurring in basalts.

390 Further, basalt clones closely related to the nitrite oxidizing *Nitrospina gracilis* (91-94%
391 similar) and *Nitrospira marina* (95-96% similar) (Watson and Waterbury, 1971; Tal et al., 2003)
392 were found (Figures 2 and 3), suggesting that nitrite-oxidation, the second step in nitrification,
393 may be occurring in basalts. This observation could not be confirmed using GeoChip; however,
394 because genes coding for nitrite-oxidation are not present on the gene chip.

395 Nitrification could provide the substrate for both denitrification and anaerobic
396 ammonium- oxidation (anammox), both of which lead to loss of nitrogen (Lam et al., 2007). In
397 fact, we found numerous genes coding for nitrate and nitrite reduction; therefore, that the genetic
398 potential for denitrification is present in this environment. Recently, Edwards et al., (2003b)
399 demonstrated that chemolithoautotrophic iron-oxidizing Bacteria are able to grow on basalt glass
400 using nitrate as the electron acceptor. Whether anaerobic ammonium-oxidation is occurring in
401 basalts remains unclear. Although Planctomycetes have been reported in basalts, including the
402 novel ocean crust clade Planctomycetes XIV delineated here, microorganisms closely related to

403 known anammox Bacteria, such as *Kuenenia stuttgartiensis*, (77% similar to basalt associated
404 microorganisms), have not been detected. Therefore, it is unclear if this process is important in
405 considering nitrogen loss from the basalt layer. Our data does suggest, however, that nitrogen
406 could be lost from marine crust by denitrification processes.

407 In addition to genes coding for denitrification processes, we also detected genes coding
408 for iron-reduction and dissimilatory sulfate reduction in basalt. Together, these genes suggest
409 that anaerobic respiration may be occurring in basalt. The presence of genes that code for
410 aerobic respiration (e.g. ammonium-oxidation) in the same sample indicates that aerobic and
411 anaerobic processes may occur simultaneously on a small spatial scale, suggesting, perhaps that
412 microniches are occupied by prokaryotes in basalt. Consistent with our findings, it was reported
413 that in upper basaltic crust redox conditions are such that aerobic and anaerobic processes are
414 likely supported (Bach and Edwards, 2003).

415

416 **Conclusion**

417 Basalts from Juan de Fuca, neighboring seamounts, and 9 °N, EPR harbor cosmopolitan
418 microorganisms that are distinct from seawater prokaryotes. Several novel ocean crust clades
419 composed only of microorganisms from basalts suggest that some Bacteria are specifically
420 adapted to this ocean crust environment. Our analysis of geochemically important functional
421 genes revealed the potential for several metabolic processes not known to be occurring in basalts,
422 particularly carbon fixation, methanogenesis, methane-oxidation, nitrogen fixation and
423 denitrification. Our data suggests that basalts not only harbor a diversity of broadly distributed
424 microbial species, but also unexpected metabolic diversity. Future studies should utilize culture-
425 dependent and -independent methods to analyze biogeochemical cycling in basalts to better

426 understand the biological processes in this vast subsurface environment and how these processes
427 ultimately affect ocean chemistry.

428

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441

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647 **Figure legends**

648

649 **Figure 1.** UPGMA analysis of T-RFLP fingerprint patterns for the Bacterial (left) and Archaeal

650 (right) communities recovered from basalts (above the dashed line) and background seawater

651 from 9 °N and JdF (below the dashed line). Older basalts (> 20 years) from JdF are underlined.

652 All 9 °N samples are less than 20 years. Sample numbers indicate Alvin dive number and

653 location: 9N is 9 °N on the East Pacific Rise and JdF is Juan de Fuca Ridge and Cobb Seamount.

654 The scale bar represents similarity.

655

656 **Figure 2.** Maximum-likelihood phylogenetic tree of proteobacterial 16S rRNA gene sequences

657 from basalt samples. The Proteobacteria tree was constructed with 25 000 puzzling steps. A

658 general Bacteria filter was used. The 16S rDNA sequence of *Aquifex pyrophilus* (M83548) was

659 used as the outgroup (not shown). The alpha-Proteobacteria ocean crust clade II, designated by

660 an (*), was delineated by Mason et al. (2007) and is included here because new basalt sequences

661 extend this clade. The scale bar indicates 0.1 nucleotide substitutions per site.

662

663 **Figure 3.** Maximum-likelihood phylogenetic tree of Actinobacteria, Cyanobacteria,

664 Bacteroidetes, Planctomycetes, Gemmatimonadetes, Acidobacteria, Nitrospirae, and SBR1093

665 16S rRNA gene sequences from basalt samples. The Bacteria tree was constructed with 25 000

666 puzzling steps. A general Bacteria filter was used. The 16S rDNA sequence of *Aquifex*

667 *pyrophilus* (M83548) was used as the outgroup (not shown). The scale bar indicates 0.1

668 nucleotide substitutions per site.

669

670 **Figure 4.** Maximum-likelihood phylogenetic tree of archaeal 16S rRNA gene sequences from
671 basalt samples. The Archaea tree was constructed with 25 000 puzzling steps. A general
672 Archaea filter was used. The 16S rDNA sequence of *Aquifex pyrophilus* (M83548) was used as
673 the outgroup (not shown). The alpha-MGI ocean crust clade VIII, designated by an (*), was
674 delineated by Mason et al. (2007) and is included here because new basalt sequences extend this
675 clade. The scale bar indicates 0.1 nucleotide substitutions per site.
676

TABLE 1. Basalt samples collected from the East Pacific Rise and the Juan de Fuca Ridge.

Alvin dive ^a	Date	Latitude	Longitude	Depth (m)	Dive feature ^b	Age ^c	Molecular analyses
East Pacific Rise (EPR) R/V Atlantis Voyage 7 Leg 3							
D3713C	10/19/01	09° 50.80' N	104° 17.63' W	2493	base of Q vent	1991 eruption	T-RFLP
D3716A	10/22/01	09° 50.30' N	104° 17.51' W	2499	axial caldera	"	T-RFLP
D3718B	10/24/01	09° 50.78' N	104° 17.58' W	2493	north of Q vent	"	T-RFLP, cloning & sequencing ^d
D3719D	10/25/01	09° 50.78' N	104° 17.58' W	2496	near M vent	"	T-RFLP
D3720R	10/26/01	09° 50.78' N	104° 17.58' W	2498	near TY vent	"	T-RFLP
D3721D	10/27/01	09° 50.79' N	104° 17.59' W	2495	near Q vent	"	T-RFLP
D3721E	10/27/01	09° 50.79' N	104° 17.59' W	2496	near Q vent	"	T-RFLP
Juan de Fuca Ridge (JdF) R/V Atlantis Voyage 7 Leg 19							
D3815F	8/05/02	45° 59.50' N	129° 56.59' W	2135	Helium Basin	<100 Ka	T-RFLP, cloning & sequencing ^e
D3816F-1,2	8/06/02	46° 31.34' N	129° 29.94' W	2653	Co-Axial Rift	10-170 Ka	T-RFLP
D3823M	8/19/02	46° 41.95' N	130° 55.94' W	1909	Cobb Seamount	3.3 Ma	T-RFLP, cloning & sequencing ^d
D3826U	8/23/02	46° 31.16' N	129° 34.92' W	2409	lava flow	1993 eruption	T-RFLP

679 ^aThe Alvin dive number and suffix is the sample identifier.

680 ^bThe 9 °N samples were collected from the area of the 1991 eruption (Haymon et al., 1993) and were 11
681 years old at the time of collection. The region of EPR vents is described in Fornari and Embley (1995).

682 ^cThe ages of Juan de Fuca samples from Helium Basin and Co-Axial Rift, were inferred from seafloor
683 spreading rate and distance from the ridge axis. Age of the Cobb Seamount sample from Desonie and
684 Duncan (1990). D3826U was collected from the 1993 lava flow (Embley et al., 2000).

685 ^dD3718B and D3823M clones were analyzed and presented in Mason et al. (2007) and are only included
686 in dendrograms in this study if they are part of novel clades delineated here, or are closely related to
687 clones from D3815F.

688 ^eD3815F clones are presented in this study.

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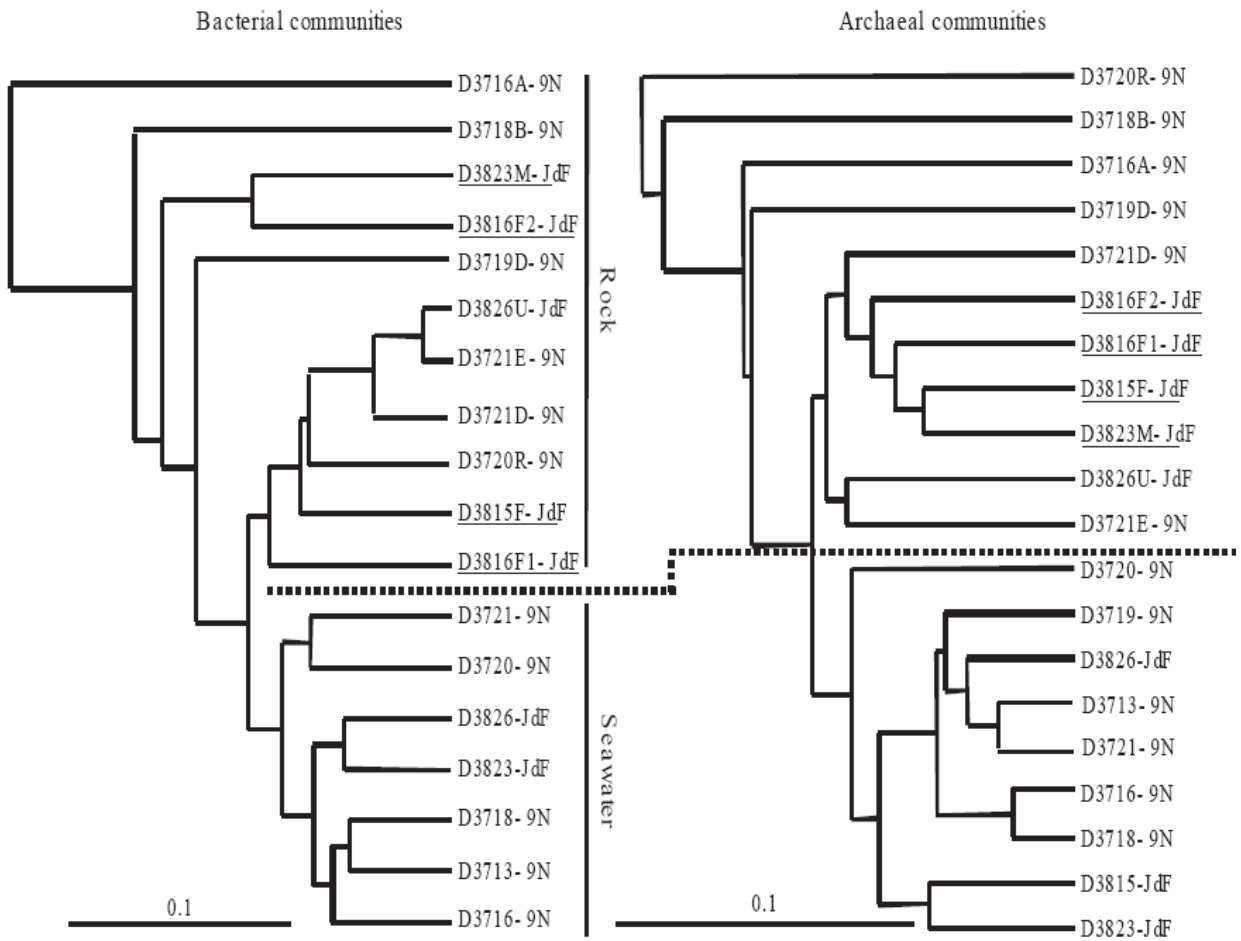
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Table 2. Clone libraries of archaeal and bacterial 16S rRNA genes from basalts and seawater.

Sample ID	Sample site	Sample type	Prokaryote domain	Clone families	Unique /total	Good's coverage (%)
D3718B	9 °N	Basalt	Archaea	20	11/94	88.3
D3718F	9 °N	Seawater	Archaea	22	13/96	86.5
D3815F	JdF	Basalt	Archaea	33	24/95	74.7
D3823M	JdF	Basalt	Archaea	20	13/95	86.3
D3718B	9 °N	Basalt	Bacteria	10	64/93	31.2
D3718F	9 °N	Seawater	Bacteria	10	44/89	50.5
D3815F	JdF	Basalt	Bacteria	43	25/96	73.4
D3823M	JdF	Basalt	Bacteria	15	38/74	48.6



698 **Figure 1.**

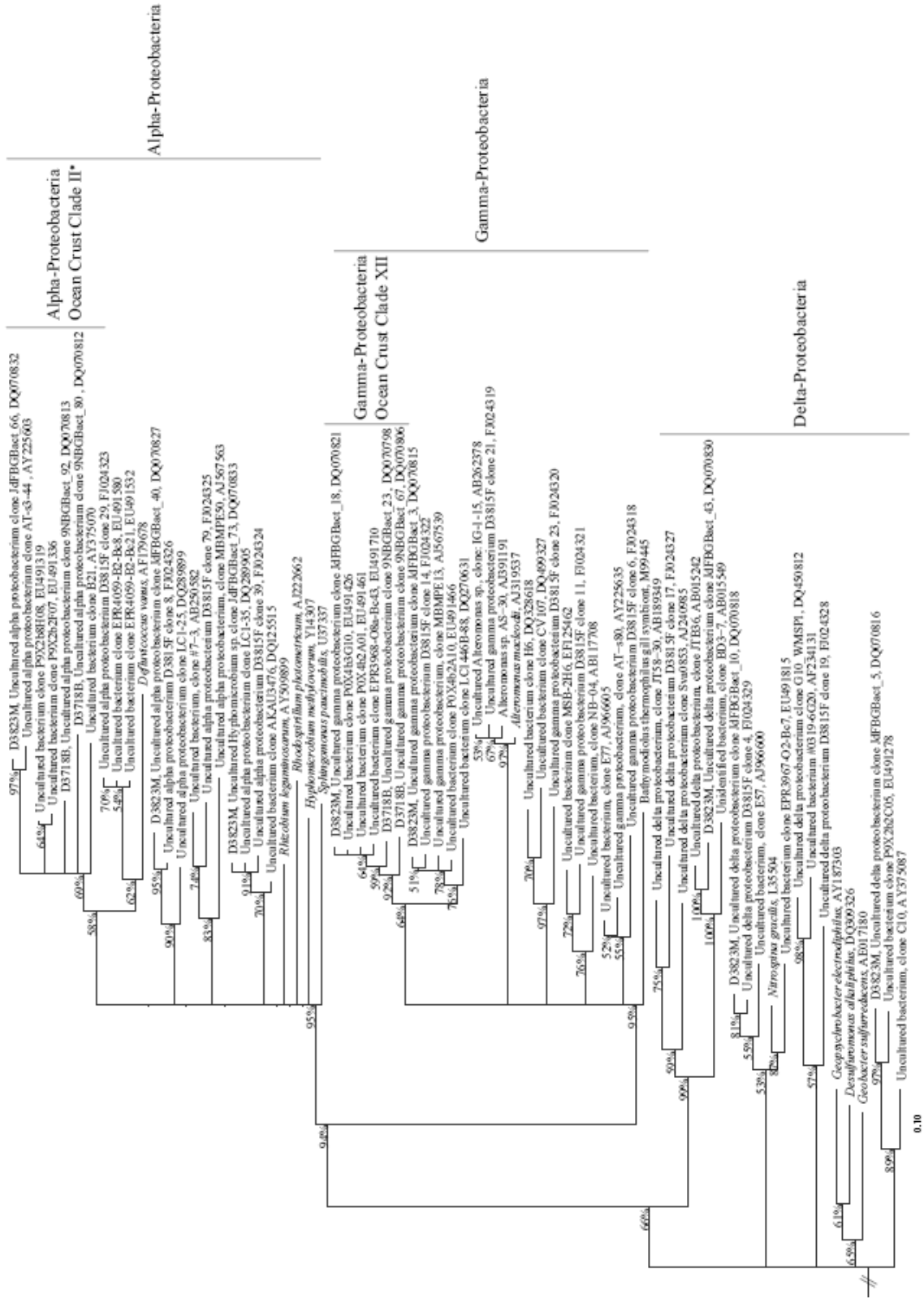


Figure 2.

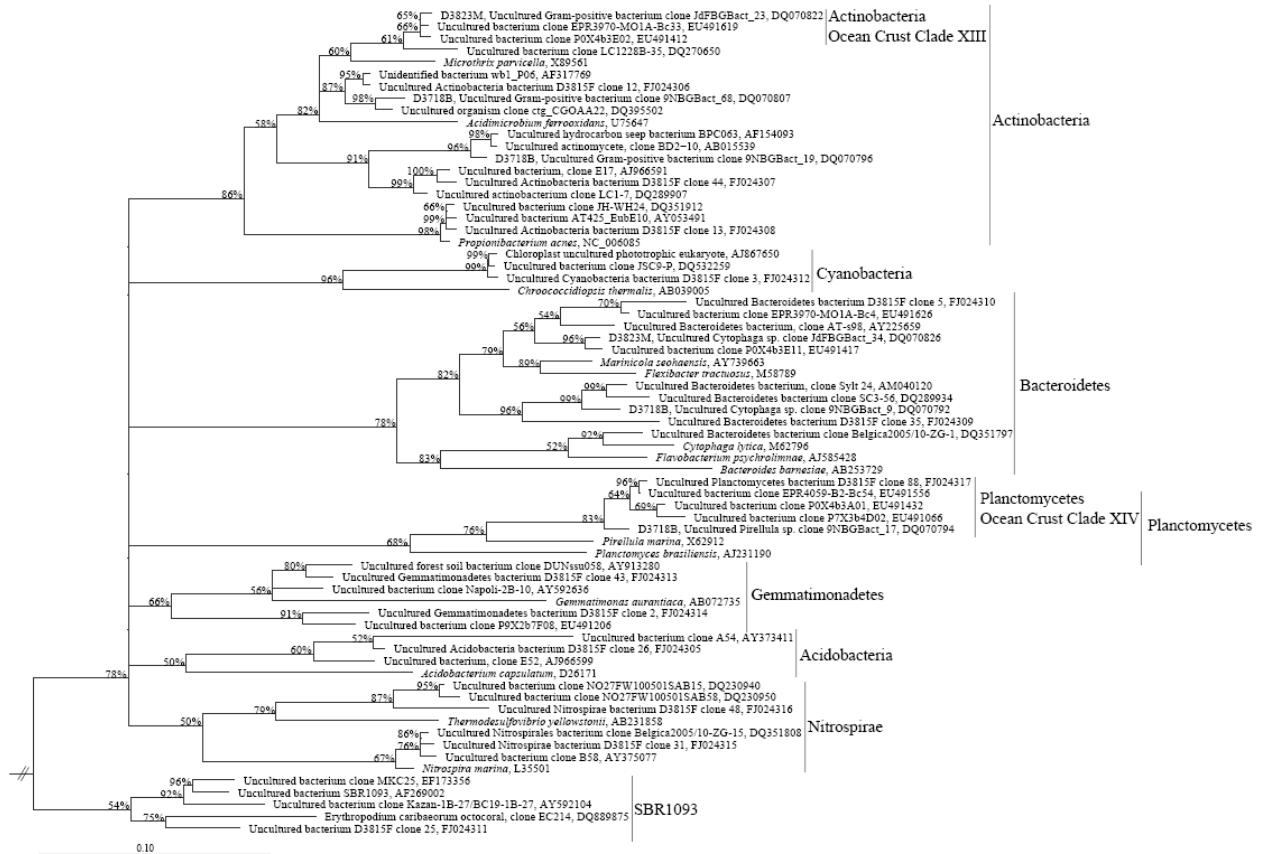


Figure 3.

700

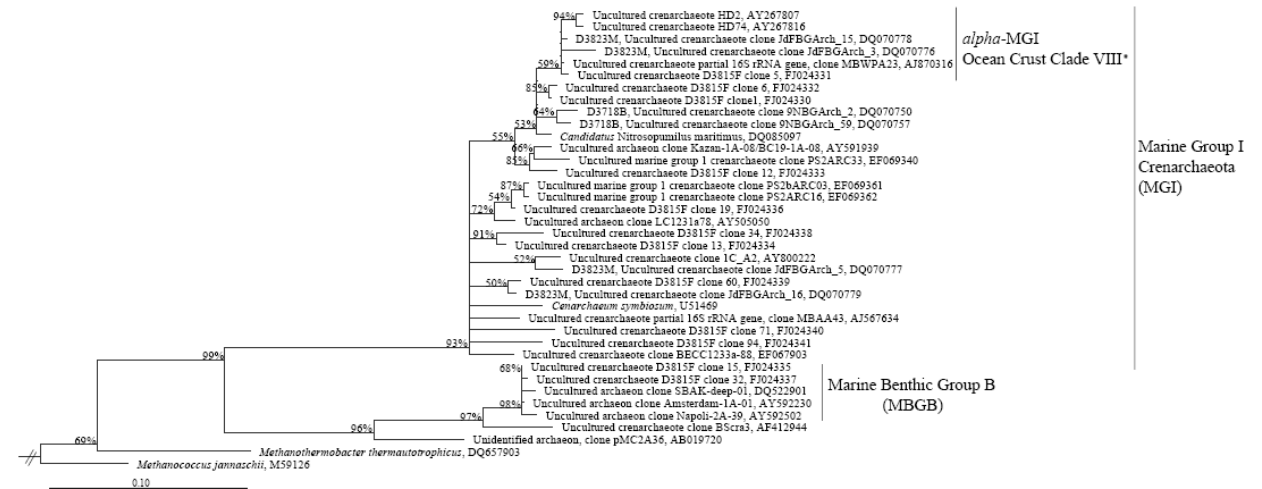


Figure 4.

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Supplementary Table 1. Functional genes present in basalt sample D3815F determined by microarray analysis.

Genbank ID	Gene short description	Gene category	Organism
23012702	0	Carbon fixation	Magnetospirillum magnetotacticum
12407235	aclB	Carbon fixation	Chlorobium limicola
30721807	FTHFS	Carbon fixation	Methylobacterium extorquens
32307749	rbcL	Carbon fixation	uncultured bacterium
505126	rbcL	Carbon fixation	Hydrogenophilus thermoluteolus
37791353	rbcL	Carbon fixation	uncultured proteobacterium
7592878	rbcL	Carbon fixation	uncultured deep-sea autotrophic bacterium ORII-2
7592885	rbcL	Carbon fixation	uncultured deep-sea autotrophic bacterium ORII-5
37791373	rbcL	Carbon fixation	uncultured proteobacterium
21217703	rbcL	Carbon fixation	uncultured bacterium
7229162	rbcL	Carbon fixation	uncultured deep-sea autotrophic bacterium SBI-5
132036	rbcL	Carbon fixation	Rhodospirillum rubrum
7592852	rbcL	Carbon fixation	uncultured deep-sea autotrophic bacterium SBII-4
6778693	dsrA	Sulfate reduction	uncultured sulfate-reducer HMS-25
40253098	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
FW015084A	dsrA	Sulfate reduction	lab clone
34017136	dsrA	Sulfate reduction	uncultured bacterium
34017154	dsrA	Sulfate reduction	uncultured bacterium
13898425	dsrA	Sulfate reduction	uncultured phenanthrene mineralizing bacterium
20501981	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
FW010274A	dsrA	Sulfate reduction	lab clone
13898437	dsrA	Sulfate reduction	uncultured phenanthrene mineralizing bacterium
12667570	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium UMTRAdsr626-27
7262420	dsrA	Sulfate reduction	uncultured sulfate-reducer HMS-54
20501993	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
34017094	dsrA	Sulfate reduction	uncultured bacterium
7262428	dsrA	Sulfate reduction	uncultured sulfate-reducer HMS-24
20502017	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
14090290	dsrA	Sulfate reduction	Desulfomicrobium escambiense
14389123	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
25990790	dsrA	Sulfate reduction	uncultured bacterium
40253034	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
14389119	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
15055587	dsrA	Sulfate reduction	Desulfococcus multivorans
22900884	dsrA	Sulfate reduction	uncultured bacterium
22999275	dsrA	Sulfate reduction	Magnetotactic cocci
14276799	dsrA	Sulfate reduction	Desulfotomaculum geothermicum

6778709	dsrA	Sulfate reduction	uncultured sulfate-reducer HMS-4
12667674	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium UMTRAdsr826-16
TPB16340A	dsrA	Sulfate reduction	lab clone
20501983	dsrA	Sulfate reduction	uncultured sulfate-reducing bacterium
FW300226A	dsrA	Sulfate reduction	lab clone
6179922	dsrB	Sulfate reduction	Solar Lake Mat Clone9065
18034323	dsrB	Sulfate reduction	Desulfacinum infernum
FW005271B	dsrB	Sulfate reduction	lab clone
13249527	dsrB	Sulfate reduction	uncultured sulfate-reducing bacterium
13561055	dsrB	Sulfate reduction	Desulfosarcina variabilis
TPB16051B	dsrB	Sulfate reduction	lab clone
15076856	dsrB	Sulfate reduction	Desulfosporosinus orientis
FW003272B	dsrB	Sulfate reduction	lab clone
14389217	dsrB	Sulfate reduction	uncultured sulfate-reducing bacterium
FW003264B	dsrB	Sulfate reduction	lab clone
28974756	dsrB	Sulfate reduction	uncultured bacterium
3892198	dsrB	Sulfate reduction	Archaeoglobus profundus
15077475	dsrB	Sulfate reduction	Desulfovibrio desulfuricans subsp. desulfuricans
FW015318B	dsrB	Sulfate reduction	lab clone
13249551	dsrB	Sulfate reduction	uncultured sulfate-reducing bacterium
FW010117B	dsrB	Sulfate reduction	lab clone
10716971	dsrB	Sulfate reduction	unidentified sulfate-reducing bacterium
TPB16055B	dsrB	Sulfate reduction	lab clone
28974734	dsrB	Sulfate reduction	uncultured bacterium
21673682	dsrB	Sulfate reduction	Chlorobium tepidum TLS
FW003269B	dsrB	Sulfate reduction	lab clone
TPB16070B	dsrB	Sulfate reduction	lab clone
13249539	dsrB	Sulfate reduction	uncultured sulfate-reducing bacterium
34017190	dsrB	Sulfate reduction	uncultured bacterium
40253070	dsrB	Sulfate reduction	uncultured sulfate-reducing bacterium
39998311	cytochrome	Metal resistance/reduction	Geobacter sulfurreducens PCA
24372200	cytochrome	Metal resistance/reduction	Shewanella oneidensis MR-1
39995722	cytochrome	Metal resistance/reduction	Geobacter sulfurreducens PCA
39996424	cytochrome	Metal resistance/reduction	Geobacter sulfurreducens PCA
24373346	cytochrome	Metal resistance/reduction	Shewanella oneidensis MR-1
39998423	cytochrome	Metal resistance/reduction	Geobacter sulfurreducens PCA
39935825	cytochrome	Metal resistance/reduction	Rhodopseudomonas palustris CGA009
2865528	cytochrome	Metal resistance/reduction	Shewanella putrefaciens
39998372	cytochrome	Metal resistance/reduction	Geobacter sulfurreducens PCA
39998004	cytochrome	Metal	Geobacter sulfurreducens PCA

39997392	cytochrome	resistance/reduction Metal	Geobacter sulfurreducens PCA
23475584	cytochrome	resistance/reduction Metal	Desulfovibrio desulfuricans G20
39996164	cytochrome	resistance/reduction Metal	Geobacter sulfurreducens PCA
39937295	cytochrome	resistance/reduction Metal	Rhodopseudomonas palustris CGA009
145083	cytochrome	resistance/reduction Metal	Desulfovibrio vulgaris
15022433	mcr	resistance/reduction Methane generation	Treponema medium
12802200	mcrA	Methane generation	uncultured archaeon 85A
38570220	mcrA	Methane generation	uncultured euryarchaeote
13259189	mcrA	Methane generation	uncultured methanogen RS-MCR04
38570178	mcrA	Methane generation	uncultured euryarchaeote
38570176	mcrA	Methane generation	uncultured euryarchaeote
799197	mcrA	Methane generation	Methanobolus oregonensis
34305116	mcrA	Methane generation	uncultured archaeon
13259303	mcrA	Methane generation	uncultured methanogen RS-ME32
7445687	mcrA	Methane generation	Methanothermobacter thermautotrophicus
34305109	mcrG	Methane generation	uncultured archaeon
20094092	mcrG	Methane generation	Methanopyrus kandleri AV19
6002398	mmo	Methane oxidation	Methylomonas sp. KSPIII
141050	mmo	Methane oxidation	Methylococcus capsulatus
6002402	mmo	Methane oxidation	Methylomonas sp. KSPIII
6002409	mmo	Methane oxidation	Methylomonas sp. KSWIII
2098700	mmo	Methane oxidation	Methylocystis sp. M
37813016	mmoA	Methane oxidation	uncultured bacterium
34915630	mmoA	Methane oxidation	uncultured methanotrophic proteobacterium
21685061	mmoA	Methane oxidation	Methylocella palustris
37813004	mmoA	Methane oxidation	uncultured bacterium
2098698	mmoA	Methane oxidation	Methylocystis sp. M
11038435	mmoA	Methane oxidation	uncultured putative methanotroph
37813020	mmoA	Methane oxidation	uncultured bacterium
7188932	pmo	Methane oxidation	Methylosinus trichosporium
7188937	pmo	Methane oxidation	Methylocystis sp. M
34733039	pmoA	Methane oxidation	uncultured bacterium
37496853	pmoA	Methane oxidation	uncultured bacterium
6424923	pmoA	Methane oxidation	uncultured eubacterium pAMC512
7578614	pmoA	Methane oxidation	uncultured bacterium FW-47
11493646	nifD	Nitrogen fixation	Azoarcus sp. BH72
29293348	nifH	Nitrogen fixation	uncultured bacterium
12001854	nifH	Nitrogen fixation	uncultured bacterium NR1611
13173333	nifH	Nitrogen fixation	uncultured bacterium
780721	nifH	Nitrogen fixation	unidentified marine eubacterium
780717	nifH	Nitrogen fixation	unidentified marine eubacterium
22449921	nifH	Nitrogen fixation	uncultured bacterium
12659182	nifH	Nitrogen fixation	Spirochaeta zuelzeriae
3157500	nifH	Nitrogen fixation	unidentified nitrogen-fixing bacteria

10863131	nifH	Nitrogen fixation	marine stromatolite eubacterium HB(0898) Z02
1236929	nifH	Nitrogen fixation	Anabaena variabilis
10863141	nifH	Nitrogen fixation	marine stromatolite eubacterium HB(0697) A100
12659198	nifH	Nitrogen fixation	Treponema azotonutricium
3157624	nifH	Nitrogen fixation	unidentified nitrogen-fixing bacteria
13173301	nifH	Nitrogen fixation	uncultured bacterium
3157594	nifH	Nitrogen fixation	unidentified nitrogen-fixing bacteria
1255464	nifH	Nitrogen fixation	unidentified bacterium
22449901	nifH	Nitrogen fixation	uncultured bacterium
19070843	nifH	Nitrogen fixation	unidentified nitrogen-fixing bacteria
13173319	nifH	Nitrogen fixation	uncultured bacterium
29649383	nifH	Nitrogen fixation	uncultured nitrogen-fixing bacterium
13173335	nifH	Nitrogen fixation	uncultured bacterium
33385573	nifH	Nitrogen fixation	uncultured bacterium
780713	nifH	Nitrogen fixation	unidentified marine eubacterium
1572591	nifH	Nitrogen fixation	Desulfovibrio gigas
29293188	nifH	Nitrogen fixation	uncultured bacterium
13173305	nifH	Nitrogen fixation	uncultured bacterium
5701924	nifH	Nitrogen fixation	Paenibacillus polymyxa
3157704	nifH	Nitrogen fixation	unidentified nitrogen-fixing bacteria
12001832	nifH	Nitrogen fixation	uncultured bacterium NR1600
12001842	nifH	Nitrogen fixation	uncultured bacterium NR1605
3157506	nifH	Nitrogen fixation	unidentified nitrogen-fixing bacteria
20804123	nifH	Nitrogen fixation	Mesorhizobium loti
22450003	nifH	Nitrogen fixation	uncultured bacterium
22988609	0	Nitrification	Rhodobacter sphaeroides
7595786	amoA	Nitrification	unidentified bacterium
7544069	amoA	Nitrification	Nitrosomonas halophila
7578632	amoA	Nitrification	uncultured bacterium WC306-54
27529221	amoA/pmoA	Nitrification	uncultured bacterium
26278794	narG	Denitrification	uncultured bacterium
29652478	narG	Denitrification	uncultured bacterium
26278922	narG	Denitrification	uncultured bacterium
26278882	narG	Denitrification	uncultured bacterium
29652532	narG	Denitrification	uncultured bacterium
38427014	narG	Denitrification	uncultured bacterium
26278770	narG	Denitrification	uncultured bacterium
32308011	narG	Denitrification	uncultured bacterium
32307889	narG	Denitrification	uncultured bacterium
38427060	narG	Denitrification	uncultured bacterium
32307981	narG	Denitrification	uncultured bacterium
26278784	narG	Denitrification	uncultured bacterium
29652590	narG	Denitrification	uncultured bacterium
29652508	narG	Denitrification	uncultured bacterium
38427022	narG	Denitrification	uncultured bacterium
29652428	narG	Denitrification	uncultured bacterium
26278684	narG	Denitrification	uncultured bacterium
32307917	narG	Denitrification	uncultured bacterium

26278870	narG	Denitrification	uncultured bacterium
17385544	narG	Denitrification	uncultured bacterium
26278676	narG	Denitrification	uncultured bacterium
30269569	nasA	Denitrification	uncultured bacterium
30269577	nasA	Denitrification	uncultured bacterium
12597209	nirK	Denitrification	Alcaligenes sp. STC1
3758830	nirK	Denitrification	Hyphomicrobium zavarzinii
NBPd1-B05	nirK	Denitrification	lab clone
27125563	nirK	Denitrification	uncultured bacterium
37999212	nirK	Denitrification	uncultured bacterium
ORA-NIRK-C01	nirK	Denitrification	lab clone
1488172	nirK	Denitrification	Rhizobium sllae
3758901	nirK	Denitrification	Rhodobacter sphaeroides f. sp. denitrificans
28542653	nirS	Denitrification	uncultured bacterium
24421455	nirS	Denitrification	uncultured organism
24421507	nirS	Denitrification	uncultured organism
37999198	nirS	Denitrification	uncultured bacterium
24421269	nirS	Denitrification	uncultured organism
24528368	nirS	Denitrification	uncultured bacterium
28542627	nirS	Denitrification	uncultured bacterium
38455926	nirS	Denitrification	uncultured bacterium
24421271	nirS	Denitrification	uncultured organism
22252866	nirS	Denitrification	uncultured bacterium
34391466	norB	Denitrification	Nitrosomonas europaea
29466090	norB	Denitrification	uncultured bacterium
29466092	norB	Denitrification	uncultured bacterium
38373207	nosZ	Denitrification	uncultured bacterium
3057083	nosZ	Denitrification	Paracoccus pantotrophus
29125972	nosZ	Denitrification	uncultured soil bacterium
13959038	nosZ	Denitrification	Azospirillum lipoferum
4633572	nosZ	Denitrification	uncultured bacterium ProR
14994626	nosZ	Denitrification	uncultured bacterium
