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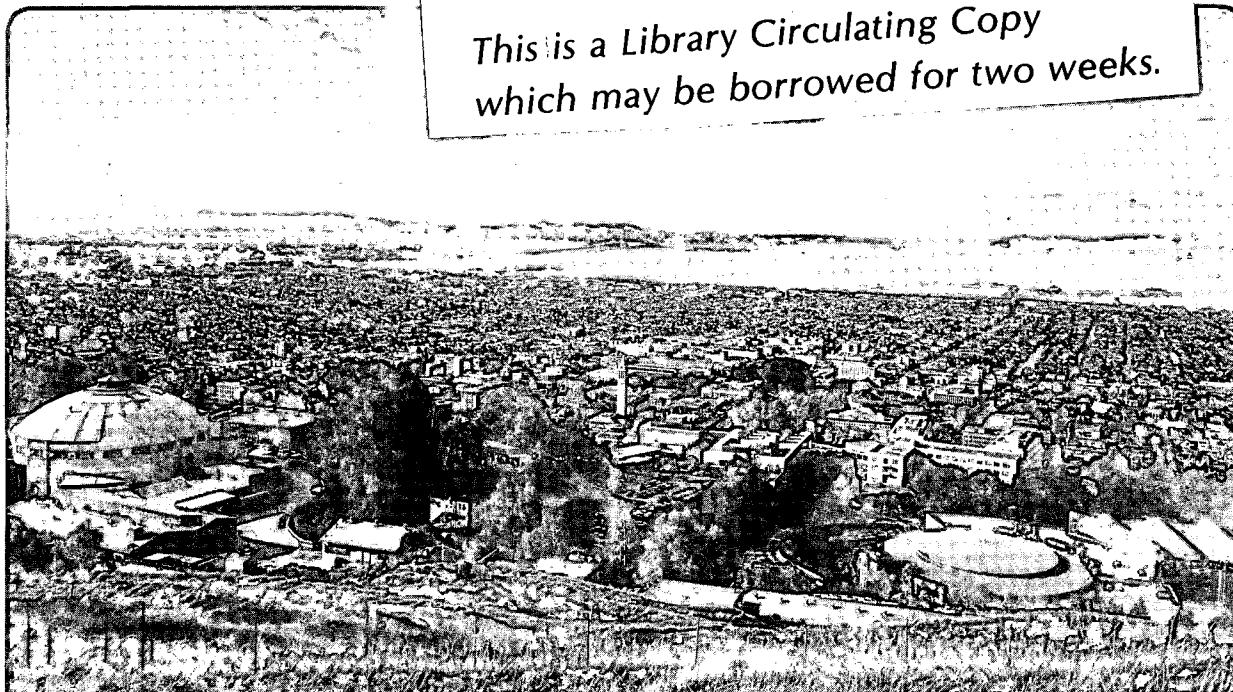
CO<sub>2</sub> TRENDS IN THE DEPLETION OF THE LARDERELLO  
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K. Pruess, R. Celati, C. Calore, and F. D'Amore

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# CO<sub>2</sub> TRENDS IN THE DEPLETION OF THE LARDERELLO VAPOR-DOMINATED RESERVOIR

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## Introduction

Most Larderello wells produce superheated steam, with a noncondensable gas content of 3-10 percent by weight, of which typically 95 percent consists of CO<sub>2</sub> (D'Amore and Truesdell, 1984). For most wells for which long-term depletion data are available the gas content of the discharge shows remarkably small temporal variations. In many cases the gas-steam ratio changes by no more than 10 to 20 percent over several decades, and there are only few wells where gas content changes by as much as a factor of 3, which is still a rather modest change. In many wells there is a tendency for gas content to increase for several years, then stabilize, and eventually decrease (Sestini, 1970; D'Amore and Truesdell, 1979; Calore et al., 1982).

The present paper is concerned with the origin of CO<sub>2</sub> in the Larderello discharges from a reservoir engineering point of view. Specifically, we consider the question of fluid reserves (water and CO<sub>2</sub>) at Larderello, and we carry out numerical simulations to obtain insight into temporal trends of CO<sub>2</sub> released from idealized models of vapor-dominated systems. Important constraints in the modeling arise from observed temperatures and enthalpies at Larderello. Discharge enthalpies are generally close to those of saturated steam near 250°C (2.8 MJ/kg), with some degree of superheat which tends to increase with time. Flow rates of most wells decrease rapidly during the first few years of production and subsequently decline very slowly (Sestini, 1970; Weres et al., 1977; D'Amore and Truesdell, 1979).

## Fluid Reserves

Total cumulative fluid production from the Larderello reservoir is large. Extrapolating data given by Sestini (1970) we estimate that approximately  $400 \times 10^9$  kg of fluid have been extracted to date from the central zone of Larderello, which covers an area of approximately 20 km<sup>2</sup>. On the basis of this large cumulative production, it has been concluded by several authors that most of the fluid reserves were originally in place in liquid form, because storage of such fluid quantities in vapor form would require an unreasonably large reservoir thickness (James, 1968; Nathenson, 1975; Weres et al., 1977). In order to calculate specific fluid depletion

(i.e., mass of fluid extracted per unit reservoir volume), it is necessary to estimate the average thickness of the productive formations, and to allow for the fact that total drainage area will be larger than the well field area. Adopting a value  $h = 1$  km for reservoir thickness, and taking a conservative factor 2 for reservoir volume to account for production from greater depth, or from outside the drilled area, we obtain a specific depletion to date of approximately 10 kg/m<sup>3</sup>. For an estimated porosity of < 5 percent, the vapor in place is < 1 kg/m<sup>3</sup> (for a vapor density of  $\rho_v \approx 20$  kg/m<sup>3</sup> at 250°C). If 10 kg/m<sup>3</sup> of fluid are extracted from boiling liquid (with a density of  $\rho_l \approx 800$  kg/m<sup>3</sup> at 250°C), then the corresponding decrease in liquid saturation is  $\Delta S_l = 0.25$ . From these considerations we conclude that initial (pre-exploitation) liquid saturation in the Larderello reservoir has to have been no less than  $S_l = 0.25$ .

Let us now turn to the question of CO<sub>2</sub> reserves. The average CO<sub>2</sub> content of the fluids produced in Larderello is  $X_c = 5$  percent by mass, with most individual wells falling in the range of  $3\% \leq X_c \leq 10\%$ . Could the produced CO<sub>2</sub> have been stored in the original reservoir fluid? In Figure 1 we have plotted CO<sub>2</sub> mass fractions of two-phase mixtures as a function of liquid saturation, with CO<sub>2</sub> partial pressure  $P_c$  as parameter (temperature 250°C). It is seen that for a given partial pressure of CO<sub>2</sub>, the CO<sub>2</sub> content of two-phase fluid decreases rapidly with increasing liquid saturation. The effects are temperature dependent. If temperature is increased the CO<sub>2</sub> content of the vapor phase diminishes, while that of the liquid phase increases. At intermediate saturations  $0.2 \leq S_l \leq 0.5$  these phase distribution effects tend to cancel out, so that CO<sub>2</sub> content is determined mostly by partial pressure, with little temperature dependence. For the lowest value of liquid saturation which is compatible with cumulative fluid production to date,  $S_l = 0.25$ , an average CO<sub>2</sub> content of  $X_c = 5$  percent requires a CO<sub>2</sub> partial pressure of  $P_c \approx 15$  bars. Such partial pressures are inconsistent with field measurements of pressures and temperatures at Larderello. Indeed, in zones which had not been heavily exploited the observed temperatures and pressures correspond rather closely to the saturated vapor pressure curve for water,  $P_v \approx P_{sat}(T)$ , indicating that CO<sub>2</sub>

partial pressure is small (perhaps  $P_c = 1-2$  bars). From this we conclude that only a small fraction of the produced  $CO_2$  at Larderello could have been originally stored in the reservoir fluids. Most of the produced  $CO_2$  has to have been supplied either by an external source (presumably located at greater depth), or by an internal source (i.e., a mineral buffer).

### Numerical Simulations

The conclusion reached above about the origin of  $CO_2$  in Larderello discharges is supported by results of numerical simulations. The simulations demonstrate conclusively that  $CO_2$  concentrations and trends in produced fluids can not be explained in terms of  $CO_2$  content in place in the original reservoir fluids. Before discussing the numerical results it is appropriate to describe the general pattern of  $CO_2$  concentration change expected from two-phase flow and phase change phenomena.

For temperature and pressure conditions relevant to vapor-dominated systems, the amount of  $CO_2$  present in a unit volume of gas phase is larger by a factor of about 3 than the amount dissolved in a unit volume of liquid phase (see Figure 2). After depletion of a block of porous rock is initiated, the liquid saturation generally tends to decrease because of boiling. This enhances gas phase mobility (relative permeability), causing  $CO_2$  concentrations in the fluids discharged from rock matrix blocks to rise. However, the boiling liquid is less rich in  $CO_2$  than the initial gas phase, so that  $CO_2$  concentrations in the gas phase decline during depletion. This effect slows down the increase in discharged  $CO_2$  concentrations brought about by enhanced gas mobility, and eventually reverses the trend, causing discharged  $CO_2$  concentrations to decline. Therefore, the combination of two-phase flow and phase change effects is expected to yield  $CO_2$  trends compatible with field observations for many Larderello wells: an initial period of increasing  $CO_2$  concentrations, followed by a period of approximately constant composition, and an eventual decline in  $CO_2$  concentrations.

The general trends described above are borne out by numerical simulations. We have used the  $H_2O/CO_2$  version of Lawrence Berkeley Laboratory's general purpose simulator MULKOM (Pruess, 1983a; O'Sullivan et al., 1983) to study fluid discharges from porous matrix blocks of low permeability. The simulations were made with the method of multiple interacting continua (MINC; Pruess and Narasimhan, 1982), and different levels of space discretization were used to verify that spatial truncation errors were negligible. Figure 3 shows simulated  $CO_2$  transients for fluid discharge from a porous cube of 50 m side length with parameters given in Table 1. At the production rate specified, total block depletion takes 51.9 years.  $CO_2$  concentrations increase by almost a factor of 3 for the first five years, and subsequently decline, as expected. Comparing these results with typical field data for

Larderello wells, it is apparent that the simulated changes in  $CO_2$  concentrations are stronger and occur more rapidly than are generally observed in the field. We do not consider this a serious shortcoming of the simulation. Observed  $CO_2$  concentrations in the field represent an average of fluid discharges from many matrix blocks. As the drainage volume of a well expands with time, matrix blocks at greater distance from the well will contribute increasing  $CO_2$  concentrations, while matrix blocks near the well will discharge fluids of declining  $CO_2$  content. The superposition of these effects is expected to diminish and slow down  $CO_2$  concentration changes in comparison to the single block response shown in Figure 3, giving a response more compatible with field data.

A more serious problem with the simulated  $CO_2$  trends is the absolute magnitude of  $CO_2$  mass fraction, which is smaller than typical field values by a factor of 20-30. In the simulation we used a  $CO_2$  partial pressure of 1 bar. In order to obtain  $CO_2$  mass fractions in the range of 5-8 percent, as observed in the field, we would require unreasonably large partial pressures. Alternatively, we could modify relative permeability parameters, or in-place liquid saturations, to enhance gas phase flow in comparison to liquid flow in the matrix. (For the conditions chosen in the simulation, the gas phase contains initially approximately 5 percent  $CO_2$  by mass.) However, this gives rise to other problems, which are discussed below.

The most serious failure of the simulation occurs with regard to flowing enthalpy. Figure 3 shows that initial discharge enthalpy is as low as 1.07 MJ/kg. Approximately 6 years of discharge are required before enthalpy approaches a value of 2.8 MJ/kg which is compatible with field observations. This enthalpy behavior was not anticipated. Rather, we had expected that because of the low matrix block permeability a strong conductive enhancement of flowing enthalpy would take place, so that superheated vapor would be discharged even as liquid is mobile and flowing in the matrix (Pruess, 1983b). Interestingly, however, the presence of  $CO_2$  suppresses the mechanism of enthalpy enhancement described by Pruess (1983b). As discharge from the matrix block is initiated, a gradient in  $CO_2$  partial pressure is established inside the matrix block, which is entirely sufficient to drive mass flux at the desired rate of fluid discharge. For a period of several years the gradient of vapor pressure remains negligibly small in comparison to the gradient of  $CO_2$  partial pressure. Therefore, temperature gradients inside the matrix block are also very small, and conductive enhancement of flowing enthalpy is negligible.

The only way to obtain discharge enthalpies near 2.8 MJ/kg is to adjust parameters in such a way that liquid is immobile in the matrix. However, in simulations with immobile liquid we invariably obtain a rapid decline of  $CO_2$  concentrations with time, caused by dilution of the gas phase from boiling liquid. Over the time period required to deplete the original gas

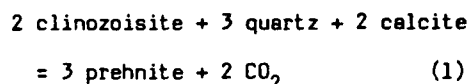
phase in place, CO<sub>2</sub> concentrations decline by a factor of approximately 30.

The conclusion from these simulation studies confirms the result from the analysis of fluid reserves, above. Namely, the concentration and trends of CO<sub>2</sub> in Larderello discharges can not be explained by assuming that produced CO<sub>2</sub> originates from the reservoir fluid.

#### Mineral Buffer for CO<sub>2</sub>

The discussion of fluid reserves and the numerical simulation results given above indicate that most of the CO<sub>2</sub> produced at Larderello originates either from external or from internal (mineral buffer) sources. We consider the external source hypothesis (CO<sub>2</sub> supplied from beneath the main reservoir) to be rather improbable. Observed temporal variations in CO<sub>2</sub> content of discharges are generally small over several decades, while flow rates of individual wells often change by an order of magnitude or more. The relative constancy of CO<sub>2</sub> content would be difficult to explain in a model with external source, while it follows naturally when a mineral buffer is assumed.

Several authors have suggested that non-condensable gas concentrations in geothermal reservoirs, and in particular CO<sub>2</sub> partial pressures, are controlled by equilibrium reactions involving mineral assemblages (Giggenbach, 1981; Cavarretta et al., 1982; Arnorsson et al., 1983; Giggenbach, 1984). Cavarretta et al. (1982) showed that the CO<sub>2</sub> partial pressures at Larderello and Serrazzano are close to equilibrium for the reaction



Following these authors we have considered a depletion model with a mineral buffer as given by Equation (1). Very little is known about the kinetic rates of reactions such as (1). We have adopted a first order rate law ad hoc, assuming that the volumetric rate of CO<sub>2</sub> release from minerals is proportional to the difference between equilibrium and actual partial pressure of CO<sub>2</sub>:

$$q_{\text{CO}_2} = \alpha (P_c^{\text{eq}} - P_c) \quad (2)$$

An internal CO<sub>2</sub> source as given by Equation (2) was incorporated into the MULKOM simulator. It is believed that the mineral buffer reaction Equation (1) can only proceed in the presence of liquid water. Therefore, the rate constant  $\alpha$  should go to zero for  $S_L \rightarrow 0$ . This is taken into account in our simulations by replacing  $\alpha \rightarrow \alpha \cdot S_L / 0.05$  for  $S_L < 0.05$ . The temperature dependence of  $P_c^{\text{eq}}$  is expressed by an empirical fit to data for the reaction (1), which were provided to us by Gianelli (private communication,

1984). The relationship is:

$$\log P_c^{\text{eq}} = 1.437 \times 10^{-2} T - 1.4 \times 10^{-5} T^2 - 2.81 \quad (3)$$

Reaction (1) and Equation (3) are considered valid in the 250–350°C temperature range and for the deeper layers of the reservoir (metamorphic basement) of Larderello. Lacking more detailed information on possible mineral buffers in different parts of the reservoir and in different temperature ranges, at present we have assumed that a relationship like Equation (3) could be considered valid for the entire reservoir and in a wider temperature range (200°C  $\leq$  T  $\leq$  350°C).

It is interesting to note that over the temperature interval 200°C  $\leq$  T  $\leq$  300°C the ratio of equilibrium partial pressure of CO<sub>2</sub>, as predicted by Equation (3), to the saturated vapor pressure is practically constant. We have

$$\frac{P_c^{\text{eq}}(T)}{P_{\text{sat}}(T)} = 0.0204 \pm 0.6\% \quad (4)$$

From this it follows that equilibrium CO<sub>2</sub> mass fractions in vapor are independent of temperature (for 200°C  $\leq$  T  $\leq$  300°C), if in fact CO<sub>2</sub> partial pressure is controlled by the reaction (1).

Using the above model of a finite rate mineral buffer, we have simulated the depletion of individual matrix blocks. In the simulations it was necessary to limit time steps to a modest saturation change per time step, because of the dependence of buffer rate upon liquid saturation. Simulation results should be considered illustrative of mechanisms rather than quantitative, because of the uncertainties involved in the parameters for the mineral buffer.

Simulation of block depletion with mineral buffer was performed with the same parameters as in Table 1, except for initial  $P_c$ , which was assumed equal to the equilibrium value at 240°C ( $0.6798 \times 10^5$  Pa). Binary diffusion was also included in the simulations. A diffusivity at standard conditions (0°C, 1 atm.) of  $D_S = 1.38 \times 10^{-5}$  m<sup>2</sup>/s and a temperature and pressure dependence of the type

$$D(T, P) = D_S \frac{P}{P_S} \left( \frac{T}{T_S} \right)^{1.91}$$

with T = absolute temperature, were assumed (Perry, 1963).

Some results for simulated block depletion are shown in Figure 4. The four curves correspond to different values of the rate constant  $\alpha$  of Equation (2).

For high values of  $\alpha$  [ $10^{-10}$  –  $10^{-12}$  kg/(m<sup>3</sup>·s·Pa)], the CO<sub>2</sub> content in the produced fluid, after an initial transient, reaches values which are comparable with those observed in the

field, but the CO<sub>2</sub> fraction remains practically constant up to the complete depletion of the block. With lower  $\alpha$  values we obtain a trend more similar to that observed in many Larderello wells, but the maximum values are significantly lower than in the field.

Enthalpy rises faster than in the case without mineral buffer, and the more rapid increases are obtained with higher values of  $\alpha$ . This is due to the fact that CO<sub>2</sub> generation tends to reduce  $P_c$ -gradient, so that conductive enthalpy enhancement is stronger. Diffusion also contributes to reduce the  $P_c$  gradient.

The initial values of CO<sub>2</sub> concentration and enthalpy appear still too low in comparison with real data due to the initial liquid flow out of the block, which in these simulations is relatively high, an initial liquid saturation of 0.8 being assumed. Higher initial values of both CO<sub>2</sub> mass fraction and enthalpy are obtained with a lower initial water saturation (Figure 5). Now the CO<sub>2</sub> content of the produced fluid does not show the fast decrease which occurred without mineral buffer. With the initial liquid saturation used in this case (0.6) the depletion time is still about 38 years.

Figure 6 shows the cumulative CO<sub>2</sub> mass generation per unit volume as a function of distance from the block surface for the four cases of Figure 4. In all cases the curves are sufficiently smooth, but the cases with higher  $\alpha$  require a finer space discretization. The curves show different trends at short distances from the block surface. With a very high  $\alpha$ , the strong CO<sub>2</sub> generation tends to maintain everywhere in the two-phase zone a  $P_c$  close to the equilibrium value, preventing the propagation of  $P_c$  drawdown inside the block. The result is that strong generation rates occur in a narrow zone close to the saturation front only. In such conditions total CO<sub>2</sub> generation depends mainly on the velocity of saturation front displacement. This velocity increases as the front moves away from the block surface, due to the decreasing liquid saturation and the consequent lower liquid flow from the inner most parts of the block. With low  $\alpha$ , on the contrary, a significant  $P_c$  gradient is established throughout the block and CO<sub>2</sub> generation begins very early also at points distant from the saturation front. CO<sub>2</sub> generation is always diffused over a relatively large volume.

Moving away from the block surface, the period of CO<sub>2</sub> generation increases as drying occurs later, but the average generation rate decreases due to the lower average  $P_c$  drawdown. The effect of the increasing period of generation prevails at intermediate distances from the block surface, while the decrease of average generation rate generally prevails in the innermost parts of the block.

### Conclusions

Overall mass balance considerations and numerical simulations of reservoir depletion show

that only a small fraction of the CO<sub>2</sub> produced in Larderello could have been originally stored in the reservoir fluid. Most of the produced CO<sub>2</sub> was supplied either by an external source, or by an internal source (i.e., a mineral buffer). The latter possibility appears more likely, because of the small temporal variations in CO<sub>2</sub> content of well discharges, even as production rates change by large amounts.

Numerical simulations were performed to study CO<sub>2</sub> discharge from rocks of low permeability, with a temperature-dependent prehnite/clinozoisite mineral buffer. Simulated discharge characteristics were found to be in approximate agreement with typical CO<sub>2</sub> trends observed for Larderello wells. However, observed CO<sub>2</sub> concentrations tend to be somewhat larger than the simulated values, suggesting that additional minerals are participating in buffering CO<sub>2</sub>.

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### References

- Arnorsson, S., Gunnlaugsson, E., and Svavarsson, H., 1983, The Chemistry of Geothermal Waters in Iceland. III. Chemical Geothermometry in Geothermal Investigations, *Geochim. Cosmochim. Acta*, Vol. 47, pp. 567-577.
- Calore, C., Celati, R., D'Amore, F. and Noto, P., 1982, Geochemical Evidence of Natural Recharge in Larderello and Castelnuovo Areas, Proc. Eighth Workshop Geothermal Reservoir Engineering, Stanford, California, pp. 323-328.
- Cavarretta, G., Gianelli, G., and Puxeddu, M., 1982, Formation of Authigenic Minerals and their Use as Indicators of the Physicochemical Parameters of the Fluid in the Larderello-Travale Geothermal Field, *Econ. Geol.*, Vol. 77, No. 5, pp. 1071-1084.
- D'Amore, F., and Truesdell, A. H., 1979, Models for Steam Chemistry at Larderello and The Geysers, Proc. Fifth Workshop Geothermal Reservoir Engineering, Stanford, California, pp. 283-297.
- D'Amore, F., and Truesdell, A. H., 1984, Helium in the Larderello Geothermal Fluid, *Geothermics*, Vol. 13, No. 3, pp. 227-239.

Giggenbach, W. F., 1981, Geothermal Mineral Equilibria, Geochim. Cosmochim. Acta, Vol 45, pp. 393-410.

Giggenbach, W. F., 1984, Mass Transfer in Hydrothermal Alteration Systems - A Conceptual Approach, unpublished manuscript (to appear in Geochim. Cosmochim. Acta).

James, R., 1968, Wairakei and Larderello: Geothermal Power Systems Compared, New Zealand Journal of Science, Vol. 11, pp. 706-719.

Nathenson, M., 1975, Some Reservoir Engineering Calculations for the Vapor-Dominated System at Larderello, Italy, U. S. Geological Survey, Open File Report 75-142.

O'Sullivan, M. J., Bodvarsson, G. S., Pruess, K., and Blakeley, M. R., 1983, Fluid and Heat Flow in Gas-Rich Geothermal Reservoirs, paper SPE-12102, presented at the Society of Petroleum Engineers 58th Annual Technical Conference and Exhibition, San Francisco, Ca., October 1983 (to appear in Society of Petroleum Engineers Journal).

Perry, J. H., 1963, Chemical Engineers' Handbook, McGraw-Hill, New York, Section 14, pp. 19-23.

Pruess, K., 1983a, Development of the General Purpose Simulator MULTOM, Annual Report 1982, Earth Sciences Division, Lawrence Berkeley Laboratory Report LBL-15500.

Pruess, K., 1983b, Heat Transfer in Fractured Geothermal Reservoirs with Boiling, Water Resources Research, Vol. 19, No. 1, pp. 201-208.

Pruess, K., and Narasimhan, T. N., 1982, On Fluid Reserves and the Production of Superheated Steam from Fractured, Vapor-Dominated Geothermal Reservoirs, Journal of Geophysical Research, Vol. 87, No. B11, pp. 9329-9339.

Sestini, G., 1970, Superheating of Geothermal Steam, Geothermics, Special Issue 2, pp. 622-648.

Weres, O., Tsao, K., and Wood, B., 1977, Resource, Technology, and Environment at The Geysers, Lawrence Berkeley Laboratory Report LBL-5231, Berkeley, California.

Table 1: Parameters for Simulation

Matrix	
rock grain density	2500 kg/m <sup>3</sup>
rock specific heat	960 J/Kg°C
rock heat conductivity	2.8 W/m°C
porosity	.05
permeability	10 <sup>-17</sup> m <sup>2</sup>
relative permeability	Corey curves S <sub>lr</sub> = 0.30 S <sub>gr</sub> = 0.05
Fractures (three perpendicular sets)	
spacing	50 m
Initial Conditions	
temperature	240°C
liquid saturation	0.80
CO <sub>2</sub> partial pressure	10 <sup>5</sup> Pa
Discharge	
Constant volumetric rate	2 x 10 <sup>-8</sup> kg/s.m <sup>3</sup>

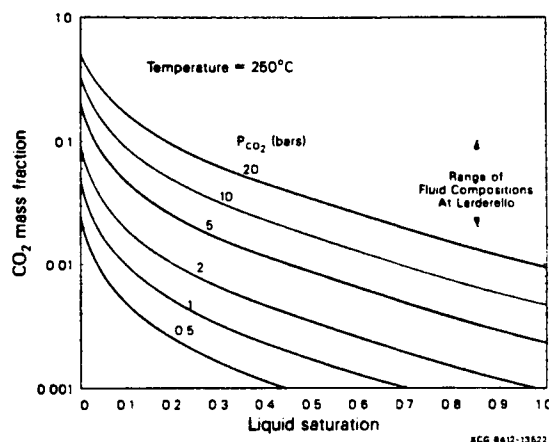


Figure 1. CO<sub>2</sub> mass fraction in two-phase fluid as a function of liquid saturation.

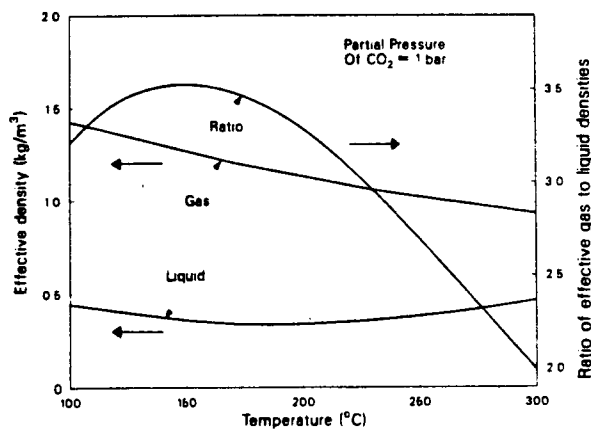


Figure 2. Effective density of CO<sub>2</sub> in gas and liquid phases at a partial pressure of P<sub>c</sub> = 1 bar.



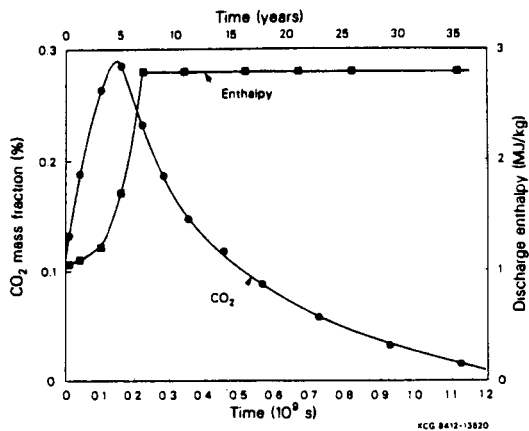
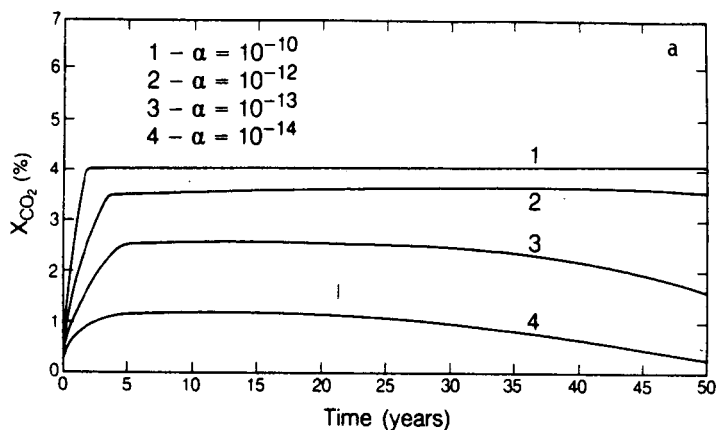
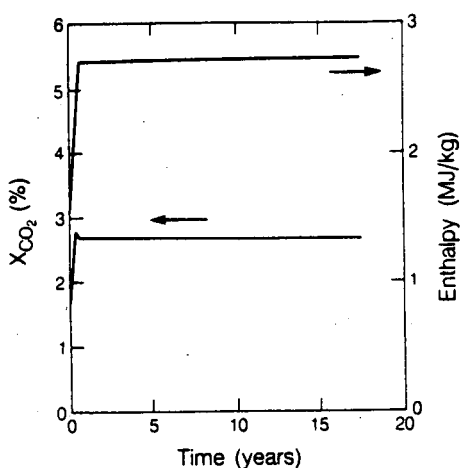


Figure 3. Simulated CO<sub>2</sub> and enthalpy transients for fluid discharge from a porous block of low permeability.

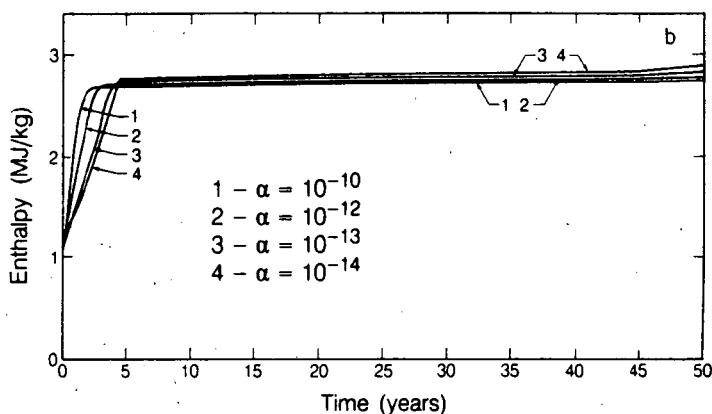


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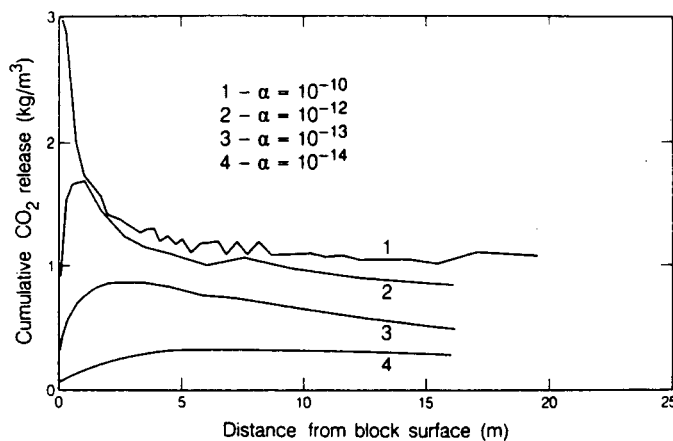
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Figure 5. Simulated CO<sub>2</sub> and enthalpy transients for fluid discharge from a porous block with CO<sub>2</sub> internal source and initial liquid saturation of 60%.



XBL 851-8803

Figure 4. Block depletion model with mineral buffer given by Equation (1). CO<sub>2</sub> mass fractions (a) and enthalpy (b) trends for the discharged fluid. Cases 1 to 4 with a generation rate constant of  $10^{-10}$ ,  $10^{-12}$ ,  $10^{-13}$  and  $10^{-14}$  kg/(m<sup>3</sup>·s·Pa) respectively.



XBL 851-8808

Figure 6. Cumulative CO<sub>2</sub> mass generation per unit volume for the four cases of Figure 4.

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