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Permalink https://escholarship.org/uc/item/1tg7t8gg

Journal ACS Earth and Space Chemistry, 6(12)

ISSN 2472-3452

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Publication Date 2022-12-15

DOI

10.1021/acsearthspacechem.2c00259

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Heterogeneous Formation of Organonitrates (ON) and Nitroxy-Organosulfates (NOS) from Adsorbed α -Pinene-Derived **Organosulfates (OS) on Mineral Surfaces**

Eshani Hettiarachchi and Vicki H. Grassian*

Cite This: ACS Earth Space Chem. 2022, 6, 3017–3030	Read Online
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ABSTRACT: Organonitrates (ON) and nitroxy-organosulfates (NOS) are important components of secondary organic aerosols (SOAs). Gas-phase reactions of α -pinene (C ₁₀ H ₁₆), a primary precursor for several ON compounds, are fairly well understood although formation pathways for NOS largely remain unknown. NOS formation may occur via reactions of ON and organic peroxides with sulfates as well as through radical-initiated photochemical processes. Despite the fact that organosulfates	Alpha-Pinene HNO ₃ Ho ho ho ho ho ho ho ho h

(OS) represent a significant portion of the organic aerosol mass, ON and NOS formation from OS is less understood, especially through nighttime heterogeneous and multiphase chemistry pathways. In the current study, surface reactions of adsorbed α pinene-derived OS with nitrogen oxides on hematite and kaolinite



surfaces, common components of mineral dust, have been investigated. α -Pinene reacts with sulfated mineral surfaces, forming a range of OS compounds on the surface. These OS compounds when adsorbed on mineral surfaces can further react with HNO₃ and NO₂, producing several ON and NOS compounds as well as several oxidation products. Overall, this study reveals the complexity of reactions of prevalent organic compounds leading to the formation of OS, ON, and NOS via heterogeneous and multiphase reaction pathways on mineral surfaces. It is also shown that this chemistry is mineralogy-specific.

KEYWORDS: α -pinene, SOA, nitroxy-organosulfates (NOS), organonitrates (ON), organosulfates (OS), iron oxide, kaolinite

INTRODUCTION

Secondary organic aerosols (SOA) are ubiquitous in the atmosphere and are known to play key roles in climate change, air quality, reduced visibility, and human health.¹⁻⁵ Organosulfates (OS, ROSO₃H), organonitrates (ON, RONO₂), and nitroxy-organosulfates (NOS, $RN_xS_yO_n$) are present within SOA and represent a structurally diverse mixture of compounds found in polluted environments.⁶⁻¹⁰ Precursors of these compounds include both biogenic and anthropogenic sources of organic compounds. Primary biogenic precursors are isoprenes, monoterpenes, sesquiterpenes, and aldehydes.¹¹

 α -Pinene (C₁₀H₁₆), a biogenic volatile organic compound, is one of the most abundant atmospheric monoterpenes with an average estimated emission of 66 Tg yr $^{-1.12-15}$ Furthermore, α -pinene-derived OS-, ON-, and NOS-SOA along with oxygenated α -pinene derivatives have been detected in ambient air.^{16–19} The formation of OS compounds in the atmosphere occurs via several reaction pathways. OS is produced from the photochemical reaction of biogenic hydrocarbons under highly acidic conditions, such as with sulfate aerosols or sulfuric acid, via nucleophilic substitution reactions of sulfate with epoxides in aqueous-phase solutions, with HO radical-involved mechanisms, and via catalytic

reactions with photocatalysts such as TiO2.^{16,17,20-25} Schmidt et al. showed increased formation of OS from methacrolein under illuminated conditions in the presence of road dust.²⁵ Numerous laboratory studies have been conducted to understand the formation of OS from α -pinene.^{22,24,26,27} Surratt et al. reported that the OS is formed when monoterpenes are oxidized in the presence of acidified sulfate aerosol.²

Despite the fact that atmospheric OS represents a significant portion of the organic aerosol mass, estimated to be between 3 and 30%,^{18,28,29} the formation of ON and NOS from reactions of OS due to nighttime heterogeneous and multiphase chemistry is less understood. Atmospheric ON formation occurs either by OH radical-initiated photochemical reactions or NO₃ radical-initiated nocturnal reactions of anthropogenic

Received: August 19, 2022 Revised: October 11, 2022 Accepted: November 18, 2022 Published: November 29, 2022





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and biogenic volatile organic compounds.^{30–35} Reactions of α pinene with strong atmospheric oxidizers such as O3 and OH radicals lead to the formation of a range of oxidized α -pinene derivatives that are known as first-generation oxidation products.^{4,12,36,37} Some of these compounds are pinonaldehyde, pinonic acid, α -pinene oxide, and structurally diverse organic peroxides. These oxidation products can also act as a reactant in generating atmospheric OS and ON compounds.³⁸ Oxidation reactions generally occur through photochemical pathways in the presence of sunlight.^{30,39} Recently, it was shown that the formation of ON from α -pinene via heterogeneous reactions on mineral surfaces under dry and dark conditions can occur.⁴⁰ In the case of NOS formation, its presence in aerosols is almost always observed in the night time. Surratt et al.²² suggested NOS formation from the reactions of ON with sulfates under acidification, whereas Iinuma et al.¹⁷ suggested the formation of some common NOS such as $C_{10}H_{17}NSO_7$ and $C_{10}H_{18}NSO_7$ occurs during night time. Similar to OS, NOS formation may occur through reactive uptake of epoxides, esterification of hydroxyl and keto groups with sulfuric acid,⁴¹ and/or through radical-initiated processes in wet aerosols.⁴² Increased NOS formation due to fireworks has also been previously observed.¹¹ However, formation pathways for NOS, in general, are less well understood, although their relative abundance and formation pathways involving heterogeneous chemistry in nighttime conditions remain scarce.

In this study, reactions of NO₂ and HNO₃ with α -pinene-OS adsorbed on mineral surfaces hematite (α -Fe₂O₃) and kaolinite $(SiO_2Al_2O_3(OH)_4)$ were studied. The adsorbed α -pinene-OS was produced by exposing α -pinene on mineral surfaces, hematite and kaolinite, first exposed to SO2. Hematite and kaolinite are major reactive components of mineral dust.⁴ Mineral surfaces can interact with gas-phase SO_2 , 5,44,455,44,45 leading to the formation of surface products.^{46,47} In fact, the interaction of SO₂ on mineral surfaces has been widely studied⁴⁸⁻⁵⁰ and it has been shown that several surfaceadsorbed species form, including sulfate, bisulfate, sulfite, and bisulfite, thereby producing a surface for α -pinene to interact with and facilitating the formation of OS. Additionally, both NO₂ and HNO₃ are found in trace quantities in the atmosphere, primarily via fossil fuel combustion, vehicle exhausts, and agricultural activities.^{45–47,51} Due to the higher correlation of atmospheric NO₂ concentrations to traffic, it is used as a traffic-related air pollution marker.⁴⁶ Therefore, interactions between α -pinene with mineral dust and trace gas pollutants during dust transport and urban dust episodes can facilitate formation of OS, ON, and NOS.

Given the importance of heterogeneous reactions of atmospheric gases with mineral dust surfaces in the atmosphere and due to the abundance of α -pinene, the role of mineral dust surface chemistry in reactions of α -pinene-OS with atmospheric NO₂ and HNO₃ on hematite and kaolinite surfaces was investigated. This study represents a step forward in understanding the complexity of reactions of volatile and semivolatile organics with trace inorganic gases on mineral surfaces to form multifunctional compounds that can modify the climate impacts of mineral dust.⁵² In the current study, experiments were conducted under dark and dry conditions. Although water vapor and adsorbed water play an important role in the chemistry of these surface-mediated reactions,⁵³ these initial studies were conducted in the absence of coadsorbed water. Both Fourier transform infrared (FTIR)

spectroscopy as an in situ probe of product formation and high-resolution mass spectrometry (HRMS) of solventextracted products were used to unravel the complex heterogeneous chemistry occurring on mineral dust surfaces.

MATERIALS AND METHODS

Transmission FTIR Experiments and Methods. Transmission FTIR spectroscopy was used to monitor reactions on mineral surfaces at 296 ± 1 K. Additional details of the infrared cell and gas handling system have been previously described.^{40,54–58} Mineral particles hematite (α -Fe₂O₃, 99+%, Fischer Scientific) or kaolinite (SiO₂Al₂O₃(OH)₄, Sigma-Aldrich) with a BET surface area of 80 \pm 10 and 8.4 \pm 0.5 m²/g, respectively, were heated in an oven at 473 \pm 1 K overnight to remove organic contaminants and then pressed onto one half of a tungsten grid (ca. 5 mg). The grid was then placed in the sample IR cell compartment held by two stainless steel jaws. Following the preparation of the mineral sample and placement in the IR cell, the system was evacuated for 4 h using a turbomolecular pump. Mineral surfaces were subsequently exposed to 50% RH water vapor for 2 h to ensure a fully hydroxylated terminated surface. Once hydroxylated, the system was evacuated for another 6 h to remove water vapor in the chamber. After evacuation, the sample was exposed to 100 mTorr of SO₂ (99+%, Sigma-Aldrich) for 20 min under dry conditions (RH < 1%) to obtain an SO₂-exposed mineral surface. The chamber was then evacuated for 4 h to remove gas-phase SO₂ and any weakly bonded SO₂ on the surface. This evacuation ensures that the incoming α -pinene interacts with only strongly adsorbed SO₂ species, and not with any weakly adsorbed SO₂ species. After evacuation, the sample was exposed to 500 mTorr of α -pinene (99+%, Sigma-Aldrich) for 20 min under dry conditions (RH < 1%), and the formation of OS on the surface was studied. The α -pinene sample was degassed at least three times with consecutive freeze-pump-thaw cycles prior to use.

Reactions of gas-phase HNO₃ from the nitric acid vapor taken from a concentrated mixture of H_2SO_4 (~96 wt $\cancel{>}$)-HNO₃(~70 wt %) 3:1 ratio⁵⁹ and NO₂ (26.5 ppm in N₂, Airgas) with adsorbed α -pinene-OS on mineral surfaces were studied. First, OS was produced on the surface as described above. The desired pressure of HNO_3 (50 mTorr) or NO_2 (7 mTorr) was then introduced into the IR cell. FTIR spectra were collected over a 4 h period through both halves of the tungsten grid to monitor the gas phase and particle phase. Following adsorption, the system was evacuated overnight. The experiments with HNO₃ were conducted in a different experimental but very similar setup, and instead of stainless steel a Teflon-coated cell was used. Relatively higher concentrations of reactants compared to their atmospheric concentrations were used in this study to determine the feasibility of these types of reactions. Additionally, mass spectroscopic analysis of surface products required these higher concentrations for definitive confirmation as lower concentrations yielded similar HRMS patterns with much lower intensity and higher mass error, suggesting the use of higher concentrations may not impact product formation beyond forming more products.

Prior to and following the exposure to gases, single-beam spectra (250 scans) of the surface and gas phase were acquired using a resolution of 4 cm⁻¹ and covering the spectral range of 600-4000 cm⁻¹. This was accomplished by the use of a linear translator with the infrared beam interrogating the portion of

the grid pressed with mineral particles followed by the portion of the grid left blank. Absorption spectra on mineral particles are reported as the difference in the mineral spectra before and after exposure to gases. Absorption bands due to gas-phase components, as measured through the blank half of the tungsten grid under identical conditions, were subtracted to obtain FTIR spectra of adsorbed species only.

HRMS Experiments and Methods. Organic products formed on mineral surfaces following reactions of α -pinene with inorganic gases (SO₂, NO₂, and HNO₃) were analyzed using a direct-injection linear ion trap (ThermoFisher Orbitrap) high-resolution mass spectrometer (HRMS). Adsorbed products were extracted from the hematite or kaolinite solid substrate using methanol (CH₃OH, Fisher Scientific, HPLC grade) as the solvent. Methanol was chosen as the solvent as it produced the best signal for α -pinene standards and OS products relative to other solvents and solvent mixtures (e.g., acetonitrile, ethanol, and their 1:1 mixtures with H_2O). The sample vial, syringe, and all other glassware used in the transfer process were cleaned prior to use with methanol, and Milli-Q water (Millipore Sigma, 18.2 M Ω), and baked in an oven at 773 \pm 1 K to further remove trace organics. Plastic vials used in sample preparation were sonicated in methanol for 60 min and washed thoroughly prior to use. All of the samples were stored at 253 ± 1 K and analyzed within 48 h of collection. Product stability was tested for up to 10 days in the freeze storage conditions used. It was determined that no transformations of products occurred. Extraction in different solvents confirmed no additional products were formed during the extraction step.

HRMS analysis in both positive electrospray ionization (ESI) ($[M + H]^+$ and $[M + Na]^+$) and negative ESI modes ([M - H]) was used. The heated electrospray ionization (HESI) source was operated at 373.15 K. The ESI capillary was set to a voltage of 3.5 kV at 623.15 K. The HESI-Orbitrap MS was calibrated prior to use. Mass spectra were acquired with a mass range of 50–2000 Da. Peaks with a mass tolerance of >5 ppm were rejected. Compositions were calculated with the following element ranges: 12C, 0–60; 1H, 0–150; 16O, 0–25; 14N, 0–5; 32S, 0–5; 23Na, 0–5; 39K, 0–5; 56Fe, 0–5. Tandem mass spectrometry (MS/MS) with a collision energy of 40 eV was used for structure determination.

RESULTS AND DISCUSSION

Formation of Organosulfates from α -Pinene on Mineral Surfaces. These experiments were initiated by preparing mineral surfaces with adsorbed OS compounds. This was done by first exposing mineral surfaces to gas-phase SO₂ followed by the exposure of α -pinene as described in the Materials and Methods section. For an SO₂-exposed hematite surface, several species and coordination modes of adsorbed SO₂ species have been observed (Figure S1 and Table S1). In the presence of gas-phase SO₂, these species include molecularly adsorbed SO₂, monodentate and bidentate adsorbed sulfite, adsorbed HSO3 and SO2·H2O complexes, and bidentate and bridging sulfate were observed.^{48,} Following evacuation of gas-phase SO₂, previous studies have shown that the adsorption of SO_2 on hematite and other metal oxides forms both adsorbed sulfite and sulfate that remain strongly bound to the surface.^{50,62} Similarly, for a sulfated kaolinite surface, adsorbed sulfate species were observed.

Upon exposure to α -pinene, the spectral features for α -pinene on these surfaces become apparent. After evacuation,

surface-adsorbed products were retained on the surface (Figure 1). As shown previously, α -pinene weakly adsorbs on hematite



Figure 1. FTIR spectra of surface-bound products from the reaction of α -pinene on (a) sulfated hematite and (b) sulfated kaolinite in the spectral regions from 1250 to 1800 and from 2750 to 3050 cm⁻¹. The absorbance scale is shown in the top left corner.

and undergoes reversible adsorption. Thus, adsorption of α pinene alone on hematite does not yield any strongly adsorbed product formation. In the case of kaolinite surfaces, it has been shown that α -pinene undergoes oxidation reactions on this surface to yield pinonaldehyde and pinonaldehyde dimer.⁴⁰ For hematite, following exposure of SO₂ and α -pinene, there remain adsorbed products following evacuation as evident by the peaks seen in the infrared spectrum. Several features were prominent in the following spectral range: from 2800-3000 cm⁻¹ (C–H stretch); 1300–1500 cm⁻¹ (C–H bond bending modes); and ~1236 and 1267 cm⁻¹ (C–O bond stretch). A concomitant loss of surface hydroxyl groups was observed between 3500 and 3700 cm⁻¹ similar to when SO₂ alone was adsorbed. For kaolinite, there were also peaks observed in the spectra albeit of weaker intensity. The most evident ones were from 1250 to 1550 and 3600 to 3900 cm⁻¹. However, given the complexity of products formed on the surface, it is difficult to assign these FTIR spectra to specific molecules. Generally, covalent organosulfates have two strong bands, one at 1415-1370 cm^{-1} and the other at 1200–1185 cm^{-1} , both of which are due to the stretching vibrations of the SO₂ group.⁶³ The peaks arising due to adsorbed sulfate and sulfite overlap in the same region as OS bond vibrations.⁶

To obtain additional molecular-level insights into these surface-bound products, compounds were solvent-extracted from the surface and further analyzed for product identification using HRMS (Figure 2). The HRMS patterns of surface products collected in neg-ESI mode from the reactions of α -pinene with adsorbed SO₂ species on hematite indicated the presence of several OS compounds (Figure 2 and Table 1).

These identified OS compounds correspond to $C_{10}H_{17}SO_3$ (compound 1, m/z = 217.09), $C_{10}H_{17}SO_4$ (compounds 2 and 3, m/z = 233.09), $C_{10}H_{15}SO_3$ (compound 4, m/z = 215.07), $C_{10}H_{17}SO_5$ (compound 5, m/z = 249.08), and $C_{10}H_{15}SO_4$ (compound 6, m/z = 231.07). The structures of these compounds were confirmed via tandem mass spectrometry. Furthermore, neither oxygenated α -pinene derivatives nor unreacted α -pinene were observed via HRMS analysis. For kaolinite, several OS compounds were identified from both neg- and pos-ESI modes. In neg-ESI mode, apart from compounds 1–6, two new OS were identified. These are $C_{20}H_{31}SO_7$ (compound 7, m/z = 415.18) and $C_{30}H_{47}SO_7$ (compound 8, m/z = 551.31). In pos-ESI mode, $C_{10}H_{17}SO_5$



Figure 2. HRMS patterns in both positive $[M + H]^+$ and negative $[M - H]^-$ ESI modes, of surface-bound products from the reaction of α -pinene with sulfated mineral surfaces on (a) hematite and (b, c) kaolinite. HRMS pattern for hematite in positive-ESI mode did not produce any analyte peaks. Confirmed peaks are color-coded according to heteroatoms. Pink box solid—CHOS; green box solid—CHO; black box solid—CH.

(compound 9, m/z = 249.09) and $C_{20}H_{32}SO_5Na$ (compound 10, m/z = 407.18) along with pinonaldehyde ($C_{10}H_{16}O_2Na$, compound 11, m/z = 191.10), pinonalehyde dimer ($C_{20}H_{32}O_4Na$, compound 12, m/z = 359.22), pinonic acid ($C_{10}H_{16}O_3Na$, compound 13, m/z = 207.10), and $C_{20}H_{32}O_5Na$ (compound 14, m/z = 375.21) were observed.

The presence of both compounds **2** and **3** was confirmed via MS/MS analysis. Similarly, the structure for **9** was determined based on its major fragments at m/z = 223.08 ($C_{10}H_{16}SO_2Na$) and 191.10 ($C_{10}H_{16}O_2Na$), which suggested the presence of a pinonaldehyde moiety. Table S2 provides detailed information on the mass fragmentation of identified compounds. Similar OS compounds were identified in various field and chamber studies, suggesting the formation of these OS compounds from α -pinene occurs in a broader environment.^{19,26,27,41,65–67}

Formation pathways of these OS compounds are facilitated on these mineral surface mechanisms (Scheme 1).^{70,73} Briefly, the adsorbed SO₂ species on mineral surfaces can interact with π bonds of α -pinene in the gas phase, producing a series of OS compounds.^{26,27,78,79} Compounds 1 and 2 can form via carbocation formation from the reaction of α -pinene with adsorbed sulfate/sulfite on mineral surfaces. The reaction of α -pinene π bonds with acid groups such as sulfate was previously proposed.¹⁷ The carbocation can react with adsorbed sulfite or sulfate species on mineral surfaces to form the isomers of compounds 1 or 2, respectively (Scheme 1a). Compounds 3 and 5 can form via the interaction of π bonds in α -pinene with the surface redox sites/surface hydroxyl groups on mineral surfaces, leading to the formation of pinene oxide-like species. A similar reaction is prominent in limonene (an isomer of α -pinene) forming limonene oxide;⁸⁰⁸⁰ here, the interactions are expected to be weaker for α -pinene given no α -pinene oxide was observed from its interactions with mineral surfaces.

The reaction of adsorbed sulfite or sulfate species on mineral surfaces with one of the two epoxide carbons on α -pinene can yield compound 3 or 5. The presence of peaks corresponding to HSO₃⁻ (m/z = 80.97), SO₃ (m/z = 79.95), and HSO₄⁻ (m/z = 233.09 in the HRMS fragmentation pattern of m/z = 233.09 further confirms the formation of both sulfate and sulfite groups containing α -pinene-derived OS compounds.

Two formation pathways are proposed for compound 4. First, the dehydration of compound 3 yields compound 4. Second, the loss of 2H atoms from compound 1 in the mass spectrometer can yield m/z = 215.07. Similarly, compound 6 can form from the dehydration of compound 5 or the removal of 2H atoms from compound 2 during the HRMS analysis (Scheme 1b). Dehydration of compounds 3 and 5 from compounds 4 and 6, respectively, can occur either on mineral surfaces or during the HRMS analysis. Duporté et al. observed an OS at 249 (compound 5 in this study) dehydrating to an OS at 231 (compound 6) during MS analysis.²⁷

On kaolinite, α -pinene derivatives pinonaldehyde (11) and pinonaldehyde dimer (12) are formed⁴⁰ and further oxidized to form their respective acids (compounds 13 and 14). Formation of pinonaldehyde and pinonaldehyde dimer upon exposure to kaolinite surfaces is explained in detail in our previous study.⁴⁰ Briefly, these are from the dihydroxylation of double bond on α -pinene followed by glycol cleavage on the kaolinite surface. Compound 7 can be expected to form from the reaction of adsorbed sulfites with compounds 11 and 12, respectively. In this case, adsorbed sulfites act as a nucleophile and attack aldehyde carbon on either compound 9 or 7. Compound 9 can also form from the olefinic cleavage of compound 4 in a similar way to pinonaldehyde formation from α -pinene. Compound 8 can be expected to form due to the reaction between compounds 14 and 1. Similarly, compound 10 may be formed due to the reaction between compounds 13 and 1 (Scheme 1c,d). Therefore, these results show the

Table 1. Identified Surface Products Using HRMS and MS/MS in Either Positive $[M + H]^+$ or Negative ESI $[M - H]^-$	⁻ Modes
from the Reactions of α -Pinene on Sulfated Mineral Surfaces ^{<i>abcd</i>}	

ID	Observed Formula,	m/z	Proposed Structure	Hematite	Kaolinite	Previously reported in
	ESI Mode (Molecular Formula)					field/ laboratory studies ^a
1	$C_{10}H_{17}SO_3, -$ ($C_{10}H_{18}SO_3$)	217.09	$\begin{array}{c} \begin{array}{c} & & \\ $	√b		N/A ^c
2	$C_{10}H_{17}SO_{4,}$ - ($C_{10}H_{18}SO_{4}$)	233.09	2(a) 2(b)			N/A
3	C ₁₀ H ₁₇ SO ₄ , - (C ₁₀ H ₁₈ SO ₄)	233.09	$\begin{array}{c} \begin{array}{c} \begin{array}{c} \begin{array}{c} \\ \\ \\ \\ \\ \\ \\ \end{array} \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} \\ \end{array} \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \end{array} $	\checkmark	\checkmark	N/A
4	$C_{10}H_{15}SO_{3,}-$ ($C_{10}H_{16}SO_{3}$)	215.07	$\begin{array}{c} \begin{array}{c} \begin{array}{c} & \\ \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ \end{array} \\ \end{array} \\ \begin{array}{c} \\ \\ \\ \end{array} \\ $			N/A
5	$C_{10}H_{17}SO_{5,}$ - ($C_{10}H_{18}SO_{5}$)	249.08	$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	\checkmark	√	22,26,27,42,65,67–69
6	$\begin{array}{c} C_{10}H_{15}SO_{4,}-\\ (C_{10}H_{16}SO_{4})\end{array}$	231.07			√	26,27
7	$C_{20}H_{31}SO_7, -$ ($C_{20}H_{32}SO_7$)	415.18	it of a transformed	^d		N/A
8	C ₃₀ H ₄₇ SO ₇ , - (C ₃₀ H ₄₈ SO ₇)	551.31	it it is int			N/A
9	$\begin{array}{c} C_{10}H_{17}SO_{5,}+\\ (C_{10}H_{16}SO_{5})\end{array}$	249.09	HOSO			22,24,41,65–67,69
10	$C_{20}H_{32}SO_5Na_5 + (C_{20}H_{32}SO_5)$	407.18	it i is the		\bigvee	N/A
11	$C_{10}H_{16}O_2Na_1 + (C_{10}H_{16}O_2)$	191.10				40,68,70–72
12	$C_{20}H_{32}O_4Na_4 + (C_{20}H_{32}O_4)$	359.22	ip of i			40,73-75
13	$C_{10}H_{16}O_3Na_3 + (C_{10}H_{16}O_3)$	207.10	H _O			22,42,68,71,72,75,76
14	$C_{20}H_{32}O_5Na, +$ ($C_{20}H_{32}O_5$)	375.21	С ОН ОН ОН ОН			74–77

^{*a*}One or more of the compound, molecular formula, and m/z has been reported. ^{*b*} $\sqrt{}$ = Detected in HRMS. ^{*c*}N/A = Not available. ^{*d*}-- = Not detected in HRMS.

formation of these multiple OS surface products from reactions of α -pinene with adsorbed SO₂ species on mineral surfaces in dark and dry conditions, in the absence of strong atmospheric oxidizers such as HO radicals and O₃, thus underscoring the unique role the surface plays. These compounds interact with adsorbed SO₂ species on surfaces, yielding a more complex mixture of SOAs. One such possible interaction is the relatively well studied sulfate ester and OS formation from pinonaldehyde.^{16,22,26,27,41,66}

Reactions of Gas-Phase Nitric Acid and Nitrogen Dioxide with Adsorbed α -Pinene-OS. In these FTIR experiments, the mineral surfaces were prepared as described above to form adsorbed OS. These surfaces were then exposed to either gas-phase HNO₃ or NO₂ for several hours (4 h). The evacuated surfaces are shown in Figure 3. Upon exposure to HNO₃, new spectral features around 1500–1550 and ~1300 cm⁻¹ started to appear on the hematite surfaces. These are attributed to different types of adsorbed nitrate species.⁵⁸ Moreover, asymmetric and symmetric stretching vibrations of the NO₂ group of ON occur at 1615–1660 and 1270–1285 cm^{-1.63} In addition to these, peaks at 2878, 2961, and 2986 cm⁻¹ were attributed to C–H stretching modes, and peaks at 1472 and 1458 cm⁻¹ were assigned to C–H bending modes. In contrast, for kaolinite surfaces, only weaker spectral features corresponding to organics at 2875, 2927, and ~1325 cm⁻¹ were observed. However, a strong band at 3233 cm⁻¹ was

Scheme 1. Proposed Formation Pathways for Several of the Compounds Observed in HRMS Compounds (a) 1 and 2, (b) 3, 4, 5, and 6 on Mineral Surfaces and (c) 9, 10, 11, and 13, (d) 7, 8, 12, and 14 on Kaolinite Surfaces



observed, suggesting the presence of hydrogen-bonded networks that could include the presence of alcohols, and/or protonated carboxylic acids. Upon exposure to NO₂, minimal changes of spectral features from OS adsorbed on hematite were observed at ~1370–1420 cm⁻¹, but a weak band at 1550 cm⁻¹ was seen.

For kaolinite, spectral features around $1300-1450 \text{ cm}^{-1}$ as well as weaker features at ~2967 cm⁻¹ were observed, suggesting the presence of C–H bonds on the surface. Furthermore, a strong band at 3488 cm⁻¹ was observed, indicating the presence of O–H groups. However, given the complexity of products formed on the surface for specific compounds, these surface-bound products were extracted and analyzed using HRMS and MS/MS to identify products and for structural determination.

Product Formation from the Reactions of Adsorbed *α*-Pinene-OS with Gas-Phase HNO₃. The HRMS patterns collected in both neg- and pos-ESI modes of surface products from reactions of HNO₃ with adsorbed OS species on hematite indicated the presence of several OS compounds, ON compounds, and NOS compounds (Figure 4a,b, Table 2). Apart from the presence of compounds 1 and 2, a new OS, Compound 15 (C₉H₁₄SO₆, sulfated pinalic-3-acid, m/z = 369.12 for C₁₂H₂₆SO₉Na) was observed. Compound 15 was observed as a sodated cluster of three methanol units and with a major fragment at 365.09 for C₁₂H₂₂SO₉Na. Furthermore, m/z = 233.09 was assigned only to compound 2, as no indication of SO₃ was shown on its MS/MS fragmentation. Two major ON products were identified. Compound 16 (C₁₀H₁₆NO₃, m/z = 198.15) was identified with its major



Figure 3. FTIR spectra of surface-bound products from the reaction of adsorbed α -pinene-OS with either HNO₃(vapor) or NO₂(g) on (a) hematite and (b) kaolinite in the spectral regions from 2800 to 3800 cm⁻¹ and from 1190 to 1800 cm⁻¹. The absorbance scale is shown in the upper left for each surface.

fragment 166.12 for C₁₀H₁₆ON. Compound 17 (C₂₀H₃₄NO₃, m/z = 336.25) was identified with major fragments at 320.26 for C₂₀H₃₄NO₂, 302.25 for C₂₀H₃₂ON, and 168.14 for C₁₀H₁₈ON. Two NOS were identified. Compound **18** (C₁₀H₁₆NSO₇, m/z = 294.06) was identified with a major fragment at 278.07 for C₁₀H₁₆NSO₆. Compound **19** (C₁₉H₃₂NSO₁₀, m/z = 595.29 for C₂₃H₄₉NSO₁₄) was identified as a 4 methanol-units cluster. Its major fragments were at m/z = 579.29 for C₂₃H₄₉NSO₁₃ and 302.14 for C₁₄H₂₄NSO₄. A smaller m/z at 467.18 for C₁₉H₃₃NSO₁₀ (compound **19** without methanol units clustered) was also observed. The nonsulfated molecule of compound **15**, pinalic-3-acid, ^{3,68,81–85} has been widely observed in the environment as well as in laboratory experiments. Similarly, compounds with

similar molar weight/formula to compounds **16**, **17**, **18**, and **19** have been previously observed in both laboratory experiments and the field.^{11,19,22,40,42,68,69,71,72,86–89}

Despite wide observations of pinalic-3-acid in the environment, the mechanistic details of how this compound forms, as well as that of compound **15**, are unclear. Pinalic-3-acid forms via the ozonolysis of α -pinene. In the current study, it can be speculated that HNO₃/adsorbed nitrates on hematite may induce the formation from the OS present on the hematite surface. Given that HSO₄⁻ is a good leaving group,⁹⁰ the formation of sulfated pinalic-3-acid (**15**) can occur either via oxidation of sulfated α -pinene, or sulfation of formed pinalic-3acid by the surface-adsorbed sulfate species. Further research studies are recommended to better understand the underlying



Figure 4. HRMS patterns in both positive $[M + H]^+$ and negative $[M - H]^-$ ESI modes of surface-bound products from the reactions of adsorbed α -pinene-OS with HNO₃(vapor) on (a, b) hematite and (c, d) kaolinite. Confirmed peaks are color-coded according to heteroatoms. Pink box solid—CHOS; blue box solid—CHON; orange box solid—CHNOS; green box solid—CHO; black box solid—CH.

mechanisms of formation of compound 15 on mineral surfaces in the presence of HNO₃.

A smaller quantity of compound 16 was observed in the study. Compound 16 was previously observed in our studies forming from reactions of α -pinene with adsorbed nitrates on hematite surfaces.⁴⁰ However, here, nitrate is likely to act as a nucleophile on either compound 1 or 2 followed by the leaving of sulfate/sulfite group as their respective acid, resulting in a C=C on compound 16 (Scheme 2). Sulfate species are known to be good leaving groups.90

Formation of compound 17 suggests that α -pinene-OS may undergo dimerization as well. Both α - and β -pinene dimerization has previously been studied.^{91,92} We speculate this is due to the interaction of surface hydroxyl groups with pi hydrogens in the α -pinene moiety of compounds 4, 6, and 16, a carbocation forms and undergoes dimerization with compound 4 or 6 on the surface followed by elimination of H_2SO_3/H_2SO_4 . However, the exact formation pathway remains unclear. Moreover, it is worth investigating whether their formation is possible at low α -pinene concentrations.

The formation of compound 18 is proposed to occur via the addition of an HNO3 molecule on the double bond of either compound 4 or 6. It is expected that the sulfite group in

compound 4 is oxidized by HNO₃ molecules in the system. Compound 19 likely forms via reactions between compound 15 (or pinalic-3-acid) and compound 18. The nitrate group in compound 18 may react with either $-COOSO_3$ (-COOH in pinalic-3-acid) or aldehyde end. However, we propose the reaction undergoes via aldehyde end due to the formation of fragment 302.14 for C14H24NSO4. This fragment may be possible if the reaction occurs via the aldehyde end. A smaller peak at m/z = 547.14 for C₁₉H₃₃NS₂O₁₃ was also observed, suggesting the presence of two sulfate groups in the same molecule. However, it was not listed as a separate compound due to the lower occurrence. Despite the fact that our experimental evidence suggests the possible formation pathway for Compound 19 is as shown in Scheme 2c, given the nitrate esters are usually considered to be weak nucleophiles, a more detailed investigation of high molecular weight NOS formation, such as Compound 19, is warranted.

Five α -pinene-OS products were observed on kaolinite surfaces after exposure to HNO3 along with three ON and two NOSs. The OS are compounds 1, 2, 3, 8, and 9. The ON, compound 20 ($C_{20}H_{32}NO_3$, m/z = 334.24, pos and $C_{20}H_{30}NO_3$, m/z = 332.22, neg), was identified. Another possible ON, compound 21 ($C_{30}H_{47}NO_3Na$, m/z = 492.34),

Table 2. Identified Surface Products Using HRMS and MS/MS in Either Positive $[M + H]$ + or Negative ESI $[M - H]^-$ Modes
from the Reactions of Adsorbed α -Pinene-OS with Either HNO ₃ or NO ₂ on Mineral Surfaces ^{<i>abcde</i>}

ID	Observed Formula,	m/z	Proposed Structure	Hematite		Kaolinite	Previously reported in
	ESI Mode (Molecular Formula)						field/ laboratory studies ^a
				AP+HNO ₃	AP+NO ₂	AP+HNO ₃	
15	$C_{12}H_{26}SO_9Na_1 + (C_9H_{14}SO_6)$	369.19	0 C C C C C C C C C C C C C C C C C C C	√ b	c		3,24,85,86,65,68,69,72,81–84
16	$C_{10}H_{16}NO_{3,}+$ ($C_{10}H_{15}NO_{3}$)	198.15	OFZ 0		\checkmark		40,71,72
17	C ₂₀ H ₃₄ NO ₃ , + (C ₂₀ H ₃₃ NO ₃)	336.25	o ^{°N} o d	√			86
18	C ₁₀ H ₁₆ NSO ₇ , - (C ₁₀ H ₁₇ NSO ₇)	294.06	о _{ї у} 20 — Добудо, н	\checkmark		\checkmark	11,19,22,42,68,69,87–89
19	C ₂₃ H ₄₉ NSO ₁₄ , + (C ₁₉ H ₃₂ NSO ₁₀)	595.29 and 467.18	H = 0 $H = 0$ $H =$	V			22
20	$\begin{array}{c} C_{20}H_{32}NO_3, + \\ C_{20}H_{30}NO_3, - \\ (C_{20}H_{31}NO_3) \end{array}$	334.24, + 332.22, -	os _N o d		\checkmark	\checkmark	86
21	C ₃₀ H ₄₇ NO ₃ Na, + (C ₃₀ H ₄₇ NO ₃)	492.34	N/A				N/A ^e
22	C ₂₀ H ₃₂ NSO ₇ , - (C ₂₀ H ₃₃ NSO ₇)	430.19				\checkmark	N/A
23	C ₃₂ H ₅₀ NSO ₆ , - (C ₃₂ H ₅₁ NSO ₆)	576.33	N/A			\checkmark	N/A
24	C ₁₀ H ₁₇ O, + (C ₁₀ H ₁₆ O)	153.13	• TA		√		26,40,71,72
25	$C_{10}H_{18}O_2Na, + (C_{10}H_{18}O_2)$	193.12	но он				72

"One or more of the compound, molecular formula, and m/z has been reported. ${}^{b}\sqrt{}$ = Detected in HRMS. "-- = Not detected in HRMS." matched structure with HRMS, MS/MS data, and molecular formula." ${}^{e}N/A$ = Not available.

with three α -pinene moieties was identified in positive-ESI mode. Compound 21 produced several key fragments. Among these are 460.35 for $C_{30}H_{47}ONNa$, 438.37 for $C_{30}H_{48}ON$, and 324.23 for C₂₀H₃₁ONNa. Compound **18** (C₁₀H₁₆NSO₇, m/z =294.06) was identified in the neg-ESI mode with a major fragment at 278.07 for C₁₀H₁₆NSO₆. Another NOS, compound 22 ($C_{20}H_{32}NSO_7$, m/z = 430.19), in the neg-ESI mode produced a series of fragments at 428.18 (C₂₀H₃₀NSO₇), 412.18 (C₂₀H₃₀NSO₆), and 396.19 (C₂₀H₃₀NSO₅). A major peak was observed at 576.33 with the formula of $C_{32}H_{50}NSO_{67}$ and its fragments at 456.20 for C₂₉H₃₀NSO₂, 440.21 for C₂₉H₃₀NSO, and 351.20 for C₂₀H₃₁SO₃ were assigned to belong to the same compound (23). The formula or the structure of compound 23 was not identified in these studies. Many studies have reported the formation of high molecular weight ON, OS, and \overline{NOS} in the environment.^{11,19,22,68,69,86–8} Similar to these studies, most interesting is that in this current study, the formation of a unique mixture of SOAs on these surfaces underscores the complexity of these reactions.

Furthermore, highlighting the importance of mineral surfaces, formation of a C₉ compound (**15**) facilitating further SOA formation of a C₁₉ compound (**19**) is observed only with hematite surfaces in the presence of HNO₃. Hematite and kaolinite surfaces represent both redox-active minerals and common clay minerals, respectively. The differences in the surface chemistry reflect the different surface properties—surface structure, surface hydroxyl group density, redox active sites, and specific surface planes. The details of these different properties and how they play a role in the reaction mechanism have yet to be fully discerned (Figure 5).

Product Formation from the Reactions of Adsorbed *α***-Pinene-OS with Gas-Phase NO**₂. Reactions of NO₂ with adsorbed *α*-pinene-OS on hematite surfaces produced a series of compounds. Some of these compounds were also observed during reactions of *α*-pinene-OS with HNO₃ and others are unique. Some of these differences may arise from the lower acidity resulting in the mineral surface upon exposure to NO₂ than when exposed to HNO₃ as acid-catalyzed reactions can

Scheme 2. Proposed Formation Pathways for Compounds Identified in HRMS: (a) Compound 15, (b) Compound 16, (c) Compound 18, and (d) Compound 19 on Mineral Surfaces



Article

induce formation of products in the presence of HNO₃ vapor. Among these were two OS, compounds 2 and 3, two ON, compounds 16 and 20, and four oxidized α -pinene derivatives, compounds 11, 13, 24 ($C_{10}H_{17}O_{1}$, m/z = 153.13 and $C_{10}H_{16}ONa$, m/z = 175.11, α -pinene oxide), and 25 $(C_{10}H_{18}O_2Na, m/z = 193.12, \alpha$ -pinene diol). Compounds 2 and 3 were identified in both negative and positive ESIs modes. Furthermore, a range of peaks containing NSO groups were identified in positive-ESI mode, suggesting the formation of NOS during these reactions. These peaks were 535.31 for C₁₉H₃₀NSO₁₃Na, 362.07 for C₁₇H₁₆NSO₆, 321.04 for C₁₆H₁₂NSO₃Na, 279.06 for C₆H₁₇NSO₉, and 113.96 for CHNSO₂Na. Although none of these peaks were assigned to a particular compound, it is important to note that both C_{17} and C_{16} compounds have been previously observed from α pinene SOAs.^{85,86,93} Similarly, reactions of adsorbed α -pinene-OS on kaolinite surfaces with NO2 produced a mixture of different SOAs. The OS 1, 2, 3, 7, 9, and 10 were observed. The oxygenated α -pinene derivatives, compounds 11, 12, and 13, were also observed. However, no new products compared to α -pinene alone with sulfated kaolinite was observed. Several peaks containing NSO were observed in both neg- and pos-ESI modes. However, similar to the case of hematite, here also no peak was assigned to a particular NOS compound due to the complexity of HRMS and MS/MS data. Overall, the current study suggests the formation of a unique mixture of ON and NOS and other oxidation products from the reactions of gasphase NO₂ with adsorbed α -pinene-OS on mineral surfaces. Future research studies are recommended for understanding the reactions of NO2 with VOC-derived OS adsorbed on mineral surfaces to better understand these reaction pathways. In particular, these studies can benefit from computational analyses. Quantification of identified surface products was not

conducted in the current study due to the lack of availability of suitable standards.

CONCLUSIONS AND ATMOSPHERIC IMPLICATIONS

The reactions of α -pinene on sulfated hematite and kaolinite surfaces at 296 K yield several different OS and oxidation products of α -pinene. Kaolinite surfaces produce pinonaldehyde and several OS compounds with a pinonaldehyde moiety, whereas all OS identified on hematite retained the α -pinene moiety, highlighting the underlying differences in the role of mineralogy and surface-specific chemistry. The adsorbed α pinene-OS reacts with HNO₃ or NO₂ to produce more OS, multiple ON, and NOS products on surfaces. The proposed formation pathways of these OS occur via an adsorbed sulfite/ sulfate species. The formation of ON and NOS compounds in the presence of HNO₃ suggests the addition of HNO₃, nucleophilic attacks by nitrates on OS, and reactions between ON and OS on the surface. Among other products, C10H17NSO7, commonly known as NOS-295, was observed on both surfaces from reactions of OS with HNO₃. On hematite, OS $(C_9H_{14}SO_6)$ and NOS $(C_{19}H_{32}NSO_{10})$ were observed, indicating the formation of C₉ compounds in the presence of HNO₃, thus underscoring the specific roles of mineral surfaces. The majority of surface products identified from the reactions of OS with NO2 on both surfaces were either OS or oxidized α -pinene derivatives, along with evidence of the formation of NOS.

Overall, this study shows how mineral dust aerosol surface can interact with inorganic and organic gas-phase compounds to yield a wide range of compounds important for secondary organic aerosol surfaces. These data show that surfaces can interact with gas-phase SO_2 that then lead to OS surface products in the presence of ubiquitous organic compounds



Figure 5. HRMS patterns in both positive $[M + H]^+$ and negative $[M - H]^-$ ESI modes of surface-bound products from the reactions of adsorbed α -pinene-OS with NO₂(gas) on (a, b) hematite and (c, d) kaolinite. Confirmed peaks are color-coded according to heteroatoms. Pink box solid—CHOS; blue box solid—CHON; orange box solid—CHNOS; green box solid—CHO; black box solid—CH.

such as α -pinene and then further can yield ON and NOS compounds when gas-phase nitrogen oxides are present. Most notable is that the specific products that form are mineral specific. Additionally, some of the SOAs observed in this study have been repeatedly observed in field studies; yet, formation pathways are poorly understood. In the environment, VOCs such as α -pinene are ubiquitous. During dust transport and urban dust episodes, these VOCs and trace gas pollutants such as NO₂ and SO₂ can interact with each other, forming more highly functional, less volatile compounds as shown here on two different mineral surfaces under dry conditions. As water plays an important role in heterogeneous chemistry,⁵³ it is important to further investigate how adsorbed water molecules impact these reactions. This study suggests heterogeneous reactions are possible pathways for the formation of these OS, ON, and NOS compounds but we also note the importance of extending this to different environmental conditions (e.g., relative humidity and solar radiation) on these formation pathways.

ASSOCIATED CONTENT

Supporting Information

The Supporting Information is available free of charge at https://pubs.acs.org/doi/10.1021/acsearthspace-chem.2c00259.

FTIR spectra of sulfated mineral surfaces (Figure S1); vibrational frequencies of different adsorbed sulfurcontaining species on hematite and kaolinite (Table S1); and MS/MS analysis of parent peaks for the identified compounds (Table S2) (PDF)

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Notes

The authors declare no competing financial interest.

ACKNOWLEDGMENTS

This study is supported by the National Science Foundation under Grant CHE-2002607. The authors acknowledge Cholaphan Deeleepojananan and Dr. Shubhrangshu Pandit for helpful discussions.

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