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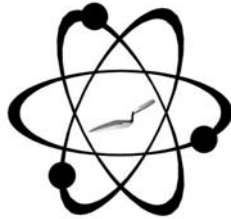
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**SOURCE PROVENANCE OF OBSIDIAN ARTIFACTS FROM THREE
SUBSURFACE TEST UNITS AT PIEDRAS MARCADAS (LA 290), MIDDLE RIO
GRANDE VALLEY, NEW MEXICO**

by

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Report Prepared for

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INTRODUCTION

The analysis here of 61 obsidian artifacts from three test units at Piedras Marcadas (LA 290) in the middle Rio Grande River valley indicates a slightly different provenance mix than the results from the general surface (Shackley 2009a). In the test unit case, all the artifact quality sources of archaeological obsidian present in the Jemez Mountains, both pre-caldera and caldera event sources occur in the assemblage. All these sources are present in the Rio Grande alluvium as far south as Albuquerque, although the Valles Rhyolite (Cerro del Medio) nodules are very small. No Mount Taylor obsidian was recovered sub-surface (Shackley 2012). Mount Taylor is not available in Rio Grande Quaternary sediments this far north.

ANALYSIS AND INSTRUMENTATION

All archaeological samples are analyzed whole. The results presented here are quantitative in that they are derived from "filtered" intensity values ratioed to the appropriate x-ray continuum regions through a least squares fitting formula rather than plotting the proportions of the net intensities in a ternary system (McCarthy and Schamber 1981; Schamber 1977). Or more essentially, these data through the analysis of international rock standards, allow for inter-instrument comparison with a predictable degree of certainty (Hampel 1984).

The trace element analyses were performed in the Geoarchaeological XRF Laboratory, Albuquerque, New Mexico, using a Thermo Scientific *Quant'X* energy dispersive x-ray fluorescence spectrometer. The spectrometer is equipped with a ultra-high flux peltier air cooled Rh x-ray target with a 125 micron beryllium (Be) window, an x-ray generator that operates from 4-50 kV/0.02-1.0 mA at 0.02 increments, using an IBM PC based microprocessor and WinTrace™ 4.1 reduction software. The spectrometer is equipped with a 2001 min⁻¹ Edwards vacuum pump for the analysis of elements below titanium (Ti). Data is acquired through a pulse processor and analog to digital converter. This is a significant improvement in analytical speed and efficiency beyond the former Spectrace 5000 and *QuanX* analog systems (see Davis et al. 2011; Shackley 2011).

For Ti-Nb, Pb, Th elements the mid-Zb condition is used operating the x-ray tube at 30 kV, using a 0.05 mm (medium) Pd primary beam filter in an air path at 200 seconds livetime to generate x-ray intensity $K\alpha_1$ -line data for elements titanium (Ti), manganese (Mn), iron (as Fe^T), cobalt (Co), nickel (Ni), copper, (Cu), zinc, (Zn), gallium (Ga), rubidium (Rb), strontium (Sr), yttrium (Y), zirconium (Zr), niobium (Nb), lead (Pb), and thorium (Th). Not all these elements are reported since their values in many volcanic rocks is very low. Trace element intensities were converted to concentration estimates by employing a least-squares calibration line ratioed to the Compton scatter established for each element from the analysis of international rock standards certified by the National Institute of Standards and Technology (NIST), the US. Geological Survey (USGS), Canadian Centre for Mineral and Energy Technology, and the Centre de Recherches Pétrographiques et Géochimiques in France (Govindaraju 1994). Line fitting is linear (XML) for all elements but Fe where a derivative fitting is used to improve the fit for iron and thus for all the other elements. When barium (Ba) is acquired, the Rh tube is operated at 50 kV and 0.5 mA in an air path at 200 seconds livetime to generate x-ray intensity $K\alpha_1$ -line data, through a 0.630 mm Cu (thick) filter ratioed to the bremsstrahlung region (see Davis et al. 2011). Further details concerning the petrological choice of these elements in North American obsidians is available in Shackley (1988, 1990, 1995, 2005; also Mahood and Stimac 1991; and Hughes and Smith 1993). A suite of 17 specific standards used for the best fit regression calibration for elements Ti- Nb, Pb, and Th, include G-2 (basalt), AGV-2 (andesite), GSP-2 (granodiorite), SY-2 (syenite), BHVO-2 (hawaiite), STM-1 (syenite), QLO-1 (quartz latite), RGM-1 (obsidian), W-2 (diabase), BIR-1 (basalt), SDC-1 (mica schist), BCR-2 (basalt), TLM-1 (tonalite), SCO-1 (shale), all US Geological Survey standards, NBS-278 (obsidian) from the National Institute of Standards and Technology, BR-1 (basalt) from the Centre de Recherches Pétrographiques et Géochimiques in France, and JR-1 and JR-2 (obsidian) from the Geological Survey of Japan (Govindaraju 1994).

The data from the WinTrace software were translated directly into Excel for Windows and into SPSS for statistical manipulation (Table 1). In order to evaluate these quantitative determinations, machine data were compared to measurements of known standards during each run (Table 1). RGM-1 is analyzed during each sample run for obsidian artifacts to check machine calibration (Table 1). Source assignments made by reference to source data at Berkeley, Baugh and Nelson (1987) and Shackley (1995, 2005).

DISCUSSION

Before a discussion of the source provenance of the samples, a short discussion of the Jemez Mountains sources is in order. Following this is a short discussion of the samples proper.

The Jemez Mountains and the Sierra de los Valles

A more complete discussion of the archaeological sources of obsidian in the Jemez Mountains is available in Shackley (2005:64-74). Distributed in archaeological contexts over as great a distance as Government Mountain in the San Francisco Volcanic Field in northern Arizona, the Quaternary sources in the Jemez Mountains, most associated with the collapse of the Valles Caldera, are distributed at least as far south as Chihuahua through secondary deposition in the Rio Grande, and east to the Oklahoma and Texas Panhandles through exchange. And like the sources in northern Arizona, the nodule sizes are up to 10 to 30 cm in diameter; El Rechuelos, Cerro Toledo Rhyolite, and Valles Rhyolite (Valles Rhyolite derived from the Cerro del Medio dome complex) glass sources are as good a media for tool production as anywhere. Until the recent land exchange of the Baca Ranch properties, the Valles Rhyolite primary domes (i.e., Cerro del Medio) have been off-limits to most research. The discussion of this source group here is based on collections by Dan Wolfman and others, facilitated by Los Alamos National Laboratory, and the Museum of New Mexico, and recent sampling of all the major sources courtesy of the Valles Caldera National Preserve (VCNP; Shackley 2005; Wolfman 1994).

There are at least four eruptive events in the last 8.7 million years that have produced the four chemical groups in the Jemez Mountains (Figure 1).

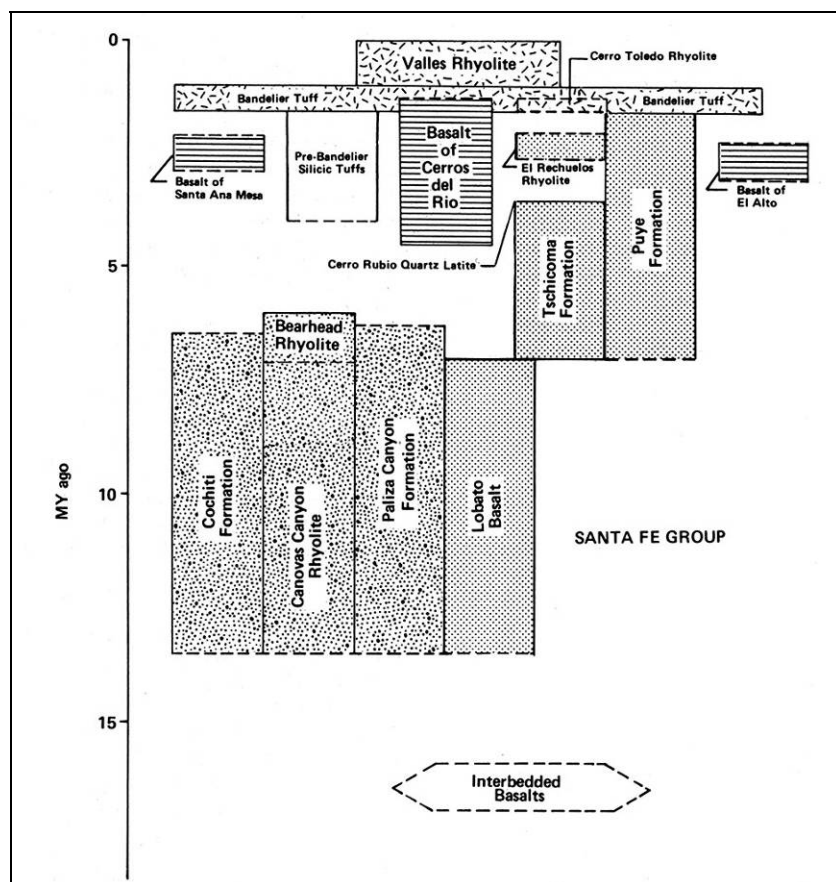


Figure 1. Generalized stratigraphic relations of the major volcanic and alluvial units in the Jemez Mountains (from Gardner et al. 1986). Note the near overlapping events at this scale for the Cerro Toledo and Valles Rhyolite members, and the position of Cerro Toledo Rhyolite at the upper termination of the Puye Formation.

The earliest is the Bear Springs Peak source, part of Canovas Canyon Rhyolite that is dated to about 8.7 mya, firmly in the Tertiary (Kempton et al. 2004; Figure 1 here). This source is a typical Tertiary marekanite source with remnant nodules embedded in a perlitic matrix. It is located in a dome complex including Bear Springs Peak on Santa Fe National Forest and radiating to the northeast through Jemez Nation land (Shackley 2009b). While the nodule sizes are small, the glass is an excellent media for tool production and has been found archaeologically at Zuni and in secondary deposits as far south as Las Cruces (Church 2000; Shackley 2012). Four of the samples were produced from this source (Table 1 and 2).

Part of the same Keres Member as Canovas Canyon Rhyolite is Paliza Canyon Rhyolite. They have similar elemental chemistry and are likely nearly contemporaneous. This source is rare in archaeological contexts, but occurs in Rio Grande alluvium, and is present as one sample here.

The second relevant eruptive event that produced artifact quality obsidian is the El Rechuelos Rhyolite. This source, present as one sample here, is what I consider the best media for tool production of the group. It dates to about 2.4 million years ago, and nodules at least 10 cm in diameter are present in a number of domes north of dacite Polvadera Peak, the incorrect vernacular name for this source. El Rechuelos has eroded through the Rio Chama into the Rio Grande and has also been found in alluvium into southern New Mexico (Church 2000; Shackley 2012).

About 1.4 mya, the first caldera collapse occurred in the Jemez Mountains, called Cerro Toledo Rhyolite. This very large event produced the Bandelier Tuffs and spread ash flows many kilometers into the area and horizontally southwest from what is now Rabbit Mountain and the Cerro Toledo domes to the east. These large ash flow sheets are responsible for the great quantity of Cerro Toledo obsidian that is present in the Quaternary Rio Grande alluvium all the way to Chihuahua (Church 2000; Shackley 2005, 2012). Cerro Toledo Rhyolite secondary deposit nodules is present relatively near to Piedras Marcadas on Quaternary terraces above the east side of the Rio Grande, including Placitas and the sands near Tijeras Wash south of the Albuquerque airport (Shackley 2005).

The second caldera collapse that produced the Valles Rhyolite member of the Tewa Formation, called Valles Rhyolite here, occurred around one million years ago and created most of the geography of the current Valles Caldera. A number of rhyolite ring domes were produced on the east side of the caldera, but only Cerro del Medio produced artifact quality obsidian. Indeed, the Cerro del Medio dome complex produced millions of tons of artifact quality glass, and is the volumetrically largest obsidian source in the North American Southwest challenged

only by the Government Mountain dome complex in the San Francisco Volcanic Field. Cerro del Medio obsidian was apparently preferred by Folsom knappers, as well as those in all periods since. While Cerro Toledo probably appears in archaeological contexts in New Mexico sites with greater frequency, it is likely because it is distributed in secondary contexts. Valles Rhyolite (Cerro del Medio), present as three samples here importantly does not erode outside the caldera, in and quantity and size and likely had to be originally procured in the caldera proper (Shackley 2005). All of the four Valles Rhyolite samples are small, although the one projectile point tip indicates a size larger than nodule sizes thus far recovered in the Rio Grande alluvium, and was probably produced from raw material from Cerro del Medio proper, suggesting procurement at the source or exchange of primary source obsidian.

Source Provenance Discussion

Most of the artifacts analyzed produced from all these sources are bipolar core or flake fragments and most appear to have waterworn cortex. This suggests that most of these raw materials were procured across the river somewhere. In the case of the Mount Taylor specimen the raw material had to be procured at Mount Taylor or in the Rio Puerco or the Rio Grande south of Socorro after the Rio Puerco joins the Rio Grande (Shackley 2005, 2012). Mount Taylor sources (Grants Ridge, Horace and La Jara Mesas) are common in historic period contexts at Zuni and the source may have been “controlled” by the Zuni (Shackley 2005; Table 1 and Figure 3). The issue of the point fragment from Valles Rhyolite raw material is discussed above. The mix of sources in the test unit assemblage mirrors the mix of sources recovered from the Rio Grande Quaternary Alluvium at Tijeras Wash almost identically (Figures 2). This is the strongest argument for local procurement of obsidian toolstone at Piedras Marcadas as indicated by the test unit assemblage.

Surface versus Subsurface Results

While the samples are relatively small, the mix of sources recovered from surface contexts versus the subsurface test unit sample is somewhat different (Figure 4). While both are

dominated by Jemez Mountains secondary deposit sources, the presence of Mount Taylor sources on the surface indicates procurement through direct access to the Zuni region or exchange with the Zuni. IF the subsurface material is earlier, and the surface material later, then one change seen is contact to the west rather than local procurement and/or contact north at an earlier period. It is possible that the Mount Taylor obsidian was procured by the Coronado Expedition knappers when they were at and around Zuni and transported the raw material to Piedras Marcadas during the siege as tool raw material. Again, the sample size is small.

REFERENCES CITED

- Baugh, T.G., and F.W. Nelson, Jr.
1987 New Mexico Obsidian Sources and Exchange on the Southern Plains. *Journal of Field Archaeology* 14:313- 329.
- Church, T.
2000 Distribution and Sources of Obsidian in the Rio Grande Gravels of New Mexico. *Geoarchaeology* 15:649-678.
- Davis, K.D., T.L. Jackson, M.S. Shackley, T. Teague, and J.H. Hampel
2011 Factors Affecting the Energy-Dispersive X-Ray Fluorescence (EDXRF) Analysis of Archaeological Obsidian. In *X-Ray Fluorescence Spectrometry (XRF) in Geoarchaeology*, edited by M.S. Shackley, pp. 45-64. Springer, New York.
- Gardner, J. N., F. Goff, S. Garcia, R. Hagan
1986 Stratigraphic Relations and Lithologic Variations in the Jemez Volcanic Field, New Mexico. *Journal of Geophysical Research* 91B2:1763-1778.
- Govindaraju, K.
1994 1994 Compilation of Working Values and Sample Description for 383 Geostandards. *Geostandards Newsletter* 18 (special issue).
- Hampel, Joachim H.
1984 Technical Considerations in X-ray Fluorescence Analysis of Obsidian. In *Obsidian Studies in the Great Basin*, edited by R.E. Hughes, pp. 21-25. Contributions of the University of California Archaeological Research Facility 45. Berkeley.
- Hildreth, W.
1981 Gradients in Silicic Magma Chambers: Implications for Lithospheric Magmatism. *Journal of Geophysical Research* 86:10153-10192.
- Hughes, Richard E., and Robert L. Smith
1993 Archaeology, Geology, and Geochemistry in Obsidian Provenance Studies. In *Scale on Archaeological and Geoscientific Perspectives*, edited by J.K. Stein and A.R. Linse, pp. 79-91. Geological Society of America Special Paper 283.
- Kempton, K., G.R. Osburn, S. Kelley, M. Rampey, C. Ferguson, and J. Gardner

2004 Geology of the Bear Springs Peak 7.5' Quadrangle, Sandoval County, New Mexico. Open File Geologic Map OF-GM 74 (Draft), New Mexico Bureau of Mineral Resources, Socorro.

Mahood, Gail A., and James A. Stimac

1990 Trace-Element Partitioning in Pantellerites and Trachytes. *Geochemica et Cosmochimica Acta* 54:2257-2276.

McCarthy, J.J., and F.H. Schamber

1981 Least-Squares Fit with Digital Filter: A Status Report. In *Energy Dispersive X-ray Spectrometry*, edited by K.F.J. Heinrich, D.E. Newbury, R.L. Myklebust, and C.E. Fiori, pp. 273-296. National Bureau of Standards Special Publication 604, Washington, D.C.

Schamber, F.H.

1977 A Modification of the Linear Least-Squares Fitting Method which Provides Continuum Suppression. In *X-ray Fluorescence Analysis of Environmental Samples*, edited by T.G. Dzubay, pp. 241-257. Ann Arbor Science Publishers.

Shackley, M. Steven

1988 Sources of Archaeological Obsidian in the Southwest: An Archaeological, Petrological, and Geochemical Study. *American Antiquity* 53(4):752-772.

1990 *Early Hunter-Gatherer Procurement Ranges in the Southwest: Evidence from Obsidian Geochemistry and Lithic Technology*. Ph.D. dissertation, Arizona State University, Tempe.

1995 Sources of Archaeological Obsidian in the Greater American Southwest: An Update and Quantitative Analysis. *American Antiquity* 60(3):531-551.

2005 *Obsidian: Geology and Archaeology in the North American Southwest*. University of Arizona Press, Tucson.

2009a Source Provenance of Obsidian Artifacts from Piedras Marcadas (LA 290), Middle Rio Grande Valley, New Mexico. Report prepared for Matt Schmader, Albuquerque Open Space, City of Albuquerque, New Mexico.

2009b Two Newly Discovered Sources of Archaeological Obsidian in the Southwest. *Kiva* 76:269-280.

2011 An Introduction to X-Ray Fluorescence (XRF) Analysis in Archaeology. In *X-Ray Fluorescence Spectrometry (XRF) in Geoarchaeology*, edited by M.S. Shackley, pp. 7-44. Springer, New York.

2012 The Secondary Distribution of Archaeological Obsidian in Rio Grande Quaternary Sediments, Jemez Mountains to San Antonito, New Mexico: Inferences for Prehistoric Procurement and the Age of Sediments. Poster presentation at the Society for American Archaeology, Annual Meeting, Memphis, Tennessee.

Wolfman, Daniel

1994 Jemez Mountains Chronology Study. Report prepared by the Office of Archaeological Studies, Museum of New Mexico for the USDA Forest Service, Contract No. 53-8379-9-14.

Table 1. Elemental concentrations for the archaeological specimens and the USGS RGM-1 standard by test unit. All measurements in parts per million (ppm).

Sample	Test Pit	Mn	Fe	Zn	Rb	Sr	Y	Zr	Nb	Ba	Pb	Th	Source
70-80-1	1	52	1152	18	22	8	65	17	98	10	35	27	Cerro Toledo Rhy.
		5	5	3	2			0					
70-80-2	1	53	1119	20	20	11	54	14	81	26	36	30	Cerro Toledo Rhy.
		1	4	3	2			7					
70-80-3	1	42	9841	15	19	10	62	17	95	0	32	20	Cerro Toledo Rhy.
		2		7	4			1					
70-80-4	1	52	1097	12	20	11	65	17	98	0	35	24	Cerro Toledo Rhy.
		4	4	0	7			3					
70-80-5	1	44	8717	60	12	43	26	10	49	415	25	26	Canovas Cnyn Rhy.
		0			2			6					
70-80-6	1	46	1019	12	19	9	59	15	89	0	32	20	Cerro Toledo Rhy.
		1	6	3	5			9					
70-80-7	1	51	1075	14	21	8	65	17	98	0	36	24	Cerro Toledo Rhy.
		2	7	0	2			8					
80-90-1	1	41	9836	88	18	10	62	17	93	0	32	23	Cerro Toledo Rhy.
		7			7			1					
80-90-2	1	50	1061	11	20	10	64	16	90	0	36	24	Cerro Toledo Rhy.
		4	8	0	8			8					
80-90-3	1	48	1052	13	20	12	64	17	94	0	35	18	Cerro Toledo Rhy.
		8	1	8	3			1					
90-100-1	1	51	1131	16	21	9	63	17	93	0	37	27	Cerro Toledo Rhy.
		7	1	9	5			4					
90-100-2	1	49	1081	14	20	8	68	17	96	0	33	24	Cerro Toledo Rhy.
		2	7	5	5			0					
110-120-1	1	48	1035	10	20	9	64	17	10	20	36	24	Cerro Toledo Rhy.
		4	8	0	4			0	0				
120-130S-1	1	52	1096	15	20	19	60	17	93	6	36	26	Cerro Toledo Rhy.
		3	5	8	4			0					
120-130S-2	1	51	1131	18	21	11	62	17	92	0	37	21	Cerro Toledo Rhy.
		0	5	6	3			0					
130-140-1	1	46	1028	12	19	9	63	16	92	0	33	17	Cerro Toledo Rhy.
		1	5	0	5			7					
140-150-1	1	48	1068	13	20	10	59	16	99	0	33	21	Cerro Toledo Rhy.
		7	5	7	4			8					
140-150-2	1	49	1104	14	21	11	67	17	95	0	36	20	Cerro Toledo Rhy.
		4	4	4	1			3					
140-150-3	1	47	1059	26	19	10	54	15	83	150	33	25	Cerro Toledo Rhy.
		3	6	6	1			3					
220-230-1	1	50	1134	14	21	8	62	17	10	0	41	34	Cerro Toledo Rhy.
		7	1	9	0			1	1				
220-230-2	1	57	1219	28	22	12	60	17	90	23	37	32	Cerro Toledo Rhy.
		3	5	8	8			0					
20-30-1	2	47	1044	13	20	9	62	17	93	38	36	26	Cerro Toledo Rhy.
		3	4	6	2			4					
20-30-2	2	54	1156	17	20	10	61	17	96	0	38	25	Cerro Toledo Rhy.
		7	4	1	7			2					
20-30-3	2	53	1211	17	21	10	64	17	92	0	35	22	Cerro Toledo Rhy.
		9	0	2	9			0					
20-30-4	2	51	1129	19	20	8	59	17	89	0	36	32	Cerro Toledo Rhy.
		5	5	3	9			0					
40-50-1	2	54	1151	22	21	11	61	16	90	0	36	18	Cerro Toledo Rhy.
		5	1	3	4			6					
50-60-1	2	49	1064	19	20	11	61	16	87	6	36	26	Cerro Toledo Rhy.
		5	4	7	3			7					
50-60-2	2	41	1069	13	15	14	43	15	52	72	28	17	Valles Rhyolite
		2	9	2	5			9					

50-60-3	2	21 5	5449	47	0	13	2	10	0	117	-2	3	not obsidian
60-70-1	2	55 8	1178 2	13 9	21 8	8	63	17 6	10 5	7	42	31	Cerro Toledo Rhy.
70-80-1	2	50 3	1098 8	13 4	21 1	10	66	17 0	96	0	36	24	Cerro Toledo Rhy.
70-80-2	2	46 2	1056 0	10 0	20 7	11	66	18 4	10 1	0	34	18	Cerro Toledo Rhy.
70-80-3	2	46 8	5343	20	0	22	1	10	0	728	2	6	not obsidian
70-80-4	2	40 7	1106 6	32 8	15 1	14	38	15 0	47	40	27	18	Valles Rhyolite
80-85-1	2	51 4	1116 1	16 0	21 8	8	62	17 5	96	19	39	30	Cerro Toledo Rhy.
85-90-1	2	48 9	1059 9	12 6	20 1	10	65	17 6	99	0	37	26	Cerro Toledo Rhy.
90-100-1	2	47 4	9089	21 2	12 3	51	26	10 3	54	477	29	32	Canovas Cnyn Rhy.
90-100-2	2	54 0	1137 5	18 5	22 0	8	64	16 9	93	0	40	30	Cerro Toledo Rhy.
90-100-3	2	47 8	1047 8	10 2	20 1	9	61	17 0	97	0	39	24	Cerro Toledo Rhy.
100-110-1	2	55 2	1181 0	13 1	20 8	9	64	17 4	97	0	37	25	Cerro Toledo Rhy.
100-110-2	2	55 3	1163 4	13 6	22 3	10	65	17 6	10 5	0	39	27	Cerro Toledo Rhy.
100-110-3	2	49 5	1093 2	19 8	20 6	8	66	17 6	96	25	35	28	Cerro Toledo Rhy.
120-130-1	2	50 4	1049 7	11 8	18 8	11	57	16 8	90	0	33	24	Cerro Toledo Rhy.
0-30-1	3	50 1	1067 1	17 9	19 6	10	58	15 8	86	0	35	25	Cerro Toledo Rhy.
0-30-2	3	56 8	1189 0	21 3	21 4	11	61	16 4	92	0	42	29	Cerro Toledo Rhy.
30-40-1	3	27 7	5390	16 2	0	13	1	13	0	17	2	3	not obsidian
30-40-2	3	46 0	1014 1	14 4	18 4	9	55	15 2	86	0	28	16	Cerro Toledo Rhy.
40-50-1	3	44 7	1076 0	21 8	18 2	13	52	14 1	84	66	33	25	Cerro Toledo Rhy.
50-60-1	3	50 6	1101 0	17 5	21 0	8	60	17 3	94	0	40	26	Cerro Toledo Rhy.
Sample	Test Pit	Mn	Fe	Zn	Rb	Sr	Y	Zr	Nb	Ba	Pb	Th	Source
60-70	3	50 4	1070 7	17 1	20 6	8	62	17 8	93	32	36	20	Cerro Toledo Rhy.
90-100-1	3	50 3	1075 9	10 4	20 9	8	68	17 4	98	33	35	25	Cerro Toledo Rhy.
90-100-2	3	43 7	8781	61 1	12	45	23	99	50	459	23	24	Canovas Cnyn Rhy.
90-100-3-1	3	54 3	1136 7	15 4	21 4	8	62	17 4	97	0	38	19	Cerro Toledo Rhy.
100-110-1-1	3	47 1	1023 1	11 7	19 6	10	64	17 2	90	0	35	28	Cerro Toledo Rhy.
110-120-1	3	54 5	1021 5	11 1	11 3	93	27	12 8	34	153	24	12	Paliza Canyon
120-130-1-1	3	45 0	7916	13 3	17 3	11	27	75	50	0	26	13	El Rechuelos
120-130-2	3	52 6	1114 5	15 9	21 7	10	60	17 2	10 2	0	35	20	Cerro Toledo Rhy.
130-140-1	3	54 8	1132 7	14 5	20 9	10	66	17 2	94	0	35	22	Cerro Toledo Rhy.
130-140-2	3	38	1018	94	15	12	42	15	53	40	25	22	Valles Rhyolite

		7	8	9				3					
140-150-1	3	49	1074	29	19	9	50	14	81	0	36	29	Cerro Toledo Rhy.
		4	9	4	1			7					
140-150-2	3	48	1055	12	20	12	64	16	99	0	35	25	Cerro Toledo Rhy.
		7	1	5	8			7					
140-150-3	3	41	9489	21	11	64	21	10	41	858	24	20	Canovas Cnyn Rhy.
		1		9	5			7					
140-150-4	3	50	1099	17	20	11	63	16	94	0	35	24	Cerro Toledo Rhy.
		7	6	3	5			4					
150-160-1	3	49	1086	35	19	10	56	16	85	0	34	29	Cerro Toledo Rhy.
		8	3	2	5			0					
RGM1-S4		27	1332	35	14	10	24	21	7	877	20	15	standard
		9	3	5	8			6					
RGM1-S4		28	1322	38	14	10	24	21	8	861	18	13	standard
		7	3	6	5			5					
RGM1-S4		27	1330	35	15	10	24	21	11	872	22	16	standard
		2	1	0	7			3					
RGM1-S4		29	1332	35	14	10	23	21	5	872	17	18	standard
		6	3	8	8			6					

Table 2. Crosstabulation of source by test unit. Non obsidian removed.

		Source					Total	
		Cerro Toledo	Valles	Canovas	El	Paliza		
		Rhy.	Rhyolite	Cnyn Rhy.	Rechuelos	Canyon		
Test Pit	1	Count	20	0	1	0	0	21
		% within Test	95.2%	0.0%	4.8%	0.0%	0.0%	100.0%
		Pit						
		% within Source	38.5%	0.0%	25.0%	0.0%	0.0%	34.4%
		% of Total	32.8%	0.0%	1.6%	0.0%	0.0%	34.4%
		Count	17	2	1	0	0	20
		% within Test	85.0%	10.0%	5.0%	0.0%	0.0%	100.0%
		Pit						
		% within Source	32.7%	66.7%	25.0%	0.0%	0.0%	32.8%
		% of Total	27.9%	3.3%	1.6%	0.0%	0.0%	32.8%
		Count	15	1	2	1	1	20
		% within Test	75.0%	5.0%	10.0%	5.0%	5.0%	100.0%
	Pit							
	% within Source	28.8%	33.3%	50.0%	100.0%	100.0%	32.8%	
	% of Total	24.6%	1.6%	3.3%	1.6%	1.6%	32.8%	
Total		Count	52	3	4	1	1	61
		% within Test	85.2%	4.9%	6.6%	1.6%	1.6%	100.0%
		Pit						
		% within Source	100.0%	100.0%	100.0%	100.0%	100.0%	100.0%
		% of Total	85.2%	4.9%	6.6%	1.6%	1.6%	100.0%

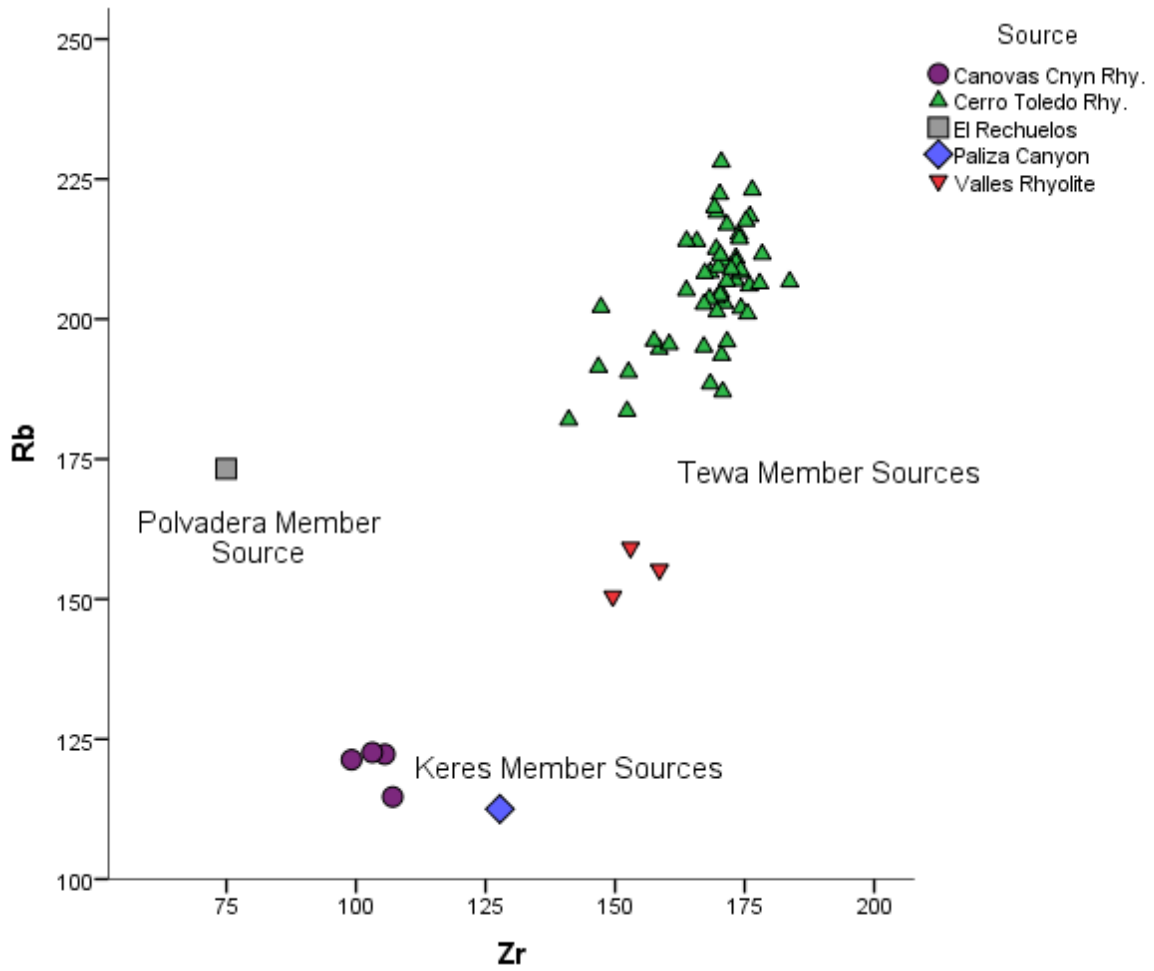


Figure 2. Zr versus Rb bivariate plot of the archaeological specimens.

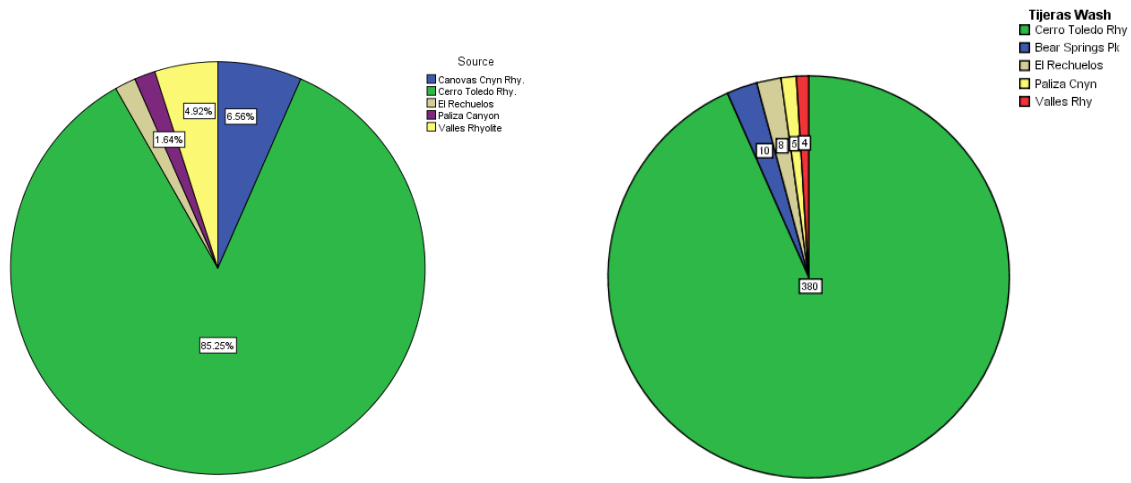


Figure 3. Frequency distribution of source provenance in the test units (%) versus the distribution recovered in Quaternary alluvium at Tijeras Wash, Albuquerque (count). Colors dissimilar.

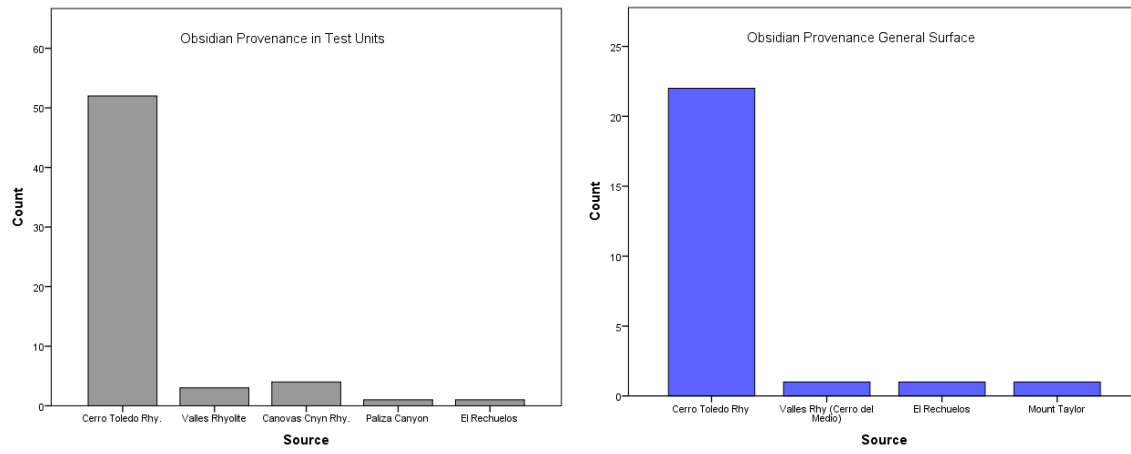


Figure 4. Frequency histograms of source provenance in the test units (left) and general surface (right)